

Dear Editor,

We thank reviewer #1 for their comments and suggestions on our paper. Below, we discuss these point-by-point and indicate how we intend to adapt the manuscript in a revised version.

Sincerely, also on behalf of my co-authors,

Carolien van de Weijst

Summary

The study by van der Weijst et al presents multiproxy-based temperature and productivity records spanning the past 15 Ma. Data are generated from an ODP core recovered from the eastern equatorial Atlantic, at a site that is under the influence of monsoon-induced upwelling. The proxies used are alkenone-based UK'37, archaeal GDGT-based TEX86 and dinocysts. Based on several lines of evidence, the authors argue that the TEX86 proxy records a mixture of surface and subsurface signal, likely that of Antarctic Intermediate Water (AAIW) as also suggested by two previous studies. Assuming that UK'37 reflects SST variability, benthic d18O records NADW variability and TEX86 reflects AAIW variability, in combination with other published temperature records and a pCO₂ record, the authors discuss possible mechanisms that lead to M2 glaciation.

General comments

*Overall I find the manuscript very clear and well-written. The topic also fits well within the remit of the journal, thus will likely be of interest to the broad readership of *Climate of the Past*. Although generally accessible to the reader, I think some arguments can be further improved and/or need further clarification. I hope the authors will find my comments and suggestions helpful in revising the manuscript. Altogether it should amount to moderate revision. Below I outline my major concerns.*

REPLY: We thank reviewer #1 for their compliments regarding our work.

(1) TEX86 reflects AAIW variability

The authors provided several lines of evidence to support their claim that TEX86 reflects subsurface temperature: high GDGT 2/3 ratio, TEX86H SSTs that are out of the modern range and unrealistically large magnitude of change over the last 15 Ma, temporal trends that are more similar to those of benthic d18O than UK'37 and Mg/Ca based on both mixed layer- and thermocline-dwelling species. The discussion is generally convincing.

What is however unclear to me is how the authors then link the TEX86 signal to AAIW variability. I find the explanation provided at Line 226-227 to be rather vague. The core-top value of their SubT TEX86 record is ~14°C, which is a lot warmer than the temperature in AAIW (~5°C according to Figure 7a) and is in fact closer to that of South Atlantic Central Water. Further, the core of depth range integrated by the HL16 calibration is 100-350 m, which is much shallower than the average depth of AAIW. As the climate interpretation hinges on this claim, the authors need to provide clearer arguments to support their statement. I wonder if it is possible to attribute the relative influence of SACW vs. AAIW using the depth distribution in the calibration used?

REPLY: South Atlantic Central Water (SACW) present in the water column at Site 959 is formed in the South Atlantic subtropical gyre. Together with the Antarctic Intermediate water (AAIW) below (characterized by relatively low salinity and an increasing component with depth; see fig 7a) this forms the permanent thermocline waters. We agree with this reviewer that our signal represents a mix including both SACW and AAIW which hence must partly track AAIW variability. In other words, our TEX86 signal tracks Southern Ocean surface temperature variability imprinted in both SACW and AAIW. But clearly, we are sampling SACW so to comply with the comment of the reviewer, we will change reference of AAIW in the remainder of the manuscript to SACW, thereby noting that both represent a Southern Ocean climate signal.

(2) Subsurface export of GDGTs

Most sedimentary GDGT 2/3 ratios along the core are > 5. The authors argue that this indicates that GDGTs in the sediment core are partially sourced from the subsurface ocean. I acknowledge that the GDGT 2/3 ratio is a routinely used indicator to cull TEX86 data, it does however worry me that the correlation between this ratio and TEX86 is quite strong for this core. The authors use the positive correlation to argue that the TEX86 variability is NOT controlled by subsurface-sourced GDGTs, as more subsurface GDGTs should result in colder temperature estimates hence also a negative correlation between TEX86 and GDGT 2/3 ratio. But alternatively this positive correlation might simply mean that the ratio is indeed reflecting temperature change instead of changes in the surface vs subsurface source of GDGT. Indeed, from Figure 7 of Taylor et al. (2013, GPC), it does look like the GDGT 2/3 ratio is not only correlated to water depth but also to SST. If true, this would mean that the authors lose one of the strongest lines of evidence for the subsurface origin of their TEX86 record. I encourage the authors to discuss this possibility in detail.

REPLY: Taylor et al. (2013) discussed this issue in their paragraph 5.1. They conclude that water depth is the dominant control of GDGT[2/3] ratios in sediments, both in the calibration data set and in the paleo-domain. They also conclude that the positive correlation between SST and GDGT[2/3] (notably the opposite relation to that expected based on homeoviscous adaptation) is an oceanographical artifact. More recent work, as described in lines 179-181, has confirmed this interpretation based on water column profiles. This work finds no evidence of GDGT[2/3] values above 4 in surface waters, making it highly unlikely that the ratio values we find at Site 959 represent solely surface-derived GDGT assemblages. Accordingly, they must integrate a larger depth range. To comply with the comment of the reviewer, we will include the conclusions that Taylor et al. present in their paragraph 5.1 in section 4.1.1.

Second, what process would be responsible for the export of GDGTs from subsurface ocean at 350 m or within the AAIW to the seafloor? I appreciate the fact that this is not a calibration study, but the lack of vehicle to export subsurface ocean GDGTs to depths is one of the main criticisms for the HL16 calibration, so I think the reader will be a lot more convinced if the authors can propose some possible mechanisms to explain/demonstrate that a subsurface export of GDGTs is indeed possible.

REPLY: This is a valid point and has been extensively discussed in previous papers, including Taylor et al. (2013) and Ho and Laepple (2017). Because it represents an issue that is outside the scope of this paper but still important, we will include a few sentences in the introduction, where the issue of deep-production is discussed.

(3) Is the temporal resolution of the records sufficiently high to assess a 5-10 ky lead/lag relationship between records?

The paleoclimate interpretation is largely based on the temporal patterns, so the choice of calibration does not matter much. But what matters here instead is the temporal resolution. Most of the discussion in the last subsection of Discussion is based on the 5-10 ky lead of TEX86 over benthic $\delta^{18}O$, and the lead/lag relationship between pCO_2 and temperature records. The lead of TEX86 over benthic $\delta^{18}O$ is, to my eyes, based on one or two data points. Having said that, I don't think that one can make such a claim given the low-resolution of the proxy records and uncertainty in proxy measurements. I urge the authors to give this some more thought and provide a more balanced discussion taking into account the limitations of their dataset.

REPLY: this is a valid point and therefore we clearly indicate that this is somewhat speculative in this section, acknowledging that data resolution is the limiting factor (283-284). As suggested below, we will include the inflection points in Figure 5 by plotting lines and will carefully reword the section to convey that we do not consider this data to represent full proof of our hypothesis.

Specific comments

Line 13-15: I would flip the arguments around. First the ecological evidence that Thaumarchaeota / GDGTs occur in the subsurface ocean, then only the circumstantial evidence of the good correlation of TEX86 to temperature from various depths.

REPLY: We will reconsider wording here to optimize the argument.

Line 17: “proved” is a bit strong. In my opinion, structural similarities in downcore records are at best circumstantial evidence, not direct proof. What about “can be best assessed in downcore studies”?

REPLY: we will replace “proved” by “determined”

Line 166-174: The authors argue that TEX86 records subsurface ocean temperature variability because the long-term trend in TEX86 differs from that of other SST proxies, including UK’37 and Mg/Ca. As the authors noted, these two proxies have their own issues: UK’37 is close to saturation and thus may be insensitive to temperature change, whereas Mg/Ca is susceptible to secular change in seawater Mg/Ca, carbonate chemistry and salinity. I note that the Mg/Ca records are from a sister paper (cp-2021-68), and likely the authors have discussed all the issues in that paper. It would still make life easier for the reader if the authors can briefly summarize these Mg/Ca issues here and why the trend is robust despite these potential caveats. As for UK’37, it might be helpful to also test the Bayspline or other tropical UK’37 calibration (e.g. Sonzogni et al. 1997 DSR II) that has a steeper slope than the one in Prahl and Wakeham (1987 Nature) or Müller et al. (1998 GCA).

REPLY: we will include a brief summary on the Mg/Ca. We are hesitant to include other tropical UK’37 calibrations. We agree with the criticism of Herbert et al. (2020) regarding the application of UK37 using the calibration suggested: “We hesitate to adopt the Tierney and Tingley (2018) approach because it potentially amplifies noise (difficulty in accurately determining the very small C37:3 peak area at high UK’37) into signal”. We will include this at the end of section 2.2 (methods).

Line 189-190: Unclear reasoning, please rephrase. Also see my general comment (2).

REPLY: this was indeed somewhat clumsily phrased. This will be corrected in the revised version.

Line 206-207: It is difficult to see this in Figure 5. It might be helpful if the authors can illustrate the lead in the figure. Also see my general comment (3). Line 223–224: Again, this is not immediately clear from the figure. Also, how is it established? How is the onset defined? See comment on Line 206-207.

REPLY: To clarify, we will include lines to indicate the interpreted age of changes discussed.

Line 240-241: I would rephrase this. Strictly speaking, HL16 calibration does not assume/target any water depths. Instead, they search for calibrations that can reconcile the variability of UK’37 and TEX86. The depth distribution indicates that the most probable depth range is 100-350 m. Thus, instead of using the entire calibration ensemble that includes all depth ranges down to 950 m, one may very well choose one that is calibrated to 100-350 m.

REPLY: This comment indicates to us that we have to further clarify how we determined the depth distribution of the calibration weights from the calibration ensemble of Ho and Laepple. The calibrations were binned based on depth ranges. In the revised version, we will include a new section in the supplementary information that will clarify our approach, including the depth ranges used.

Line 242-244: Thaumarchaeotal cell counts and GDGT concentrations vary a lot in space. Are studies from the South Atlantic and Arabian Sea the suitable choice of reference here? If no water column studies are available in the study area, at least tell the reader why it is reasonable to assume that what happens in the water column elsewhere might be applicable to eastern equatorial Atlantic.

REPLY: We are slightly confused regarding this comment as this sentence summarizes inferences from the globally available data sets of the vertical distribution of Thaumarchaeotal cells counts and GDGT concentrations, extensively discussed in the literature. This is used as a model for Site 959, as is the TEX86 calibration from a global set of surface sediment data. To address the concerns of the reviewer, we will reference the Atlantic Ocean study of Sintès et al. (2016, <https://doi.org/10.3389/fmicb.2016.00077>).

Line 251-257: Given the hydrography at the study site which is under the influence of episodic monsoon-induced upwelling, I would expect the temperature variability here to be larger than that at pelagic sites like the warm pool. So I think it unlikely that the 1:1 relationship between the surface and subsurface ocean would hold over millions of years, nor is it reasonable to assume that the subsurface temperature variability here would be comparable to global SST change in the tropics.

REPLY: Here, we merely assess the implications of assuming the 1:1 relationship (tested by Ho & Laepple (their figure S5) but indeed not over million-year time scales). This indicates a SST drop that is consistent with estimates of the global average cooling (so indeed slightly larger than elsewhere in the tropics as the reviewer suggests). We indicate quite clearly that this assumption needs further testing (final 5 lines of this paragraph).

Line 277-285: See general comment (3).

REPLY: See our reply to general comment 3