

Dear editor,

Thank you for editing our manuscript!

We really appreciate the careful scrutiny that professors Helmut Weissert and Dan Breecker took on the manuscript. The constructive comments and suggestions have pushed us to greatly improve the article.

Besides of one-by-one reply (see PS: Replies to the reviewers), we have made additions and corrections following the reviewers. Below are the main changes.

- 1) According to the suggestion of reviewer #1, we have added a paleogeographic map showing the study area and the paleoclimatic zone, and numbered it the Fig. 1a
- 2) We have deleted the supplementary data Note S2 as the reviewer #1 suggested and placed this part in the end of section 3.3. Totally 26 lines (new lines 159 to 184) were added in text for the change.
- 3) We assembled the Early Jurassic carbon isotope ratio curve of organic matter (Fig. 8d) following the suggestion of the reviewer #1.
- 4) By the suggestion of reviewer #2, we increased a paragraph in the text (subsection 3.2) to make notes of the cathodoluminescence images (CL) of calcretes, and combined CL with the observation results of petrography and field occurrence we explained the validity of carbon isotope for the estimate of $p\text{CO}_2$ (lines 135-143).
- 5) We have made the error propagation for the $p\text{CO}_2$ results with procedures and formulas by Breecker (2013) and Breecker and Retallack (2014). $p\text{CO}_2$ error curve is added in figure 8a and a paragraph of uncertainty assessment is complemented in the end of the subsection 4.3 (lines 242-248). The procedure and parameters are added as the supplementary data Table S7.
- 6) In addition to the replies to the two reviewers, we have done lots of correction and amendment for the whole text, comparing to the previous version. English language is further polished as possible as we can.

Thank you for time!

Yours sincerely,

Xianghui Li, Jingyu Wang, Troy Rasbury, et al.

PS: Replies to the reviewers.

1 Replies to Prof. Helmut Weissert

1) Introduction The authors start their study with a rather general sentence: “The Jurassic was a typical greenhouse period”, what is a “typical greenhouse” ? The Jurassic was an interesting time because of major plate tectonic changes affecting global climate. The fragmentation of Pangea resulted in the collapse of a Monsoonal-type climate in the earliest Jurassic and by the end of Jurassic a more zonal climate was established. Jurassic was, no surprise, a time with cooler and warmer climate, possibly even with Ice Age episodes and the

new data in this study confirm that changes in Early Jurassic pCO₂ were considerable. In their study, the authors compare data collected from the Sichuan Basin with data from the Colorado Plateau and from North China and northern Gondwanaland. As a reader I am, of course, interested in the paleogeography of the Early Jurassic and in the relative paleopositions of the study site and of other localities mentioned in the text. I recommend that the authors add a figure showing the plate tectonics of the Early Jurassic and showing the locations mentioned in their study. In addition to paleogeography, paleoclimate models could help to better understand climate trends in SW China and in North America, both discussed in this study. However, there are almost no models available on Early Jurassic climate. One of the few simulations providing information on precipitation pattern and aridity in the Early Jurassic can be found in Robertson et al. (Sedimentology, 2017). The authors may refer to this study. Models may help to better formulate hypotheses and questions addressed in this study. And, clear hypotheses and questions facilitate reading of the paper . » The authors may considerably improve their introductory part. Time window chosen: Your time window is defined by post-mass extinction time (Hettangian) at the base and the Toarcian OAE at the top. If you extend part of your discussion to the Mesozoic-Cenozoic (terrestrial proxies have begun to provide information: line 41) then you soon are moving on slippery grounds: Many, many studies on terrestrial climate exist for the Late Mesozoic and even more for the Cenozoic. Focus on your time window, even in your figures (fig 8)

We deleted the sentence “The Jurassic was a typical greenhouse period”. This gets rid of greenhouse, but still expresses that the Jurassic was hotter than today. More works (almost rewriting) were made in the paragraph, with suggestions of the referee, to enhance expression of the Early Jurassic changeable and oscillating climate.

Following the suggestion, we have added a global map of climatic paleogeography, on which locations of the study area and correlative Colorado Plateau are marked. We replaced the previous Fig. 1 as Fig. 1b and the new paleogeographic map as Fig. 1a. Please see the new Fig.1. We laid the global Early-Middle Jurassic climate zones (Boucot et al., 2013) on the Early Jurassic (~193 Ma, Sinemurian) paleogeographic map (Scotese, 2014), instead of a GCM map selection in the paper by Robinson (not Robertson) et al. (2017. Fig. 5, page 223, 64(4), Sedimentology). This is because oceans and continents (plates) are not distinct and climatic zones are not shown in the Early Jurassic GCM maps.

We further deleted the Gondwanaland term in the text, because it was a superland (composed of Antarctic, Africa, South America, India, Australia, New Zealand, etc.) of the South Hemisphere during the Paleozoic, and began to disassemble after the Pangea break-up since the Late Triassic, and it was not an independent supercontinent when the Pangea formed.

For the time window, we focused the interval of the Sinemurian to the early Toarcian in section 5.2 (pCO₂ discussion) and figure 8 although we still keep the Hettangian in section 5.1 (climate evolution) as the Hettangian sediments can be observed in some investigated sections, and they show the climate transition from more humid to arid into the Jurassic.

2) Material and methods The methods you apply are up-to-date and they provide useful information on past climate. I recommend to add a paragraph on your selection procedure of CO₂ concentration in soils. You have this discussion in your notes, that is fine. However, the S(z) value is a crucial value for your pCO₂ estimates, please include your selection arguments in your main text. See also your line 162, where you may discuss the S(z) selection procedure if you did not do it in your method paragraph.

We have deleted the supplementary data Note S2, and placed the main part in the section 3.3

(Calculation of atmospheric CO₂ concentration). Then we added the notes for the procedure of CO₂ concentration estimate and S(z) value selection in the revision. Totally 26 lines (new lines 161 to 186) were added.

3) Discussion of your data In your discussion, you present first your sediment proxy data stage by stage, then, in a next chapter, you discuss your pCO₂ data through the Early Jurassic. This structure of the text makes the reading of the discussion, at least for me, quite difficult. I prefer to see a climate discussion starting with the drivers of climate change, in this case changes in carbon cycling and in pCO₂, which are both global signals (See, for example, the new and detailed C-isotope curve for the Hettangian-Pliensbachian in Storm et al., PNAS, 2020). You can, if you take a general reference curve as a start, project your data into this curve and test, if pCO₂ trends coincide, for example, with C-isotope trends, you can see if regional climate pattern (China, N-America etc) reflects some of the global trends. This may make the paper much easier to read and you can easily show the regional pattern as part of a global climate curve. You may refer in this discussion, if possible, to the few available model data. Of course, you also will use literature data you chose for your study your figure 8 (e.g. Dera et al. 2011 and many others, as you cited correctly). A discussion starting with the global pattern also can make it easier to understand your comparison with regional data from N. America.

In the section 5 of the revision, we start with an introduction of the globally oscillated Early Jurassic climate with causes. Then we discuss the climatic change of the GSB based on sedimentary proxies and compare to other places of the world. In section 5.2, the pCO₂ record and rapid change (event), and correlation to the global climate (sea water temperature) and carbon cycle are discussed. This contexture aims to test the hypotheses from the marine climate records.

In figure 8, we assembled the Early Jurassic carbon isotope ratio curve of organic matter (Fig. 8d) from the Mochras borehole, Cardigan Bay Basin, UK (Xu et al., 2018; Storm et al., 2020) with seven-point average smoothing against depth (mbs). We did not combine the carbon isotope ratio curves of organic matter from the Paris Basin with that of Cardigan Bay Basin into one because the former is mainly against depth, and the latter against time.

4) Details and corrections, including some comments on language Abstract : : terrestrial sediments show more complicated environment and climate: : ” I assume that you want to say that climate proxies in terrestrial sediments are more complex and more difficult to interpret than many marine proxies. Please do not write “carbon-oxygen isotopes”, but carbon and oxygen isotopes. line 42 You are rather imprecise in your wording, when you write about a “negative feedback : : . has been hypothesized to account for: : : in the carbon cycle” > what to you exactly mean? 105 : : : descriptions for sedimentary facies analysis were executed: : : 116 the description how to distinguish dolomite from a calcite the field is not really needed. It should be basic knowledge for students in geology: : : 216 : : : a distinct transfer of climate.. ??? 239 : : : a warm-humid climate followed the Late Triassic: : : > Late Triassic climate 244-248 You compare North American climate with GSB climate throughout your study, this is ok, however, I like to know why you chose North America, what are hypotheses on climate similarities and differences between N. America and SW China. 248 » here it will be interesting to discuss climate pattern in the Early Jurassic, why is which regional climate similar/ different from another regional climate? 305 Interesting is the observation that lake facies was widespread during a dry period in the Toarcian. You may further comment on this. 318 you use dolomite formation as a climate proxy, this is fine. You may refer to literature on dolomitization along the Persian Gulf (“Arabian Gulf”) , eg. by McKenzie and others. 120 chichen-wire > chicken-wire (anhydrite) 253 That is the reddish rocks developed through the whole member: : : ..., but it started: : : revise 257 : : : . Calcisols were also interpreted with the description of

abundant calcretes: : : (??) revise style 282 The Ma'an-shan Member is likely the Pliensbachian: : : 302 In other hand,: : : 321 : : : can serve the determination: : : I did not mark the many additional small language inconsistencies » please revise text carefully.

In the abstract, we have deleted the sentence “Unlike marine archives, terrestrial sediments show more complicated and dynamic environment and climate”, added a new sentence to introduce the climate of the Early Jurassic, and changed all the “carbon-oxygen” as “carbon and oxygen” in the text, tables, and figures.

Line 42. The questionable sentence was partly cited from the original paper (Xu et al., 2017). We deleted the sentence that is not distinctly relevant to the topic, and added another sentence to complement the subject of this paragraph.

Line 105. We revised the sentence as “We have made observations and descriptions of sedimentary characteristics for lithofacies analysis at six outcrop sections.....”.

Line 116. Yes, we agree that the dolomite recognition is basic for any students in geology. We have deleted the sentence relevant (lines 115-123).

Line 120. We have corrected “chichen-wire” as the “chicken-wire”.

Line 216. Taking the $S_{(z)}=2000$ ppmV for pCO_2 estimate originates from the calcisol determination and (semi-) arid climate indication (details are in the added parameter conditions in section 3.3). This climate condition is harmonized with the results indicated by climate-sensitive sediments, and may not be the representative of a climate transfer.

Line 239. Yes, the climate of the Hettangian (early Early Jurassic) is similar with that of the Late Triassic, but different from other ages of the Early Jurassic. The evidence of the Late Triassic warm-humid is provided in the first sentence of the first paragraph in section 5.1. And the warm-humid climate continued in the Hettangian age (verified in subsection 5.1.1). It is noted that we have stressed a total climate condition of the Early Jurassic as a (semi-) arid in the whole article even though the Hettangian was different in climate from the main Early Jurassic.

Line 244-248. Yes, we compared the climate between GSB and (North) America through the section 5.1. From the comparison, we can see that: in the Hettangian, warm-humid climate in GSB was different from the arid climate in North America, but in the Sinemurian-Toarcian, both had the similar (semi-) arid climate. There are two reasons for us to compare the GSB with America in climate. One is that they were close in paleolatitude (~15°-30°N. Fig. 1a) and similar climate zone (tropical) of the North Hemisphere in the Early Jurassic. The other one is that relatively complete climate-sensitive records of the Early Jurassic have only been published and are available in North America. In North (-west) China, the Early Jurassic succession is also complete, but the warm-humid climate is totally different from the GSB although we sometime compare them in the text. The comparison of the GSB and North America illustrates that the secular Early Jurassic (semi-) arid climate in the two regions could not have been responded to the (e.g., CAMP and Karro-Ferrar LIP) volcanism, which may have led to the global oscillation climate. This hypothesis is added in the last paragraph of section 5.1. The secular (semi-) arid climate could be related to the paleogeography. The aridity for the two regions does not obscure the total trend and eventful change of pCO_2 demonstrated in section 5.2 as pCO_2 is a global signal of climate.

Line 253. Yes, this sentence was confusing. We rewrote as: “Differences in the color appearance show that the reddish color started in the middle member in the central basin (Location A6. Fig. S2) but almost developed through the whole member in the western basin (Location A4. Fig. 6)” (new lines 282-284).

Line 257. We have revised the sentence to “We also interpret the reddish muddy sediments with abundant calcretes as the calcisol at sections of Dafang (Location A8. Zhang et al., 2016), Tianzhu (Location A9. Li and Chen, 2010), and Weiyuan (Location A10. SBG, 1980a) (new lines 287-289).

Line 282. The sentence was deleted in the new version as the age assignment is introduced in section 2. It is actually repetitious.

Line 302. We have composed this sentence with the first sentence in the paragraph.

Line 305. Yes, arid climate indication seems not consistent with the large lacustrine transgression. We then added some phrases in the next sentence (together composing a composite sentence), to show the possibility of arid records. Actually, the next couple of paragraphs are the process that verifies the arid climate by proxies of climate-sensitive sediments (lacustrine micritic dolomites, gypsum) and carbon and oxygen isotopes of lacustrine carbonates as well as (lacustrine or land side) reddish mudrocks.

Line 318. As we know, there are a number of hypotheses for the formation of dolomite mineral, such as primary authigenic origin, diagenetic replacement, microbial mediation. Yes, McKenzie J. and her team have much contributed to the microbial dolomite. We have referred to her papers in the revision, while we preferred to interpret the dolomites formed in an arid/evaporate climate condition especially when it is associated with gypsum (references added).

Line 321. We have considered the issue, and reorganized the sentence with others in the same paragraph. Please see the new version.

Other similar issues were carefully checked and corrected in revision.

Figures I like your figure 6 which serves as a very good baseline for your discussion (you may also add data from Xu et al. from the same study area in figure 6 or elsewhere?)

Fig. 6. Adding the $\delta^{13}\text{C}_{\text{om}}$ curve (Xu et al., 2017) in western GSB in figure 6 is a good idea. But the T-OAE $\delta^{13}\text{C}_{\text{om}}$ curve (Xu et al., 2017) is too short (thin strata) to occur in the figure, i.e., the height of the curve is less than 1 cm if it is placed in the corresponding position. The other problem is that it is difficult to in age correlate between the two sections.

Fig 8 You plot changes in temperature not temperature, please correct this in the figure. The global paleosol curve is, in this case not very helpful. I also wonder, if you should start your curve after the T-J boundary (?). The high-resolution data are not really discussed in your work and you did not include any T-J boundary data.

In figure 8, the seawater temperature was corrected!

Yes, the global paleosol curve is not very helpful for the observation and comparison of the Early Jurassic global climate, but we have to keep it in the figure as it shows the present situation of the global paleosol $p\text{CO}_2$ reconstruction except for this work.

Yes, we agree and delete the high-resolution latest Triassic data in the figure.

2 Replies to Prof. Dan Breecker

My concerns are primarily related to the new pCO₂ determinations made. First, luminescent calcite (lines 141 and 142) is probably not a good material to use for paleoCO₂ determinations because **luminescent** pedogenic carbonate is thought to form under anoxic conditions, associated with water-saturation when there is a poor connection between soil pore spaces and the atmosphere (Mintz, J. S., Driese, S. G., Breecker, D. O., & Ludvigson, G. A. (2011). Influence of changing hydrology on pedogenic calcite precipitation in Vertisols, Dance Bayou, Brazoria County, Texas, USA: implications for estimating paleoatmospheric pCO₂ (Journal of Sedimentary Research, 81(6), 394-400.). The paleosol carbonates studied here might be 'weakly' luminescent, but it is hard to tell without any quantification/standardization. It is also possible that there are other factors that influence luminescence. But all of this needs to be discussed so readers can evaluate the selection of materials. I will say, however, that the careful petrography and drilling of dense micritic zones is a plus.

We thank for the valuable comment and providing the reference.

Mintz et al. (2011) provides a good example that hydrology influences the luminescence of calcretes, that we now cite and use to refine our discussion of the luminescent quality and justification for sampling the Jurassic calcretes we present in this paper.

All of the samples studied were screened for cathodoluminescence (CL), and only a few are shown in the manuscript. The samples shown have the brightest luminescence of all of the studied samples, and they were chosen because the quality of the images is much better. Most are calcites and dull to non-luminescent, with little (~ 5-10%) light orange or brownish red luminescence. The luminescence is almost certainly due to a relatively high Mn/Fe ratio, and we expect that seasonally the soils may have been water logged and disoxic.

More importantly, the key to distinguish the pedogenic calcretes from other geneses is the identification of both field occurrence and micro-texture. Our field and microscopic observations demonstrate that the calcretes have typical pedogenic features. The ginger-like calcretes are discrete within the Bk horizons, and do not form linear/tabular limestone. Slickensides and vertical rhizoliths can be often / sometimes seen in paleosols. Petrographically, we see that predominant micritic calcites occupy the dense areas of the calcretes (Fig. 2). Some calcretes have areas that were cracked and filled by secondary / diagenetic spar-calcites, that were avoided when micro-sampling.

Based on field and petrographic observations, we drilled powder samples for stable isotope analysis in dense micritic zones as the referee suggested. Thus, the carbon isotope value of carbonates can be used to estimate pCO₂.

In the new figure 2, we inserted the scanned photos of the thin-sections and marked the cathodoluminescent and drilling dense areas and added plane light photos (Fig 2c and 2d), which roughly correspond to the CL image positions. We also added sentences in the text (3.2) to make notes on the observation results of petrography and CL images (new lines 135-143).

I am concerned that the CO₂ changes the authors interpret here may not be statistically significant changes. This is impossible to evaluate without **uncertainty** quantification. The authors do consider the effect of using different input values for the pCO₂ calculation, but my guess is that they have nonetheless largely

underestimated the error associated with their approach. For instance, the authors calculate $\delta^{13}\text{C}_r$ values from $\delta^{13}\text{C}$ values of OM measured in different locations (across the globe) from the carbonate nodules. What magnitude of uncertainty might this introduce? Furthermore, $\delta^{13}\text{C}_a$ is calculated from $\delta^{13}\text{C}_r$. Given the effects of CO_2 and water stress on $\delta^{13}\text{C}$ values of C3 plants, this approach is associated with substantial uncertainty that is not addressed in this manuscript. The authors recognize that there is uncertainty associated with the value of $S(z)$. However, their consideration of $S(z) = 2000$ and 2500 ppmV is not an accurate representation of the uncertainty. I suggest error propagation that includes uncertainty associated with each input to the equation on line 159 and the results shown as error bars on each CO_2 determination.

Following the suggestion, we have made the error propagation for the $p\text{CO}_2$ results with procedures and formulas by Breecker (2013) and Breecker and Retallack (2014), and also considered parameters from Zhang et al. (2018). We have added the errors in figure 8a and a paragraph of uncertainty assessment in the end of the subsection 4.3 (new lines 242-248).

It is noted that parameters of temperature, $\delta^{13}\text{C}_r$, $\delta^{13}\text{C}_a$, $\delta^{13}\text{C}_s$, and $S(z)$, remain the same as the calculation of $p\text{CO}_2$ in the supplementary Table S4 in order to be consistent. We did not use variable temperatures and $S(z)$ because the clumped isotopes analyses is currently not available in China. Additionally, and the depth (m) to the Bk is not known due to the disappearance or erosion of the top boundaries of the observed paleosols, thus making our $S(z)$ values minimum estimates. 2°C , 0.10%, 0.1%, and 0.1% are selected for the standard errors of temperature, $\delta^{13}\text{C}_c$, $\delta^{13}\text{C}_{\text{om}}$, $\delta^{13}\text{C}_a$, and 788 ppmV is adopted for the standard error of soil carbonate transfer function with $S(z)$ as suggested by Breecker and Retallack (2014). For details please see Table S5.

Results of error propagation show that the largest source of the uncertainty is the $S(z)$ standard error 766 ppmV of modern soil carbonate (Breecker and Retallack, 2014). The second largest source of error is the $S(z)$ value selection. Details of these errors are now discussed in a new paragraph in text (new lines 250-259, subsection 4.3). Other errors such as those for temperature, $\delta^{13}\text{C}_r$, $\delta^{13}\text{C}_a$, $\delta^{13}\text{C}_s$, exert far less to the uncertainty of $p\text{CO}_2$ estimates.

The descriptions of the sediments and paleosols reported here will be useful. I'm not sure I would call these **Aridisols**, though, because redoximorphic features are prominent (at least in some of the soils, e.g., Fig 3 a,b,c). Are you sure these are not Vertisols? Are there wedge-shaped peds? The authors mention abundant slickensides- a feature common in Vertisols.

We thank the reviewer for asking this question. There is much confusion from the literature with multiple classifications, and discussions of features such as slickensides that do seem to suggest more moisture. The aridisol (calcisol and gypsisol) and vertisol are distinguished with clay-heave structure by Retallack (1993, 1998, 2001). For calcisols, a high ratio of clay can produce clay-heave and slickenside structures due to hydrological changes. Additionally, some angular, subangular, and platy peds are common in the calcisols, but few are wedge-shaped. We concur with the reviewer that the Jurassic soils we have studied are likely vertisols.