Model and proxy evidence for coordinated changes in the hydroclimate of distant regions over the Last Millennium

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Abstract. The Medieval Climate Anomaly (MCA; ca. 950-1250 CE) and the Little Ice Age (LIA; ca. 1450-1850 CE)~~, were~~ periods generally characterised by respectively higher and lower temperatures in many regions~~,~~. However, they have also been associated with drier and wetter conditions in areas around the Intertropical Convergence Zone (ITCZ), the Asian Monsoon region, and in areas impacted by large-scale climatic modes like the Northern and Southern Annular Modes (NAM and SAM, respectively). To analyse coordinated changes in large-scale hydroclimate patterns, and whether similar changes also extend to other periods of the Last Millennium (LM) outside the MCA and the LIA, reconstruction-based products have been analysed~~, including~~. This includes the collection of tree-ring based Drought Atlases (DA), the Paleo Hydrodynamics Data Assimilation product (PHYDA) and the Last Millennium Reanalysis (LMR). These analyses have shown coherent changes in the hydroclimate of tropical and extratropical regions, such as northern and central South America, East Africa, western North America, Western Europe, the Middle East, Southeast Asia and the Indo-Pacific, during the MCA, the LIA and other periods of the LM. Comparisons with model simulations from the Community Earth System Model - Last Millennium Ensemble (CESM-LME) and phases 5 and 6 of the Coupled Model Intercomparison Project (CMIP5 and CMIP6) show that both external forcing and internal variability contributed to these changes, with the contribution of internal variability being particularly important in the Indo-Pacific basin and that of external forcing in the Atlantic basin. These results may help to identify not only those areas showing coordinated changes, but also those regions ~~are able to successfully reproduce the evolution of hydroclimate during the LM~~more impacted by the internal variability, where forced model simulations would not be expected to successfully reproduce the evolution of past actual hydroclimate changes.

1 Introduction

The climate of the Last Millennium (LM) was characterised in many regions by periods of warmer and cooler conditions with respect to the pre-industrial mean climate (Diaz et al., 2011; Graham et al., 2010), two well-known examples being the Medieval Climate Anomaly (MCA; ca. 950-1250 CE) and the Little Ice Age (LIA; ca. 1450-1850 CE; Graham et al., 2007; Laird et al., 2012; Ledru et al., 2013). Such temperature changes are often associated with changes in external forcing factors like solar variability and volcanic activity (Mann et al., 2009; Schurer et al., 2013; Fernández-Donado et al., 2013). The impact of these changes on hydroclimate is not straightforward, but reconstructions show that episodes of severe drought also took

25 place during the LM in many regions (Cook et al., 2022). Periods of wetter and drier conditions existed in large areas of North America (Cook et al., 2010b), Europe (Luterbacher et al., 2012), South America (Vuille et al., 2012) and Monsoon Asia (Hu et al., 2008), with occurrences that coincide with the MCA and the LIA.

~~To survey the available evidence, we assessed 92 reconstructions reporting changes from wetter to drier or from drier to wetter conditions during the transition from the MCA to LIA (Fig. 1, with detailed references in Table 1). Some of these~~
30 ~~reconstructions show consistent drier or wetter conditions during the whole LIA (dark green and dark brown), while others show drier or wetter conditions until late 1500s CE, during the early LIA (light green and light brown), as reported for areas of Pakistan, western India and southern China by Graham et al. (2010).~~ These similarities in the evolution of the hydroclimate of distant regions suggest coordinated changes at a global scale (Graham et al., 2010; Atwood et al., 2021). For example, reconstructions from southwestern North America (Graham et al., 2007; Meko et al., 2001; Cook et al., 2004, 2010b; Hughes
35 and Funkhouser, 1998; Anderson, 2011) and the Mediterranean basin (Luterbacher et al., 2012; Martín-Puertas et al., 2010; Morellón et al., 2009) show drier conditions during the MCA and wetter conditions during the LIA~~, while~~. On the contrary, reconstructions from northwestern North America (Steinman et al., 2013; Stevens and Dean, 2008), central Europe (Büntgen et al., 2010, 2011) and the British Isles (Wilson et al., 2012; Proctor et al., 2000) tend to show wetter conditions during the MCA and drier conditions during the LIA.

40 An antiphased relationship between the MCA and LIA is also observed in tropical areas of South America~~, with~~. Wetter conditions are found during the MCA in areas of northern South America, like the Cariaco Basin in Venezuela (Haug et al., 2001), and drier conditions in areas of central South America, like Peru (Bird et al., 2011; Reuter et al., 2009; Vuille et al., 2012; Thompson et al., 1986; Apaéstegui et al., 2014) and Eastern Brazil (Vuille et al., 2012; Taylor, 2010; Novello et al., 2012). In East Africa, Anchukaitis and Tierney (2013) showed coordinated changes for lake Challa (Verschuren et al., 2009;
45 Wolff et al., 2011), Naivasha (Verschuren et al., 2000; Verschuren, 2001), Tanganyika (Tierney et al., 2010a), Masoko (Gilbert et al., 2002; Garcin et al., 2006, 2007) and Malawi (Johnson et al., 2001; Brown and Johnson, 2005; Johnson and McCave, 2008), with an opposite behavior between the lakes in the south and the north of the Intertropical Convergence Zone (ITCZ). Changes in the ITCZ have been also associated with coordinated changes in the hydroclimate of the Indian Monsoon region~~, with~~. Wetter conditions during the MCA and drier conditions during the LIA are found for wide areas of Pakistan (von Rad
50 et al., 1999; Treydte et al., 2006), India (Sinha et al., 2007; Berkelhammer et al., 2010; Bryson and Swain, 1981; Phadtare and Pant, 2006) and the Arabian Sea (Burns et al., 2002; Fleitmann et al., 2003; Anderson et al., 2002; Gupta et al., 2003). Alterations of the ITCZ have been also associated with changes in the hydroclimate of the Indo-Pacific basin (Atwood et al., 2021)~~, with~~. A marked transition between the MCA and LIA is found in areas of China (Wang et al., 2005; Zhang et al., 2008; Hu et al., 2008), Southeast Asia (Buckley et al., 2010, 2007), Indonesia (Griffiths et al., 2016; Newton et al., 2006; Tierney
55 et al., 2010b), northern Australia (Denniston et al., 2016) and the eastern Pacific islands (Sachs et al., 2009; Higley et al., 2018).

~~Most reconstructions in Fig. 1 are located in extratropical areas~~Most of the impacted regions are located in the extratropics, in the boundary between low and high pressures within the regions of the Northern and Southern Annular Modes (NAM and SAM; Thompson and Wallace, 2001; Jones et al., 2009; Fogt et al., 2009), or in tropical areas around the ITCZ ~~(Table 1)~~. Analyses based on simulated data (Roldán-Gómez et al., 2020) link transitions from the MCA to LIA in extratropical

areas with alterations in the variability of modes like the NAM and SAM (Ortega et al., 2015; Jones et al., 2009; Fogt et al., 2009), mainly driven by changes in external forcing factors (Roldán-Gómez et al., 2020). In tropical areas, the transitions from the MCA to LIA can be linked to alterations of the ITCZ (Atwood et al., 2021), like a contraction over the Western Pacific (Denniston et al., 2016; Yan et al., 2015) and shifts over East Africa (Anchukaitis and Tierney, 2013), the Indian Monsoon region (Fleitmann et al., 2003) and the Eastern Pacific (Higley et al., 2018). The analysis of the simulated hydroclimate from the Community Earth System Model - Last Millennium Ensemble (CESM-LME; Otto-Bliesner et al., 2015) performed by Roldán-Gómez et al. (2022) shows that these alterations of the ITCZ may be impacted by both external forcing factors and internal variability.

Even if individual reconstructions from different regions suggest coordinated changes in the hydroclimate during the MCA and LIA, global analyses based on proxy data are scarce and limited to certain reconstructions and products. The analysis of simulated data provides an insight into the mechanisms explaining coordinated changes in the hydroclimate of the LM, but analyses of LM simulations must account for both forced and internal variability (Roldán-Gómez et al., 2022). This work addresses these challenges by performing a comparative assessement of the LM hydroclimate at global and large-continental scales from different sources, including reconstructions, model simulations, and hybrid products based on data assimilation and reanalysis. The use of several reconstruction-based products allows the identification with a robustness assessment of those regions more impacted by hydroclimatic changes during the MCA and LIA, and more generally during the LM. The comparison between reconstruction-based products and model simulations from different ensembles allows an assessment of the contributions of external forcing and internal variability on the hydroclimate of each region, showing those regions where model simulations can provide representative information about changes in the hydroclimate.

2 Evidences from the literature

To survey the available evidence, we assessed 96 reconstructions reporting changes from wetter to drier or from drier to wetter conditions during the transition from the MCA to LIA (Fig. 1, with detailed references in Table 2). The reconstructions used in this compilation include proxy data from tree rings, marine and lake sediments, speleothems, ice cores and documentary information. These sources provide information about precipitation, moisture, level of lakes and river flows, which can be linked to drier and wetter conditions. Some reconstructions represent the annual hydroclimate while some others are sensitive to a particular season, as described in Table 2.

All the reconstructions included in the compilation cover the MCA and LIA. Some of them show consistent drier or wetter conditions during the whole LIA (dark green and dark brown), while some others show drier or wetter conditions until late 1500s CE, during the early LIA (light green and light brown), as reported for areas of Pakistan, western India and southern China by Graham et al. (2010).

The reconstructions from Fig. 1 are used as reference for the analysis of reconstruction-based and model-based products included in the following sections.
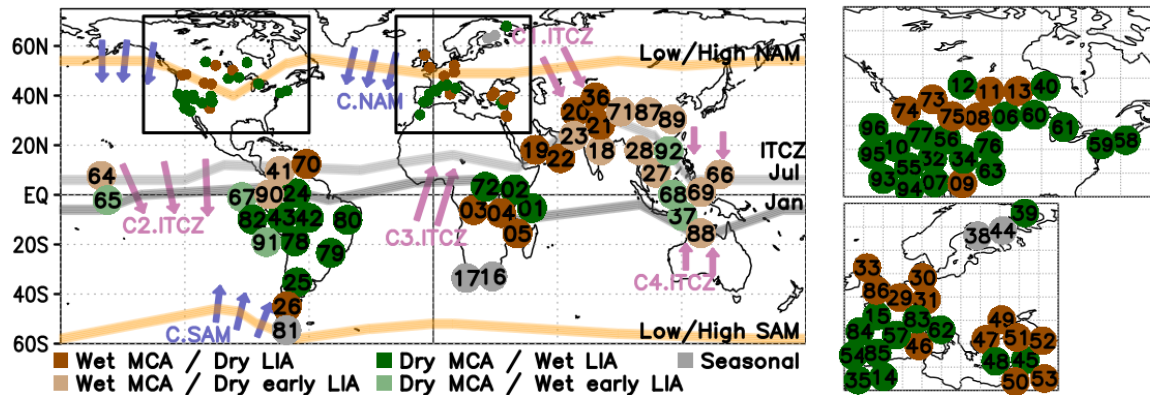
Figure 1. Hydroclimate reconstructions showing **(dark brown)** drier conditions during MCA and wetter conditions during LIA, **(light brown)** drier conditions during MCA and wetter conditions during early LIA (until late 1500s CE), **(dark green)** wetter conditions during MCA and drier conditions during LIA, **(light green)** wetter conditions during MCA and drier conditions during early LIA (until late 1500s CE), and **(gray)** drier or wetter conditions in MCA and LIA depending on the season. ~~Detailed references are included in Table 1. As per references of Table 2, A1 and A2 respectively indicate the current position of the ITCZ in July and January, B1 to B4 the ITCZ changes in the transition from MCA to LIA, C1 and C2 the boundary between low and high pressures in the pattern of NAM and SAM, and D1 and D2 the changes in this boundary in the transition from MCA to LIA.~~ The current position of the ITCZ in July and January, the ITCZ changes in the transition from MCA to LIA in the Indian Monsoon region (C1.ITCZ), Eastern Pacific (C2.ITCZ), East Africa (C3.ITCZ) and Western Pacific (C4.ITCZ), the boundary between low and high pressures in the pattern of NAM and SAM, and the changes in this boundary in the transition from MCA to LIA (C.NAM and C.SAM) are shown within the map, according to the references included in Table 1. Regional maps over North America and Europe are shown on the right side of the panel. Detailed references are included in Table 2.

3 Data and methods

The following reconstruction-based products are employed: the 100-member ensemble of the Paleo Hydrodynamics Data Assimilation product (PHYDA; Steiger et al., 2018b), the 20-member ensemble of the Last Millennium Reanalysis (LMR; Tardif et al., 2019) and the tree-ring based Drought Atlases (DA), including the Old World Drought Atlas (OWDA; Cook et al., 2015), the North American Drought Atlas (NADA; Cook et al., 2010b), and the Monsoon Asia Drought Atlas (MADA; Cook et al., 2010a) ~~and the Mexican Drought Atlas (MXDA; Stahle et al., 2016)~~. Due to the fact that they do not extend back into the MCA and LIA, other DAs have not been included in our analyses, and since the MADA does not extend back into the MCA, it has been only considered for the analysis of time series of Sect. 3.3. Even if the PHYDA and LMR provide data for the whole millennium, some of the proxies used for the assimilation do not extend so far in the past. Figure 2 shows the distribution of proxies from Steiger et al. (2018a), including those that cover the MCA and LIA. There is a large amount of proxies covering these periods in regions like North America and Europe, so for these regions the products would be expected to follow the behavior of the proxies. However, for other regions like central Africa and eastern South America, the lack of proxies could increase the contribution of the model used for the assimilation, and thus the uncertainty of the products.

Table 1. References for the position of ITCZ, the boundary between low and high pressures in the pattern of NAM and SAM, and the changes in these two elements during MCA and LIA, according to the codes included in Fig. 1.

Code	Mode	References
Jul ITCZ	Modern position of July ITCZ	Newton et al. (2006)
Jan ITCZ	Modern position of January ITCZ	Newton et al. (2006)
C1.ITCZ	ITCZ changes in the Indian Monsoon region	Fleitmann et al. (2003)
C2.ITCZ	ITCZ changes in the Eastern Pacific	Higley et al. (2018)
C3.ITCZ	ITCZ changes in East Africa	Anchukaitis and Tierney (2013)
C4.ITCZ	ITCZ changes in the Western Pacific	Denniston et al. (2016); Yan et al. (2015)
Low/High NAM	Boundary of low/high in the pattern of NAM	Li and Wang (2003)
Low/High SAM	Boundary of low/high in the pattern of SAM	Gong and Wang (1999)
C.NAM	Changes in NAM/NAO	Ortega et al. (2015)
C.SAM	Changes in SAM	Jones et al. (2009); Fogt et al. (2009)

105   Regarding the model-based products, we use the 13 all-forcing simulations of the CESM-LME, together with 10 LM simulations from the Coupled Model Intercomparison Project Phase 5 / Paleoclimate Modelling Intercomparison Project Phase 3 (CMIP5/PMIP3; Taylor et al., 2007; Schmidt et al., 2011, 2012; Stocker et al., 2013) and 4 LM simulations from the Coupled Model Intercomparison Project Phase 6 / Paleoclimate Modelling Intercomparison Project Phase 4 (CMIP6/PMIP4; Eyring et al., 2016; Jungclaus et al., 2017), all of which are described in Table 3. All these simulations have been interpolated to a

110   common grid resolution of 6º x 6º, the coarsest among the analysed simulations, to allow for the calculation of ensemble averages across the CESM-LME, CMIP5 and CMIP6 products. The CESM-LME and CMIP5 simulations are forced according to the CMIP5/PMIP3 protocol (Schmidt et al., 2011, 2012), while the CMIP6 simulations consider the CMIP6/PMIP4 protocol (Jungclaus et al., 2017). Agreement between reconstruction-based and model-based products would be only expected for those areas dominated by the external forcing, as may be the case for extratropical areas like North America and Europe

115   (Roldán-Gómez et al., 2020) and tropical areas in the Atlantic basin like South America and Central Africa (Roldán-Gómez et al., 2022), but not in those areas dominated by the internal variability, like the tropical Indo-Pacific (Roldán-Gómez et al., 2022).

Considering the wide bibliography of single-point reconstructions showing wetter or drier conditions during MCA and LIA (Fig. 1), most analyses have been based on these two periods. Following the approach of Masson-Delmotte et al. (2013), the

120   MCA has been considered from 950 to 1250 CE, and the LIA from 1450 to 1850 CE, even if these temporal intervals might not be suitable for all the regions and for all products (Neukom et al., 2014, 2019). To assess the robustness of our results to the specifics of this definition, we have performed additional analyses based on the trend during the whole period from 950 to 1750 CE, and report our findings in the Appendix A.

Table 2. Location, references, season and type of reconstruction (sediments, tree rings, speleothems, documentary, ice cores or multi-proxy) for the reconstructions represented in Fig. 1. The first column indicates the code used in Fig. 1 for each reconstruction.

Code	Location	References	Season	Type
01	Lake Challa (Tanzania)	Verschuren et al. (2009); Wolff et al. (2011)	Annual	Sediments
02	Lake Naivasha (Kenya)	Verschuren et al. (2000); Verschuren (2001)	Annual	Sediments
03	Lake Tanganyika (Tanzania, DRC)	Tierney et al. (2010a)	Annual	Sediments
04	Lake Masoko (Tanzania)	Gilbert et al. (2002); Garcin et al. (2006, 2007)	Annual	Sediments
05	Lake Malawi (Malawi)	Johnson et al. (2001); Brown and Johnson (2005); Johnson and McCave (2008)	Annual	Sediments
06	Moon Lake (North Dakota, USA)	Laird et al. (1996)	Annual	Sediments
07	California (USA)	Graham et al. (2007)	Jun-Aug, Dec-Feb	Multi-proxy
08	Bighorn (Wyoming, USA)	Gray et al. (2004)	Annual	Tree rings
09	Malpais (New Mexico, USA)	Grissino-Mayer (1995)	Annual	Tree rings
10	Sacramento River (California, USA)	Meko et al. (2001)	Annual	Tree rings
11	Humboldt Lake (Canada)	Michels et al. (2007)	Annual	Sediments
12	North Saskatchewan River (Canada)	Case and MacDonald (2003)	Jun-Aug, Dec-Feb	Tree rings
13	Red River (Canada)	St.George and Nielsen (2002)	Annual	Tree rings
14	Guadalentín River (Spain)	Benito et al. (2010)	Annual	Sediments
15	Gardon River (France)	Sheffer et al. (2007)	Annual	Sediments
16	Orange River (South Africa)	Tyson and Lindesay (1992)	Jun-Aug, Dec-Feb	Sediments
17	Buffels River (W South Africa)	Benito et al. (2011)	Jun-Aug, Dec-Feb	Sediments
18	Dandak Cave (India)	Sinha et al. (2007); Berkelhammer et al. (2010)	Jun-Aug	Speleothems
19	S Oman	Burns et al. (2002)	Annual	Speleothems
20	Lunkaransar Lake (India)	Bryson and Swain (1981)	Jun-Oct, Annual	Sediments
21	Pindar Valley (India)	Phadtare and Pant (2006)	Annual	Sediments
22	Arabian Sea (Oman)	Anderson et al. (2002); Gupta et al. (2003)	Jun-Aug	Sediments
23	Karachi (Pakistan)	von Rad et al. (1999)	Aug-Oct	Sediments
24	Lake Pumacocha (Peru)	Bird et al. (2011)	Dec-Mar	Sediments
25	N Argentina	Boucher et al. (2011)	Dec-Feb	Multi-proxy
26	Patagonia	Boucher et al. (2011)	Dec-Feb	Multi-proxy
27	S Vietnam	Buckley et al. (2010)	Jun-Aug	Tree rings
28	NW Thailand	Buckley et al. (2007)	May-Sep	Tree rings
29	NE France	Büntgen et al. (2011)	Jun-Aug	Tree rings
30	N Germany	Büntgen et al. (2010, 2011)	Jun-Aug	Tree rings
31	S Germany	Büntgen et al. (2010, 2011)	Jun-Aug	Tree rings

Code	Location	References	Season	Type
32	SW North America	Cook et al. (2004, 2010b)	Jun-Aug	Tree rings
33	Uamh Cave (NW Scotland)	Proctor et al. (2000)	Annual	Speleothems
34	Arizona and Utah (USA)	Ely et al. (1993)	Annual	Sediments
35	Morocco	Esper et al. (2007)	Feb-Jun	Tree rings
36	Karakorum Mountains (Pakistan)	Treydte et al. (2006)	Annual	Tree rings
37	Liang Luar Cave (Indonesia)	Griffiths et al. (2016)	Annual	Speleothems
38	Finland	Helama et al. (2009)	Jun-Aug, Dec-Feb	Tree rings
39	Kola Peninsula (Russia)	Kremenetski et al. (2004)	Jul-Aug	Multi-proxy
40	NW Ontario (Canada)	Laird et al. (2012)	Annual	Sediments
41	Papallacta (Ecuador)	Ledru et al. (2013)	Annual	Sediments
42	Marcacocha (S Peru)	Chepstow-Lusty et al. (2009)	Annual	Sediments
43	Cascayunga (Peru)	Reuter et al. (2009)	Annual	Speleothems
44	Saavanjoki River (Finland)	Luoto et al. (2013)	Jun-Aug, Dec-Feb	Sediments
45	Antalya and Mersin (SW Turkey)	Touchan et al. (2007)	May-Jun	Tree rings
46	Grotta Verde (Italy)	Antonioli et al. (2003)	Annual	Speleothems
47	Uzuntarla Cave (NW Turkey)	Göktürk (2011)	Annual	Speleothems
48	Kocain Cave (S Turkey)	Göktürk (2011)	Annual	Speleothems
49	Sofular Cave (N Turkey)	Göktürk et al. (2011)	Annual	Speleothems
50	Nahal Zin (Israel)	Greenbaum et al. (2000)	Annual	Sediments
51	Nar Gölü (C Turkey)	Jones et al. (2006); Woodbridge and Roberts (2011)	Annual	Sediments
52	Tercer Lake (C Turkey)	Kuzucuoglu et al. (2011)	Annual	Sediments
53	Dead Sea	Bookman et al. (2004); Neumann et al. (2007)	Annual	Sediments
54	Zoñar Lake (S Spain)	Martín-Puertas et al. (2008, 2010)	Annual	Sediments
55	Sierra Nevada (California, USA)	Stine (1994)	Annual	Tree rings
56	Great Basin (W North America)	Hughes and Funkhouser (1998)	Jun-Jul	Tree rings
57	NE Spain	Llasat et al. (2003)	Annual	Documentary
58	Little Pond (Massachusetts, USA)	Oswald and Foster (2011)	Annual	Sediments
59	Piermont Marsh (New York, USA)	Pederson et al. (2005)	Annual	Sediments
60	Hole Bog (Minnesota, USA)	Booth et al. (2006)	Annual	Sediments
61	Minden Bog (Michigan, USA)	Booth et al. (2006)	Annual	Sediments
62	Lake Accesa (Italy)	Magny et al. (2007)	Annual	Sediments
63	San Juan Mountains (Colorado, USA)	Routson et al. (2011)	Mar-Jul	Tree rings
64	Teraina (Kiribati)	Sachs et al. (2009)	Annual	Sediments

Code	Location	References	Season	Type
65	Kiritimati (Kiribati)	Sachs et al. (2009); Higley et al. (2018)	Annual	Sediments
66	Mecherchar Island (Palau)	Sachs et al. (2009)	Annual	Sediments
67	San Cristóbal Island (Ecuador)	Sachs et al. (2009)	Annual	Sediments
68	Makassar Strait (Indonesia)	Newton et al. (2006); Tierney et al. (2010b)	Annual	Sediments
69	Kau Bay (Indonesia)	Langton et al. (2008)	Annual	Sediments
70	Cariaco Basin (Venezuela)	Haug et al. (2001)	Annual	Sediments
71	Dongge Cave (S China)	Wang et al. (2005)	Annual	Speleothems
72	Lake Victoria (Tanzania, Uganda, Kenya)	Stager et al. (2005)	Annual	Sediments
73	Castor Lake (Washington, USA)	Steinman et al. (2013)	Dec-Feb	Sediments
74	Lime Lake (Washington, USA)	Steinman et al. (2013)	Dec-Feb	Sediments
75	Crevice Lake (Montana, USA)	Stevens and Dean (2008)	Annual	Sediments
76	Bison Lake (Colorado, USA)	Anderson (2011)	Annual	Sediments
77	Pyramid Lake (Nevada, USA)	Benson et al. (2002)	Annual	Sediments
78	Quelccaya Ice Cap (Peru)	Vuille et al. (2012); Thompson et al. (1986)	Dec-Feb	Ice cores
79	Cristal Cave (SE Brazil)	Vuille et al. (2012); Taylor (2010)	Dec-Feb	Speleothems
80	Diva & Torrinha Caves (NE Brazil)	Vuille et al. (2012); Novello et al. (2012)	Dec-Feb	Speleothems
81	S South America	Neukom et al. (2010)	Jun-Aug, Dec-Feb	Multi-proxy
82	Palestina Cave (NW Peru)	Apaéstegui et al. (2014)	Dec-Feb	Speleothems
83	Lake Allos (France)	Wilhelm et al. (2012)	Jul-Oct	Sediments
84	Lake Taravilla (Spain)	Moreno et al. (2008)	Annual	Sediments
85	Lake Estanya (NE Spain)	Morellón et al. (2009, 2011)	Annual	Sediments
86	SC England	Wilson et al. (2012)	Mar-Jul	Tree rings
87	Wanxiang Cave (China)	Zhang et al. (2008)	Annual	Speleothems
88	Cave KNI-51 (N Australia)	Denniston et al. (2016)	Annual	Speleothems
89	Heshang Cave (China)	Hu et al. (2008)	Annual	Speleothems
90	Laguna Pallcacocha (S Ecuador)	Moy et al. (2002)	Annual	Sediments
91	Peruvian Shelf (Peru)	Rein et al. (2004)	Annual	Sediments
92	Dongdao Island (S China Sea)	Yan et al. (2011)	Annual	Sediments
93	Zaca Lake (California, USA)	Kirby et al. (2014)	Annual	Sediments
94	Lake Elsinore (California, USA)	Kirby et al. (2019)	Annual	Sediments
95	Abbott Lake (California, USA)	Hiner et al. (2016)	Annual	Sediments
96	Maddox Lake (California, USA)	Kirby et al. (2023)	Annual	Sediments

Our analyses focus on reconstructions of the Palmer Drought Severity Index (PDSI; Palmer, 1965). Even if other variables
125   like the soil moisture, the precipitation or the difference between precipitation and evaporation may also provide information

Table 3. Reconstruction-based and model products considered in this work. In case of being refered to an ensemble of experiments, the acronym used for the ensemble and the number of members (N) are provided. The starting year, the available seasons and references describing each product are provided in the last ~~column~~columns.

Ensemble	Product/Model	N	Start	Seasons	References
DA	Old World Drought Atlas (OWDA)	1	0	JJA	Cook et al. (2015)
DA	North American Drought Atlas (NADA)	1	0	JJA	Cook et al. (2010b)
DA	Monsoon Asia Drought Atlas (MADA)	1	1300	JJA	Cook et al. (2010a)
LMR	Last Millennium Reanalysis (LMR)	20	850	Annual	Tardif et al. (2019)
PHYDA	Paleo Hydrodynamics Data Assimilation (PHYDA)	100	1	JJA, DJF, Annual	Steiger et al. (2018b)
CESM-LME	Community Earth System Model (CESM)	13	850	Monthly	Otto-Bliesner et al. (2015)
CMIP5	Goddard Institute for Space Studies (GISS)	3	850	Monthly	Schmidt et al. (2006, 2014)
CMIP5	Model for Interdisciplinary Research on Climate (MIROC-ESM)	1	850	Monthly	Watanabe et al. (2011)
CMIP5	Hadley Centre Coupled Model (HadCM3)	1	800	Monthly	Schurer et al. (2014)
CMIP5	Community Climate System Model (CCSM)	1	850	Monthly	Landrum et al. (2013)
CMIP5	Institut Pierre Simon Laplace (IPSL)	1	850	Monthly	Dufresne et al. (2013)
CMIP5	Commonwealth Scientific and Industrial Research Organization (CSIRO)	1	850	Monthly	Phipps et al. (2012)
CMIP5	Meteorological Research Institute (MRI)	1	850	Monthly	Yukimoto et al. (2011); Adachi et al. (2013)
CMIP5	Max Planck Institut fur Meteorologie (MPI)	1	850	Monthly	Giorgetta et al. (2013)
CMIP6	MIROC Earth System version 2 for Long-term simulations (MIROC-ES2L)	1	850	Monthly	Hajima et al. (2020)
CMIP6	Meteorological Research Institute (MRI)	1	850	Monthly	Yukimoto et al. (2019)
CMIP6	Max Planck Institut fur Meteorologie (MPI)	2	1	Monthly	Mauritsen et al. (2019)

about the hydroclimate conditions, the use of a standardised index like the PDSI is generally preferred when combining different products (Cook et al., 2014). The PDSI takes into account the soil moisture conditions from the previous months, estimated based on temperature and precipitation and considering the water storage capacity of the soil. The cumulative behavior of the PDSI makes its variations smoother than those of atmospheric variables, making it a suitable variable for

130 analyses on decennial to centennial scales, while the direct impact of temperature on the PDSI makes it more sensitive to changes in external forcing than variables exclusively based on precipitation (Roldán-Gómez et al., 2020).

The PDSI is directly provided by the DAs, the PHYDA and the LMR. For the CMIP5, CMIP6 and CESM-LME simulations, PDSI has been computed based on atmospheric variables and soil parameters, following the approach of a self-calibrating PDSI (scPDSI; Wells et al., 2004). The potential evapotranspiration is calculated using the Thornthwaite's method (Thornthwaite,
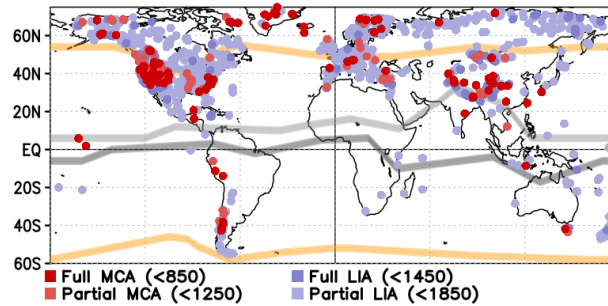
Figure 2. Coverage of the MCA and LIA by the proxies from Steiger et al. (2018a), considered for the generation of PHYDA and LMR. Positions of the modern ITCZ and NAM and SAM boundaries are also shown, as described in Fig. 1.

135 1948). While the Thorntwaite formulation has been shown to have limitations for projections of 21st-century soil moisture conditions because of its strong dependence on temperature, assessments of the Thornthwaite formulation in model simulations of the LM have shown that such limitations are not a concern for the LM period (Smerdon et al., 2015). For the 20th Century, the PDSI from the different sources have been compared to the instrumental observation based PDSI from the Climatic Research Unit (CRU; Barichivich et al., 2021).

140 The analysis of hydroclimate based on proxy data is sensitive to the seasonality of the proxies. As shown in Table 2, most proxies provide annual values, but certain proxies of the Northern Hemisphere are linked to the hydroclimate of the boreal summer (June, July and August; JJA), while certain proxies in the Southern Hemisphere are associated with the boreal winter (December, January, and February; DJF). The selection of the season considered in the different products is then a critical point for the analyses. As shown in Table 3, the PHYDA provides water-year annual averages and seasonal averages for ~~the~~

145 ~~boreal summer (June, July and August; JJA) and winter (December, January, and February; DJF)~~JJA and DJF. The LMR only provides annual averages, while the Northern Hemisphere DAs exclusively target the JJA period. For all the analyses, JJA values are considered for all the products except for the LMR, for which the annual values are used. To confirm that the annual results from the LMR can be meaningfully compared to the seasonal results from the other products, the annual PDSI from the PHYDA and the ensembles of CMIP5, CMIP6 and CESM-LME are analysed in Appendix B.

150 Coordinated changes during the MCA and the LIA are characterized by computing the difference between each variable and product during the LIA and MCA periods. The agreement between products is also assessed by computing the global spatial correlation between PDSI LIA-MCA maps from each product (Coats et al., 2013a), and by counting the number of products showing positive or negative PDSI differences for each location. Finally, time series of PDSI for the whole millennium are extracted from different locations and different products, to analyse the correlations between them beyond the periods of the

155 MCA and LIA.
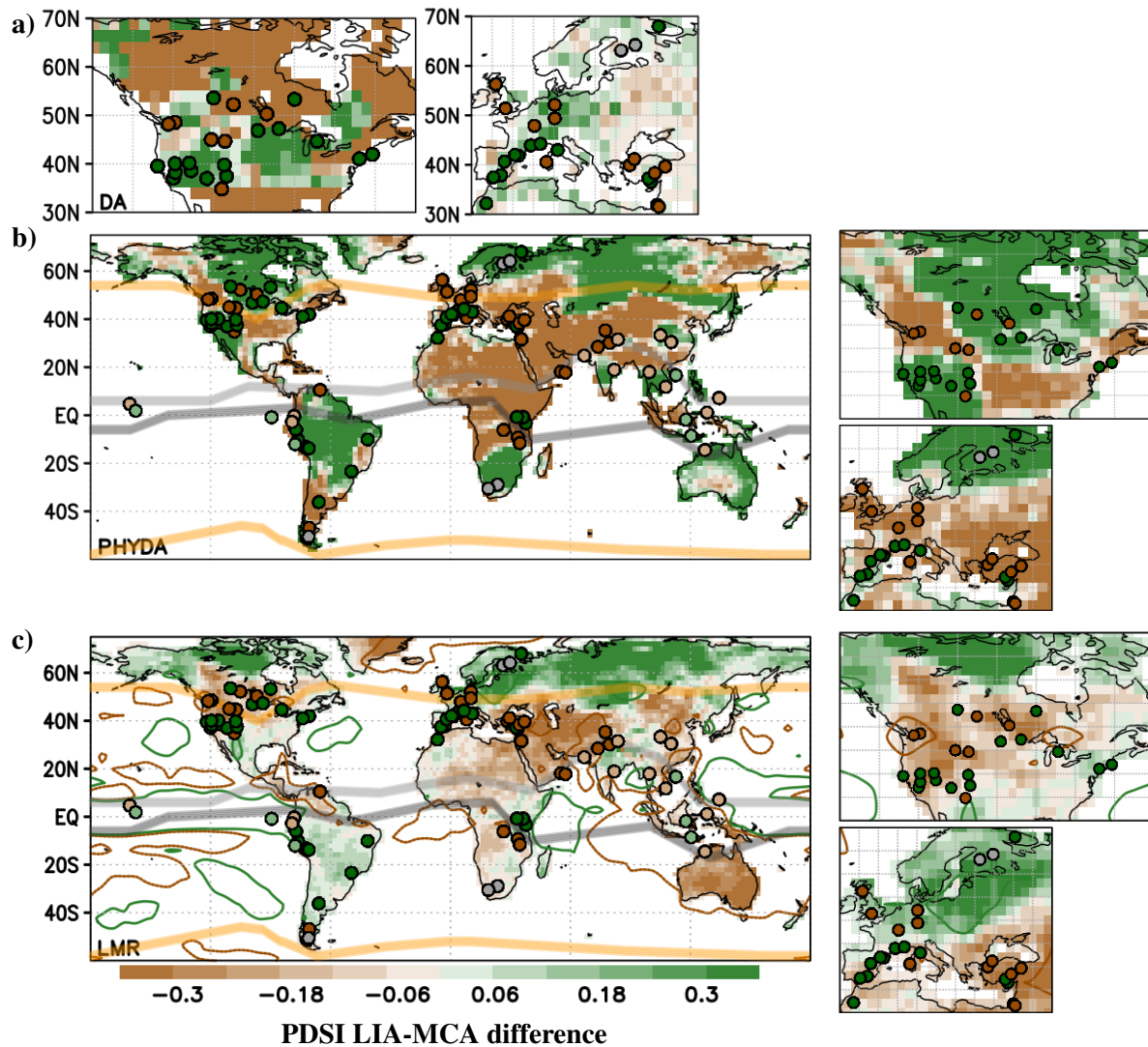
Figure 3. LIA-MCA differences in PDSI (shading) or precipitation (contours) from **(a)** DAs, **(b)** PHYDA, and **(c)** LMR. Dots represent the locations of Fig. 1, with the same color code considered in that figure; positions of the modern ITCZ and NAM and SAM boundaries are also shown, as described in Fig. 1. For the precipitation, contours of -1 (brown) and 1 (green) mm/month are shown. Average of JJA is considered for the PHYDA and the DAs, while the annual average is considered for the LMR. For the PHYDA and the LMR, regional maps over North America and Europe are shown on the right side of each corresponding panel.
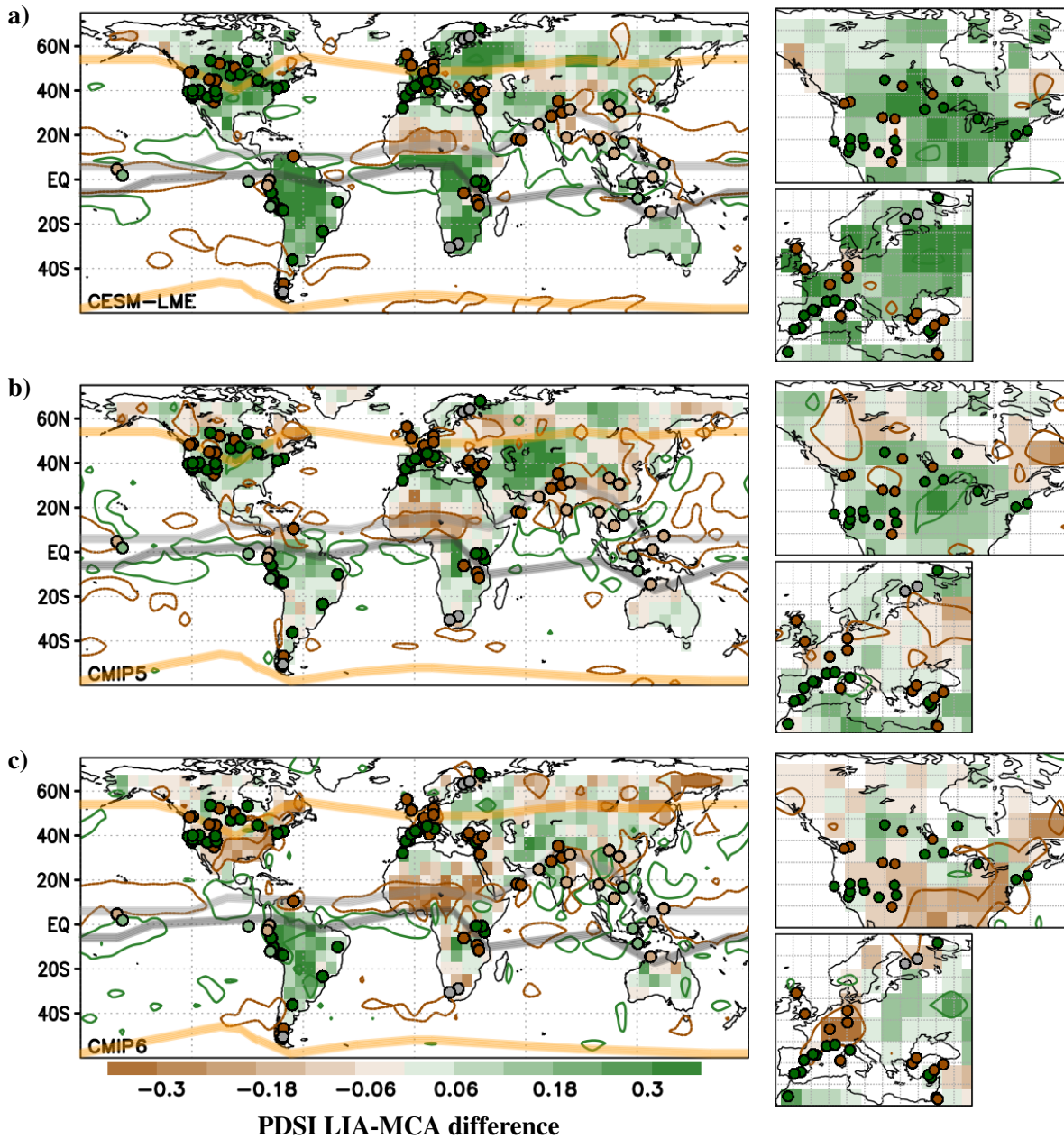
Figure 4. LIA-MCA differences in PDSI (shading) or precipitation (contours) from the ensemble average of **(a)** CESM-LME, **(b)** CMIP5 and **(c)** CMIP6 LM simulations. Dots represent the locations of Fig. 1, with the same color code considered in that figure; positions of the modern ITCZ and NAM and SAM boundaries are also shown, as described in Fig. 1. For the precipitation, contours of -1 (brown) and 1 (green) mm/month are shown. In all the cases, average of JJA is considered. For each product, regional maps over North America and Europe are shown on the right side of each corresponding panel.

4 Results

4.1 Reconstructions of hydroclimate during the MCA and LIA

PDSI differences between the LIA and MCA are shown for the DAs, the PHYDA and the LMR (Fig. 3). Some wetting and drying patterns are consistent across the DAs (Fig. 3a), the PHYDA (Fig. 3b), and the LMR (Fig. 3c). For example, all products
160 estimate drier conditions during the LIA and wetter conditions during the MCA over the Middle East. ~~In the PHYDA and LMR these conditions are associated with a warmer LIA in these regions (Fig. 3).~~ There is also good agreement across the PHYDA and LMR in South America, with drier conditions in the north and wetter conditions in the central continent during the LIA, as well as in northern Canada, Scandinavia and northern Eurasia, with wetter conditions during the LIA~~, associated all of them with lower temperatures (Fig. 3)~~. The PHYDA and the LMR agree well with the reconstructions of Fig. 1, which also
165 show drier conditions during the LIA in southeastern Europe and Anatolia (Göktürk, 2011; Jones et al., 2006; Woodbridge and Roberts, 2011; Kuzucuoglu et al., 2011), the Indian Monsoon region (Burns et al., 2002; Bryson and Swain, 1981; Phadtare and Pant, 2006; Anderson et al., 2002; Gupta et al., 2003) and South America north of the ITCZ (Haug et al., 2001). The LMR also agree with the reconstructions of Fig. 1 in regions of East Africa (Anchukaitis and Tierney, 2013).

    In western North America there is a good agreement between the DA and PHYDA, and between these two products and
170 the reconstructions of Fig. 1, most of them showing wetter conditions in the southwest and drier conditions in the northwest during the LIA (Steinman et al., 2013; Cook et al., 2004, 2010b). Each of these products are nevertheless in contrast to the LMR, which shows drier conditions during the LIA for most of the American West, in sharp contrast with a broad range of literature on megadroughts during the MCA period (Cook et al., 2010b; Coats et al., 2013b). For the case of central and western Europe, the PHYDA estimates drier conditions during the LIA, in agreement with the reconstructions of Fig. 1 (~~Luterbacher~~
175 ~~et al., 2012~~), while the DA and LMR estimate wetter conditions. The same happens for areas of central and eastern North America, especially for those areas with a small number of proxies covering the MCA (Fig. 2). Differences between products are particularly relevant in the Indo-Pacific basin, where the PHYDA and LMR show opposite behavior in Australia, Southeast Asia and Patagonia, and in high latitudes, with opposite behavior in areas of Alaska and northeast Asia. ~~These differences coincide with opposite estimates in temperature differences as well, with the PHYDA (Fig. 3a) and LMR (Fig. 3b) respectively~~
180 ~~showing warmer or cooler conditions in Patagonia and cooler or warmer conditions in Australia during the LIA.~~

4.2 Simulated hydroclimate during the MCA and LIA

The analysis of the model ensembles allows an estimate of those areas that are potentially more impacted by external forcing as estimated by the models, because individual simulations include different realisations of internal variability but share the external forcing changes imposed as boundary conditions (Fernández-Donado et al., 2013; Jungclaus et al., 2017).
185   Figure 4 shows the mean PDSI differences between the LIA and MCA for the CESM-LME, CMIP5 and CMIP6 LM ensembles. For the areas of the Middle East and southern Asia the agreement between the DAs, the PHYDA, the LMR and the reconstructions of Fig. 1 extends also to the simulations from CESM-LME (Fig. 4a), CMIP5 (Fig. 4b) and CMIP6 (Fig. 4c). There is also a good agreement in central South America, where the CESM-LME (Fig. 4a), CMIP5 (Fig. 4b) and CMIP6 (Fig.

4c) simulations all show a wet LIA for the areas south of the ITCZ~~, also linked to cooler temperatures during this period (Fig.~~
190 ~~5)~~. For northern South America, the dry LIA observed in reconstructions (Fig. 3b,c) is also found, with a limited extension, in CMIP5 (Fig. 4b) and CMIP6 (Fig. 4c) simulations. The agreement found for these areas is consistent with an impact of external forcing on the position of the Atlantic ITCZ (Roldán-Gómez et al., 2022) and the intensity of the monsoon system (Roldán-Gómez et al., 2020).

~~In western North America and Western Europe, the CMIP5 simulations (Fig. 4b) show a good agreement with~~The CMIP5
195 simulations (Fig. 4b) show a wetter LIA in most areas of western North America and Western Europe, which is also found with a more limited extension in the reconstructions of Fig. 1 and ~~with~~ the PHYDA (Fig. 3b). However, this consistency is not so clear in the simulations from CMIP6 (Fig. 4c), for which a dry LIA is found in most regions of North America~~, associated with limited cooling during that period (Fig. 5c)~~. Disagreement also exists in the CESM-LME (Fig. 4a), which simulates a dry LIA in some areas of southwestern North America and a wet LIA in Western and Northern Europe. ~~These differences can be~~
200 ~~linked to different atmospheric dynamics, with the CMIP5 ensemble showing important differences in sea level pressure (SLP) between the MCA and the LIA in most regions of North America and Europe, while the CMIP6 ensemble and the CESM-LME limit these changes to some areas of Southern and Eastern Europe and northern North America. Despite these differences, all the models show relevant changes in these regions during the MCA and LIA, in line with the changes in the North Atlantic Oscillation (NAO) shown in Roldan-Gomez et al. (2020).~~As shown in Appendix C, this disagreement can be explained by a
205 different behavior in terms of temperatures and atmospheric dynamics.

In the Indo-Pacific basin, the ensemble average of CESM-LME (Fig. 4a), CMIP5 (Fig. 4b) and CMIP6 (Fig. 4c) LM simulations show small PDSI differences between the MCA and LIA, with no clear agreement with the PHYDA (Fig. 3b) and the LMR (Fig. 3c). ~~The CESM-LME (Fig. 5a) tends to show cooler conditions during the LIA in these regions, but the temperature difference is small compared to that of North America and Europe. For the case of CMIP5 (Fig. 5b) and CMIP6 (Fig. 5c),~~
210 ~~warmer conditions are found during the LIA for some areas of India and northern Australia. Regarding the atmospheric dynamics, only the LMR shows relevant changes in SLP for the Indo-Pacific, with positive anomalies in the Indian and western Pacific basin and negative anomalies in the eastern Pacific basin during the LIA~~ This behavior is consistent with a larger impact of internal variability in these areas (Roldán-Gómez et al., 2022), with simulation-dependent changes that are filtered out when working with ensemble averages.

215 **4.3 Agreement between products**

To assess the agreement between different pairs of products, spatial pattern correlations for the global maps of Fig. 3 and Fig. 4 have been included in Table 4. Despite the regional differences found in areas of North America and the Indo-Pacific basin, the ensembles of CESM-LME, CMIP5 and CMIP6 show global spatial correlations between 0.29 and 0.39. The global correlation of reconstruction-based products is strongly impacted by low correlations in some areas of the southern hemisphere
220 and at high latitudes, with a lower density of proxy data. If these regions are excluded (Table 5), correlations larger than 0.29 are also found between DAs and PHYDA and between PHYDA and LMR. The disagreement between LMR and the other reconstruction-based products in areas of Western Europe and western North America explains the small global correlation
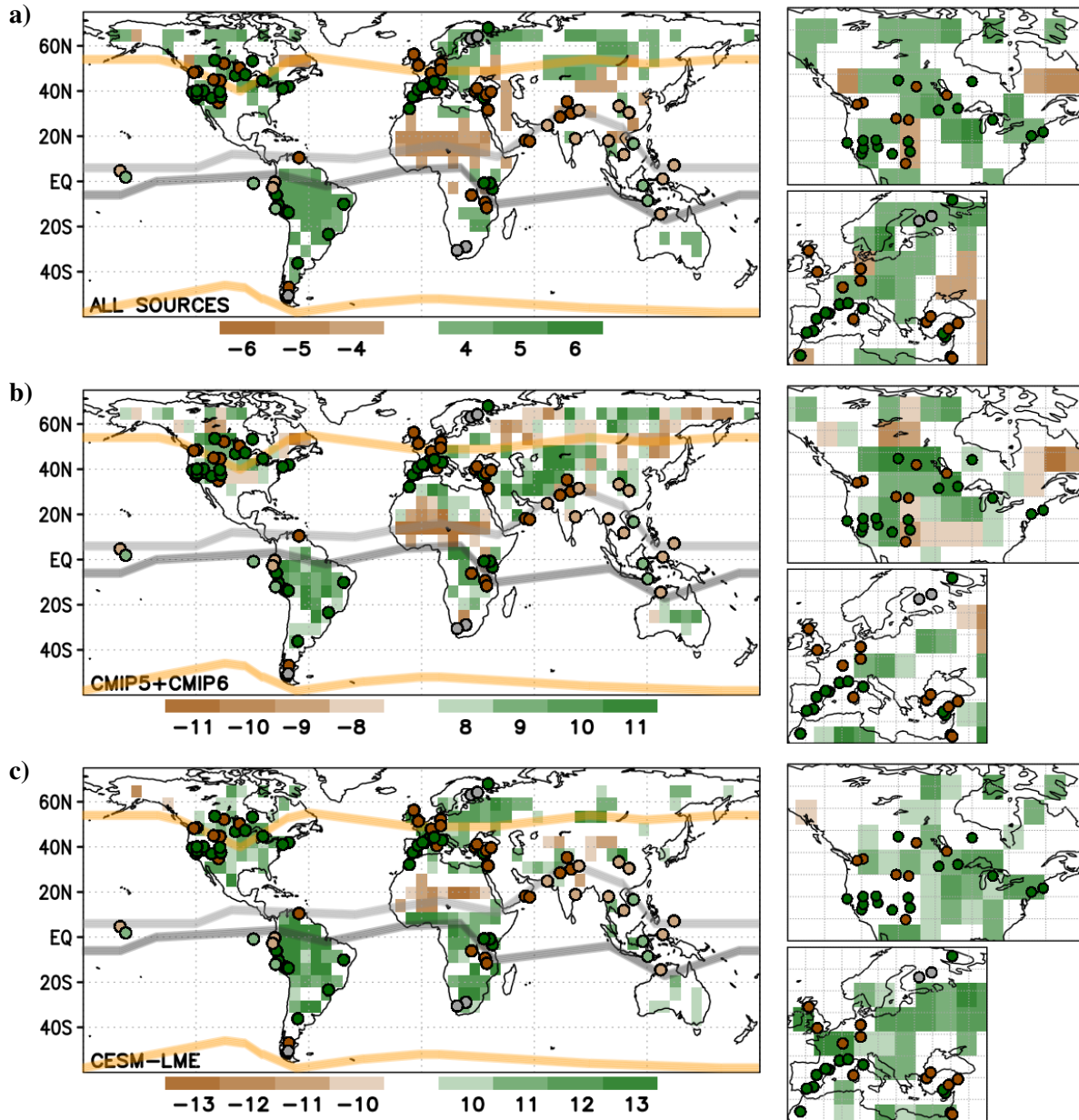
Figure 5. (a) Number of products, from the 6 analysed products (DAs, PHYDA, LMR, and the ensemble averages of CESM-LME, CMIP5 and CMIP6), showing positive (green) or negative (brown) PDSI differences LIA-MCA. **(b)** Number of simulations, from the 14 simulations in the ensembles of CMIP5 and CMIP6, showing positive (green) or negative (brown) PDSI differences LIA-MCA. **(c)** Number of simulations, from the 13 simulations of the CESM-LME, showing positive (green) or negative (brown) PDSI differences LIA-MCA. Dots represent the locations of Fig. 1, with the same color code considered in that figure; positions of the modern ITCZ and NAM and SAM boundaries are also shown, as described in Fig. 1. Regional maps over North America and Europe are shown on the right side of each corresponding panel.

Table 4. Correlation between the maps of differences LIA-MCA of PDSI from DAs (Fig. 3a), PHYDA (Fig. 3b), LMR (Fig. 3c), CESM-LME (Fig. 4a), CMIP5 (Fig. 4b) and CMIP6 (Fig. 4c). Significant correlations (p < 0.05) accounting for autocorrelation are shown in bold.

	PHYDA	LMR	CESM	CMIP5	CMIP6
DAs	-0.05	-0.06	-0.11	0.08	-0.08
PHYDA		**0.32**	**0.24**	0.10	**0.18**
LMR			0.05	-0.08	0.05
CESM				**0.33**	**0.29**
CMIP5					**0.39**

Table 5. Same as Table 4, but considering only the data between 20ºS and 50ºN.

	PHYDA	LMR	CESM	CMIP5	CMIP6
DAs	**0.35**	-0.08	-0.04	**0.30**	0.08
PHYDA		**0.29**	**0.35**	0.11	**0.27**
LMR			**0.24**	-0.10	0.16
CESM				**0.36**	**0.32**
CMIP5					**0.45**

between the DAs and LMR. It is important to note that the spatial correlation between reconstruction-based and model-based products is impacted by internal variability. The use of ensemble averages removes the contribution of internal variability as it was manifest in the climate trajectory from the simulated data. For the reconstruction-based products the contribution of internal variability is not removed, even if the multi-century averages for the LIA and MCA periods emphasize the contribution of the external forcing.

The agreement between sources is particularly poor in areas of the Indo-Pacific basin. As shown in Fig. 5a, there is no clear agreement regarding the sign of the PDSI LIA-MCA differences for areas of Southeast Asia and Australia. There is also no clear agreement for these areas among the individual simulations comprising the CMIP5 and CMIP6 ensembles (Fig. 5b), suggesting that the hydroclimate of these areas is strongly impacted by simulation-dependent processes, likely those associated with internal variability. The link to the internal variability is clear when analysing the simulations of the CESM-LME (Fig. 5c), which share the same forcing conditions and differ only in the internal variability. These results are in line with the results obtained with the CESM-LME by Roldán-Gómez et al. (2022), showing a relevant impact of internal variability on the position and extension of the Indian and western Pacific ITCZ.

In other areas, most products (Fig. 5a) and most individual simulations (Fig. 5b,c) show the same sign of PDSI differences such as in central South America, with a drier MCA and a wetter LIA, or the Sahel, extending from Western Africa to the

Arabian peninsula, for which most products and simulations show a wetter MCA and a drier LIA. Consistency also extends to areas of western North America, with at least 4 products showing wetter conditions in the south and drier conditions in the

240   north during the LIA, or to northern India, where most products show a wetter MCA and a drier LIA. The agreement for these areas suggest a contribution of external forcing that is more important than that of internal variability.

The MCA and the LIA are periods extensively addressed in the literature, but they are not the only periods showing relevant droughts (Cook et al., 2022). The analyses have been then extended to other periods, by extracting time series of PDSI for the whole millennium from the same regions that show coordinated changes during the MCA and LIA (Fig. 6a-e and 7a-e),

245   including western North America, southwestern Europe, northern and central South America, East Africa, Pakistan, India, Southeast Asia and Indonesia. For each region, the correlations between the DAs, the PHYDA, the LMR, and the ensemble average of CESM-LME, CMIP5 and CMIP6 LM simulations have been computed (Fig. 6f-j and 7f-j).

The correlation between the DAs and PHYDA is significant for western North America, while the correlation between the PHYDA and LMR is significant for central South America, northwestern North America, southwestern Europe, Pakistan,

250   Southeast Asia and Indonesia. Some of these correlations exceed 0.7 and are clearly visible in the time series of PDSI from these regions. This is the case for the correlation between the DA and the PHYDA in southwestern North America (Fig. 6a,f), and the correlation between PHYDA and LMR in Pakistan (Fig. 6d,i), Indonesia (Fig. 7a,f) and central South America (Fig. 7e,j). Regions showing high correlations between reconstruction-based products are in general those regions showing the largest hydroclimate changes during the LM, like those around the ITCZ (Yan et al., 2015; Anchukaitis and Tierney, 2013),

255   those in the Monsoon region (Fleitmann et al., 2003) and in the area of influence of the NAM (Ortega et al., 2015).

Significant correlations are also found between the ensemble averages of CESM-LME, CMIP5 and CMIP6 for regions of South America, India and Southeast Asia, between CMIP5 and CMIP6 ensembles for areas of southeastern Europe, East Africa and Pakistan, and between CESM-LME and CMIP5 ensembles in areas of southwestern Europe and Indonesia. As for the case of reconstruction-based products, some of these correlations exceed 0.7 and are visible in the time series of PDSI,

260   like those of northern South America (Fig. 7d,i) and Southeast Asia (Fig. 7c,h) for CESM-LME, CMIP5 and CMIP6, and those of East Africa (Fig. 6e,j) and India (Fig. 7b,g) for CMIP5 and CMIP6. In these cases, similarities are found during the whole millennium, including a trend in the 20th century not found in the reconstructions (Ljungqvist et al., 2016). If only the pre-industrial period is considered, the correlations are smaller but still significant for most of these areas. The correlation between simulated products is particularly high for those areas with a larger impact of external forcing, like the Atlantic ITCZ

265   (Roldán-Gómez et al., 2020), consistent with the fact that the use of the ensemble average emphasizes the forcing signal (Fernández-Donado et al., 2013; Roldán-Gómez et al., 2020).

Considering this, the comparison between time series from reconstruction-based products and from model simulations would be only relevant for those areas where the contribution of the forcing dominates. Significant correlations are found between the ensemble average of CESM-LME, CMIP5 and CMIP6 LM simulations and the PHYDA and LMR in northern and central South

270   America and, to a lower extent, in regions of southwestern Europe, northern North America and Southeast Asia, suggesting a relevant contribution of the forcing to the hydroclimate of these areas. Some of these correlations between reconstruction-based products and model simulations exceed 0.5 and are visible in the time series of PDSI, like those in central South America (Fig.
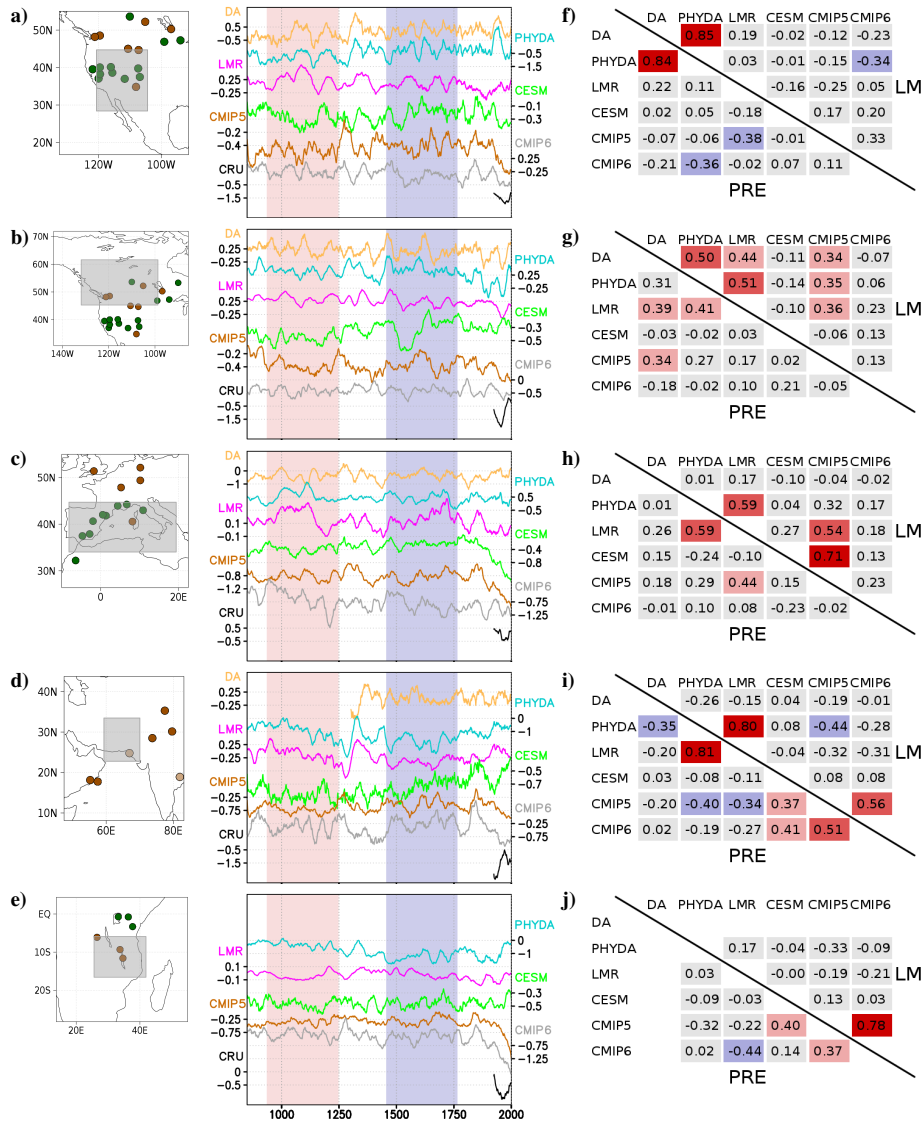
Figure 6. (a-e) Time series of PDSI for **(a)** southwestern North America, **(b)** northwestern North America, **(c)** southwestern Europe, **(d)** Pakistan and **(e)** East Africa, obtained from DAs, PHYDA, LMR, the ensemble average of CESM-LME, CMIP5, and CMIP6 LM simulations. Time series from the CRU PDSI data set for the 20th Century have been also included for comparison. Time series of DAs are not available for those regions not considered in OWDA, NADA, MADA and MXDA. Periods of MCA and LIA are highlighted in red and blue in the middle panels. Dots in the map of the left side represent the locations of Fig. 1, with the same color code considered in that figure. **(f-j)** Correlations between time series of PDSI for each region obtained from DAs, PHYDA, LMR, and the ensemble average of CESM-LME, CMIP5, and CMIP6 LM simulations, both for the full millennium (LM) and the pre-industrial period (PRE). Color code for the correlations is the same as included in Fig. 8.
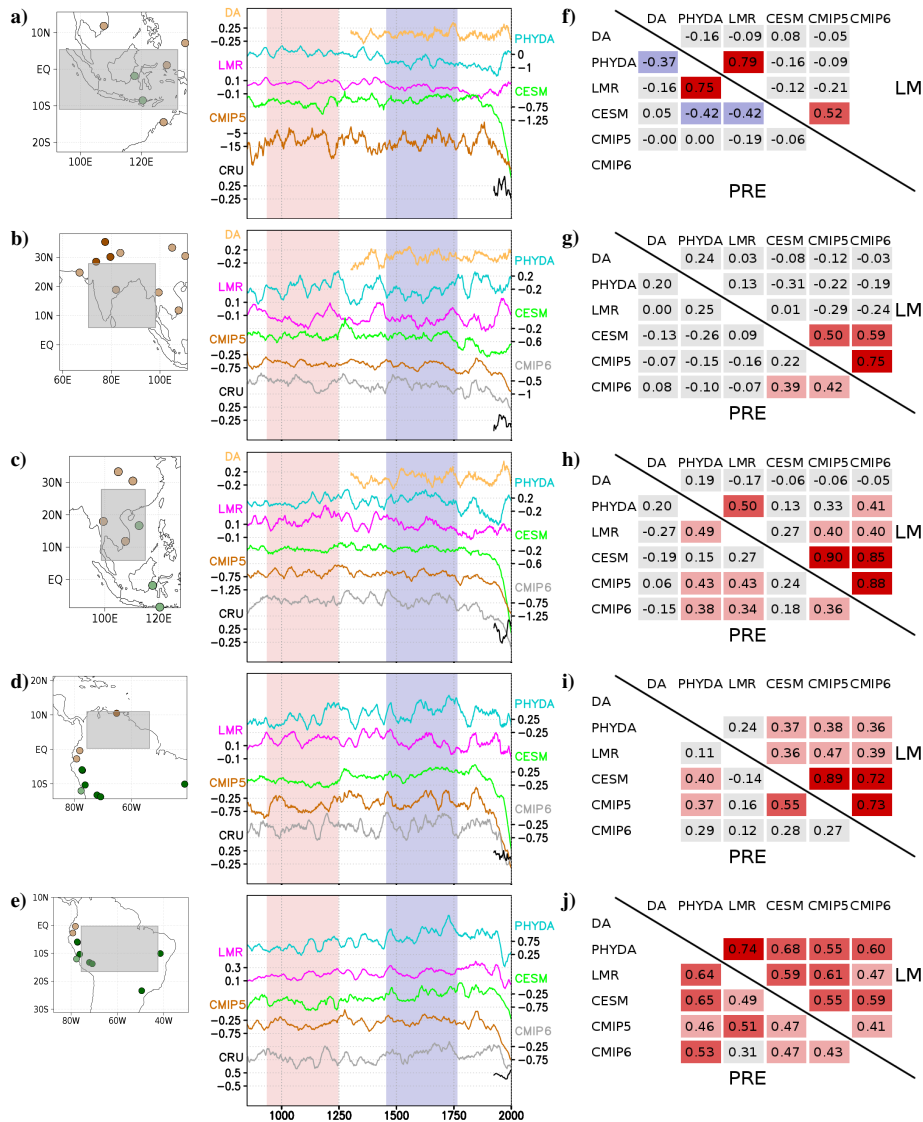
Figure 7. Same as Fig. 6 for **(a,f)** Indonesia, **(b,g)** India, **(c,h)** Southeast Asia, **(d,i)** northern South America and **(e,j)** central South America. Time series of DAs are not available for those regions not considered in OWDA, NADA, MADA and MXDA. Time series of CMIP6 are not available for Indonesia, since most CMIP6 simulations do not provide the necessary variables to compute the PDSI for this region.

7e,j) and those between the LMR and the ensemble average of CMIP5 in southwestern Europe (Fig. 6c,h). These results can be then linked to a particularly important contribution of the forcing in these regions, present in both proxy data and model simulations.
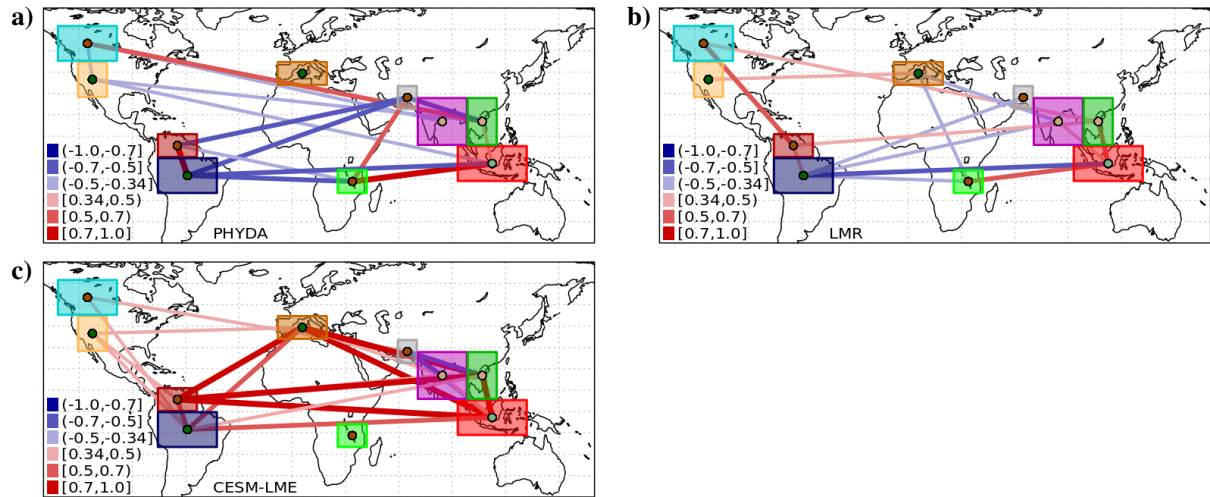
Figure 8. Correlations between time series of PDSI for different regions (solid lines), obtained from **(a)** PHYDA, **(b)** LMR and , **(c)** CESM-LME. Only significant correlations (p < 0.05; |r| > 0.34) accounting for autocorrelation are shown, following the color code in the bottom left side of the pannels. The color of the points for each region indicate whether the LIA was wetter (green) or drier (brown) than the MCA in the reconstructions of Fig. 1.

4.4 Correlation between distant regions

As described in the previous sections, distant regions showed similar hydroclimate behaviors during the MCA and LIA. These similarities suggest the presence of coordinated changes between distant regions that go beyond these two periods. To analyse these changes, correlations between the time series of different regions from Fig. 6 and 7 have been computed for the
280 PHYDA (Fig. 8a) and the LMR (Fig. 8b). Both reconstruction-based products show significant positive and negative correlations between distant regions, some of them in agreement with the behavior of reconstructions during the MCA and LIA (Fig. 1).

In particular, both the PHYDA and the LMR show significant negative correlations between central South America and the regions of East Africa and Pakistan. These correlations can be linked to coordinated changes of the Atlantic ITCZ and
285 the monsoon system, in line with those observed in model simulations and associated with shifts of the Atlantic ITCZ in response to changes in the external forcing (Roldán-Gómez et al., 2020, 2022). Both reconstruction-based products also show significant positive correlations across the Indian and Pacific basins, such as those between Indonesia and East Africa, and between Southeast Asia and northwestern North America, consistent with the changes in the Pacific Walker circulation found in simulations (Yan et al., 2015).

290 The correlations across the Pacific basin are emphasized in the PHYDA, which also shows significant negative correlations between southwestern North America and the regions in India and Indonesia. In contrast, the LMR tends to emphasize the variability in the Atlantic basin, with positive correlations between southwestern North America and southwestern Europe

following the pattern of the ~~NAO~~North Atlantic Oscillation (NAO). The spatial correlations found in the PHYDA and LMR are mainly driven by the model used for the data assimilation. However, if the correlations are computed for the CESM-LME (Fig. 8c) without any data assimilation, only positive correlations associated to the impact of temperature in the PDSI are obtained, and the fact that the PHYDA and LMR respectively ephasize the variability in the Pacific and Atlantic basin shows that the regional variability given by the proxies is consistent with the spatial patterns from the models. This increases the confidence on the ability of the models to reproduce the behavior of the hydroclimate at a global scale, in particular for those areas and those mechanisms more impacted by the external forcing.

5 Conclusions

The analysis of different reconstruction-based and model-based products shows consistent coordinated changes during the LM in the hydroclimate of distant regions. The areas more impacted by these changes are those around the ITCZ, like northern and central South America, East Africa and the Indo-Pacific, those in the area of influence of the NAM/NAO, like western North America and Western Europe, and those of the Indian Monsoon region, extending from the Middle East to Southeast Asia. Even if these changes are particularly important during the MCA and the LIA, for which a high number of reconstructions show drier or wetter conditions, significant correlations between distant regions are also found when considering the whole millennium.

The agreement between reconstruction-based and model-based products for areas of South America and East Africa is indicative of a relevant contribution of the forcing in these regions, consistent with the shifts of the Atlantic ITCZ obtained in simulations in response to the forcing (Roldán-Gómez et al., 2022). Agreement is also high in the Indian Monsoon region, consistent with an alteration of the monsoon system as a consequence of changes in the forcing (Roldán-Gómez et al., 2020). A certain agreement is also found for areas of western North America and Western Europe, consistent with an alteration of the NAM/NAO in response to changes in the forcing (Roldán-Gómez et al., 2020), even if the different products show discrepancies at regional scales. These discrepancies may be linked to the contribution of internal variability in these areas, but also to uncertainties in the reconstructed data, limitations in the physics of the models and inaccuracies in the assessment of external forcing considered for the simulations. The largest discrepancies between reconstructed and simulated products are found in the Indo-Pacific basin, suggesting that this area is more impacted by internal variability processes than by changes in the forcing.

The combination of reconstructions, model simulations and hybrid products is a powerful technique to overcome the limitations of each individual source. On one hand, the sparse hydroclimate records available from proxy data can be completed with gridded information of atmospheric dynamics and hydroclimate variables provided by model simulations and hybrid products to extend the analysis to areas not fully covered by the proxies. On the other hand, the representativity of the mechanisms found in simulations, often involving temperatures, atmospheric dynamics and hydroclimate, like the shifts of the ITCZ, the alteration of variability modes like the NAM/NAO and the changes in the Walker circulation during the MCA and LIA, could be confirmed with the hydroclimate information obtained from proxy data and hybrid products.

In this work, we considered an exhaustive and up-to-date compilation of all the available reconstruction-based and model-based sources to analyse the hydroclimate of the LM at a global scale, with a particular focus on the MCA and LIA as periods with particular forcing conditions. This exercise, perhaps more typical in the analysis of temperatures but systematically applied for the first time to the global hydroclimate, provided a novel assessment of global hydroclimate changes. The focus on the

330 MCA and LIA periods and the comparison with model simulations allowed the isolation of those mechanisms potentially impacted by the forcing, while the novel approach based on combining reconstruction-based and model-based sources showed a coherence in the large-scale changes of the hydroclimate of the LM.

Appendix A: Impact of the definition of the MCA and LIA

For the analyses described in the previous sections, the periods of the MCA and LIA have been respectively defined from 950

335 to 1250 CE and from 1450 to 1850 CE. This definition may depend on the product and is not applicable at a global scale (Neukom et al., 2014, 2019). For example, reconstructions of the Indo-Pacific basin (Fig. 1, light brown and light green) tend to show stable conditions only during the early LIA, until late 1500s CE, rather than during the whole LIA as found in regions of Europe and North America (Fig. 1, dark brown and dark green). To make the analyses independent from the particular selection of periods for the MCA and LIA, the differences LIA-MCA of Fig. 3 and Fig. 4 have been replaced in Fig. A1 and

340 Fig. A2 by the slope of PDSI computed over the the whole period from 950 to 1750 CE. The results obtained with this method show no major differences with respect to those shown in Fig. 3 and Fig. 4, confirming that the conclusions obtained in the paper are independent from the particular definition of the MCA and LIA.

Appendix B: Seasonal changes in the PDSI

In the results presented in the previous sections, the JJA PDSI values from the DAs, the PHYDA and the simulations from

345 CESM-LME, CMIP5 and CMIP6 are compared with the annual PDSI provided by the LMR. To confirm that the use of different averaging periods does not have a significant impact on the results, the annual PDSI from the PHYDA and the different ensembles of simulations have been also analysed. The difference between the LIA and MCA for the annual PDSI of the PHYDA and the ensembles of CESM-LME, CMIP5 and CMIP6 is respectively shown in Fig. B1a, Fig. B1b, Fig. B2a and Fig. B2b. When comparing these results with those obtained with the average of JJA (Fig. 3b and Fig. 4) only minor differences

350 are observed, mostly limited to regional scales. This shows that despite the relevant seasonal changes found in the atmospheric variables, the cumulative behavior of the PDSI makes it less sensitive to the season.

Appendix C: Link to temperature and atmospheric dynamics

For many regions, the agreement in the PDSI differences between the MCA and LIA from different products could be linked to an agreement in temperatures. This is the case for central South America, where all the products agree to show a wet LIA (Fig.

355 3 and Fig. 4), linked to cooler temperatures during this period (Fig. C1 and Fig. C2), or the Middle East, where the PHYDA and
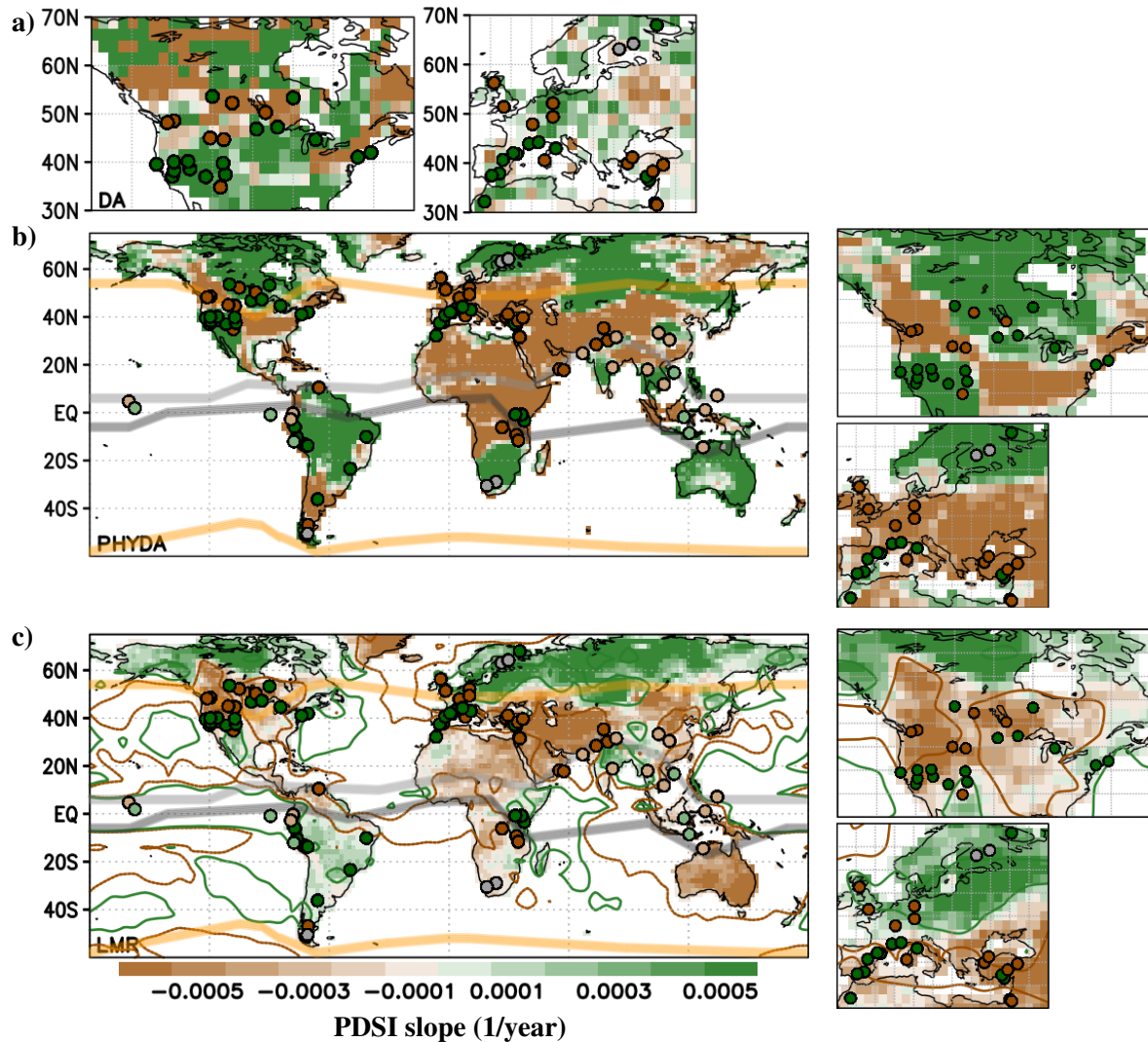
Figure A1. Slope of PDSI (shading) or precipitation (contours) from **(a)** DAs, **(b)** PHYDA, and **(c)** LMR, obtained with a linear regression for the period 950-1750 CE. Dots represent the locations of Fig. 1, with the same color code considered in that figure; positions of the modern ITCZ and NAM and SAM boundaries are also shown, as described in Fig. 1. For the slope of precipitation, contours of -0.001 (brown) and 0.001 (green) mm/month/year are shown. Average of JJA is considered for the PHYDA and the DAs, while the annual average is considered for the LMR. For the PHYDA and the LMR, regional maps over North America and Europe are shown on the right side of each corresponding panel.

LMR show a drier LIA and a wetter MCA (Fig. 3) consistent with a cooler MCA and a warmer LIA shown by these products (Fig. C1).
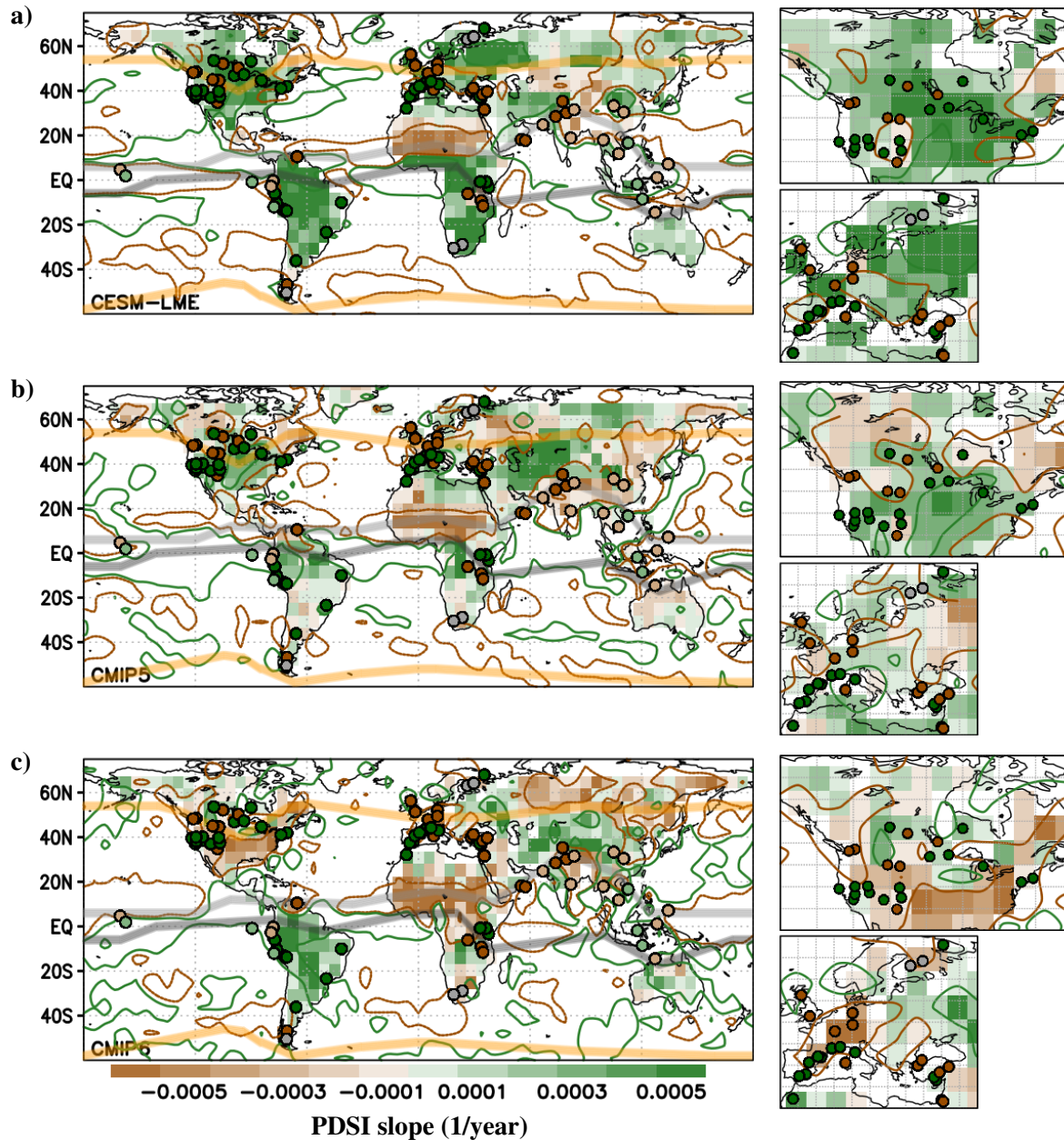
Figure A2. Slope of PDSI (shading) or precipitation (contours) from the ensemble average of **(a)** CESM-LME, **(b)** CMIP5 and **(c)** CMIP6 LM simulations, obtained with a linear regression for the period 950-1750 CE. Dots represent the locations of Fig. 1, with the same color code considered in that figure; positions of the modern ITCZ and NAM and SAM boundaries are also shown, as described in Fig. 1. For the slope of precipitation, contours of -0.001 (brown) and 0.001 (green) mm/month/year are shown. In all the cases, average of JJA is considered. For each product, regional maps over North America and Europe are shown on the right side of each corresponding panel.
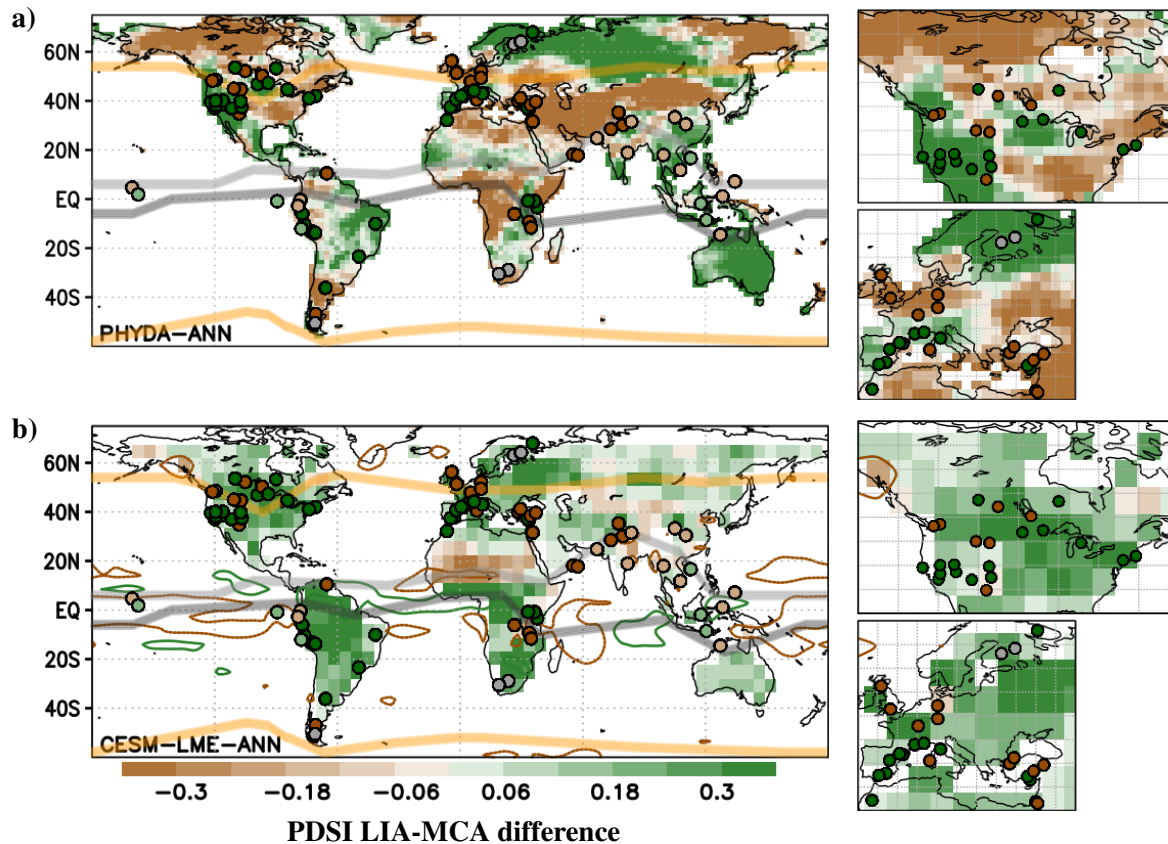
Figure B1. LIA-MCA differences in PDSI (shading) or precipitation (contours) from **(a)** PHYDA and **(b)** the ensemble average of CESM-LME, considering annual averages. Dots represent the locations of Fig. 1, with the same color code considered in that figure; positions of the modern ITCZ and NAM and SAM boundaries are also shown, as described in Fig. 1. For the precipitation, contours of -1 (brown) and 1 (green) mm/month are shown. For each case, regional maps over North America and Europe are shown on the right side of each corresponding panel.

The analysis of temperatures can also explain some of the discrepancies found in the analysis of PDSI. For example, the different behavior obtained from PHYDA and LMR for areas of Australia and Patagonia (Fig. 3) can be explained by an opposite behavior in terms of temperature, with the PHYDA (Fig. C1a) and LMR (Fig. C1b) respectively showing warmer or cooler conditions in Patagonia and cooler or warmer conditions in Australia during the LIA. The drier LIA shown by CMIP6 simulations for regions of North America (Fig. 4c) could be also associated with a limited cooling during that period in the CMIP6 ensemble (Fig. C2c) compared to that shown in CMIP5 (Fig. C2b) and CESM-LME (Fig. C2a), while the wetter LIA shown by CESM-LME simulations in Western and Northern Europe (Fig. 4a) could be associated with the fact that CESM-LME shows cooler conditions during the LIA for these regions (Fig. C2a) compared to those simulated in CMIP5 (Fig. C2b) and CMIP6 (Fig. C2c) ensembles.
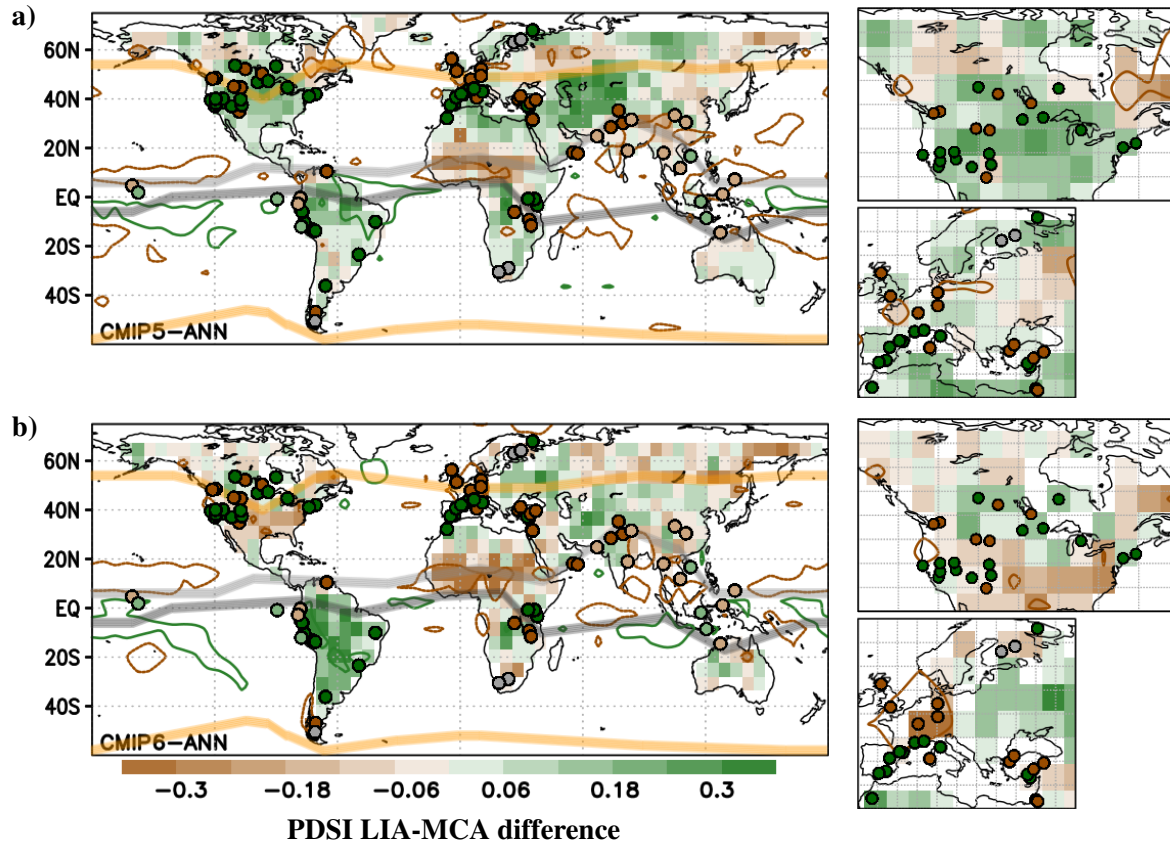
Figure B2. LIA-MCA differences in PDSI (shading) or precipitation (contours) from the ensemble average of **(a)** CMIP5 and **(b)** CMIP6 LM simulations, considering annual averages. Dots represent the locations of Fig. 1, with the same color code considered in that figure; positions of the modern ITCZ and NAM and SAM boundaries are also shown, as described in Fig. 1. For the precipitation, contours of -1 (brown) and 1 (green) mm/month are shown. For each case, regional maps over North America and Europe are shown on the right side of each corresponding panel.

These differences can be linked to differences in atmospheric dynamics, with the CMIP5 ensemble showing important differences in sea level pressure (SLP) between the MCA and the LIA in most regions of North America and Europe (Fig. C2b), while the CMIP6 ensemble (Fig. C2c) and the CESM-LME (Fig. C2a) limit these changes to some areas of Southern
370 and Eastern Europe and northern North America. Despite these differences, all the models show relevant changes in these regions during the MCA and LIA, in line with the changes in the NAO shown in Roldán-Gómez et al. (2020).

For the Indo-Pacific basin, the CESM-LME (Fig. C2a) tends to show cooler conditions during the LIA for most regions, but the temperature difference is small compared to that of North America and Europe. For the case of CMIP5 (Fig. C2b) and CMIP6 (Fig. C2c), warmer conditions are found during the LIA for some areas of India and northern Australia. Regarding
375 the atmospheric dynamics, only the LMR shows relevant changes in SLP for the Indo-Pacific, with positive anomalies in the
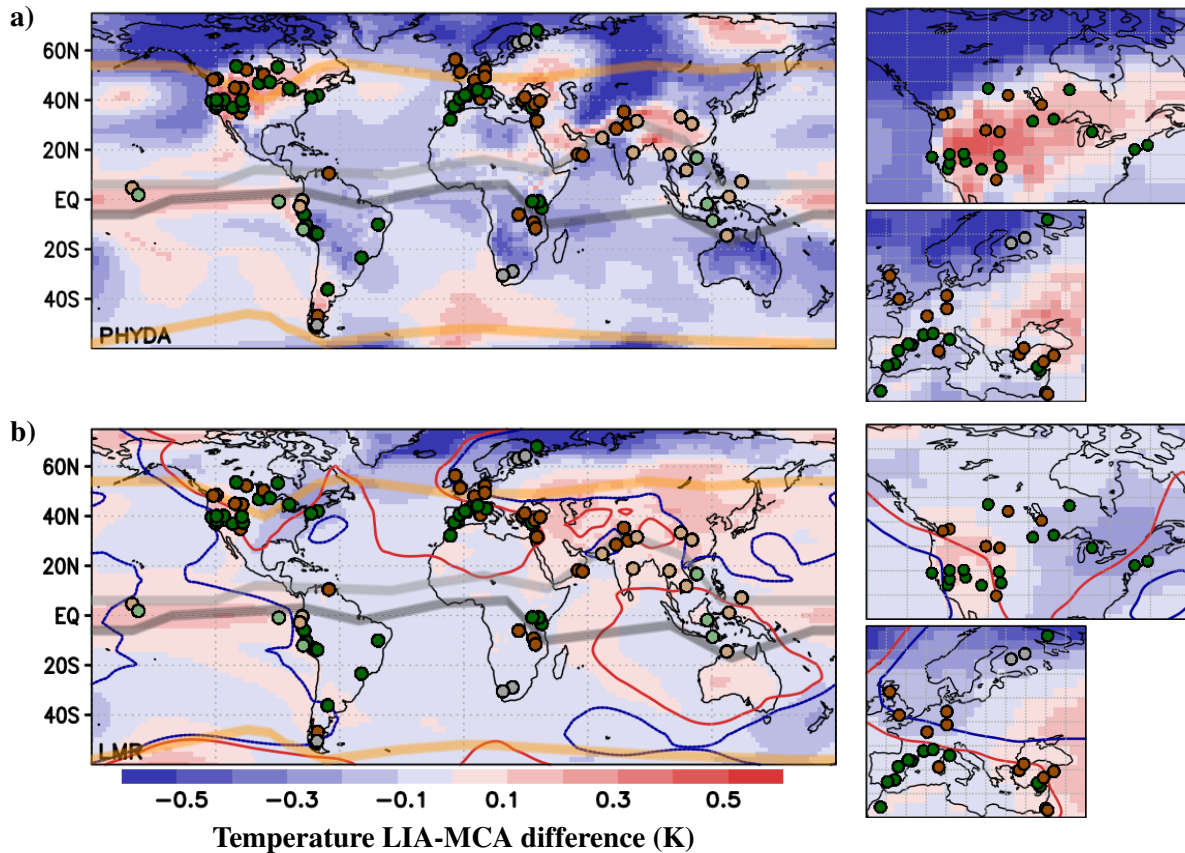
Figure C1. LIA-MCA differences in temperature from **(a)** PHYDA, and **(b)** LMR. Dots represent the locations of Fig. 1, with the same color code considered in that figure; positions of the modern ITCZ and NAM and SAM boundaries are also shown, as described in Fig. 1. For the LMR, differences in SLP are also shown with contours of -5 (blue) and 5 (red) Pa. Average of JJA is considered for the PHYDA, while the annual average is considered for the LMR. For each product, regional maps over North America and Europe are shown on the right side of each corresponding panel.

Indian and western Pacific basin and negative anomalies in the eastern Pacific basin during the LIA. These relatively small changes found in temperature and SLP could explain the limited impact observed in the analysis of PDSI for this area.

Author contributions. This study is part of PJRG's PhD. PJRG contributed with data processing, analysis of results and writing of the paper. JFGR, JES and FGP contributed to the analysis and discussion of results and to writing the paper.

380 *Competing interests.* The authors declare that they have no conflict of interest.
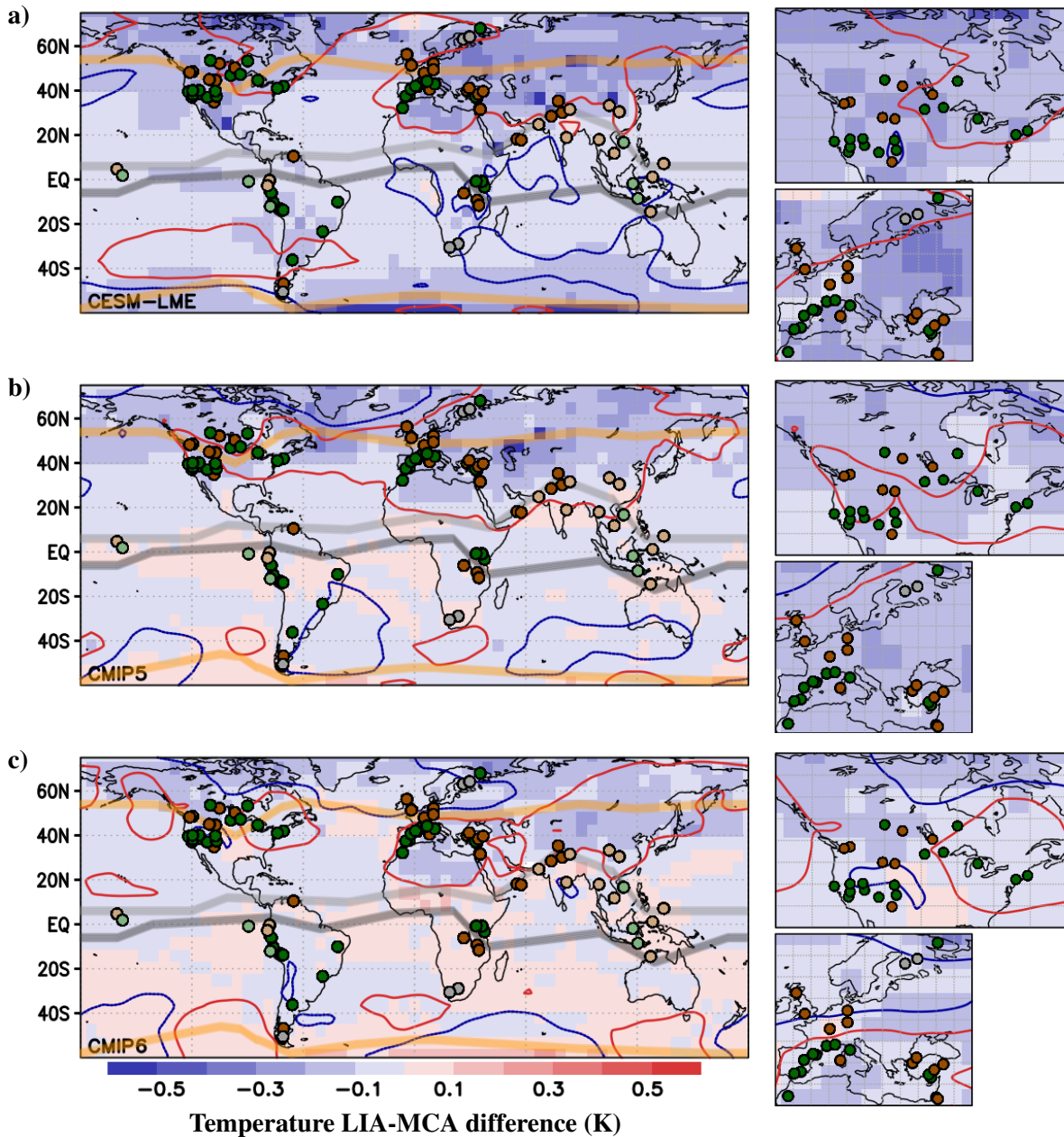
Figure C2. LIA-MCA differences in temperature (shading) or SLP (contours) from the ensemble average of **(a)** CESM-LME, **(b)** CMIP5 and **(c)** CMIP6 LM simulations. Dots represent the locations of Fig. 1, with the same color code considered in that figure; positions of the modern ITCZ and NAM and SAM boundaries are also shown, as described in Fig. 1. For the SLP, contours of -5 (blue) and 5 (red) Pa are shown. In all the cases, average of JJA is considered. For each product, regional maps over North America and Europe are shown on the right side of each corresponding panel.

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