

1 **Late Aptian paleoclimate reconstruction of Brazilian**  
2 **equatorial margin: inferences from palynology**

3 Michelle Cardoso da Silva Giannerini<sup>1</sup>, Marcelo de Araujo Carvalho<sup>1</sup>,

4 Cecília Cunha Lana<sup>1</sup>, Gustavo Santiago<sup>1</sup>, Natália de Paula Sá<sup>1</sup>, Gabriel da Cunha

5 Correia<sup>1</sup>

6 <sup>1</sup>Laboratorio de Paleoecologia Vegetal (LAPAV), Departamento de Geologia e Paleontologia, Museu  
7 Nacional Universidade Federal do Rio de Janeiro; 20940-040, Rio de Janeiro, Brazil.

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9 *Corresponding author:* giannerini@gmail.com

10 **Abstract.**

11 This study conducted high-resolution paleoclimatic analyses based on the identification  
12 of palynological groups from the late Aptian age (Biozone *Sergipea veriverrucata*) in  
13 the Bragança and Codó formations within the Bragança-Viseu, São Luís, and Parnaíba  
14 basins. The analysis comprised 40 palynological samples, with 200 palynomorphs per  
15 slide counted when possible. Bioclimatic analysis was mainly supported by the  
16 identification of botanical affinities, and ecological and climatic parameters such as  
17 wet/arid trend (Fs/X), Shannon-Wiener diversity, and indicator species analysis  
18 (IndVal) were used. Statistical analyses such as principal component and cluster  
19 analyses were employed to support the paleoclimatic interpretations. The study  
20 recognized 69 genera distributed among the main groups of living plants, including  
21 bryophytes, ferns, lycophytes, gymnosperms, and angiosperms. It was possible to  
22 attribute botanical affinity in 94.2% of the taxa, and nine genera occurred in all sections  
23 studied: *Afropollis*, *Araucariacites*, *Callialasporites*, *Cicatricosisporites*, *Classopollis*,  
24 *Cyathidites*, *Deltoidospora*, *Equisetosporites*, and *Verrucosisporites*, with *Classopollis*

25 being the most abundant. The stratigraphic distribution of the bioclimatic groups  
26 (hydrophytes, hygrophytes, lowland tropical flora, upland flora, and xerophytes)  
27 allowed for the identification of climatic phases: pre-evaporitic, evaporites, and post-  
28 evaporites. In the pre-evaporitic phase, the most significant abundances were between  
29 the hygrophytes and upland flora, indicating a certain level of humidity. Xerophytes  
30 were the most abundant in all phases, with a conspicuous increase in the evaporitic  
31 phase, reflecting an increase in aridity. In the post-evaporitic phase, there was a  
32 significant increase in the upland flora with the return of wetter conditions. This study  
33 confirmed an increasing humidity trend in the analyzed sections, probably owing to the  
34 influence of the Intertropical Convergence Zone that already operated during the late  
35 Aptian.

36

## 37 **1. Introduction**

38 The palynoflora preserved in the upper Aptian rocks of South America and Africa  
39 is typical of hot conditions and is commonly associated with arid climates (Chumakov  
40 et al., 1995, Hay and Floegel, 2012). However, because biodiversity tends to be higher  
41 in wetter climates, the high diversity observed during the Aptian raises the possibility  
42 that this arid phase fluctuated during that ~~period~~age. The palynoflora related to hot and  
43 humid climates exhibits a growing trend toward these conditions, even during the  
44 Aptian. This trend may be linked to the shifting and strengthening of a humid belt  
45 associated with the Intertropical Convergence Zone (ITCZ) (Hay and Floegel, 2012;  
46 Scotese, 2016; Carvalho et al., 2022; Santos et al., 2022), as well as to the establishment  
47 of the South Atlantic, which affected the marine current system.

48 Palynology plays an important role in paleoclimate studies, because analyzing the  
49 assemblages of palynomorphs (e.g., spores, pollen grains), it is possible determine from  
50 the botanical affinities the types of plants that existed in the past and infer their climatic  
51 preferences. For instance, certain pollen types are indicative of wetter climates, while  
52 others are associated with drier conditions. Palynological analysis can also provide  
53 evidence of seasonal changes, temperature variations, moisture levels, and extreme  
54 climatic events.~~Palynology is a valuable tool for inferring paleoclimatic conditions~~  
55 ~~based on palynomorph studies.~~ Late-Upper Aptian rocks from Brazilian sedimentary  
56 basins, including the Bragança and Codó formations, contain a significant  
57 representation of conifers from the family Cheirolepidiaceae and their pollen grains,  
58 such as *Classopollis* (Regali et al., 1974; Carvalho et al., 2017, 2019, 2022).  
59 *Classopollis* is typically associated with arid conditions, often found in lagoons and  
60 coastal areas, and frequently associated with evaporites (Batten, 1975; Vakhrameev,  
61 1970, 1981; Doyle et al., 1982; Hashimoto, 1995; Heimhofer et al., 2008, Carvalho et  
62 al., 2019). However, studies of the Sergipe Basin (northwestern Brazil), suggest strong  
63 fluctuations in the abundance of *Classopollis* and other xerophytic flora, with a  
64 decreasing trend toward the late Aptian accompanied by an increase in fern spores that  
65 require water for reproduction (Carvalho et al., 2017, 2019). The geographic extent of  
66 these trends remains controversial, and further investigation is required to identify  
67 possible climatic oscillations in other sedimentary basins in Brazil. Analysis of the  
68 Codó and Bragança formations, located in the Cretaceous section of the São Luís,  
69 Bragança-Viseu, and Parnaíba basins near the paleoequator, where the Intertropical  
70 Convergence Zone (ITCZ) occurs, has great potential to provide insights into this topic.

71 The objective of this study ~~was-is~~ to infer the paleoclimate of the late Aptian  
72 ~~period-age~~ in the Bragança-Viseu, São Luís, and Parnaíba basins, all located in

73 equatorial areas (Fig. 1), by examining the relationships ~~between~~among groups of  
74 palynomorphs that are sensitive climatic taxa. Furthermore, this study aimed to  
75 investigate how variations in the composition of paleofloras and indicator species are  
76 linked to climatic changes such as shifts in humidity and temperature, as well as other  
77 paleoenvironmental forcings.

78

## 79 **2. Geological settings**

80 According to Milani et al. (2007), the three sedimentary basins considered in this  
81 study are grouped into large assemblies based mainly on the tectonic context in which  
82 they developed: Mesozoic aborted rift basins (Bragança-Viseu and São Luís basins) and  
83 Paleozoic Synclises (Parnaíba Basin).

84 The Bragança-Viseu, São Luís, and Parnaíba basins show a similar stratigraphic  
85 evolution. The Bragança-Viseu and São Luís basins are located on the equatorial margin  
86 and the Parnaíba Basin in north-central Brazil (Fig. 1). The basins constitute a rift  
87 system (graben and semi-graben) located between the terrain~~ess~~ of the folding belt.  
88 Together, these cover an area of approximately 645,000 km<sup>2</sup>. The sedimentary  
89 succession of the basins consists of Paleozoic, Mesozoic, and Cenozoic rocks. The  
90 Cretaceous strata are represented by the Bragança (Bragança-Viseu and São Luís  
91 basins), Grajaú, Codó, and Itapecuru Formations. ~~In the studied sections, the Bragança~~  
92 ~~(EGST-01 and VN-01 wells) and Codó (CI-01, PR-01, PE-01, RL-01 wells) formations~~  
93 ~~were recognized.~~

94 The Bragança Formation consists of gray medium- to coarse-grained sandstones  
95 and conglomerates, with subordinate medium-grained sandstones supported by

96 conglomerates and greenish siltstones. This formation is interpreted as an alluvial fan  
97 deposit.

98 The Codó Formation is composed of dark shales, anhydrite, and calcilutites, with  
99 sandstone intercalations. These deposits were assigned to a lagoonal environment.  
100 Marine incursions are indicated by fossil contents and the occurrence of evaporitic  
101 deposits.

102

### 103 3. Late Aptian climatic evolution

104 The pre-evaporitic, evaporitic, and post-evaporitic phases are recognized for the  
105 late Aptian (Petri et al., 1983; Milani et al., 2007). These phases occur within the K40-  
106 K50 supersequences, and show an average maximum thickness of approximately 650 m  
107 in the studied basins. The pre-evaporitic phase is represented by carbonate and  
108 siliciclastic deposits formed in fluvial and lacustrine deltaic environments within a large  
109 proto-oceanic gulf (Petri et al., 1983; Milani et al., 2007). The peak of the evaporitic  
110 deposition is recorded in the K50 supersequence, with widespread occurrences in the  
111 Brazilian equatorial margin. The origin of these deposits is heat intensification  
112 associated with the widening of the Atlantic Ocean. These conditions caused strong  
113 evaporation, leading to a wide distribution of evaporites (mainly halite and anhydrite  
114 gypsum) in the South Atlantic basins. The post-evaporitic phase is characterized by  
115 fully marine conditions, evidenced by the rich assemblages of marine fossils. During  
116 this phase, carbonates were deposited, followed by muddy and sandy sediments, in  
117 shallow marine to slope environments (Petri et al., 1983; Milani et al., 2007).

118 The Bragança and Codó formations are inserted within the K40-K50 Supersequence.  
119 However, in the Bragança Formation, only the pre-evaporitic phase is recognized. On

120 the other hand, the Codó Formation has recorded the three climatic phases (pre-  
121 evaporitic, evaporitic, and post-evaporitic) (Milani et al., 2007).

122

123

## 124 **4. Material and methods**

125

### 126 **4.1. Studied sections**

127 This study was based on core samples from three basins: Bragança-Viseu and São  
128 Luís located in the equatorial margin, and the Parnaíba Basin in north-central Brazil.  
129 All of the core samples were obtained from Petrobras (the Brazilian oil company)  
130 drilling. (Fig. 1).

131 The stratigraphic succession studied comprises parts of the Bragança and Codó  
132 formations. The Bragança Formation (Bragança-Viseu Basin) includes wells EGST-1  
133 (676-1872.1 m), consisting of sandstones, siltstones, and conglomerates, and VN-1  
134 (1287.6-1317.69 m), consisting only of sandstones (Fig. 2) (Table 1). The Codó  
135 Formation includes three section from the São Luís Basin: PR-1 (1507.6-1513.1 m),  
136 composed of sandstones and siltstones, and PE-1 (1562-1776.8 m), which has a lithology  
137 similar to that of the previous one, with the addition of calcarenites. RL-1 (1157.3-1240.3  
138 m) is composed of sandstones, siltstones, calcarenites, and anhydrites. The fourth section,  
139 CI-1 (768-907.1 m), is from the Parnaíba Basin. CI-1 has a lithology similar to that of  
140 RL-1, but the former has a more pronounced package of anhydrites than the latter does  
141 (Fig. 2) (Table 1).

142 ~~This study was based on core samples from three basins: Bragança Viseu and São~~  
143 ~~Luís located in the equatorial margin, and the Parnaíba Basin in north-central Brazil.~~

144 ~~The material for this study was derived from the cores of wells EGST-1 and VN-1 from~~  
145 ~~the Bragança Viseu Basin; PR-1, PE-1, and RL-1 from the São Luís Basin; and CI-1~~  
146 ~~from the Parnaíba Basin (Table 1), all drilled by Petrobras (the Brazilian oil company)~~  
147 ~~(Fig. 1). Detailed information on the studied sections is presented in Table 1.~~

148 ~~The stratigraphic succession studied comprises parts of the Bragança Viseu, São Luís,~~  
149 ~~and Parnaíba Basins. The Bragança Viseu Basin includes wells EGST-1 (676–1872.1 m),~~  
150 ~~consisting of sandstones, siltstones, and conglomerates, and VN-1 (1287.6–1317.69 m),~~  
151 ~~consisting only of sandstones (Fig. 2) (Table 1).~~

152 ~~Three sections are from the São Luís Basin: PR-1 (1507.6–1513.1 m), composed of~~  
153 ~~sandstones and siltstones, and PE-1 (1562–1776.8 m), which has a lithology similar to~~  
154 ~~that of the previous one, with the addition of calcarenites. RL-1 (1157.3–1240.3 m) is~~  
155 ~~composed of sandstones, siltstones, calcarenites, and anhydrites. The fourth section, CI-~~  
156 ~~1 (768–907.1 m), is from the Parnaíba Basin. CI-1 has a lithology similar to that of RL-1,~~  
157 ~~but the former has a more pronounced package of anhydrites than the latter does (Fig. 2)~~  
158 ~~(Table 1).~~

159 The late Aptian age of the samples is based on the *Sergipea variverrucata*  
160 Biozone recognized in two studied drill cores (PR-1 and CI-1), which is correlated with  
161 part of the upper Aptian *Globigerinelloides algerianus* Zone (Carvalho et al., 2016). In  
162 the other four sections (EGST-1, VN-1, PE-1, and RL-1), *Sergipea variverrucata* was  
163 not recognized. However, the identified floristic associations (e.g., *Afropollis jardinus*,  
164 *Araucariacites australis*, *Bennettittapollenites regaliae*, *Equisetosporites maculosus*,  
165 *Klukisporites foveolatus*, *Sergipea simplex*) are attributed to the late Aptian of Brazil  
166 (Regali and Santos, 1999; Carvalho et al., 2017, 2019).

## 169 4.2. Sample preparation

170 The samples were prepared at the Research and Development Center of Petrobras  
171 (CENPES) in Rio de Janeiro, ~~applying. The method used was~~ the standard Petrobras  
172 method of palynological, ~~preparation~~ compiled by Uesugui (1979) ~~on the basis of~~  
173 ~~methods developed~~ based on by Erdtman (1943, 1969) and Faegri et al. (1966). Thus, in  
174 this study, most mineral constituents were dissolved by hydrochloric and hydrofluoric  
175 acids before heavy-liquid separation, and the remaining organic matter was sieved  
176 through a 10 µm mesh before mounting on slides. The cores are stored at CENPES (Rio  
177 de Janeiro, RJ).

## 179 4.3. Palynological analyses

180 The samples were analyzed using a transmitted light microscope. Analysis was  
181 based on the first 200 palynomorphs counted on each slide. The marine elements  
182 (dinoflagellate cysts and microforaminiferal linings) were counted separately.  
183 Taxonomic identification was based on the methods of Regali et al. (1974), Lima  
184 (1978), Dino (1992, 1994), and Carvalho et al. (2019, 2022).

## 186 4.4. Bioclimatic analysis

187 Palynomorphs are useful climatic indicators (bioclimatic groups) because of their  
188 botanical affinities that allow the application of the ecological preferences of taxa.  
189 However, identifying the spores and pollen grains of the parent plant classified at the  
190 family level is often challenging. This study referred to the literature (e.g., Dino, 1994;



191 Carvalho, 2004; Souza-Lima and Silva, 2018; Jansonius et al., 1976-1996) to identify  
192 the botanical affinities of the indicator species.

193 On the basis of botanical affinities and inferred paleoenvironmental conditions  
194 ([e.g., Dino 1992, Balme 1995; Antonioli, 2001; Carvalho et al., 2017, Carvalho et al.,](#)  
195 [2019, Carvalho et al., 2022](#)), this study proposes five bioclimatic groups: hydrophytes,  
196 hygrophytes, tropical lowland flora, upland flora, and xerophytes. These groups provide  
197 valuable insights into the climate and vegetation of the study area.

198

#### 199 **4.5. Wet-dry trend**

200 To support the bioclimatic group distribution, we used the Fs/X (fern spores  
201 *versus* xerophytes) ratio. This ratio is based on the co-occurrence of fern spores and  
202 xerophytic palynomorphs (*Classopollis* and polyplicate gnetalean pollen); therefore, it  
203 can be used as an indicator of dry-wet trends (Carvalho et al., 2019). The ratio of fern  
204 spores to xerophytic palynomorphs (Fs/X) was calculated as  $Fs/X = nFs / (nFs + nX)$ ,  
205 where n is the number of specimens counted, Fs is the number of fern spores (non-  
206 reworked), and X is the number of xerophytic pollen grains. In summary, Fs/X  
207 approaching 1 implies high humidity, and that approaching -1 indicates low humidity.

208

#### 209 **4.6. Diversity**

210 Shannon-Weaver diversity indices H (S) were calculated for all samples by using  
211 PAST software (Hammer et al., 2001) to provide information for interpreting  
212 paleoclimatic trends. Diversity H(S) considers the abundance of each species and is  
213 used to characterize the diversity of the assemblages.

214

215

216

#### 217 **4.7. Indicator species**

218 To characterize the climate changes during the late Aptian based on paleoflora, we  
219 employed the indicator species analysis (IndVal) method. The IndVal is a widely used  
220 measure in ecological studies to evaluate the association between a particular species  
221 and a specific habitat or environmental condition. The IndVal index quantifies the level  
222 of association between a species and a habitat by considering two components:  
223 specificity and fidelity. Specificity refers to the extent to which a species is associated  
224 with a particular habitat or condition, while fidelity represents the probability of finding  
225 a species in a habitat given its occurrence in the overall study area. The IndVal index  
226 has demonstrated successful applications in palynological studies (Caron and Jackson,  
227 2007; Roucoux et al., 2013; Carvalho et al., 2017; Trindade and Carvalho, 2018,  
228 Leandro et al. 2020). In our study, the IndVal index was employed to assess the degree  
229 of association between taxa and specific sample groups corresponding to different  
230 paleoclimatic phases. To reconstruct the late Aptian vegetation in the studied basins, we  
231 employed the indicator species analysis (IndVal) method. The IndVal is a measure used  
232 in ecological studies to assess the association between a particular species and a specific  
233 habitat or environmental condition. The IndVal index quantifies the association between  
234 a species and a habitat by considering two components: specificity and fidelity.  
235 Specificity refers to the degree to which a species is associated with a particular habitat  
236 or condition, while fidelity represents the probability that a species is present in a

237 ~~habitat given its occurrence in the overall study area. This index was applied for the first~~  
238 ~~time in pre-Quaternary samples in the der Sergipe Basin (Carvalho et al., 2017).~~

239 ~~The IndVal index determines the taxa strongly associated with particular groups~~  
240 ~~of samples and is assumed to reflect the climatic conditions of those groups.~~ It was  
241 calculated using the formula proposed by Dufrière and Legendre (1997):  $\text{IndValGroup } k, \text{ Species } j = 100 \times A_{k,j} \times B_{k,j}$ , where  $A_{k,j}$  represents specificity, and  $B_{k,j}$  represents  
242 fidelity. We used PAST software (Hammer et al., 2001) to calculate these values.

244 To ensure that our IndVal analysis fulfilled the criteria of ordination and climate-  
245 focused approach, we grouped the samples according to three climatic phases: pre-  
246 evaporitic, evaporitic, and post-evaporitic. This allowed us to identify the specific  
247 indicator species associated with each climatic phase and gain insights into the  
248 vegetation that existed during the late Aptian ~~periodage~~.

## 250 5. Results

251 Sixty-nine genera were identified in the 40 samples and were distributed into five  
252 plant groups: bryophytes (four genera), ferns (17 genera), lycophytes (10 genera),  
253 ~~pteridosperms (one genus)~~, gymnosperms (23-24 genera), and angiosperms (14 genera)  
254 (Appendix 1) (Table 2). Twenty indeterminate morphotypes were found in ferns and 10  
255 in angiosperms. Of the 69 genera identified, nine occurred in all the wells studied:  
256 *Afropollis*, *Araucariacites*, *Callialasporites*, *Cicatricosisporites*, *Classopollis*,  
257 *Cyathidites*, *Deltoidospora*, *Equisetosporites*, *Verrucosisporites*. The suggested  
258 botanical affinity of the 69 genera was 94.2%. The 5.8% without botanical affinity  
259 refers to the group of angiosperms.

261 ~~The recognition of botanical affinities of the 69 genera was based on the literature~~  
262 ~~(e.g., Dino 1992, Balme 1995; Antonioli, 2001; Carvalho et al., 2017, Carvalho et al.,~~  
263 ~~2019, Carvalho et al., 2022) and the database of fossilworks.org. The suggested~~  
264 ~~botanical affinity of the 69 genera was 94.2%. The 5.8% without botanical affinity~~  
265 ~~refers to the group of angiosperms.~~

266 All bioclimatic groups were present in the studied sections (Table 3, Appendix 2).

267 In general, the palynological assemblage is predominantly composed of the xerophytic  
268 bioclimatic group, characterized by a high abundance of *Classopollis*. The average  
269 abundance of xerophytes was 55.7%, ranging from 46.3% to 63.6% in the sections  
270 studied (Table 4). In sequence, the upland flora had an overall average abundance of  
271 18.9% (ranging from 7.8% to 26%), with *Araucariacites* being the dominant taxon. The  
272 hygrophyte bioclimatic group is characterized by the presence of *Cicatricosisporites*,  
273 which had an average abundance of 18.6% (ranging from 11.4% to 28.4%). By contrast,  
274 the hydrophyte bioclimatic group is the least abundant, with an overall average of 0.7%,  
275 and is dominated by the genus *Crybelosporites*. Regarding diversity, the Shannon-  
276 Wiener indices (H') obtained for the 40 samples showed an overall average of H'= 2.0,  
277 which ranged from H'=1.6 in the VN-1 section to H'= 2 .2 in section PE-1 (Table 4).  
278 The values of the wet-dry trend (Fs/X ratio) ranged from 0.19 (dry) in section CI-1 to  
279 0.39 (wet) in EGST-1(wet) (Table 4).

### 281 **5.1. Stratigraphic distribution of bioclimatic groups in EGST-1 well**

282 Although xerophytes are dominant overall, EGST-1 well exhibits a higher  
283 abundance of hygrophytes (24.9%) due to moderate to high occurrences of  
284 *Cicatricosisporites* and *trilete psilate* (e.g., *Cyathidites*), especially at the base of the  
285 well (Fig. 3). Additionally, the abundance of hygrophytes, tropical lowland flora, and

286 upland flora increases toward the upper sections, whereas the abundance of xerophytes  
287 decreases (Fig. 4). The Shannon-Wiener indices ( $H'$ ) showed an overall average of  $H'=$   
288 2.1, slightly above the general average ( $H'= 2.0$ ). The Fs/X ratio had the highest value  
289 for all sections (0.38), above the overall average (0.28), indicating more humid  
290 conditions (Table 4).

291

## 292 **5.2. Stratigraphic distribution of bioclimatic groups in VN-1 well**

293 Similar to the EGST-1 well, the VN-1 well is composed of four samples from the  
294 Bragança Formation, in which xerophytes dominate. However, unlike the former well,  
295 hygrophytes exhibit the highest average abundance (28.4%) among all studied wells,  
296 primarily because of the abundance of trilete psilate. Despite few samples, an increasing  
297 trend of hygrophytes, tropical lowland flora, and upland flora was observed, with a  
298 significant peak in hygrophytes (Fig. 4). The average diversity of  $H'=1.6$  is the lowest  
299 for the studied basins, below the overall average ( $H'= 2.0$ ). The Fs/X ratio was 0.31,  
300 above the overall average (0.28).

301

## 302 **5.3. Stratigraphic distribution of bioclimatic groups in PR-1 well**

303 The section comprises four samples from the Codó Formation. Notably, the PR-1  
304 well exhibits the lowest average abundance of xerophytes (46.3%) (Table 4). However,  
305 it shows the highest average abundance in the tropical lowland flora group (20.4%) of  
306 all the wells studied, driven by the presence of the genus *Afropollis*. In general, an  
307 increasing trend toward hygrophytes, upland flora, and mainly tropical lowland flora  
308 was observed (Fig. 5). The average diversity was  $H'= 2.1$  in this well. This value is one  
309 of the highest values among all the wells studied. This high diversity is mainly

310 attributed to the significant number of species belonging to the tropical lowland flora  
311 group. The Fs/X ratio was 0.25, slightly below the overall average (0.28) (Table 4).

312

#### 313 **5.4. Stratigraphic distribution of bioclimatic groups in PE-1 well**

314 The PE-1 well shows a clear decreasing trend upward of the xerophytes, which  
315 did not exceed 20% (Fig. 6). By contrast, hygrophytes and upland flora show a  
316 conspicuous increase. Highlight for the upland flora group show an average of 26%  
317 driven by the genus *Araucariacites*. The average diversity of  $H'=2.2$  is the highest for  
318 the basins. This average diversity is due to the many species of upland flora and  
319 hygrophytes. The Fs/X ratio was 0.28, the same as the overall average (0.28) (Table 4).

320

#### 321 **5.5. Stratigraphic distribution of bioclimatic groups in RL-1 well**

322 The section consists of seven samples from the Codó Formation. The xerophytic  
323 bioclimatic group dominated the entire section, with no ~~sudden-abrupt~~ changes in the  
324 abundance curve observed, except at the base of the section, where the hygrophytes,  
325 tropical plain flora, and upland flora groups together reached almost 40% (Fig. 7). The  
326 average diversity of  $H'=1.9$  is the second lowest for the studied basins. The Fs/X ratio  
327 was 0.24, slightly below the overall average (0.28) (Table 4).

328

#### 329 **5.6. Stratigraphic distribution of bioclimatic groups in CI-1 well**

330 The Parnaíba Basin is represented by one well, which comprises 13 samples from  
331 the Codó Formation. The palynological assemblage of this section was dominated by  
332 the xerophytic bioclimatic group, with a high average of 63.6%, largely because of the  
333 abundance of *Classopollis* and *Equisetosporites*. [The abundance curves of bioclimatic](#)  
334 [groups show that in the base occurs a balance between the xerophytes and the other](#)

335 groups, especially the upland flora, and in the top a clear dominance of xerophyte  
336 group~~The dendrogram shown in Fig. 8 reveals two intervals: the base with a greater~~  
337 ~~balance between xerophytes and the other groups, especially the upland flora (15.9%)~~  
338 (Table 4). The Fs/X ratio recorded the lowest value in all sections (0.19), which was  
339 below the overall average (0.28), indicating drier conditions (Table 4). However,  
340 despite this, the average diversity of  $H'=2.0$  was one of the highest, with the same value  
341 as the overall average of 2.0.

342

### 343 **5.7. Climatic phases**

344 All six sections were individually analyzed for palynology. However, a composite  
345 section was constructed (Table 5) based on the stratigraphically evident chronological  
346 distribution of the climatic phases in each studied section. The composite section of the  
347 Bragança-Viseu, São Luís, and Parnaíba basins consists of 40 samples, with 24 samples  
348 from the pre-evaporitic phase, eight from the evaporitic phase, and eight from the post-  
349 evaporitic phase (Table 5). In general, the composite section highlights the bioclimatic  
350 groups of hygrophytes (18.8%) and tropical lowland flora. The diversity and Fs/X ratio  
351 curves showed strong synchrony, indicating a relation between diversity and humidity  
352 (Fig. 9). No marine elements were recorded in these sections.

353 During the pre-evaporitic phase, there is a higher abundance of xerophytes,  
354 hygrophytes, and upland flora, but with strong oscillations observed in their respective  
355 curves. The dendrogram in Fig. 9 identifies two intervals within this phase: with  
356 significant values of xerophytes at the base but with a decreasing trend toward the top.  
357 The interval above, the xerophyte curve exhibits an upward trend. The diversity and  
358 Fs/X ratio curves show synchrony but with a decreasing trend toward the top. The  
359 indicator species (IndVal) identified for the pre-evaporitic phase, *Deltoidospora* spp.

360 (Cyatheaceae-Dicksoniaceae) is related to the montane rainforest, suggesting more  
361 humid conditions (Table 5).

362 The evaporitic phase, which corresponds to the gypsum layers of the Codó  
363 Formation, is characterized by the highest average of the xerophytic bioclimatic group  
364 in the composite section (Table 5). Additionally, the average abundance of the tropical  
365 lowland flora group is also high, driven by the genus *Afropollis*. Surprisingly, the mean  
366 diversity is high during this phase, but the mean Fs/X ratio is the lowest. The high  
367 diversity in arid conditions is due to the great diversity of species in the xerophytic  
368 group, such as *Classopollis classoides*, *Equisetosporites maculosus*, and  
369 *Gnetaceaepollenites jansonius*. The IndVal for the evaporitic phase is *Afropollis* spp.  
370 related to tropical lowland flora (Table 5).

371 The post-evaporitic phase, which includes part of a section of the Codó  
372 Formation, is characterized by a significant decrease in the dominance of the xerophytic  
373 bioclimatic group; lower average abundance (47%) in PR-1; and the clear dominance of  
374 hygrophyte groups, including tropical lowland flora and upland flora. The dendrogram  
375 reveals a break between this phase and the evaporitic phase (Fig. 9). In general, this  
376 reflects an inversion in abundance between groups related to humidity (hygrophytes,  
377 hydrophytes, tropical flora, and upland flora) and groups related to drier conditions  
378 (xerophytes) (Fig. 9). In this phase, the indicator species is *Deltoidospora* spp.,  
379 suggesting more humid conditions for pre- and ~~pre~~post-evaporitic phases.

380

## 381 **6. Discussion**

382 The data obtained from these sections provide clear evidence of the dominance of  
383 the xerophytic bioclimatic group during the late Aptian in Brazilian sedimentary basins.



384 This information supports that in the literature that suggests an essentially arid climate  
385 during this period-age (e.g., Lima, 1983; Suguio and Barcelos, 1983; Petri, 1983;  
386 Rossetti et al., 2003; Hay and Floegel, 2012, Carvalho et al., 2017, 2019, 2022). These  
387 authors attributed this aridity to the predominance of conifers from the  
388 Cheirolepidiaceae family and their *Classopollis* pollen grains. However, climatic  
389 oscillations were identified during this period-age, indicated by bioclimatic groups  
390 related to the humid conditions: hydrophytes, hygrophytes, tropical lowland flora, and  
391 upland flora. A relationship between these groups has been suggested (e.g., Carvalho et  
392 al., 2017, 2019, 2022). In this study, principal component analysis (PCA) was  
393 conducted between bioclimatic groups that exhibited patterns similar to those observed  
394 in the literature (e.g., Carvalho et al., 2017, 2019, 2022). The PCA revealed a positive  
395 correlation among hygrophytes, hydrophytes, tropical lowland flora, and upland flora,  
396 whereas xerophytes show a negative relationship on the first axis (Component 1) (Fig.  
397 10), explaining more than 70% of the variation. Component 1 characterizes the wet-dry  
398 trend.

399 The sections of the São Luís Basin (PE-1, RL-1, and PR-1) showed the lowest  
400 abundance of xerophytic flora, followed by the sections of the Bragança-Viseu Basin  
401 (VN-1 and EGST-1) and the CI-1 section (Parnaíba Basin) farther south (Fig. 11A).  
402 More humid conditions also were suggested by Santos et al. (2022) for the São Luís  
403 Basin. This study utilized palynological data and PCA analysis to propose the existence  
404 of a wet phase during the late Aptian in the São Luis Basin. Through the analysis of the  
405 abundance of *Araucariacites* and fern spores, as well as the presence of the genus  
406 *Classopollis* associated with carbonate sedimentation in two semi-arid intervals, an  
407 intermediate humid interval was identified. According to Santos et al. (2022), ~~the~~  
408 authors suggested that the data were sufficient to identify a pre-Albian humid belt,

409 which challenges the view of exclusively arid Gondwana during the Aptian and  
410 supports the presence of a wet phase.

411 As also suggested by Carvalho et al. (2022), we compared the studied sections  
412 with sections in the Espírito Santo Basin, located much farther south (at 20°S). We  
413 found that the studied basins had a lower abundance of xerophytic flora than the  
414 Espírito Santo Basin did (Fig. 11B-C). The decreasing trend in aridity observed from  
415 the southeast (Espírito Santo Basin) to the northeast (Fig. 11B-C) coincides with the  
416 location of the hot and humid belt attributed to the ITCZ (Ohba et al., 2010, Chaboureau  
417 et al., 2012, 2014; Scotese, 2016). Notably, the approach to the ITCZ belt, where xeric  
418 restrictions are milder, reflected even in the most aridity phase, the evaporitic phase,  
419 whose indicator species was the *Afropollis* spp. of the lowland tropical flora. This  
420 indicates that the ITCZZC must have had diminished aridity. The genus *Afropollis* has  
421 been associated with hot, humid climates. According to Carvalho et al. (2022), this  
422 genus exhibits the weakest negative correlation with xerophytic flora (e.g.,  
423 *Classopollis*).

424 The ITCZ belt proposed by Scotese (2016) for the Aptian covers the entire  
425 African continental paleoequator. However, although very close, it did not reach South  
426 America (Fig. 11B). Palynological analyses conducted by Deaf et al. (2020) on the late  
427 Aptian material of the Dahab Formation (Mathruh Basin, Egypt) indicated a  
428 predominance of fern spores from the hygrophyte bioclimatic group (e.g.,  
429 *Triplanosporites*, *Cicatricosisporites*) and uplands (e.g. *Deltoidospora*, *Araucariacites*),  
430 accounting for approximately 60% on average. This finding suggests that the Dahab  
431 Fformation is characterized by humid conditions.

432 The xerophytic flora (*Classopollis* and *Equisetosporites*) in the Dahab Formation  
433 averaged approximately 25%. Considering the climatic belts proposed by Scotese  
434 (2016, 2021), this formation occurred “inside” the ITCZ, which is reflected in the  
435 prevalence of bioclimatic groups associated with more humid conditions. The  
436 abundance of xerophytic flora in the Dahab Formation (~~Mathrub Basin~~) was lower than  
437 that in the sections studied. This difference was particularly significant when compared  
438 with the Espírito Santo Basin, where the abundance of xerophytic flora was 87.3%, as  
439 opposed to 25% in the Dahab Formation (Fig. 11C). Notably, a significant contributor  
440 to the humidity in the Dahab Formation was likely a marine influence, which was not  
441 present in the sections studied.

442

## 443 7. Conclusion

444 The Aptian sections studied have well-preserved palynological diversity  
445 dominated by the genera *Classopollis* (Cheirolepidiaceae) and *Araucariacites*  
446 (*Araucariaceae*). Some genera of ferns are also abundant such as *Cicatricosisporites*  
447 (*Anemiaceae*), *Verrucosisporites* (*Osmundaceae*), and *Deltoidospora* (*Cyatheaceae*).

448 Five bioclimatic groups were identified and proposed for interpretation:  
449 hydrophytes, hygrophytes, tropical lowland flora, upland flora, and xerophytes. The  
450 bioclimatic groups provide evidence that the climate during the late Aptian was arid.  
451 However, when considering the distribution curves of bioclimatic groups, as well as the  
452 indicator species (IndVal) and diversity, a clear upward trend toward increased humidity  
453 was observed.

454 The late Aptian period-age was characterized by three distinct climatic phases:  
455 pre-evaporitic, evaporitic, and post-evaporitic. During the pre-evaporitic phase, despite

456 the dominance of xerophytic flora, there were episodes of humidity, evidenced by  
457 indicator species such as *Deltoidospora* spp. The evaporitic phase was dominated by  
458 xeric elements, although the moderate to high abundance of lowland tropical flora,  
459 confirmed by *Afropollis* spp. as an indicator species, indicated some periods of  
460 humidity. The post-evaporitic phase was marked by a lower abundance of xerophytic  
461 elements and a clear dominance of groups associated with wet conditions, mainly the  
462 upland flora, suggesting a wetter climate during this phase.

463 The climatic variation during the late Aptian is reflected in the palynological  
464 assemblages, with the arid phase being dominated by the genus *Classopollis* and other  
465 xerophytic bioclimatic group indicators. The wet phase is marked by a significant  
466 decrease in xerophytes and a high abundance and diversity of *Araucariacites*, fern  
467 spores, and other genera related to highland, hydrophytic, and hygrophytic bioclimatic  
468 groups. The "mirror effect" observed in the frequency curves highlights the ecological  
469 differences between the arid and humid trend groups.

470 According to our findings, vegetation dynamics were affected by a combination of  
471 the Intertropical Convergence Zone (ITCZ) and the opening of the South Atlantic Ocean  
472 during the late Aptian ~~periodage~~. The influence of the ITCZ is currently stronger in the  
473 north-central region of South America. Notably, climate evolution during the late  
474 Aptian in the South Atlantic led to increased humidity, which was closely linked to  
475 plant diversity and marine influences.

476

477 **Appendices**

478

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488

489 **Data availability**

490 The data and code used in this paper are deposited at CENPES, PETROBRAS, Rio de  
491 Janeiro, RJ, Brazil (wells VN-1, EGST-1, RL-1, PE-1, CI-1, and PR-1). Additional  
492 information on samples (wells VN-1, EGST-1, RL-1, PE-1, CI-1 and PR-1) can be  
493 accessed in [www.anp.gov.br](http://www.anp.gov.br).

494

495 **Author contributions**

496 M.C.S.G and M.A.C. led the writing with contributions of all coauthors; M.C.S.G.,  
497 C.C.L, G.S., N.P.S. and G.C.C. collected the palynological data and M.C.S.G. and  
498 M.A.C. carried out the pollen data analysis.

499

500 **Competing interests**

501 The authors declare no competing interests.

502

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