East Antarctic Ice Sheet Variability In The Central Transantarctic Mountains Since The Mid Miocene

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Abstract. The response of the East Antarctic Ice Sheet to warmer-than-present climate conditions has direct implications for projections of future sea level, ocean circulation, and global radiative forcing. Nonetheless, it remains uncertain whether the ice sheet is likely to undergo net loss due to amplified melting coupled with dynamic instabilities, or whether such losses will be balanced, or even offset, by enhanced accumulation under a higher precipitation regime. The glacial-depositional record from the central Transantarctic Mountains (TAM) provides a robust geologic means to reconstruct the past behaviour of the East Antarctic Ice Sheet, including during periods thought to have been warmer than today, such as the Mid Pliocene Warm Period (~3.3–3.0 Ma). This study describes a new surface-exposure-dated moraine record from Otway Massif in the central TAM spanning the last ~9 Myr, and synthesises these data in the context of previously published moraine chronologies constrained with cosmogenic nuclides. The resulting record, although fragmentary, represents the majority of direct and unambiguous terrestrial evidence for the existence and size of the East Antarctic Ice Sheet during the last 14 Myr, and thus provides new insight into the long-term relationship between the ice sheet and global climate. At face value, the existing TAM moraine record does not exhibit a clear signature of the Mid Pliocene Warm Period, thus precluding a definitive verdict on the East Antarctic Ice Sheet's response to this event. In contrast, an apparent hiatus in moraine deposition both at Otway Massif and neighbouring Roberts Massif suggests that the ice sheet surface in the central TAM potentially was lower than present during the Late Miocene and earliest Pliocene.

1   Introduction

Relict glacial deposits preserved in the Transantarctic Mountains (TAM) constitute direct geologic evidence for the East Antarctic Ice Sheet (EAIS) having been present at certain times in the past, and thus provide a terrestrial means for reconstructing former ice sheet configuration. In contrast, marine-based proxy records from the seas surrounding Antarctica, while more continuous and often better dated, are intrinsically indirect; inferences of former ice sheet extent and thickness based on such proxies are therefore speculative. To assess the long-term behaviour of the EAIS directly and from a terrestrial

vantage, we describe the ages of relict glacial deposits in the central Transantarctic Mountains (TAM), constrained by cosmogenic nuclide surface-exposure dating. As the world's largest remaining ice sheet, the configuration of the EAIS and its sensitivity to radiative forcing have direct implications for sea level and ocean circulation, and thus global climate, under amplified greenhouse conditions. In recent decades, observational data have revealed a considerable degree of ice marginal variability (e.g., changes in ice-flow velocity, grounding line position, and ice shelf thickness) in East Antarctica, potentially reflecting the influence of rising ocean temperature on the stability of submarine grounding lines (Li et al., 2015; Rintoul et al., 2016; Konrad et al., 2018; Brancato et al., 2020; Noble et al., 2020). Inland, however, evidence for an EAIS response to climate forcing is less conclusive, in part due to the continental scale of the ice sheet and the relatively short observational periods involved (Stokes et al., 2022). Consequently, it remains uncertain whether the EAIS is likely to undergo net loss due to amplified melting at the margins coupled with dynamic instabilities (Golledge et al., 2015; DeConto and Pollard, 2016), or whether such losses will be balanced, or even offset, by enhanced accumulation under higher precipitation conditions (e.g., Davis et al., 2005; Gregory and Huybrechts, 2006; Rapley, 2006; Boening et al., 2012; King et al., 2012; Winkelmann et al., 2012; Zwally et al., 2015; The IMBIE Team, 2018).

To explore this question further, a growing body of palaeoclimate research has sought to establish how the EAIS responded to past periods of relative warmth, including previous interglacials (DeConto and Pollard, 2016), the Mid Pliocene Warm Period (MPWP: ~3.3–3.0 Ma), and the late Miocene. Of these, the MPWP has received particular attention as it affords, arguably, a closer analogue for our future than any other period since (e.g., Zubakov & Borzenkova, 1988; Dowsett & Cronin, 1990; Bart, 2001; Huybrechts, 2009; Yamane et al., 2015; Blasco et al., 2023, Burke et al, 2018): average MPWP temperatures were elevated relative to today (e.g., Budyko, 1982; Dowsett et al., 2009; Lunt et al., 2010; Haywood et al., 2013) and sea level generally was higher (Raymo et al., 2006), although reconstructions of the latter remain highly variable (+15–60 m) (Miller et al., 2012; Levermann et al., 2013; Rovere et al., 2014; Dumitru et al., 2019; Hearty et al., 2020; Richards et al., 2022). The contribution of the EAIS to past ocean volume, and ultimately ocean circulation, is thus an outstanding question in climate science.

To date, the majority of proxy-based assessments have inferred long-term EAIS variability from marine-geologic data. Yet, these offshore records display a considerable, and sometimes conflicting, range of potential ice sheet configuration. For instance, whereas sedimentologic data from McMurdo Sound indicate persistent cold-based conditions along the ice-sheet margins throughout the Pliocene (Barrett & Hambrey, 1992; McKay et al., 2009), despite generally warmer sea-surface temperatures (Scherer et al., 2007; Naish et al., 2009), ice-rafted debris (IRD) provenance studies posit a dynamically unstable EAIS and widespread collapse of marine-based sectors, such as the Aurora and Wilkes Basins (Fig. 1), during warmer periods (Williams et al., 2010; Pierce et al., 2011; Cook et al., 2013; Tauxe et al., 2015; Valletta et al., 2018). Meanwhile, interpretations of sea-floor morphology on the Antarctic continental shelf have invoked *expansion* of the EAIS during the Pliocene due to enhanced accumulation (Bart & Anderson, 2000; Bart, 2001; Cowan, 2002), a scenario supported

by some model simulations (Hill et al., 2007) and evidence for accumulation-induced thickening of EAIS outlet glaciers during Pleistocene interglacials (Higgins et al., 2000; Todd et al., 2010).
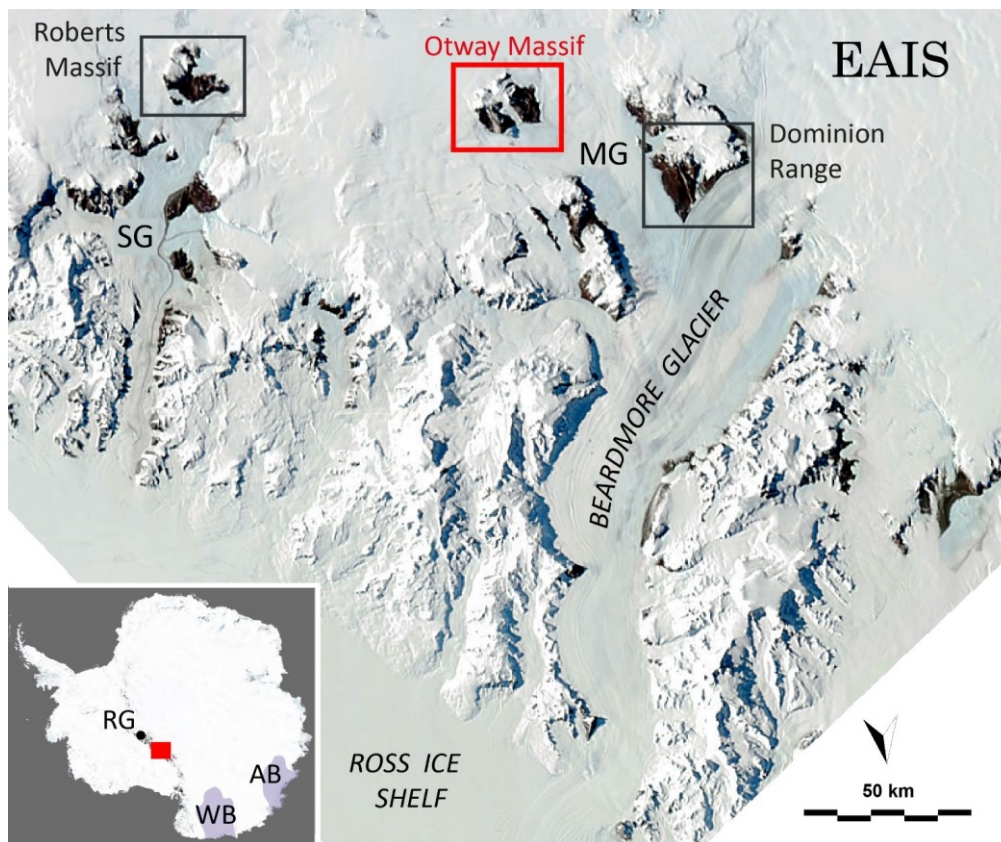
Figure 1. Location map showing position of Otway Massif and other Transantarctic Mountain sites mentioned in the text: SG - Shackleton Glacier; MG - Mill Glacier; RG (inset map) - Reedy Glacier. Also shown are the respective positions of the Aurora (AB) and Wilkes (WB) subglacial basins (inset map).

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75   The glacial-geologic record from Antarctica itself, although fragmentary due to the limited exposure of ice-free land, provides direct terrestrial evidence for past EAIS behaviour. Pioneering work by J. Mercer on the glacial geology of the TAM identified two principal types of deposits, the first comprising lodgement tills emplaced by wet-based glaciation, the second characterised by loose, generally thin ablation tills (Mercer, 1968, 1972). The former – collectively termed the Sirius Group – have since been observed throughout the upper TAM (Barrett & Powell, 1982; McKelvey et al., 1991; Webb et al.,

80   1987; Wilson et al., 1998; Balter-Kennedy et al., 2020), commonly overlying moulded, striated bedrock, and have been interpreted as representing overriding of the mountains by an enlarged EAIS (Mayewski, 1975; Denton et al., 1984, 1989; Webb et al., 1984; Marchant et al., 1993). While direct dating control for the Sirius Group is lacking, and recognising that these wet-based deposits likely represent multiple glacial episodes~~are unlikely to have been emplaced simultaneously~~, a degree of minimum-limiting age constraint is afforded by overlying deposits at several sites (Marchant et al., 1996; Bruno et

85   al., 1997; Schäfer et al., 1999; Ackert and~~&~~ Kurz, 2004), with the most recent data indicating emplacement at Roberts Massif prior to the late Miocene (Balter-Kennedy et al., 2020).

      In contrast to the Sirius Group, the physical characteristics of the overlying deposits are more suggestive of cold-based ice (Mercer, 1972; Bockheim et al., 1986; Prentice et al., 1986), suggesting that the EAIS is frozen to the bed throughout much

90   of the high TAM. Moreover, the spatial distribution of relict cold-based deposits within the current landscape suggests an ice configuration similar to today: a large EAIS draining via TAM outlet glaciers into the Ross Sea Embayment (Bromley et al., 2010). These geologic observations, coupled with a bracketing $^{40}Ar/^{39}Ar$ age control from interbedded volcanic deposits in the McMurdo Dry Valleys, were used to argue for enduring cold polar conditions in East Antarctica potentially since the mid Miocene and, correspondingly, general stability of the EAIS since that time (Denton et al., 1993; Sugden et al., 1995;

95   Marchant and Denton, 1996). In subsequent decades, the application of *in situ* cosmogenic nuclides to glacial deposits and buried ice in the central TAM has corroborated that model at least regionally, revealing that this sector of the EAIS has maintained a configuration broadly similar to today for much of the Pleistocene–Pliocene (Bibby et al., 2016; Bader et al., 2017; Kaplan et al., 2017; Bergelin et al., 2022) and latter Miocene (Balter-Kennedy et al., 2020). Environmental conditions vary considerably in East Antarctica, however, and it is possible the EAIS has undergone greater variability in extent and/or

100 volume in sectors far from the TAM. For instance, sedimentologic data from Lambert Glacier – the catchment of which comprises 7% of the EAIS – have been interpreted as indicating fjord conditions there until the ~~mid Pleistocene~~middle Pleistocene (Hambrey and McKelvey, 2000). Those authors posited that the EAIS was less extensive at that time, a scenario potentially supported by marine evidence from neighbouring Prydz Bay (Mahood and Barron, 1996; Quilty et al., 2000; Whitehead et al., 2005).

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      Taken together, the existing Antarctic dataset presents a variable view of long-term ice sheet behaviour, particularly during key periods of relative warmth. Most, but not all, direct terrestrial evidence suggests the EAIS maintained a configuration similar to present during some periods of apparently elevated atmospheric CO_2 and climatic warmth. In contrast, the indirect

marine record leaves open the possibility of larger-scale variability, including widespread retreat of marine-based sectors. To help address this knowledge gap, we synthesise the existing glacial-geologic dataset in the central and southern TAM spanning the last 15 Myr, focusing specifically on directly dated records constrained with cosmogenic nuclide surface-exposure dating. This synthesis comprises previously published data from Dominion Range (Ackert and Kurz, 2004) and Reedy Glacier (Bromley et al., 2010), the extensive Roberts Massif record (Balter-Kennedy et al., 2020), and our new dataset from Otway Massif in the central TAM (Fig. 1).

2 Otway Massif: Geographical setting and methodology

Otway Massif (-85.415, 172.780; Fig. 1) overlooks the heads of Mill and Mill Stream Glaciers, two tributaries of the Beardmore Glacier, and comprises ~100 km^2 of mountainous polar desert terrain underlain primarily by Ferrar dolerite. The massif exhibits a high-relief relict alpine topography of horn peaks separated by aretes and glacially carved cirques and valleys, and ranges in elevation from 2000 m at the modern ice surface to 3260 m. Previous work by Denton et al. (1989) mapped glacial deposits in the north-west sector of Otway Massif. Our investigation focuses on the range's north-eastern sector (Fig. 1), where the land surface adjacent Mill Stream Glacier forms a broad, shallow valley running parallel to the ice margin. Above this feature, low-angled slopes rise gradually for ~2–3 km in a south-easterly direction towards the base of the Mt. Petlock–Mom Peak ridgeline. The undulating surface of these lower-elevation slopes (~8 km^2) is characterised by glacially scoured, moulded, and deeply oxidised bedrock knolls separated by shallow valleys; overlying this landscape are thin, poorly developed, and potentially sterile soils similar to those described by Dragone et al. (2021) for the high central TAM.

Assuming past ice surface profiles were broadly similar to today, we posit that the EAIS margin here has remained relatively thin for the duration of our record due to the low-gradient landscape over which it fluctuated. Together with the prevailing cold-based regime, this minimally erosive glacial configuration has no doubt contributed to the widespread deposition and long-term preservation of the ablation tills and moraine ridges forming the basis for our cosmogenic surface-exposure chronology.

During the 2015-16 austral summer, we mapped the distribution, elevation, and morphology of moraines and drift sheets onto 2-m resolution vertical satellite imagery provided by the Polar Geospatial Center, University of Minnesota. We collected samples for cosmogenic surface-exposure dating from the upper surfaces of stable boulders located on moraine crests and drift sheets, using a hammer and chisel and/or drill and wedges to remove samples of 1–5 cm thickness. To minimise the impacts of post-depositional erosion or snow burial, we specifically targeted boulders with ≥ 0.5 m relief and lacking visible signs of anomalously extreme surface weathering. We avoided boulders exhibiting obvious signs of subglacial modification (e.g., moulding, polish, and striations) due to the risk of such clasts having originated from remobilised Sirius Group tills, outcrops of which occur locally at Otway Massif, and thus containing inherited cosmogenic

nuclides (Balter-Kennedy et al., 2020). For each sample, we photographed, drew, and described each boulder and measured horizon shielding with a clinometer and compass; shielding values were calculated using the method of Balco et al. (2008). Boulder positions were recorded with a handheld GPS (~6 m horizontal precision) and elevations measured with a barometer calibrated against temporary GNSS benchmarks and corrected to orthometric heights relative to the EGM96 geoid (full procedure described by Balter-Kennedy et al., 2020).

We measured cosmogenic helium-3 concentrations in pyroxene isolated from the dolerite using the protocol described by Balter-Kennedy (2020; modified from Bromley et al., 2014). Following crushing and sieving, the 125–250-μm size fraction was boiled in 10% HNO_3 to remove metal oxides, after which we used heavy liquids (sodium polytungstate) to separate the heavier pyroxene from lighter components. Pyroxene aliquots were then leached in weak (5%) hydrofluoric acid for 2 hours to remove (1) adhering remnants of feldspar and glass and (2) the outer rims of pyroxene grains, which can contain elevated concentrations of radiogenic helium (Blard and Farley, 2008; Bromley et al., 2014). Finally, leached pyroxenes were passed through a Frantz magnetic separator before being hand-picked under a microscope to remove any remaining impurities. We measured helium concentrations in clean pyroxene separates at the Berkeley Geochronology Center using the 'Ohio' MAP 215-50 sector field mass spectrometer; we also periodically analysed aliquots of the CRONUS-P pyroxene standard (Blard et al., 2015) alongside our samples during each measurement period. Full details of the helium measurement procedure, including analysis of procedural blanks, replicates, and lab standards, are provided by Balter-Kennedy et al. (2020). Between 2016 and 2019, we measured helium in samples from 52 boulders, 28 of which were analysed in replicate. As measured concentrations of non-cosmogenic helium in Ferrar dolerite (<$7x10^6$ atoms g^{-1}: Ackert, 2000; Margerison et al., 2005; Kaplan et al., 2017) pyroxenes are typically <1% of the average observed concentrations in samples from Otway Massif, we assume all measured ^{3}He in our samples is cosmogenic and thus make no correction for non-cosmogenic contributions.

Alongside the ^{3}He analyses, we also also measured cosmogenic ^{10}Be and ^{21}Ne in three sandstone clasts. We purified quartz via established physical and chemical procedures (Schaefer et al., 2009) at Lawrence Livermore National Laboratory (LLNL), whereupon beryllium was extracted by anion chromatography and converted to BeO. Ratios of ^{10}Be/^{9}Be in the three samples and one procedural blank were measured at LLNL relative to the 07KNSTD standard (Nishiizumi et al., 2007). Reported uncertainties incorporate uncertainties in AMS measurements, process blanks, and carrier concentrations. We measured cosmogenic ^{21}Ne in the same quartz separates used for ^{10}Be at BGC using the procedure described by Balter-Kennedy et al. (2020).

We calculated surface-exposure ages using Version 3 of the University of Washington online calculator (https://hess.ess.washington.edu) as described by Balco et al. (2008) and subsequently updated, with the 'primary' production rate callibrationcalibration data sets for ^{10}Be and ^{3}He from Borchers et al. (2016). The production rate for ^{21}Ne is calculated assuming a ^{21}Ne/^{10}Be production ratio of 4.03, after Balco et al. (2019), Balco and Shuster (2009), and Kober et al. (2011).

175   Production rates were scaled to each sample site using the time-variable 'LSDn' model (Lifton et al., 2014) and the Antarctic atmosphere model of Stone (2000). A full description of our protocol for identifying likely outliers is given in Section 3.2.

~~3~~ Results

180   ### 3.1 Otway Massif surficial geomorphology

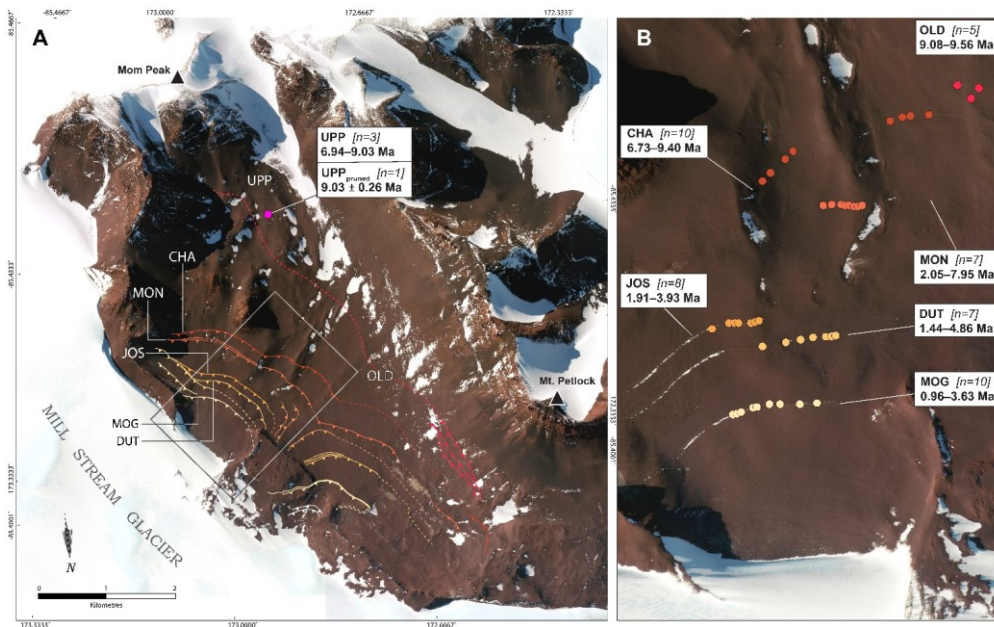
The lower reaches of Otway Massif are mantled by a thin and patchy cover of loose, coarse-grained till characterised by angular and sub-angular clasts, predominantly of dolerite lithology. Evidence for subglacial modification of the clasts (e.g., moulding, striations) is absent, suggesting deposition as ablation till by cold-based ice. Punctuating this surface, numerous moraine ridges, some as much as 1.5 m in relief, are composed of boulders, cobbles, and gravel and mark different marginal
185   positions of the EAIS. In all, we identified six prominent moraines (see Fig. 2 for moraine nomenclature), which are compositionally similar to the adjacent till but differ from one another in the degree of relative weathering, reflecting the time-transgressive nature of deposition. We describe both tills and moraines below in order from lowest to highest, which we interpret as youngest to oldest based on relative weathering criteria and exposure-age data discussed below. Underlying the surficial polar ablation tills, we also observed conspicuous patches of a light-coloured, semi-lithified, clay-supported
190   diamicton (Fig. 3) containing abundant sub-angular and sub-rounded clasts; many clasts bear striations, glacial polish, and

clear evidence of glacial moulding. Reflecting the compositional and morphological contrast with the overlying ablation tills, we interpret this lower unit as lodgement till deposited under more temperate glaciologic conditions and potentially correlated with the Sirius Group (Mercer, 1972).

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Figure 2. Glacial-geomorphic map and cosmogenic nuclide surface-exposure chronology of the Otway Massif study area. (A) Relative positions of moraines and drift edges marking the former extent of the Mill Stream Glacier (EAIS) margin. Dashed lines denote inferred drift edges. White box indicates the area covered in panel B. Also shown are the cosmogenic ^{3}He surface-exposure ages for the Upper (UPP) drift unit, both as single (n=3) and pruned (n=1)
200 **populations; bold type denotes age ranges; italics denote mean ages. (B) Inset map of Mogul (MOG)–Old (OLD) units and their corresponding ^{3}He surface-exposure ages. Note: ^{10}Be and ^{21}Ne ages for the Mogul moraine are not shown. The base map is derived from WorldView-2 satellite imagery (© 2013, DigitalGlobe, Inc.).**

Within a few hundred metres of the modern ice margin, a thin veneer of fresh, unweathered dolerite and sandstone boulders
205   represents the most recent advance of the EAIS along the north-eastern margin of Otway Massif (Fig. 3). While this relatively young unit is not associated with moraine landforms in this sector of Otway Massif, we note that in the north-western part of the range, Denton et al. (1989) reported two distinct units of thick, similarly fresh drift that are bounded by conspicuous moraine ridges. The lower of the two extends ~8 m above the ice margin and was correlated by those authors with the Holocene-age Plunket drift (Denton et al., 1989), whereas the second, higher drift extends up to 40 m above the
210   glacier surface and was correlated with the Late Pleistocene (marine-isotope stage 2) Beardmore drift. We sampled sandstone boulders from the fresh unit for ^{10}Be dating; those data will be the focus of a separate paper covering the most recent EAIS behaviour in the central TAM.

Beyond the spread of fresh drift, the moraines and drift surfaces mapped in Fig. 2 appear generally more weathered with
215   distance from the ice margin. Dolerite boulders on the Mogul moraine (Figs. 2, 3), which demarks two separate lobes extending as much as a kilometre from the ice edge, exhibit moderate orange-brown staining to a depth of one millimetre, with shallow pitting on windward surfaces and light polish on lee surfaces. Sandstone boulders on this same landform are lightly exfoliated, exhibit minor pitting, and are lightly stained except for where flakes have been lost to weathering. Outboard of this unit, dolerite boulders on the Dutch moraine (Fig. 2) are marginally more weathered – exhibiting thicker (2
220   mm) weathering rinds, surface pits up to 1 cm in diameter, and well-developed desert varnish – while those associated with the Joshua moraine (Fig. 2) show advanced flaking and occasional cavernous weathering (Fig. 3); sandstone boulders on the latter moraine are highly disintegrated remnants. Farther from the modern ice margin, dolerite boulders associated with the Montana and Charlie moraines and the Old drift surface (Fig. 2) are highly weathered and many have developed a thick iron-oxide crusts (Fig. 3), which, although being resistant to further surface weathering (e.g., pitting), are prone to break off in

225 thick plates. By far the most aggressively weathered dolerite boulders we sampled, however, are those perched upon the Upper drift surface, one of which (15-OTW-058) exhibits fist-size caverns, while the remaining two (15-OTW-059 and 060) have undergone structural collapse (Fig. 3). For the Montana moraine and older deposits, centimetre-thick accumulations of salt/calcium carbonate occur on the undersides of most dolerite boulders. Moreover, we note that sandstone boulders are absent from these older deposits. As sandstone weathers faster than dolerite, its absence from these surfaces probably reflects

230 the destructive impact of protracted exposure to weathering, and as such is -a common characteristic of very old surfaces.

Figure 3. Examples of glacial geomorphology and relative weathering at Otway Massif. (A) Unweathered sandstone boulder perched on weathered dolerite-rich till, representing the most recent expansion of the EAIS at Otway Massif. (B) Southern section of the Mogul moraine. (C) Sandstone and dolerite boulders on the crest of the Mogul moraine,

illustrating the moderate degree of surface staining on both lithologies. (D) Weathered boulders making up the
260   **Charlie moraine, the relief of which has been progressively lowered by inflation of the surrounding desert pavement (see Section 3.1). (E) Sub-surface salt accumulations revealed in a shallow excavation in the Charlie drift unit. (F) Severely weathered dolerite boulder on the surface of the Upper drift.**

In arid polar environments such as the central TAM, the development of desert pavement provides an additional measure of
265   relative age. Shallow (~40 cm) excavations in the Montana moraine surface revealed a 2–3 cm thick surficial layer of loose medium- and coarse-grained sand, below which is a compacted unit of oxidised coarse sand, gravel, and cobbles, the latter exhibiting moderate desert varnish. Thick salt layers are common throughout this compacted unit. Within the higher Old drift, we observed that the appearance of cobble-sized surface clasts exhibits a grain-size dependence; coarse-grained dolerite clasts are visibly exfoliated and crumbly, whereas fine-grained dolerite clasts tend to be highly desert-varnished.
270   Immediately beneath these surface clasts, a unit of unconsolidated medium- and coarse-grained sand extends to a depth of 5 cm and overlies a compacted horizon of deeply oxidised coarse sand and gravel and interbedded salt layers. The highest unit identified at Otway massif, the Upper drift, displays similar sub-surface sedimentologic characteristics to the Old drift.

3.2   Correlative deposits in the Central and Southern TAM

275   Overall, the stratigraphy and morphology of glacial deposits at Otway Massif are nearly identical to those described by Balter-Kennedy et al. (2020) for Roberts Massif, which is located ~100 km east of Otway Massif (Fig. 1) and underlain by the same dolerite-sandstone lithology. The last 15 Myr of ice sheet deposition at Roberts Massif is represented by at least 24 individual moraine/drift units, the oldest being positioned farthest from – and generally highest above – the modern ice margin (Balter-Kennedy et al., 2020). As summarised in Fig. 6, exposure ages for the highest Roberts Massif deposits range
280   between ~5 Ma and 13 Ma (Misery Platform); stratigraphically and geomorphically, these deposits are analogous to the uppermost units observed at Otway Massif (Charlie, Old, Upper deposits). Indeed, the only significant disparity between the two sites is the relative abundance at Roberts Massif of highly weathered sandstone clasts, even within the oldest deposits on SW Col; this feature reflects the extensive outcropping of Beacon sandstone at the head of Shackleton Glacier (Balter-Kennedy et al., 2020). Below the oldest Roberts Massif deposits, mid-elevation (Arena and Ringleader moraines) and low-
285   elevation (e.g., northern and southern transects) landforms give age ranges of ~0.5–3.2 Ma, respectively. Moraine ridges at both Roberts Massif and Otway Massif are nearly identical in relief, composition, relative weathering, and development of sub-surface salt layers.

Closer to Otway Massif, pre-Pleistocene deposits at Dominion Range (Fig. 1) also bear a strong similarity to those at our
290   field site. Multiple studies have reported lateral moraines extending up to 2500 m elevation on the Oliver and Mercer Platforms (Mercer, 1972; Denton et al., 1989; Prentice et al., 1989; Ackert, 2000; Ackert and Kurz, 2004), with Ackert (2000) describing the highest landforms – collectively termed 'Dominion drift' – as cold-based deposits in which boulders

exhibit extreme structural degradation; cavernous weathering is well-developed in dolerite clasts while sandstone boulders have been destroyed. Cosmogenic ^{3}He ages constraining the Oliver Platform moraine (Dominion drift) range from 1.7 Ma to 4.6 Ma and are included in Fig. 6. Between the Dominion drift and modern surface of Beardmore Glacier, similar yet undated deposits are grouped into three broad units that, on the basis of relative weathering, get progressively younger with decreasing elevation (Mercer, 1972; Denton et al., 1989). Farther afield, at the southern end of the TAM range, Bromley et al. (2010) reported at least five distinct cold-based till units alongside Reedy Glacier in the Wisconsin Range (Fig. 1). As at Otway Massif, the distribution of landforms at Quartz Hills, Polygon Spur, and Tillite Spur indicates deposition by north-flowing EAIS ice conforming to the current landscape, while relative weathering characteristics suggest that deposits increase in age with elevation. The highest cold-based tills at Reedy Glacier – 'Reedy E' drift – comprise highly deflated surfaces of exfoliated granite debris, into which are set granitic boulders exhibiting extreme cavernous weathering. Cosmogenic ^{10}Be dates constrain the Reedy E deposits to 2.0–7.0 Ma at Quartz Hills, 1.2–6.6 Ma at Polygon Spur, and 1.7–5.2 Ma at Tillite Spur (see Section 4). Together, the geomorphologic and weathering characteristics of deposits at Reedy Glacier, Dominion Range, and Roberts Massif are sufficiently similar to those at Otway Massif to permit a first-order regional synthesis (Section 4).

3.3   Otway Massif cosmogenic surface-exposure ages

The application of surface-exposure dating to glacial deposits relies on the premise that glacially transported clasts originate from subglacial erosion, and thus contain negligible concentrations of cosmogenic nuclides prior to deposition on a moraine ridge or drift surface. Upon exposure to the cosmic ray flux, nuclide accumulation begins within the lattices of constituent minerals; concentrations of cosmogenic ^{3}He, ^{10}Be, and ^{21}Ne in glacial erratics thus yield exposure ages for the final occupation of the moraine or the deglaciation of a land surface. Including replicates, the new dataset from Otway Massif comprises 80 measurements of cosmogenic ^{3}He in samples from 52 boulders located on the crests of five moraine ridges (Mogul, Dutch, Joshua, Montana, Charlie) and perched on two drift surfaces (Old, Upper).

Individual ages are given in Table S1 and calculated landform ages in Table 1. Complete step-degassing results and nuclide concentrations are given in Tables S2–S4; all boulder information and metadata are archived online in the ICE-D:ANTARCTICA database (www.ice-d.org/antarctica). To resolve likely depositional ages for individual units and landforms, we measured at least three (and typically > 5) samples per feature in order to perform statistics on each population. Specifically, we employed the coupled statistical-stratigraphic protocol detailed by Balter-Kennedy et al. (2020) to identify outlier samples: assuming that a moraine's true age lies within the range of measured exposure ages from that landform, any exposure age on a moraine that is older (younger) than *all* exposure ages on a stratigraphically older (younger) landform is deemed erroneous and rejected (Balter-Kennedy et al., 2020). We also pruned as likely geomorphic outliers any samples returning ages > 2σ beyond the main age population of that moraine. While the Otway dataset as a whole includes few outliers under this elimination method, we note that this primarily reflects the broad and/or non-normal

distributions of many of the moraines studied. Indeed, recognising that some landforms exhibit relatively tight distributions (i.e., Charlie moraine and Old drift), the majority of dated landforms at Otway Massif exhibit sufficient scatter (coefficient of variance > 15%) to warrant using the range as the most conservative and best estimate of deposition age. Since production rate uncertainty is minor by comparison to the range of exposure ages for any given moraine, we do not consider this in our treatment of the chronology. All ages are shown as box plots in Fig. 4.

Table 1. Otway Massif moraine and drift ^{3}He ages and statistics.

Site	Elevation at sample site (m)	Count (samples pruned)	Age range of raw data (Ma)	Mean age (Ma)[1]	**Age range (Ma)[2]**	Coefficient of Variance (%)	Reduced χ^2
Moraines							
MOG	2262	10	0.96–5.36	1.96 ± 0.86	**0.96–3.63**	44	362
DUT	2262	7 (2)	1.44–4.86	2.54 ± 0.92	**1.44–3.66**	36	242
JOS	2278	8	1.91–3.93	2.85 ± 0.68	**1.91–3.93**	24	84
MON	2275	7	2.05–7.95	4.49 ± 2.56	**2.05–7.95**	57	617
CHA	2277	10	6.73–9.40	7.94 ± 0.88	**6.73–9.40**	11	15
Drift							
OLD	2290	5	9.08–9.56	9.33 ± 0.19	**9.08–9.56**	2	1
UPP	2503	1 (2)	6.94–9.03	9.03 ± 0.26	**-**	-	-

Ten dolerite samples from the crest of the Mogul moraine range in age from 0.98 to 3.63 Ma (Table 1; Figs. 2 and 4), while three small sandstone boulders on the same crest (15-OTW-011, 012, and 013) returned apparent beryllium-10 ages ranging from 1.09 to 1.92 Ma (Table 1; Fig. 4). The same three sandstone samples also gave apparent neon-21 ages ranging from 1.65 to 5.36 Ma (Table 1; Fig. 4). As illustrated in the two-nuclide diagram in Fig. 5, these paired nuclide concentrations plot below the line of simple exposure, indicating that the boulders have experienced either significant post-depositional erosion or a complex exposure history. Because none of the boulders shows evidence of substantial weathering (being fresh in appearance and exhibiting only minor surficial alteration (see Section 3.1), and their apparent ^{10}Be and ^{21}Ne ages are not significantly younger than the ^{3}He ages of adjacent dolerite boulders (as would result if the less-resistant sandstone clasts experienced greater erosion than the dolerite), we consider the scenario of multiple periods of exposure as the most plausible cause for these complex exposure histories. Recognising that all three sandstone boulders apparently experienced complex exposure, it would nonetheless be incorrect to assume that all 52 dolerite in our dataset must therefore have undergone similarly complex exposure histories: first, it is not justified statistically to draw such a conclusion based on only three

samples, and second, the source and transport trajectories of sandstone and dolerite boulders at Otway Massif likely differ considerably based on the relative rarity of sandstone outcrops versus dolerite outcrops. Thus, we continue to interpret the

350   ^{3}He data on the basis that the sample set probably includes boulders both with and without inheritance. ~~Acknowledging that all boulders sampled at Otway Massif – both sandstone and dolerite – might have experienced a degree of complex exposure~~ For the Mogul moraine, ~~(see Discussion), the demonstrably non-simple exposure histories of the three Mogul sandstones precludes using the ^{10}Be or ^{21}Ne data for direct age control in this instance. Nonetheless,~~ we note that these apparently inheritance-influenced measurements on sandstone clasts nonetheless provide a minimum-limiting constraint for the Mogul

355   moraine of ~1 Mya, in broad agreement with the helium data.
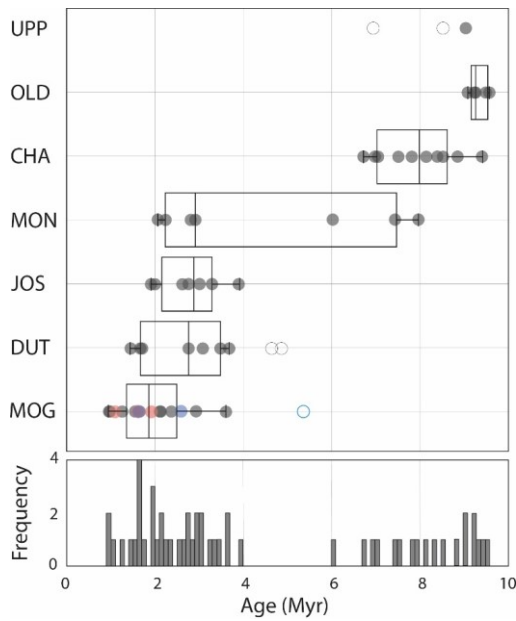
Figure 4. Box plots illustrating the surface-exposure ages for the Plio-Pleistocene moraines and drift units preserved

360   at Otway Massif. Units are arranged stratigraphically, with names corresponding to those given in Fig. 2. Filled grey, pink, and blue circles denote ^{3}He, ^{10}Be, and ^{21}Ne ages, respectively; empty circles denote outliers, with colours

corresponding to nuclide. Mean landform/unit age is shown by the red crosses. The histogram panel depicts all apparent exposure ages from Otway Massif and illustrates the potential hiatus in deposition discussed in Section 4.

365 Seven samples from boulders on the Dutch moraine return a strongly multimodal distribution, ranging in age from 1.44 to 3.66 Ma (Table 1; Figs. 2 and 4). Two additional samples from this landform (15-OTW-018 and 019) returned ages older than all eight samples from the next, stratigraphically older Joshua moraine and thus were excluded as stratigraphic outliers. Farther upslope, eight boulders on the Joshua moraine give a range of 1.91–3.93 Ma, while the Montana moraine dates to 2.05–7.95 Ma (Table 1; Figs. 2 and 4). Ten samples from the Charlie moraine give a relatively tight dataset ranging from

370 6.73 to 9.40 Ma (Table 1; Figs. 2 and 4), while five boulders embedded in the Old drift surface, within 200 m of the Charlie moraine, provide an internally consistent distribution of 9.08–9.56 Ma (Table 1; Figs. 2 and 4). Finally, at the upper end of our transect and close to the northern flank of Mom Peak, three samples from the Upper drift (Fig. 2) give a bimodal age distribution ranging from 6.94 to 9.03 Ma (Table 1; Fig. 4). Being younger than all five ages from the stratigraphically younger Old drift, we excluded samples 15-OTW-059 and 60 as young outliers. Consequently, our best ~~age~~ estimate for the

375 deposition of Upper drift is based on a single age of 9.03 ± 0.26 Ma afforded by sample 15-OTW-58.

The presence of dolerites boulders in the TAM with > 10 Ma exposure ages confirms that erosion rates are very low, thus limiting the possible impact of erosion on the younger part of the Otway chronology. Nonetheless, since erosion rates are greater than zero, we sought to evaluate the impact of erosion on our dataset. If we assume that the oldest sample in our

380 record (15-OTW-055) is no older than 14.5 Ma, which is the oldest material reported from neighbouring Roberts Massif (Balter-Kennedy et al., 2020), the erosion rate at our site cannot be greater than 3.3 cm/Myr. Applying that rate to a sample of Mid-Pliocene age (e.g., 15-OTW-026) changes the apparent exposure age from 3 to 3.3 Ma, which is not significant relative to the age ranges exhibited by each moraine reported here. Therefore, we do correct our age calculations for erosion.
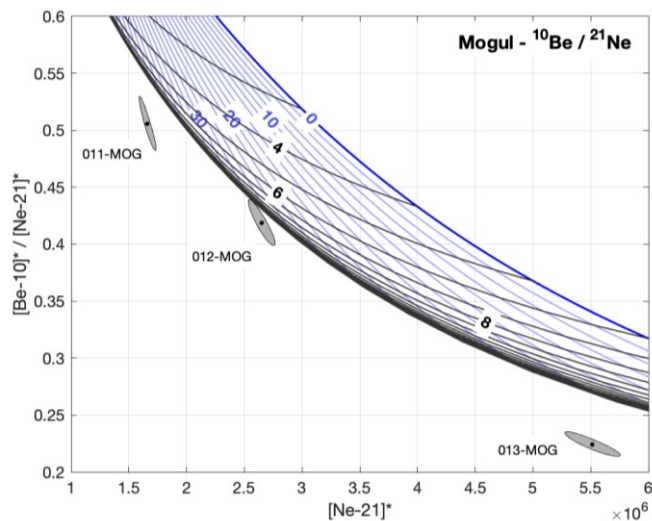
Figure 5. [10]Be–[21]Ne normalised two-nuclide diagram for three sandstone erratics sampled on the Mogul moraine (15-OTW-011, 012, and 013). Black and blue numbered lines are isolines of constant exposure age (Ma) and steady erosion (cm Myr[-1]), respectively.

390

4   Discussion

The glacial-geologic record from Otway Massif spans the latter half of the Miocene (during which the Charlie, Old, and Upper deposits were emplaced) to at least the middle Pleistocene (i.e., Mogul moraine). Together, the morphology and sedimentology of the Otway glacial deposits, like those described from neighbouring sites (Ackert, 2000; Bader et al., 2017; Balter-Kennedy et al., 2020; Bergelin et al., 2022), are typical of cold-based ablation tills reflecting deposition under glaciologic conditions similar to today. Coupled with the dearth of outwash or fluvial landforms, these glacial-geologic characteristics indicate that the central TAM have been dominated by a cold, arid climate for the duration of our record, with no evidence for even brief interruptions during that period. This is significant because it requires that the high-elevation Antarctic interior remained a polar desert environment even during the relative warmth of the late Miocene epoch and the subsequent MPWP, as proposed by previous glacial-geologic research in the central TAM (Barrett, 2013; Bibby et al., 2016; Balter-Kennedy et al., 2020; Spector and Balco, 2021; Bergelin et al., 2022). Our record also provides additional age control for the transition from warm- to cold-based glaciation at Otway Massif. According to our ages from the Old drift, both the shaping of the alpine topography by temperate ice and emplacement of wet-based Sirius Group lodgement tills occurred prior to 9 Ma, a result that is in general agreement with previous minimum estimates from elsewhere in the TAM based upon [40]Ar/[39]Ar ages (Sugden and Denton, 2004, and references therein; Lewis et al., 2007; Smellie et al., 2011) and cosmogenic surface-exposure ages (Bruno et al., 1997; Schäfer et al., 1999; Di Nicola et al., 2012; Balter-Kennedy et al. 2020).

As shown in Figs. 2 and 4, the ages reported here for each drift unit fall broadly in stratigraphic order, becoming increasingly older with distance from the modern EAIS margin. Not only does it support our observations of relative weathering and desert pavement development, this stratigraphic relationship also highlights a persistent first-order pattern of relative ice-surface lowering at Otway Massif, which aligns with similar patterns documented elsewhere in the central TAM (Denton et al., 1989; Hagan 1995; Ackert, 2020; Ackert and Kurz, 2004; Bader et al., 2017; Balter-Kennedy et al., 2020; Bergelin et al., 2022) and southern TAM (Bromley et al., 2010, 2012). Since this process both post-dates the end of significant orographic uplift (Fitzgerald, 1994; Miller et al., 2010; He et al., 2021) and is comparatively minor in scale (tens of metres rather than kilometres), we posit that a likely driver is localised isostatic uplift of mountain blocks by selective linear erosion beneath wet-based portions of EAIS outlet glaciers since the middle Miocene (Stern et al., 2005), as described in detail by Balter-Kennedy et al. (2020) for the upper Shackleton Glacier. Fluctuations in the volume of the WAIS, into which EAIS

outlet glaciers drain during ice ages, may also influence the elevational offset between relict moraines and the modern ice surface at Otway Massif, though the magnitude of that effect diminishes with increasing distance and elevation from the

420 WAIS (Bromley et al., 2010, 2012; Todd et al., 2011). Assuming that the long-term relative drop in ice sheet surface observed in the TAM is unlikely to have been monotonic, it is possible – if not probable – that the Otway moraines have been overridden by the EAIS on multiple occasions, in which case relict moraines might contain material accrued over several glacial cycles. While we cannot test this scenario with the current dataset, such an 'accretionary' process might account for at least some of the scatter exhibited by moraines with broad age ranges. In the same vein, the tight age

425 distribution of the higher-elevation Old drift (Table 1; Figs. 2 and 4) might suggest that the EAIS margin only occupied this position once. The absence of this pattern from neighbouring Shackleton Glacier, however, where the lowest/youngest moraines at Roberts Massif give relatively tight age distributions despite the likelihood of repeated overriding (Balter-Kennedy et al., 2020), leads us to infer that the generally broader age distributions at Otway Massif primarily reflect a greater degree of inheritance and/or exhumation at this site.

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~~Acknowledging~~ Beyond the overall stratigraphic coherence of our dataset, we recognise that the age distributions for ~~some~~ these glacial landforms are too broad and/or complex to pinpoint the timing of deposition beyond a range. The Mogul moraine, for example, includes ten ages spanning ~2.6 Ma, none of which can be rejected as an unequivocal outlier according to our protocol. Assuming that occupation of the Mogul moraine for the entire 2.6 Myr range is unlikely (given the

435 typical duration of Pleistocene glacial maxima), the most plausible explanation for this – and other – broad distribution is the incorporation of boulders with nuclide concentrations ~~of ³He~~ inherited from past periods of exposure, which skews the landform's apparent age towards being anomalously old. In this case, the Mogul moraine's true age would be closer to 1 Ma than to 3.5 Ma, a scenario supported by our use of the three ^{10}Be measurements as maximum-limiting ages (Fig. 4). ~~Yet~~ to simply correlate deposition with the youngest age in a population would ~~this approach can~~ ignore the potential influence of

440 geomorphic factors (e.g., burial by snow, erosion, exhumation or rolling due to moraine slumping, etc.) that might affect boulders following deposition and which, along with the addition of younger material during subsequent overriding episodes (see above), can skew the apparent age towards being anomalously *young*. Therefore, in the absence of obvious physical indicators either for young or old anomalies, we conclude that the accuracy ~~precision~~ of single-nuclide-based age estimates in cold-based glacial regimes remains ~~variable~~ poor and that in most ~~some~~ instances ranges are the only reliable means for

445 presenting ages.

For the Otway record specifically, and the existing TAM dataset generally, this fact presents a challenge to assessments of ice sheet response to climate events such as the MPWP; dated moraines typically yield age ranges exceeding the ~300 kyr duration of the MPWP and, with the exception of the Charlie moraine, all Otway moraines overlap with the event within

450 error (Fig. 4). The Mogul moraine illustrates this matter well: if we take the young end of the range, the moraine postdates the MPWP by 2 million years, whereas the older end of the range falls squarely within the MPWP. Further, at the

neighbouring sites within the central and southern TAM where mean moraine ages fall within the MPWP (e.g., Reedy E and Oliver Platform moraines), their respective uncertainties remain too great to be comparable with climate events of this scale (Fig. 6). Compounding this issue is uncertainty over (1) the applicability of long-term erosion rates (see Section 3.3) and (2) uncertainty in extrapolating production rates inferred from calibration data with 10–100 ka exposure durations to exposure~~time~~ periods in the 1–10 Ma range. While some constraints (e.g., Balter-Kennedy et al., 2020) indicate that the calibrated production rates are probably accurate over the longer term, it remains challenging to verify this with the available data, highlighting the growing need for production rate calibrations on the continent itself.
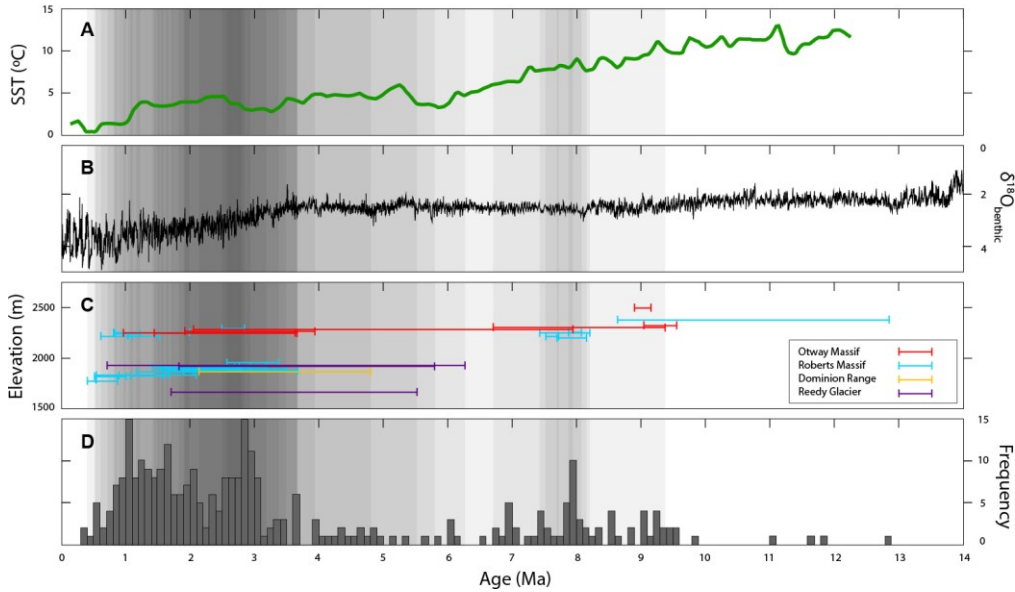
Figure 6. Temporal comparison of long-term global climate conditions and frequency of moraine deposition in the central-southern TAM, adapted from Balter-Kennedy et al. (2020) and including additional data from the present study and other published work (Ackert and Kurz, 2004; Bromley et al., 2010). A - Southern Hemisphere average sea-surface temperature reconstructed from alkenones (Herbert et al., 2016); B - Benthic oxygen isotope curve (De Vleeschouwer et al., 2017); C - Moraine age plotted against elevation for landforms at Otway Massif (red circles), Roberts Massif (turquoise circles), Dominion Range (yellow circle), and Reedy Glacier (purple circles); D -

Histogram depicting all apparent exposure ages (all samples from each of the four TAM locations discussed) shown in panel C. Grey bars denote moraine ages (including uncertainty). Deeper grey shades occur where error bars overlap and therefore represent a higher frequency of landforms with that age.

Bounded stratigraphically by the younger Joshua moraine (1.9–3.9 Ma; Figs. 2 and 4) and older Charlie moraine (~~7.94 ± 0.88~~6.7–9.4 Ma; Figs. 2 and 4), the Montana moraine exhibits both an unusually broad age range (~6 Myr) and a conspicuous 3-Myr 'hiatus' during which there are no ages in our dataset (Fig. 4). From the data on hand, and assuming that the Montana moraine represents a single period of deposition, the moraine's true age could equally fall at the older end of the range – as suggested by the strong similarity to Charlie and Old deposits in relative weathering and desert pavement development (see below) – or at the younger end, bringing the Montana moraine closer in age to the Joshua moraine (Fig. 4). There are no data for the intervening ~3 Myr. Alternatively, recognising that cold-based ice is minimally erosive (passing over landscapes with little or no impact on pre-existing landforms), it is conceivable that the Montana dataset represents not one but multiple periods of deposition at this site, the first during the late Miocene and the second during the Pliocene (see above). By either model, the apparent gap in moraine deposition at Otway Massif is similar to that documented by Balter-Kennedy et al. (2020) at the head of Shackleton Glacier (Figs. 1 and 6), where deposition of the Misery moraines and the next stratigraphically youngest moraine (Ringleader) was separated by ~3.5 Myr. Together, these datasets raise the possibility that the EAIS margin was absent from the upper TAM for a protracted period (Fig. 4), potentially reflecting a smaller ice sheet at that time. If true, resolving the climatic and/or physical drivers of a reduced EAIS during the late-Miocene–early-Pliocene, which is not generally viewed as a period of significant warmth or enhanced radiative forcing (Fig. 6), poses a compelling challenge for Antarctic research.

5   Conclusions

Relict glacial deposits at Otway Massif provide a direct terrestrial record of EAIS behaviour in the central TAM spanning the last ~9 million years. Coupled with our geomorphic observations, cosmogenic ^{3}He, ^{21}Ne, and ^{10}Be surface-exposure ages from five separate moraine limits and two drift units confirm the presence of a cold-based ice sheet in the upper central TAM since at least the Late Miocene, and thus add to the growing terrestrial consensus that the transition from temperate to polar climate conditions in Antarctica occurred prior to the late Miocene, as did emplacement of wet-based Sirius Group tills. Gradual lowering of the ice surface, relative to Otway Massif, over the same period mirrors that observed elsewhere in the TAM and is suggestive of gradual isostatic uplift in response to long-term trough erosion beneath wet-based portions of EAIS outlet glaciers.

According to both the new Otway dataset and our regional synthesis of published moraine ages (Fig. 6), there is no compelling geologic evidence from the central and southern TAM for a significant change in EAIS configuration during the

MPWP. While this might signify that interior portions of the ice sheet are perhaps insensitive to the magnitude of climate warming proposed for the middle Pliocene, we (1) caution that the uncertainties associated with this dating method in Antarctica remain too large to capture such short-lived events as the MPWP and (2) highlight that possible episodes of smaller-than-present ice extent cannot be ruled out from the available terrestrial geologic record.

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Of potentially greater climatic-glaciologic significance, an apparent gap in moraine deposition between ~4–7 Ma mirrors the glacial-geologic record from the head of Shackleton Glacier (Fig. 6) and is not inconsistent with the hypothesis that the ice sheet surface in the central TAM was lower during that period than subsequently. Recognising that this hiatus overlaps broadly with early-mid Pliocene episodes of EAIS ice-marginal variability inferred from marine proxies, clear disparities in
510   the timing and duration of the terrestrial and marine events mean these events cannot be correlated using existing data. Nonetheless, the Otway hiatus, if replicated, has clear implications for our understanding of EAIS stability and climate interactions, and thus affords a potentially valuable new target for terrestrial palaeoclimatology.

Data Availability

515   All data reported here that are not already published elsewhere are available in the supplement.

Supplement

The supplement for this article is available online at: URL

Author Contributions

520   GRMB and GB designed the research. GRMB, GB, MSJ, and HT conducted fieldwork and sample collection. AB-K, GB, and GRMB prepared and analysed samples. GRMB and GB wrote the paper with input from all authors.

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Competing Interests

The contact author has declared that none of the authors has any competing interests.

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