1 2	Magnetic properties and geochemistry of loess/paleosol sequences at Nowdeh section northeastern of Iran
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11 Abstract

The loess-paleosol sequences in the northeastern part of Iran serve as a high-12 resolution natural archive documenting climate and environmental changes. These 13 sequences offer evidence of the interaction between the accumulation and erosion 14 of aeolian and fluvial sediments during the Middle and Late Pleistocene periods. In 15 this study, the Azadshar (Nowdeh Loess Section) site was chosen to reconstruct 16 Late Quaternary climate shifts. The 24-meter thick Nowdeh loess/paleosol 17 sequence was sampled for magnetic and geochemical analysis. The sampling 18 involved 237 samples taken systematically at high resolution (10 cm intervals 19 Selected samples, corresponding to peaks in magnetic susceptibility, underwent 20 geochemical analysis to aid in the interpretation of paleoclimatic changes indicated 21 by the magnetic signals. The magnetic susceptibility results of the loess/paleosol 22 deposits revealed low values during cold and dry climate periods (loess) and high 23 values during warm and humid climate periods (paleosol). The magnetic 24 susceptibility at a depth of 22.1 meters (approximately 130 Ka) has significantly 25 decreased, suggesting cold climate conditions at this time. The most substantial 26 changes in magnetic susceptibility occur at depths between 18.6 to 21.3 meters 27 (approximately 100 to 120 Ka). During this period, there are four phases of 28 decrease (indicating cold and dry conditions) interspersed with three phases of 29 increase (signifying warm and humid conditions) in magnetic susceptibility. The 30 comparison of magnetic and geochemical data showed that variations in 31 geochemical weathering ratios corresponded to changes in magnetic parameters. A 32 high level of correlation was observed between the magnetic susceptibility 33 intensity and ratios such as Rb/Sr, Mn/Ti, Zr/Ti, and Mn/Sr. The findings from this 34 research indicate that the sedimentary section of Nowdeh has experienced six 35 distinct climate periods over the last 160,000 years. Notably, three cold and dry 36

periods occurred between three warm and humid periods. Additionally, during these climate phases, short-term cold (stadial) and warm (interstadial) intervals were also observed.

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41 Keyword: Loess/paleosols sequences, Climate, Magnetic parameters, Geochemical proxies,

42 Northeastern of Iran.

43 Introduction

Reconstruction of the Quaternary climate is important for the development of climate models that lead to a better understanding of past and present and prediction of future climate development. Loess–paleosol sequences are now recognized as one of the most complete terrestrial archives of glacial–interglacial climate change (Porter, 2001; Muhs and Bettis, 2003, Pierce et al, 2011, Guo et al, 2002) and have been used to reconstruct climate and geomorphological changes during the Quaternary (Karimi et al., 2011; Frechen et al., 2003; Prins et al., 2007).

Loess deposits occur in large areas of the northeast, east central, north and central 51 parts of Iran which is part of the loess belt that covers the Middle East and extends 52 further northward into Turkmenistan, Kazakhstan and Tajikistan (Okhravi and 53 Amini, 2001). The extensive and thick loess deposits in northern Iran have been 54 recently studied in detail establishing a more reliable chronological framework for 55 the last interglacial/glacial cycle (Lateef, 1988; Pashaee, 1996; Kehl et al., 2006; 56 Frechen et al., 2009, Karimi et al, 2009, Karimi et al, 2013, Okhravi and Amini, 57 2001, Mehdipour et al, 2012). 58

Paleoclimate studies of loess deposits based on rock magnetism and combined 59 analyses of rock magnetism and geochemistry around the world have attained 60 appreciable advances in the past few decades (Bader et al,2024; Jordanova and 61 Jordanova, 2024; Heller and Liu, 1984; Forster et al., 1996; Ding et al., 2002; Guo 62 et al., 2002; Chlachula, 2011; Bronger, 2003; Baumgart et al., 2013, Guanhua, et al, 63 2014). These studies comprise loess-paleosol records that cover Chinas loess 64 plateaus, Germany, Poland, Tajikestan, Austrian, Ukraine, and the Danube 65 catchment (Hosek et al, 2015, Ahmad and Chandra, 2013, Chen, 2010; Jordanova 66 et al., 2011; Buggle et al., 2009; Fitzsimmons et al., 2012; Fischer et al., 2012; Jary 67 and Ciszek, 2013; Baumgart et al., 2013; Schatz et al., 2014; Gocke et al., 2014). 68

Despite its suitable geographical location there is only a limited number of studies of loess deposits from the North of Iran. In this work we explore the potential of loess deposits in northern Iran for reconstructing late quaternary
 climate/environmental change.

73

74 Study area

The Nowdeh section is exposed at about 20 km southeast of Gonbad-e Kavus and east of Azadshahr city. The Nowdeh river dissects a more than 24 m thick sequence of yellowish brown (10 YR 5/4) loess covering northeast dipping weathered limestone.

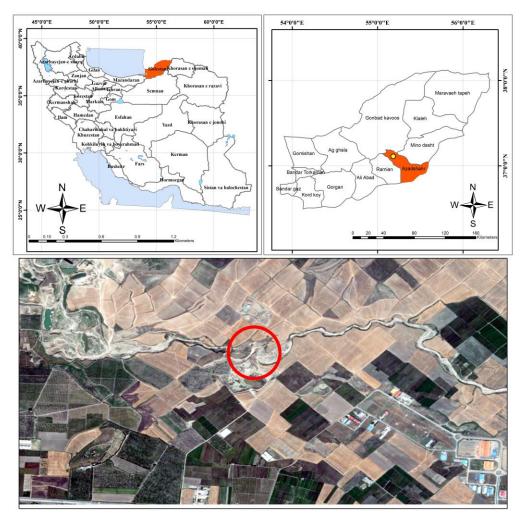
The study area (37° 05' 50" N and 55° 12' 58" E) is part of the Alborz structure and this structure continues beneath the Caspian Sea. This zone includes regions north of the Alborz fault and south of the Caspian Sea. Toward the east, the

⁸² Gorgan-Rasht zone is covered with thick layers of loess.

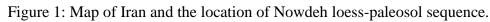
83 The Nowdeh section was selected for this work due to earlier soil studies by Kehl

et al (2005) and Frichen et al (2009) combined with the existence 12 dating for this

ss section (Figure 1).

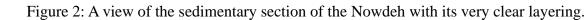












90 Methodology

The Nowdeh loess section is approximately 24 meters thick and was sampled at 10 91 cm intervals for magnetometry and geochemical analysis. The sampling location 92 and method were determined following a detailed study of the area. Magnetic 93 susceptibility measurements of all samples were conducted at the Environmental 94 and Paleomagnetic Laboratory of the Geological Survey of Iran in Tehran. 95 Magnetic susceptibility is indicative of the collective response of diamagnetic, 96 paramagnetic, ferrimagnetic, and imperfect antiferromagnetic minerals present in 97 the samples. Each sample was placed in a 11 cm³ plastic cylinder for use in 98 magnetic measurement devices. The measurement of magnetic susceptibility was 99 performed using the AGICO Kappabridge model MFK1-A. To ensure the 100 reproducibility of our results, we have meticulously documented all experimental 101 procedures, including the setup, equipment used, and analytical methods. Our 102 findings have been validated by testing multiple independent samples and 103 conducting experiments repeatedly under controlled conditions. 104

The determination of the Saturation Isothermal Remanent Magnetization (SIRM) was carried out to assess the concentration of ferromagnetic and imperfect antiferromagnetic minerals in the samples. The calculation of the Hard Isothermal Remanence (HIRM) magnetization was performed to identify magnetically

- significant components such as hematite in the samples using the followingformula:
- 111 HIRM = 0.5(SIRM + IRM 0.3T)

112 Where IRM–0.3T is the remanence after application of a reversed field of 0.3 T

after growth and measurement of SIRM. The HIRM reflects the contribution

specifically of the imperfect antiferromagnetic minerals hematite and goethite

- 115 (Bloemendal et al., 2008).
- 116 The S-0.3T value, or S-ratio, is calculated as
- 117 S=0.3T = 0.5[(-IRM=0.3T/SIRM) + 1]

and it ranges from 0 and 100%. It reflects the ratio of ferrimagnetic to imperfectantiferromagnetic minerals (Bloemendal et al., 2008).

Based on the magnetic susceptibility results, 70 samples were selected for 120 geochemical analyses (trace elements) to assist the paleoclimatic interpretation of 121 the magnetic signals. Each sample was washed using a sieve with a mesh size of 122 400 and then dried in an oven. Once dried, the samples were further sieved with a 123 325-mesh sieve. The very fine sediments were collected, packed, and labeled as the 124 tested material in special containers. A 0.2-gram portion of the powder from each 125 sample was then placed in a 1 molar hydrochloric acid solution. After two hours, 126 the samples were analyzed using an ICP device in the laboratory. The 127 concentrations of the main elements were measured as a percentage, while the 128 minor elements were quantified in milligrams per kilogram. To ensure the 129 reproducibility of the results, we meticulously document all experimental steps, 130 including the setup, equipment used, and analysis methods. The findings of this 131 research have been validated through the use of multiple independent samples and 132 by conducting experiments under controlled conditions. 133

As explained, the studied area was previously studied by Frichen et al. (2009) and 134 Kehl et al (2005). Therefore, we chose this sedimentary section to investigate 135 climate changes and used their dating data. The infrared stimulated luminescence 136 (IRSL) technique is utilized for this dating. Forty-five samples were taken in light-137 tight tubes for the IRSL dating study. About 250 g of sediment was sampled. 138 Polymineral fine-graine material (4-11 mm) was prepared for the measurements. 139 The sediment material brought on disc was irradiated by a ⁹⁰Sr/⁹⁰Y source in at 140 least seven dose steps with five discs each and a radiation dose up to 750 Gy. All 141 discs were stored at room temperature for at least 4 weeks after irradiation. The 142

irradiated samples were preheated for 1 min at 230 °C. De values were obtained by 143 integrating the 1–10 s region of the IRSL decay curves. An exponential growth 144 curve was fitted to the data and compared with the natural luminescence signal to 145 estimate the De value. Alpha efficiency was estimated to 0.08 ± 0.02 for all samples. 146 Dose rates were calculated from potassium, uranium and thorium contents, as 147 measured by gamma spectrometry (Germanium detector) in the laboratory, 148 assuming radioactive equilibrium for the decay chains. The IRSL ages gradually 149 increase with depth from 20.5 ± 2.0 to 103 ± 10 ka. The stratigraphically oldest 150 sample was collected below the lowermost exposed strongly developed palaeosol 151 (PC2) at a depth of 16.10m below surface. The about 10.50 m thick loss covering 152 the uppermost strong paleosol (PC1) likely accumulated between about 61.9±6.7 153 and 20±2.0 ka (Frechen et al, 2009). 154

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- 156

157 **Results Magnetic properties**

In Figure 3, the relationship between susceptibility, NRM (Natural Remanent 158 Magnetization), SIRM, HIRM, and S-0.3T in the Nowdeh section is illustrated. 159 The variability in the magnetic susceptibility signal within the Nowdeh section 160 indicates fluctuations in climate conditions and associated mechanisms during the 161 Late Quaternary period. The values of magnetic susceptibility (χ) in the Nowdeh 162 section range from 28.17 to 203.13 (in units of 10⁻⁸ m³ kg⁻¹). The maximum χ 163 values (203.13) are found in the lower paleosol layer at 19.4 meters depth, while 164 the minimum values are observed in the uppermost loess layer at 7.4 meters depth. 165 The rock magnetic records exhibit a strong correlation with the lithology observed 166 in the Nowdeh section. Generally, the paleosol layers exhibit higher magnetic 167 signal intensities compared to the loess layers. 168

In figure 3 The paleosols exhibit higher magnetic susceptibility (χ) values 169 compared to the loesses, with magnified magnetic enhancement observed in the 170 Bw, Bt, and Btk horizons, while the underlying C (loess) horizon displays lower χ 171 values. This difference probably reflects precipitation of iron oxides in the Bw 172 horizons, resulting in a higher concentration of pedogenetic magnetite in 173 comparison to the C horizons (Jordanova et al., 2013; Hosek et al., 2015). As 174 illustrated in Figure 3 The χ values in the lower and middle sections of the Nowdeh 175 profile, approximately 53-80 and 120-140 thousand years ago (Ka), (respectively at 176

depths of 9 to 15 and 18 to 23 meters), represent intermediate values between
unweathered loesses and weathered paleosols.

The results indicate that the Natural Remanent Magnetization (NRM) is consistent 179 with the variance in magnetic susceptibility, particularly notable at lower depths, 180 with the highest recorded value of this parameter observed at 13.1 meters depth in 181 the BW, BWK horizon (figure 3). Variations and discrepancies in magnetic 182 susceptibility align closely with the SIRM values of the Loess sequence. As 183 magnetic susceptibility decreases, SIRM also shows a corresponding decrease. In 184 the interval between 20 to 50 thousand years ago (ka (Depth 2.1 to 8.4 meters), 185 during which much of the upper Loess formation occurred, magnetic susceptibility 186 shows minimal variation, a pattern mirrored in the SIRM diagram for this period. 187 The elevated HIRM values in Figure 3 suggest an increase in the concentration and 188 frequency of magnetic deterrent minerals such as Goethite, maghemite, or 189 hematite. 190

In figure 3 The comparison between the lower values of saturation (S) (-0.3 T) 191 (between 0.6 to 0.12 Am/m) and the higher values of Hard Isothermal Remanent 192 Magnetization (HIRM) (between 2 to 5 Am/m) indicates that the proportion of 193 minerals with lower saturation, such as magnetite, is significantly lower than the 194 proportion of minerals with higher saturation in paleosols. This pattern contrasts 195 with the composition of loess deposits. As illustrated in Figure 3, This can be 196 clearly seen in loess sediments, for example, at a depth of 1 to 9 meters, 197 representing the time interval between 18 and 52 ka. 198 199

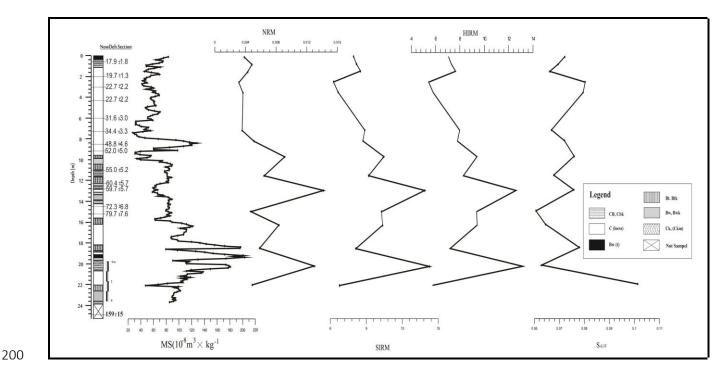




Figure 3: Basic magnetic parameters for Nowdeh section.

202 Element stratigraphy

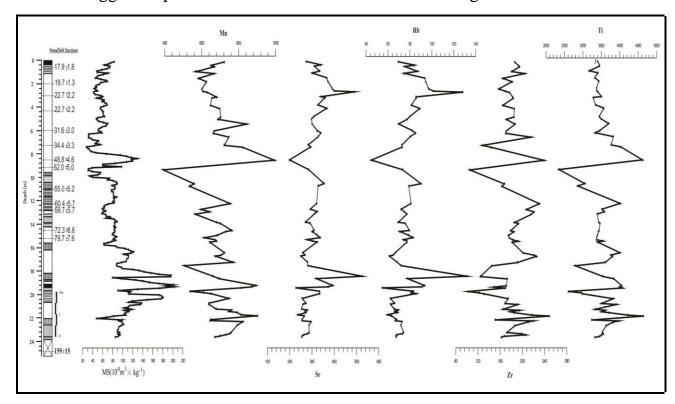
Figure 4 illustrates the correlation between the concentration of selected elements (Sr, Rb, Zr, Ti, and Mn) and magnetite susceptibility in the Nowdeh section. The figure indicates significant variations in the concentration of these elements with noticeable differences between them. Sr and Rb exhibit similar trends along the Nowdeh section. At a depth of 2.9 meters, there is a notable increase in the concentration of these two elements, corresponding to an age of 22 thousand years ago (ka). Higher in the section, the concentration of Sr and Rb decreases.

In figure 4 At a depth of 18 meters, the Nowdeh sedimentary section recorded the highest concentrations of elements such as (Sr, Rb, Zr, Ti and Mn). Conversely, the lowest concentrations of these elements were observed at a depth of 8.5 meters,

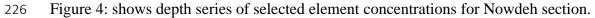
which dates back approximately 48, Ka.

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In figure 4Ti, Zr, and Mn exhibit approximately similar trends in the diagram. These elements show little variation in concentration at the beginning of the section. The changes in element concentrations from the end of the sedimentary section down to a depth of 16.7 meters (approximately 90 ka) display a zigzag pattern. Between the depths of 16.7 and 9.3 meters, the fluctuations in element concentrations are minimal. At a depth of 9.3 meters, corresponding to roughly the
last 52 ka, the research indicates the lowest concentrations of the elements
measured. However, from this point onward, the concentration of elements begins
to rise, peaking at a depth of 8.5 meters, which dates back about 48 ka. This
increase suggests a period of hot and humid conditions during that time.



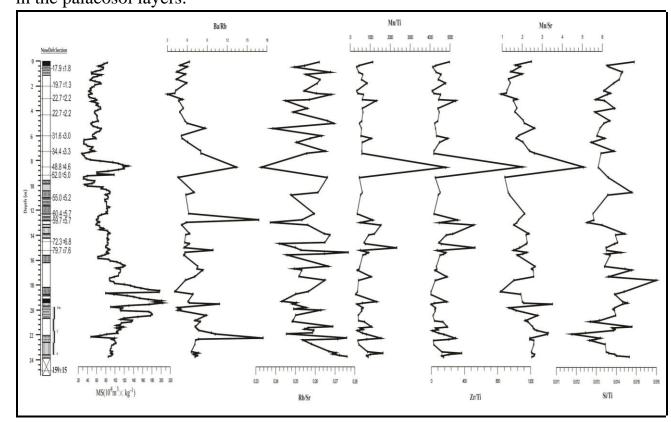
225



227 Trace element ratio

The variation of the Si/Ti ratio in figure 5 generally follows the magnetic 228 susceptibility pattern (figure 5), except for the lower part of the section (23-24 229 meters). The ratios of Mn/Sr, Zr/Ti, and Mn/Ti in figure 5 show almost no long-230 term change, except for at a depth of 8.5 meters, corresponding to an age of 48.8 231 thousand years. These changes suggest hot and humid climatic conditions, which 232 can be correlated with the high level of magnetic susceptibility. The Rb/Sr ratio 233 exhibits an opposite pattern to the magnetic susceptibility, especially at the depths 234 of 8.5, 16, 19, and 22 meters. The Ba/Rb ratio generally follows the magnetic 235 susceptibility pattern, except at depths of 13, 15, 19, and 22.8 meters where they 236 vary oppositely. 237

The variation in the Si/Ti ratio does not exhibit a consistent relationship to the sequence of loess/palaeosol layers, as defined by the magnetic susceptibility, in the Nowdeh section. On the other hand, the Mn/Ti ratios tend to show elevated values in the palaeosols, likely due to the concentration of Mn oxide in the finer sediment fraction (Bloemendal et al., 2008). This suggests that the presence of Mn oxide plays a significant role in influencing the Mn/Ti ratio in the sediments, particularly in the palaeosol layers.



245 246

Figure 5: show selected element ratios in Nowdeh section

247 Discussion

Over the entire 159 Ka sequence at the Nowdeh site, there appears to be a 248 reasonable first-order co-variation between the magnetic and geochemical 249 indicators of weathering and soil formation, particularly with magnetic parameters 250 reflecting variations in ferrimagnetic content and Sr-based ratios. However, upon 251 closer detailed examination based on individual loess and palaeosol layers, an 252 253 inconsistent relationship is observed between the amplitudes of individual peaks and troughs of magnetic and geochemical parameters (Fig 4 and 5). This suggests 254 that while there is an overall correlation between these indicators at a broader 255 scale, at a finer resolution within specific layers, the relationship becomes more 256

complex and inconsistent. Additional factors or processes may be influencing the 257 variations in magnetic and geochemical parameters within the individual 258 stratigraphic units. This issue can be seen clearly in Figures 4 and 5. As noted by 259 Hosek et al (2015) and Makeey et al (2024), there is a significant relationship 260 between magnetic receptivity, chemical elements, and climatic conditions. Our 261 study reinforces this finding, as indicated by the results obtained. Recent studies 262 have highlighted a significant relationship between climate change and the 263 magnetic properties of sediments, indicating that alterations in sediment 264 composition and depositional environments can reflect past climate conditions. For 265 instance, Huang et al. (2022) demonstrated that variations in magnetic 266 susceptibility in loess deposits are closely linked to fluctuations in moisture levels 267 and temperature, underscoring the sensitivity of magnetic minerals to climatic 268 changes. These findings suggest that analyzing the magnetic properties of 269 sediments can provide valuable insights into historical climate dynamics and 270 contribute to our understanding of how magnetic signatures may evolve in 271 response to ongoing climate change. To investigate the relationship between 272 climate change and the magnetic properties of sediments, we conducted magnetic 273 susceptibility measurements on loess sediments from the Nowdeh section. The 274 results of the magnetic susceptibility analysis demonstrated clear patterns that 275 correlate with historical climatic fluctuations. The observed distinct sequences in 276 magnetic susceptibility reveal insights into past environmental conditions, with 277 lower values indicating cold and dry periods typical of loess deposition, while 278 higher values correspond to warmer and more humid phases associated with 279 paleosol development. These correlations between magnetic susceptibility and 280 alternating loess-paleosol sequences provide a compelling illustration of how 281 sediment magnetic properties reflect climatic changes over time in the Nowdeh 282 region. The implications of these findings are significant, as they suggest that 283 magnetic susceptibility can serve as a reliable proxy for reconstructing 284 paleoclimatic conditions. By establishing this relationship, we enhance our 285 understanding of how the sedimentary environment responded to climate shifts 286 during the Pleistocene. Furthermore, these patterns of magnetic susceptibility not 287 only inform us about specific climatic conditions but also help elucidate the 288 feedback mechanisms between climate and weathering processes in the region. In 289 essence, our analysis underscores the potential of magnetic susceptibility as a key 290 indicator of past climate change, linking geological records with broader climatic 291 events. This enhanced understanding is crucial for interpreting how similar climatic 292

fluctuations may affect sediment dynamics in other regions with comparable loessdeposits.

According to Song et al. (2008), sediment loess is typically formed under cold and 295 dry climate conditions, leading to lower magnetic susceptibility values due to 296 minimal weathering processes. This observation aligns with the findings in the 297 Nowdeh section, where our data indicates that the loess layers are characterized by 298 significantly lower magnetic susceptibility than the overlying paleosols. In the 299 paleosols of the Nowdeh section, formed through extensive pedogenic processes, 300 we observe elevated magnetic susceptibility attributed to increased levels of 301 oxidation and enhanced concentrations of key elements such as iron and aluminum 302 oxides. Our magnetic susceptibility data corroborate the assertion by Song et al. 303 (2008) that paleosols generally exhibit higher magnetic susceptibility compared to 304 adjacent loess layers. Specifically, our measurements indicate a distinct rise in 305 magnetic susceptibility values within these paleosols, further supporting the notion 306 that the formation of strong magnetic minerals, such as iron oxides (Fe3O4, y-307 Fe2O3, and Fe2O3), occurs through pedogenesis. Moreover, the composition of 308 magnetic minerals in our loess samples is influenced by the grain composition 309 from the aeolian sources of sedimentation. This distinction is crucial, as our data 310 illustrates a clear contrast in magnetic mineral content between the well-developed 311 paleosols and the less altered loess layers. In our analysis, we found that the 312 paleosols not only have higher magnetic susceptibility but also exhibit a more 313 diverse range of magnetic mineral types, indicating a more complex pedogenic 314 history. This relationship highlights the importance of understanding the processes 315 of magnetic mineral formation and alteration in determining the magnetic 316 susceptibility profiles of sediment sequences. Our own findings reinforce the 317 established understanding from previous studies, such as those by Song et al. 318 (2008), while providing new insights into the specific conditions and processes at 319 play in the Nowdeh section. 320

In Fig. 3, the brown layer sequences of dark and light paleosols within the loess deposits illustrate distinct weathering processes that reflect the climatic patterns observed during glacial and interglacial periods of the middle and late Pleistocene. The higher magnetic susceptibility of the paleosols compared to the surrounding loess layers indicates a significant degree of pedogenesis and oxidation, consistent with findings by Maher (2011) and Spassov (2002). This difference is particularly pronounced at lower depths, suggesting that these layers experienced greater

weathering variability during those periods. At a depth of 21 meters (approximately 328 110 ka), the notable decrease in magnetic susceptibility suggests a transition to 329 colder and drier conditions, which aligns with the known climatic shifts of that era. 330 This finding is crucial because it highlights how magnetic susceptibility can serve 331 as a proxy for past climatic conditions, providing insights into the environmental 332 changes that influenced soil development (Thompson & Oldfield, 2021; Liu et al., 333 2022). The magnetic susceptibility chart for the Nowdeh section reveals around 334 eight distinct periods of increasing magnetic susceptibility, indicative of elevated 335 temperatures and humidity, consistent with findings from other regions that link 336 magnetic properties to climatic variations (Dearing et al., 2023; Heller & Liu, 337 2020). The magnetic susceptibility chart for the Nowdeh section reveals around 338 eight distinct periods of increasing magnetic susceptibility, indicative of elevated 339 temperatures and humidity. This pattern suggests a correlation between magnetic 340 susceptibility and climatic conditions, where periods of increased susceptibility 341 correspond to warmer, more humid phases conducive to soil formation. 342

In accordance with the standard global loess characteristics, paleosols consistently 343 exhibit higher magnetic susceptibility values compared to adjacent loess layers due 344 to pedogenesis and oxidation processes, as highlighted by Maher (2011) and 345 Spassov (2002). The NRM results suggest a decrease during loess formation and an 346 increase during paleosol formation (figure 3). This pattern suggests a relationship 347 between NRM and magnetic susceptibility figure 3 (Bloemendal et la, 2008,). A 348 decrease in NRM indicates dry and cold Climate conditions in figure 3 (The depth 349 of 7.2 meters, which is approximately equal to 34 Ka), consistent with the 350 deposition of loess layers, while an increase in NRM represents warmer and more 351 humid climate conditions, corresponding to paleosol formation. The results 352 presented in the previous section illuminate the significant relationship between 353 NRM (Natural Remanent Magnetization) and magnetic susceptibility, contributing 354 to our understanding of past climatic conditions. The observed decrease in NRM at 355 a depth of 7.2 meters (approximately 34 ka) is indicative of drier and colder 356 climate conditions, consistent with the processes associated with loess deposition. 357 This finding reinforces the notion that periods of extensive loess accumulation 358 correspond with colder climatic phases, as seen in other studies of similar 359 stratigraphic sequences. Conversely, the increase in NRM at depths of 18.6 to 21.3 360 meters marks a shift toward warmer and more humid conditions, which correlates 361 with paleosol formation. This transitional phase highlights the dynamic interplay 362

between climate and soil development, suggesting that optimum conditions for soil 363 formation fostered the development of paleosols during this time. The highest 364 magnetic susceptibility values in this layer, documented in Fig. 3, further 365 corroborate the enhancing environmental conditions during this interval. The peak 366 alignment of NRM and magnetic susceptibility at a depth of 19.4 meters, and 367 approximately 120 ka, is particularly noteworthy. This correlation may signify a 368 period of climatic stability that allowed for the establishment of rich soil profiles, 369 crucial for understanding the ecological dynamics at play during the late 370 Pleistocene. Such findings align with previous research by Bloemendal et al. 371 (2008), which also emphasized the relevance of magnetic properties in interpreting 372 paleoclimatic conditions. 373

The probable justifications for the low alteration in magnetic susceptibility and isothermal remnant magnetization between 20 to 50 Ka (Figure 3) can be attributed to two main factors:

- 1- Decreased Pedogenesis due to cold and dry periods.
- ³⁷⁸ 2- Reduction in the influx of magnetic particles into loess layers.

During the last 20 ka, there seems to be a correlation between magnetic 379 susceptibility variations in the surface soil layer and climatic conditions. This 380 period coincides with the transition from cold Climate to the current warm and 381 humid climate in the northern region of Iran (Frichen et al, 2009). As a result, the 382 soil's magnetic properties, specifically the saturation isothermal remanent 383 magnetization (SIRM), have likely increased during this time frame. However, 384 since the SIRM samples were only collected at magnetic susceptibility peak points, 385 they may not capture the full extent of variations. Comparing these findings with 386 the research by Antoine et al. (2013) on loess/paleosol sediments in Central Europe 387 reveals a close relationship, particularly around 32 ka. The close relationship 388 observed in our findings and those of Antoine et al. around 32 ka suggests a 389 synchronous response of the magnetic susceptibility of sediments to environmental 390 changes during this period. This synchrony reinforces the idea that widespread 391 climatic or geological factors were influencing sediment formation across regions. 392 Also, Antoine et al. provided critical insights into the paleoenvironmental 393 conditions in Central Europe during the Late Pleistocene. 394

395 Geochemical charts can serve as useful indicators of Climate patterns, as they can 396 highlight different levels of weathering severity. In the study of loess deposits,

certain chemical ratios can be utilized to reconstruct variations in paleoclimate 397 (Ding et al., 2001). The Zr/Ti, Mn/Ti, Rb/Sr, and Mn/Sr records from the Nowdeh 398 section exhibit a clear pattern of higher values prevailing in the palaeosols, and 399 their high degree of similarity is noteworthy. Rb/Sr has been suggested by several 400 researchers as an indicator of pedogenic intensity in loess, based on the differential 401 weathering of the major host minerals, specifically K-feldspar for Rb and 402 carbonates for Sr (Hosek et al, 2015, makeey et al, 2024). In the case of Mn/Sr, the 403 higher values observed in the palaeosols are likely a result of the combined effects 404 of grain size on Mn concentration, as well as the loss of Sr through solution 405 processes. This indicates that these ratios can serve as important indicators of 406 pedogenic processes and weathering dynamics in the sedimentary record of the 407 Nowdeh section. Rubidium is derived from K-feldspar, while strontium comes 408 from carbonate minerals. As soils weather, an increase in Rb/Sr often indicates 409 selective retention of Rb while Sr is leached away, reflecting greater weathering 410 intensity (Bai et al., 2022). Climate-driven changes, such as increased 411 precipitation, can enhance Sr leaching, leading to higher Rb/Sr ratios in wetter 412 conditions. This relationship highlights how climate influences elemental cycling 413 in soils (Zhang et al., 2021). 414

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The magnetic susceptibility record shows high values at a depth of 19.4 meters, 416 indicating hot and humid climate conditions prevailing around 120 ka. Variations 417 in the concentrations of manganese (Mn), zirconium (Zr), and titanium (Ti) in the 418 soil reflect a clear stratigraphic pattern (figure 4), with higher values seen in 419 paleosols and lower values in the loess layers (Bloemendal et al., 2008). This 420 pattern is influenced, in part, by carbonate dilution/concentration effects, as a 421 significant portion of the variability in these elements disappears when expressed 422 on a carbonate-corrected basis. In a study by Chen et al. (1999), a comparison was 423 made between the Rb/Sr ratios and magnetic susceptibility values in the uppermost 424 (last glacial/interglacial) sections of the Luochuan and Huanxian regions. The 425 researchers noted a remarkable correspondence between the amplitudes of 426 variation in magnetic susceptibility and Rb/Sr ratios. This finding suggests a close 427 relationship between magnetic susceptibility variations and the Rb/Sr ratios in 428 these regions during the last glacial and interglacial periods. 429

In the Nowdeh section, the amount of rubidium (Rb) in paleosols was lower compared to its concentration in loess layers (Figure 5). This discrepancy can be

attributed to the higher solubility of Rb in warm and humid climates, typical of 432 interglacial periods. Gallet et al. (1996) observed significant depletion of Rb in the 433 paleosols, supporting this interpretation. Recent research by Zhu et al. (2021) 434 demonstrated that climate change alters the biogeochemical cycling of nutrients, 435 including Rb, leading to increased leaching in saturated soils during periods of 436 heavy rainfall. Additionally, studies by Arias-Ortiz et al. (2020) highlighted that 437 changing climate conditions contribute to shifts in soil chemistry and nutrient 438 availability, affecting soil formation and stability. Their findings underline the 439 complex interactions between climate variations and elemental behavior in soil 440 profiles. Furthermore, a comprehensive review by Jiang et al. (2020) emphasized 441 the impacts of climate change on soil properties, particularly focusing on leaching 442 processes and the resultant changes in nutrient concentration due to increased 443 precipitation and temperature. 444

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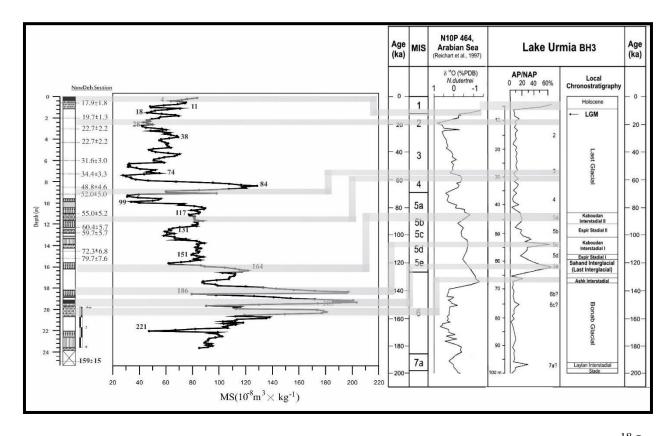
Our results indicate that the Mn/Ti, Zr/Ti, and Mn/Sr ratios tend to exhibit higher 446 values in the paleosols. According to Ding et al. (2001), elevated Mn/Ti values in 447 paleosols may result from the concentration of iron (Fe) and manganese (Mn) 448 oxides in the finer sediment fractions. They also noted that the Rb/Sr and Mn/Sr 449 ratios show a clear pattern of elevation in the paleosols, which aligns with the 450 findings of our study (figure 5). The Rb/Sr ratio has been proposed by various 451 researchers as an indicator of pedogenic intensity in loess deposits, based on the 452 differential weathering of major host minerals such as K-feldspar for Rb and 453 carbonates for Sr. The higher Mn/Sr values in paleosols may be attributed to grain-454 size effects on Mn concentrations and the solubilization loss of Sr. 455

- Chen et al. (1999) compared Rb/Sr and magnetic susceptibility in the uppermost 456 parts of the Luochuan and Huanxian sections, revealing a significant 457 458 correspondence between the variations in magnetic susceptibility and Rb/Sr ratios. This suggests a link between weathering intensity and magnetic properties in these 459 sediments. In the context of the Nowdeh sedimentary section, the magnetic 460 parameters were compared with those from other studies conducted in various 461 regions of the world, further contributing to our understanding of paleoclimatic 462 variations and weathering processes in loess deposits. 463
- In figure 6 The comparison of magnetic receptivity results from the Nowdeh sedimentary section with the palynological data from sedimentary cores of Urmia Lake (Djamali et al., 2008) and the ¹⁸O analysis from Arabian Sea sedimentary cores (Tzedakis, 1994) has provided valuable insights into past climate conditions

(Figure 6). In the analysis, an increase in the AP/NAP index (Arboreal Pollen
grains (AP) to that of the Non-Arboreal Pollen grains (NAP)) in the lakes
corresponded with the presence of ancient soil layers in the seedling sedimentary
section.

This increase signifies warmer temperatures and higher humidity levels, conducive 472 to the growth of trees and shrubs (Harrison et al., 2020; Zhang et al., 2021). 473 Conversely, a decrease in the AP/NAP index indicates a decline in temperature and 474 humidity, leading to the disappearance of trees and shrubs and changes in surface 475 vegetation cover (Chen et al., 2022). This correlation suggests that the climate 476 conditions and their fluctuations in western Iran align with the sedimentary 477 deposition at Nowdeh, consistent with findings from similar studies in the region 478 (Fuchs et al., 2013; Hosek et al., 2015). 479

Moreover, in figure 6 the ¹⁸O analysis of the Arabian Sea exhibited a strong 480 agreement with magnetic receptivity data. A decrease in the ¹⁸O indexes points to 481 warmer climate conditions, while an increase indicates colder conditions (Diamali 482 et al, 2008). The relationship between magnetic susceptibility and ¹⁸O levels in the 483 Arabian Sea sediments, as shown in Figure 6, Verify that an increase in magnetic 484 susceptibility corresponds with a decrease in ¹⁸O levels, indicating warmer climate 485 conditions. This alignment further supports the connection between the recorded 486 palynology data of Lake Urmia, ¹⁸O data from the Arabian Sea, and the sequence 487 of ancient loess-soil sediments in the Nowdeh sedimentary section. 488



489

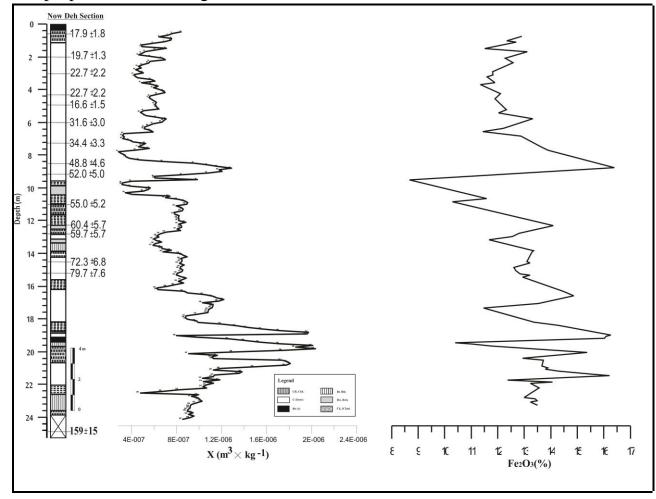
Figure 6: Correlation between recorded palynological data of Lake Urmia (Djamali et al, 2008) and ¹⁸O
 of Arabian Sea sediments (Tzedakis, 1994) with the Loess-Paleosol sediment sequence of Nowdeh
 sedimentary section.

The results of our current research demonstrate a significant correlation with the 493 findings of Fuchs et al. (2013) and Hosek et al. (2015), which examined ancient 494 loess and paleosol deposits in Central Europe, indicating that similar magnetic 495 properties can be observed across different regions in response to climate change. 496 Figure 8 depicts consistent patterns in the magnetic receptivity parameter over the 497 past 45, 73, 90, 104, and 108 ka across the study sections. Around 45 and 73 ka, 498 there is a clear increasing trend in magnetic receptivity observed in all analyzed 499 layers, indicating a shift towards warmer and more humid climate conditions 500 compared to earlier periods. This increase in magnetic susceptibility can be 501 attributed to the higher presence of iron oxides in the soil resulting from increased 502 chemical weathering. Conversely, during the periods of 90, 104, and 108 ka, a 503 decrease in magnetic susceptibility is evident across all regions, signifying colder 504 and drier climatic conditions during these time intervals. This issue can also be 505 seen in the amount of $F_{e_2}O_3$ in Nowdeh sediments (Figure 7). While the older 506 sediments also show a significant association with climate variations in Central 507

508 Europe and the Nowdeh area, the absence of radiometric dating in these older 509 sediments introduces some uncertainty when interpreting these findings. 510 Nonetheless, the consistent patterns in magnetic susceptibility across different time 511 periods provide valuable insights into past climate fluctuations and their impact on

512 soil properties in these regions.

513



514 Figure 7: The relationship between magnetic susceptibility and $F_{e2}O_3$

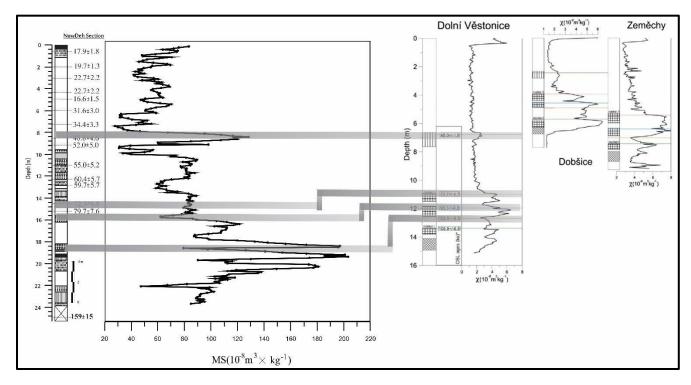




Figure 8: Comparison of changes in magnetic receptivity of Dolní Věstonice sedimentary section, Fuchs
et al, 2013, Dobsice and Zemechy section, Hošek et al, 2015, with Nowdehh sedimentary section

518

The comparison of magnetic receptivity trends as recorded in sedimentary sections of Beiyuan, Heimugou, Biampo, and the ¹⁸O records by Imbrie et al. (1984) in Figure 8 reveals a high agreement with the Nowdeh sedimentary section. This alignment indicates similar Climate conditions across different locations in the Northern Hemisphere.

The consistency in magnetic receptivity trends among these various sites suggests a commonality in the climatic conditions experienced during the corresponding time periods. This synchronization in magnetic susceptibility patterns further supports the notion that these regions were subjected to comparable environmental changes and fluctuations in the past.

Additionally, the correlation observed between the magnetic receptivity data and the ¹⁸O records underscore the close relationship between climatic factors and sedimentary deposition patterns across these sites (Figure 9). By examining these geological proxies, researchers can gain valuable insights into the past climate dynamics and variations that have affected the Northern Hemisphere over time.

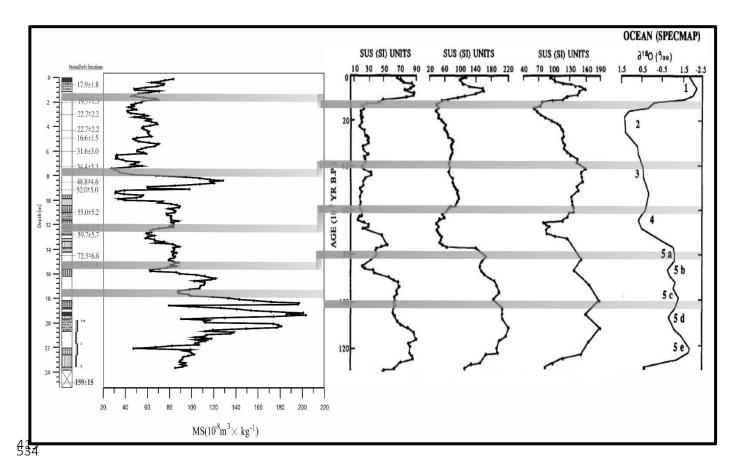
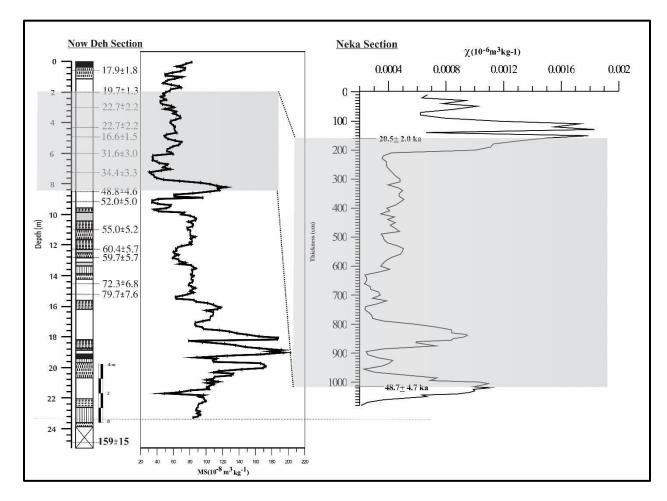


Figure 9: Comparison of magnetic receptivity changes of Beiyuan, Heimugou, Biampo, An et al, 1991,
 records of ¹⁸O Imbrie et al, 1984 with Nowdeh sedimentary section

The findings of Mehdipour et al. in 2012 in the realm of fine loess exhibit a close 537 resemblance to the results presented in our research, as illustrated in Figure 9. In 538 their study, they employed both magnetic and geochemical approaches to assess 539 different climatic periods, and the outcomes align significantly with the findings of 540 Our research. The comparison in Figure 10 reveals a strong consistency in the 541 magnetic receptivity trends between the Nowdeh section and the Neka sedimentary 542 section analyzed by Mehdipour et al (2012). Between 48 and 20 thousand years 543 ago, notable similarities are observed in the fluctuations of magnetic receptivity in 544 both sedimentary sections. Whenever there is an increase in magnetic receptivity, it 545 indicates a warm and humid period with the formation of ancient soil layers. This 546 shared pattern implies a synchrony in climatic conditions between the two regions 547 during this time frame, showcasing the utility of magnetic susceptibility as a proxy 548 for understanding past environmental changes and soil development processes. 549 550



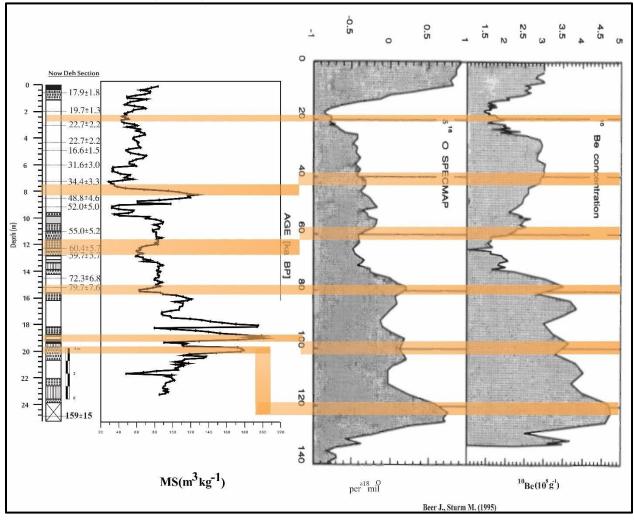
551

552 Figure 10: Comparison of magnetic receptivity diagram of Nowdeh sedimentary section with Neka 553 sedimentary section (Mahdi et al., 2012)

The results of this research exhibit a strong consistency with the findings of Beer and Sturm (1995) regarding beryllium saturation in the Zaifang sedimentary section and ¹⁸O in marine sediments. In both cases, there is a clear correlation between the fluctuations in beryllium saturation, ¹⁸O, and magnetic receptivity.

558 When beryllium saturation and ¹⁸O decrease, there is a corresponding decrease in 559 magnetic receptivity, indicating colder and drier climate conditions. Conversely, an 560 increase in beryllium saturation and ¹⁸O is accompanied by an increase in magnetic 561 receptivity, signifying warmer and more humid periods.

The high agreement between the climatic periods identified based on these parameters in the Zaifang sedimentary section and marine sediments, and the magnetic receptivity trends observed in the Nowdeh sedimentary section, highlights the synchrony of similar climate events in the past across different locations. This consistency further supports the robustness of magnetic 567 susceptibility as a proxy for understanding past climate variations and 568 environmental changes (figure 11).



569

Figure 11: Comparison of magnetic receptivity results of Nowdeh sedimentary section in comparison with 18 O and Be 10 isotope results of Xifeng sedimentary section (Beer and Sturm, 1995).

572

573 **Conclusion**

In conclusion, the loess/paleosol sequences from Northeastern Iran serve as a 574 valuable archive for studying the paleoenvironmental changes during the Upper 575 employing By multi-proxy approach Pleistocene. a that integrates 576 sedimentological, magnetic, and geochemical methods, the following key insights 577 have been revealed: 578

The stratigraphy of the studied section aligns well with the typical pattern of
 Upper Pleistocene loess/paleosol successions in the region, providing
 valuable insights into the past environmental conditions.

- Magnetic parameters show a strong correlation with climate conditions,
 making them effective variables for reconstructing climate change patterns
 in the region.
- 3. Comparisons between magnetic and geochemical data indicate that
 variations in geochemical weathering ratios mirror changes in magnetic
 weathering parameters, such as magnetic susceptibility, further enhancing
 our understanding of past environmental dynamics.
- The high degree of coherence observed between the amplitudes of magnetic susceptibility and various geochemical ratios, including Rb/Sr, Mn/Ti, Zr/Ti, and Mn/Sr, reinforces the reliability of magnetic susceptibility as a proxy for tracking environmental changes and provides additional insights into the interplay between magnetic and geochemical processes.
- 594 Overall, this comprehensive multi-proxy analysis enhances our understanding of 595 the paleoenvironmental changes in Northeastern Iran during the Upper Pleistocene 596 period and emphasizes the importance of integrating sedimentological, magnetic, 597 and geochemical data to unravel past climatic fluctuations and environmental 598 dynamics.
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