

1 *Supplement of*

2

3 **Extreme warming rates affecting alpine areas in SW Europe deduced**  
4 **from algal lipids**

5

6

7 Antonio García-Alix, et al.

8

9 *Correspondence to:* Antonio García-Alix (agalix@ugr.es)

10

11

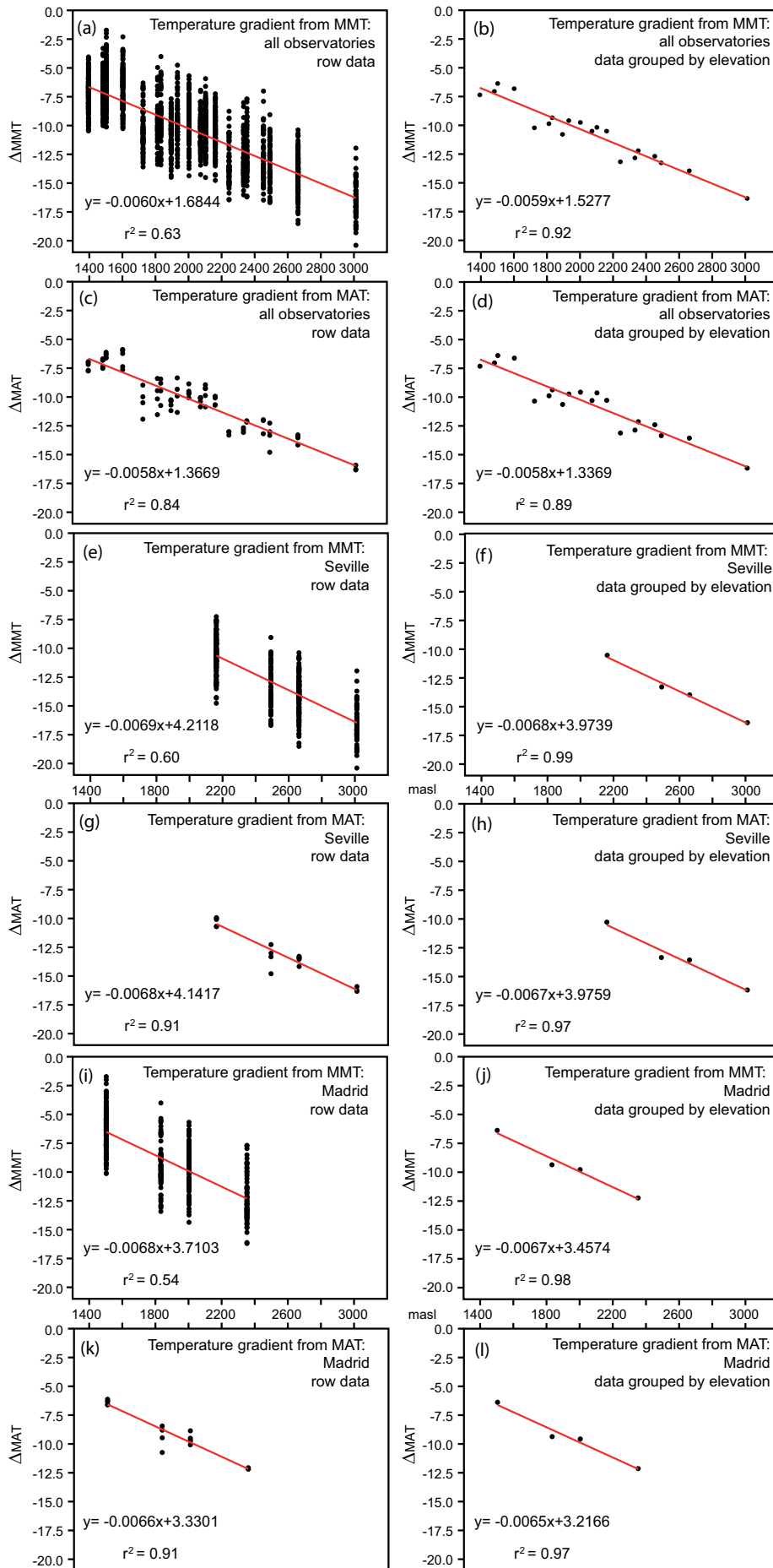
12 **This Supplement includes the following data:**

13

14       Supplementary Figures S1-S3

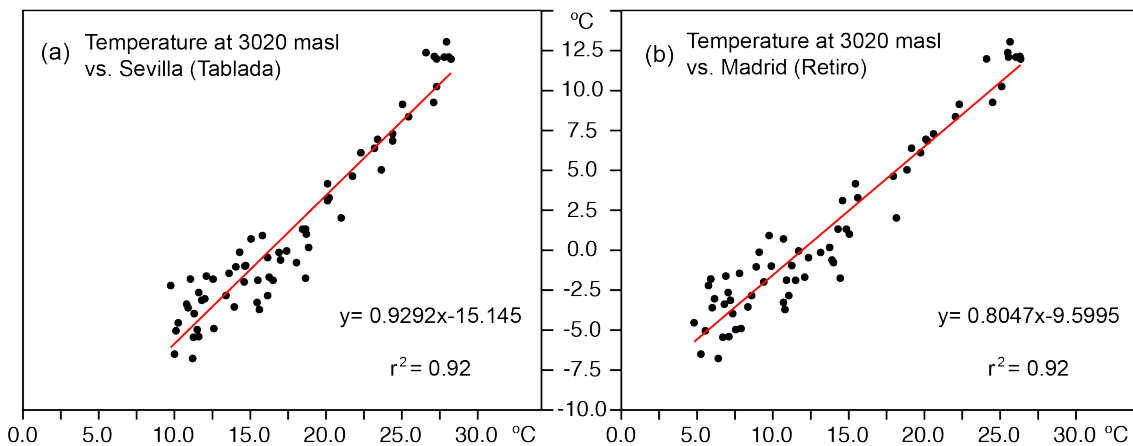
15       Supplementary Tables S1-S7

16       Supplementary References



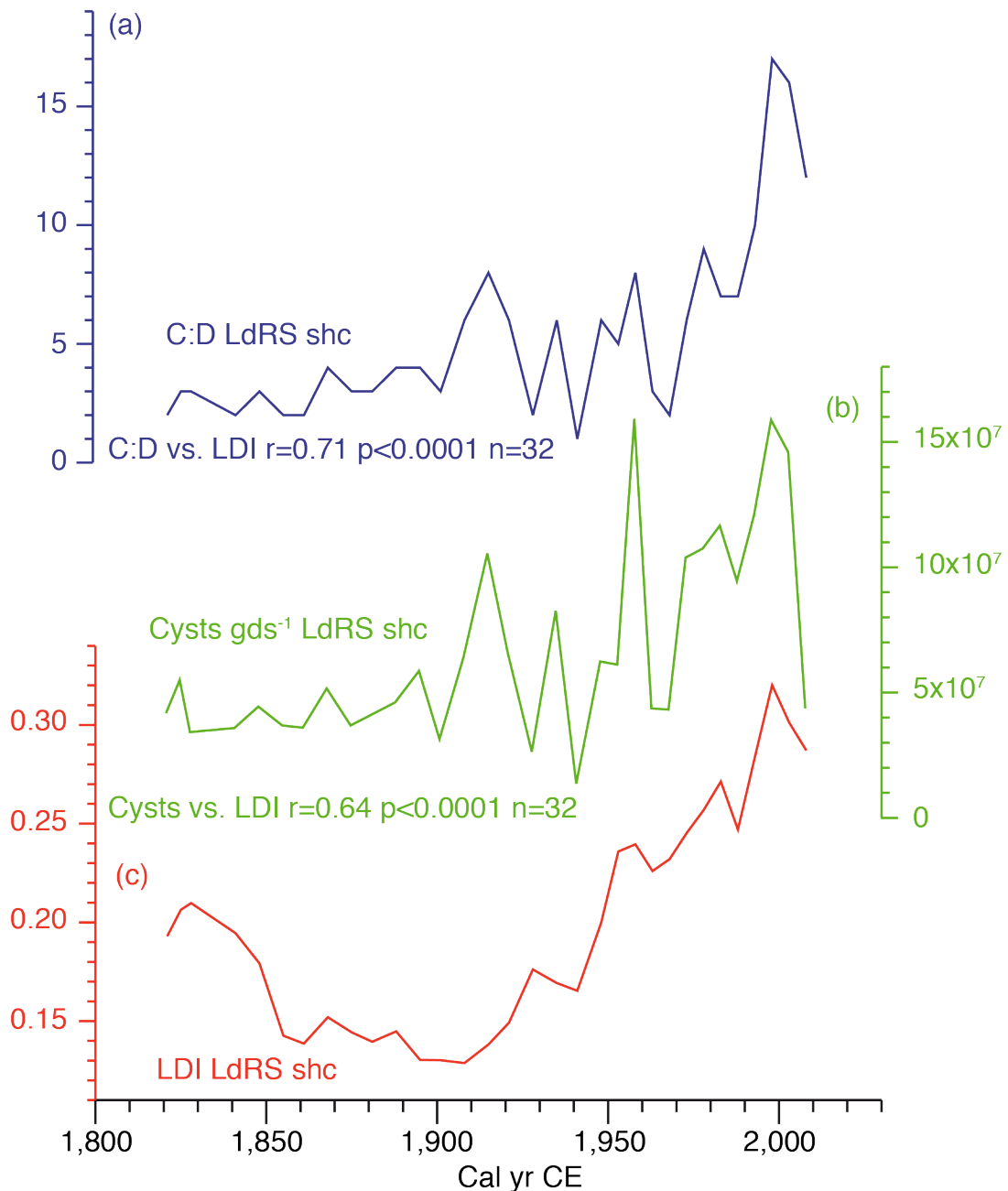
18 **Supplementary Figure S1.** Calculation of the Environmental Lapse Rate (ELR, °C/m) by means of  
 19 Ordinary Least Square regressions from temperature and elevation variations ( $\Delta_{\text{elevation}}$  and  $\Delta_{\text{MMT}}$  or  $\Delta_{\text{MAT}}$ )  
 20 between low and high elevation observatories listed in Table S1. Data from (Spanish National Weather  
 21 Agency - AEMet Open Data, 2017;Gonzalez-Hidalgo et al., 2015;Observatorio del cambio global de Sierra  
 22 Nevada, 2016). MMT (Monthly Mean Temperature) MAT (Mean Annual Temperature). (a) raw  $\Delta_{\text{MMT}}$  vs  
 23  $\Delta_{\text{elevation}}$  data (all observatories vs Sierra Nevada observatories); (b) mean  $\Delta_{\text{MMT}}$  vs  $\Delta_{\text{elevation}}$  data grouped by  
 24 elevation (all observatories vs Sierra Nevada observatories); (c) raw  $\Delta_{\text{MAT}}$  vs  $\Delta_{\text{elevation}}$  data (all observatories  
 25 vs Sierra Nevada observatories); (d) mean  $\Delta_{\text{MAT}}$  vs  $\Delta_{\text{elevation}}$  data grouped by elevation (all observatories vs  
 26 Sierra Nevada observatories); (e) raw  $\Delta_{\text{MMT}}$  vs  $\Delta_{\text{elevation}}$  (Sevilla observatory vs. Sierra Nevada  
 27 observatories); (f) mean  $\Delta_{\text{MMT}}$  vs  $\Delta_{\text{elevation}}$  (mean data grouped by elevation: Sevilla observatory vs. Sierra  
 28 Nevada observatories); (g) raw  $\Delta_{\text{MAT}}$  vs  $\Delta_{\text{elevation}}$  (Sevilla observatory vs. Sierra Nevada observatories); (h)  
 29 mean  $\Delta_{\text{MAT}}$  vs  $\Delta_{\text{elevation}}$  (mean data grouped by elevation: Sevilla observatory vs. Sierra Nevada  
 30 observatories); (i) raw  $\Delta_{\text{MMT}}$  vs  $\Delta_{\text{elevation}}$  from (Madrid observatory vs. Sierra Nevada observatories); (j)  
 31 mean  $\Delta_{\text{MMT}}$  vs  $\Delta_{\text{elevation}}$  (mean data grouped by elevation: Madrid observatory vs. Sierra Nevada  
 32 observatories); (k) raw  $\Delta_{\text{MAT}}$  vs  $\Delta_{\text{elevation}}$  (Madrid observatory vs. Sierra Nevada observatories); (l) mean  
 33  $\Delta_{\text{MAT}}$  vs  $\Delta_{\text{elevation}}$  (mean data grouped by elevation: Madrid observatory vs. Sierra Nevada observatories).  
 34 The obtained ERLs (ranging from 0.0058°C/m to 0.0069°C/m) are certainly close to the global mean ERL  
 35 (~0.0065°C/m) (Organization, 1993), showing that the different calculations are in agreement with the  
 36 global temperature-elevation gradients.

37  
 38  
 39



40  
 41  
 42  
 43  
 44  
 45  
 46  
 47

**Supplementary Figure S2.** (a) Correlation by means of Ordinary Least Square regression between Sevilla monthly temperatures and those from Cetursa 5 (3020 masl). (b) Correlations by means of Ordinary Least Square regression between Madrid monthly temperatures and those from Cetursa 5 (3020 masl). Data from (Spanish National Weather Agency - AEMet Open Data, 2017;Gonzalez-Hidalgo et al., 2015;Observatorio del cambio global de Sierra Nevada, 2016).



48  
49

50 **Supplementary Figure S3.** Comparison between (a) the ratio of Chrysophyceae cysts (including  
51 *Chromulina* spp.) against the cysts + diatoms frustules (C:D), (b) the concentration of Chrysophyceae cysts  
52 per gram of dry sediment (cysts  $\text{gds}^{-1}$ ), and (c) LDI in LdRS short core (LdRS shc), as well as their Pearson  
53 correlation (unpublished data from C. Pérez-Martínez). Interestingly, some Chrysophyceae algae produce  
54 resting siliceous cysts, more specifically *Chromulina nevadensis* and *Ochromonas* sp., two of the most  
55 abundant planktonic algae in Sierra Nevada lakes (Carrillo et al., 1991; Barea-Arco et al., 2001). There is a  
56 direct relationship between the amount of cysts produced and the number of live chrysophyte cells  
57 (Sandgren, 1988). The ratio between the cysts and the sum of cysts + diatoms frustules, as well as the  
58 number of cysts per gram of dry sediment (cysts  $\text{gds}^{-1}$ ) in LdRS show a significant long-term Pearson  
59 correlation with the LDI ( $r=0.71$ - $0.72$   $p<0.0001$ ) and ( $r=0.64$   $p<0.0001$ ) respectively. Since a statistically  
60 significant increasing trend has been observed in all the variables ( $p<0.05$  in Mann-Kendall test), variables  
61 were transformed to the squares of the z-scores ( $p<0.05$  in Mann-Kendall test for the LDI) and detrended,  
62 showing a significant Pearson correlation ( $r=0.76$   $p<0.0001$ ) and ( $r=0.45$   $p<0.01$ ) for the ratio between the  
63 cysts and the sum of cysts + diatoms frustules vs LDI and the number of cysts per gram of dry sediment  
64 (cysts  $\text{gds}^{-1}$ ) vs LDI, respectively.

Observatory	Observatory elevation (masl)	Sierra Nevada (SN) observatory	SN Observatory elevation (masl)	$\Delta$ elevation (masl)
Sevilla (Tablada)	8	Albergue	2500	2492
Sevilla (Tablada)	8	Cetursa 1	2170	2162
Sevilla (Tablada)	8	Cetursa 3	2670	2662
Sevilla (Tablada)	8	Cetursa 5	3020	3012
Granada Airport	567	Albergue	2500	1933
Granada Airport	567	Cetursa 1	2170	1603
Granada Airport	567	Cetursa 3	2670	2103
Granada Airport	567	Cetursa 5	3020	2453
Madrid (Retiro)	667	Albergue	2500	1733
Madrid (Retiro)	667	Cetursa 1	2170	1503
Madrid (Retiro)	667	Cetursa 3	2670	2003
Madrid (Retiro)	667	Cetursa 5	3020	2353
Granada Armilla	687	Albergue	2500	1813
Granada Armilla	687	Cetursa 1	2170	1483
Granada Armilla	687	Cetursa 3	2670	1983
Granada Armilla	687	Cetursa 5	3020	2333
Granada Cartuja	775	Albergue	2500	1725
Granada Cartuja	775	Cetursa 1	2170	1395
Granada Cartuja	775	Cetursa 3	2670	1895
Granada Cartuja	775	Cetursa 5	3020	2245

65  
66  
67  
68  
69  
70

**Supplementary Table S1.** Elevational difference between low and high elevation observatories used in this study

Low vs high elevation observatories	Cetursa 5 3020 masl (n=67)		Cetursa 3 2670 masl (n=113)		Cetursa 1 2170 masl (n=121)		Albergue 2500 masl (n=72)	
	<i>r</i>	<i>p</i>	<i>r</i>	<i>p</i>	<i>r</i>	<i>p</i>	<i>r</i>	<i>p</i>
<b>Madrid 667 masl</b>	0.95920	7.99E-37	0.9671	8.28E-68	0.9641	2.44E-70	0.9598	2.40E-40
<b>Sevilla 8 masl</b>	0.95790	2.14E-36	0.9635	2.24E-65	0.9620	5.85E-69	0.9641	4.86E-42
<b>Gr-Airport 567 masl</b>	0.95430	2.80E-35	0.9655	4.18E-66	0.9626	8.66E-69	0.9591	4.37E-40
<b>Gr-Cartuja 775 masl</b>	0.96887	1.61E-40	0.9769	3.39E-76	0.9751	1.04E-79	0.9680	9.63E-44
<b>Gr-Armilla 687 masl</b>	0.96821	3.11E-40	0.9783	9.84E-78	0.9767	2.01E-81	0.9645	3.22E-42

71  
72  
73  
74  
75  
76  
77

**Supplementary Table S2.** Pearson correlations between MMT from low and high elevation (Sierra Nevada) observatories Data from (Spanish National Weather Agency - AEMet Open Data, 2017;Gonzalez-Hidalgo et al., 2015;Observatorio del cambio global de Sierra Nevada, 2016).The Mann-Kendall test performed using PAST software (Hammer et al., 2001) in the raw data of these variables showed a  $p > 0.05$  pointing towards no linear trends; so, no transformation has been applied.

LDI vs Low elevation observatories	n	Normal correlation		(z-score) <sup>2</sup> correlation		Mann-Kendall	Detrended correlation	
		<i>r</i>	<i>p</i>	<i>r</i>	<i>p</i>	<i>no trend p&gt;0.05</i>	<i>r</i>	<i>p</i>
Sevilla MATA	19	0.9092	7E-08	0.8926	2E-07	0.6746	0.9462	9E-10
Sevilla MSTA	19	0.7198	0.0005	0.2655	0.2719	0.0424	0.3957	0.0935
Madrid MATA	19	0.9074	8E-08	0.8383	7E-06	0.6746	0.9291	9E-09
Madrid MSTA	19	0.7662	0.0001	0.5071	0.0267	0.0863	0.6276	0.0040
Gr-Airport MATA	19	0.7056	0.0007	0.6338	0.0036	0.9998	0.7015	0.0008
Gr-Airport MSTA	19	0.5788	0.0094	0.3381	0.1568	0.1837	0.4247	0.0699
Gr-Cartuja MATA	19	0.7327	0.0004	0.6606	0.0021	0.5756	0.6927	0.0010
Gr-Cartuja MSTA	19	0.6477	0.0027	0.4080	0.0829	0.2939	0.4501	0.0532
Gr-Armilla MATA	19	0.6971	0.0009	0.6853	0.0012	0.8337	0.6934	0.0010
Gr-Armilla MSTA	19	0.6113	0.0054	0.3849	0.1037	0.3446	0.4067	0.0840

78  
79  
80  
81  
82  
83  
84  
85  
86  
87  
88  
89  
90  
91  
92  
93  
94  
95  
96  
97  
98  
99  
100  
101  
102

**Supplementary Table S3.** Pearson correlations (normal and detrended) for the last ~100 years among LdRS shc LDI and temperature time-series of different observatories. Normal correlations show the relationship between long-term trends. Data from (Spanish National Weather Agency - AEMet Open Data, 2017;Gonzalez-Hidalgo et al., 2015;Observatorio del cambio global de Sierra Nevada, 2016). Data were standardised (z-scores), normalised (squares) and a Mann-Kendall trend test was performed using PAST software (Hammer et al., 2001) in order to assess the existence of any trend over time in the data series. Afterwards, data were detrended by fitting a linear regression versus time, and a Pearson correlation was worked out with the residuals. Observatories: Granada city (~600-700 masl and 30km from LdRS), Madrid (Retiro: 667 masl and ~360km from LdRS) and Sevilla (Tablada: 8 masl and ~230km from LdRS but almost similar latitude as LdRS). Mean annual temperature anomaly (MATA) and mean warm season (May-September) temperature anomaly (MSTA) have been tested. MSTA have also been included in the comparison because warm season temperature (May-September) influences algae growth in the studied area (Sánchez-Castillo, 1988;Carrillo et al., 1991). The three time-series available from Granada city only have reliable data from the 1970s onwards (AEMet) (Spanish National Weather Agency - AEMet Open Data, 2017), what is too short for a proper proxy calibration. Longer temperature time-series have been obtained from these Granada series using different approaches to fill in gaps and correct outlying data; i.e. MOTEDAS approach (Gonzalez-Hidalgo et al., 2015). Even though these reconstructed temperatures from Granada for the last ~100 years show a good correlation with LDI (Gr-Airport  $r>0.70$  Gr-Cartuja  $r> 0.69$ ; Gr-Armilla  $r>0.69$   $p<0.001$ ), they are likely biased by the quality of the reconstructed data. Besides, the quality of the records (presence of gaps), this discrepancy between LdRS LDI and Granada record series could be linked to the strong control of Granada basin geomorphology in local rain and temperature patterns at low elevations, with local clouds and frequent thermal inversion phenomena and specific microclimate. Madrid and Sevilla area are not influenced by this effect (Dogniaux, 1994).

$\Delta_{\text{temperature}}$ vs $\Delta_{\text{elevation}}$	$\Delta_{\text{temp}}$ from <i>eq b</i> (global $\Delta_{\text{MMT}}$ )	$\Delta_{\text{temp}}$ from <i>eq d</i> (global $\Delta_{\text{MAT}}$ )	$\Delta_{\text{temp}}$ from <i>eqs f</i> <i>and j</i> (MMT)	$\Delta_{\text{temp}}$ from <i>eqs</i> <i>h and l</i> (MAT)	Real $\Delta_{\text{temp}}$ from MMT	Real $\Delta_{\text{temp}}$ from MAT
<b>Sevilla</b> <b>8 masl</b>	16.24 °C	16.13 °C	16.51 °C	16.20 °C	16.39 °C	16.17 °C
<b>Madrid</b> <b>667 masl</b>	12.36 °C	12.31 °C	12.31 °C	12.08 °C	12.23 °C	12.14 °C

103

104

105

106

107

108

109

110

111

112

**Supplementary Table S4.**  $\Delta_{\text{temperature}}$  between Madrid and Sevilla observatories and Cetursa5 (at the same elevation as LdRS, 3020 masl) worked out using different approaches: 1) using equation from Fig. S1b from the mean values between  $\Delta_{\text{MMT}}$  and  $\Delta_{\text{elev}}$  among all the studied low elevation observatories vs those from high elevation; 2) the same as the previous one but with the  $\Delta_{\text{MAT}}$  (equation from Fig. S1d); 3) using equations from Fig. S1f and j (Sevilla-Madrid respectively) from the mean values between  $\Delta_{\text{MMT}}$  and  $\Delta_{\text{elev}}$  of Sevilla or Madrid observatories respectively vs those from high elevation; 4) the same as the previous one but with the  $\Delta_{\text{MAT}}$  (equations from Fig. S1h and l); 5) real  $\Delta_{\text{MMT}}$  between Sevilla or Madrid observatories and Cetursa 5 observatory (3020 masl); 6) the same as the previous one but with the  $\Delta_{\text{MAT}}$ .

LdRS shc LDI vs		n	Normal correlation		(z-score) <sup>2</sup> correlation		Mann-Kendall <i>no trend</i> <i>p&gt;0.05</i>	Detrended correlation	
			<i>r</i>	<i>p</i>	<i>r</i>	<i>p</i>		<i>r</i>	<i>p</i>
Solar	ΔTSI	32	0.5850	0.0004	-0.0992	0.5891	0.9611	-0.0256	0.8893
	TSI	32	0.5570	0.0009	0.1050	0.5675	0.8711	0.0204	0.9120
Volcanic	NH volcanic aerosol	27	-0.0200	0.9181	-0.0561	0.7724	0.7196	0.0892	0.7724
	Global volcanic aerosol	27	0.0247	0.8987	0.1156	0.5505	0.9085	0.2062	0.5505
	*Global volcanic forcing	27	0.0843	0.7165	0.2099	0.2745	0.4365	0.3001	0.2745
Atmospheric	NAO	30	-0.0281	0.8830	0.3276	0.0772	3E-05	0.0542	0.7760
	AMO	32	0.6097	0.0002	0.4063	0.0210	0.3724	0.3233	0.0711
Green house gases	CO <sub>2</sub> (ppm)	31	0.8328	6E-09	0.7655	5E-07	0.3587	0.7039	0.0000
	NO <sub>2</sub> (ppm)	31	0.8533	1E-09	0.7776	2E-07	0.4646	0.7134	0.0000
	CH <sub>4</sub> (ppm)	31	0.8610	5E-10	0.7346	2E-06	0.5633	0.6493	0.0001
Temperatures	CPS Summer temperatures	32	0.5775	0.0007	0.4406	0.0131	0.2998	0.3991	0.0261
	SST uk37 Gol-Ho1B	32	0.7601	4E-07	0.2386	0.1884	0.8330	0.1626	0.3739
	Global Temperatures (GLSS)	19	0.8898	3E-07	0.6920	0.0010	0.6243	0.7435	0.0003

113  
114  
115  
116  
117  
118  
119  
120  
121  
122  
123  
124  
125  
126  
127  
128  
129  
130  
131  
132  
133  
134  
135  
136  
137

**Supplementary Table S5.** Pearson correlations (normal and detrended) between LdRS short core LDI record and different proxies for solar and volcanic forcing, North Atlantic modes, greenhouse gases, and temperatures. Normal correlations show long-term trends. Data were standardised (z-scores), normalised (squares) and a Mann-Kendall trend test was performed using PAST software (Hammer et al., 2001) in order to assess the existence of any trend over time in the data series. Afterwards, data were detrended by fitting a linear regression versus time, and a Pearson correlation was worked out with the residuals. \*: Note that inverse global volcanic forcing values have been used in order to show the same trends as in Fig. 5.

**Solar Proxies:** ΔTSI, reconstruction of the difference of the total solar irradiance from the value of the PMOD composite series during the solar cycle minimum of the year 1986 CE (1365.57 W m<sup>-2</sup>) (Steinhilber et al., 2009); TSI, total solar Irradiance (Coddington et al., 2016).

**Volcanic proxies:** Annual stratospheric volcanic sulfate aerosol injection for the past 1500 years in the North Hemisphere and worldwide (Gao et al., 2008); global volcanic aerosol forcing (W m<sup>-2</sup>) (Sigl et al., 2015).

**North Atlantic modes:** NAO, North Atlantic Oscillation reconstruction (Trouet et al., 2009); AMO, Atlantic Multidecadal Oscillation reconstruction (Mann et al., 2009).

**Greenhouse gases:** reconstructed concentrations of atmospheric CO<sub>2</sub>, NO<sub>2</sub>, and CH<sub>4</sub> (ppm) (Schmidt et al., 2011).

**Temperatures:** Composite-plus-scaling (CPS) mean summer temperature anomaly reconstruction from tree rings records in Europe with respect to 1974-2003 CE (MSTA °C) (Luterbacher et al., 2016); Alkenone-Sea Surface Temperatures (SST °C) of the core Gol-Ho1B\_KSGC-31 (Gulf of Lion: NW Mediterranean Sea) (Sire et al., 2016), and global land and sea surface (GLSS) mean annual temperature anomalies with respect to 1979-2008 CE (Hansen et al., 2010).



LdRS lge LDI vs		n	Normal correlation		(z-score) <sup>2</sup> correlation		Mann-Kendall	Detrended correlation	
			r	p	r	p	no trend p>0.05	r	p
Solar	* $\Delta^{14}\text{C}$	16	-0.7260	0.0015	0.3949	0.1301	0.1917	0.3810	0.1454
	$\Delta\text{TSI}$	20	0.6916	0.0007	0.3759	0.1024	0.0231	0.3668	0.1117
	TSI	13	0.8779	0.0001	0.1301	0.0863	0.8548	0.5630	0.0451
Volcanic	NH volcanic aerosol	19	-0.0700	0.7758	0.1305	0.5945	0.1595	0.0963	0.6951
	Global volcanic aerosol	19	0.0647	0.7924	0.277	0.2507	0.0478	0.2528	0.2965
	*Global volcanic forcing	16	0.3170	0.2316	0.2549	0.2923	0.1945	0.1358	0.5795
North Atlantic modes	NAO	16	0.0999	0.7127	0.3315	0.2098	0.3004	0.2312	0.3890
	AMO	19	0.6022	0.0064	0.3732	0.1156	0.4841	0.2547	0.2928
Greenhouse gases	CO <sub>2</sub> (ppm)	20	0.7364	0.0002	0.7293	0.0003	0.9741	0.6529	0.0018
	NO <sub>2</sub> (ppm)	20	0.6538	0.0018	0.6846	0.0009	0.9225	0.6102	0.0043
	CH <sub>4</sub> (ppm)	20	0.7285	0.0003	0.7737	0.0001	0.3468	0.7071	0.0005
Temperatures	CPS Summer temperatures	20	0.7071	0.0005	0.4586	0.0420	0.1835	0.3403	0.1421
	SST uk37 Gol-Ho1B	20	0.6097	0.0043	0.7254	0.0003	0.1192	0.6519	0.0018
	SST uk37 TTR-17-1-384B	17	0.2378	0.3581	0.0206	0.9374	0.7731	0.0152	0.9537
	SST TEX86 TTR-17-384B	17	0.4489	0.0707	0.5265	0.0299	0.9671	0.4754	0.0538
	SST uk37 TTR-17-1-436B	18	0.3383	0.1697	0.6774	0.0020	0.3247	0.6989	0.0013
	SST TEX86 TTR-17-436B	18	0.4338	0.0721	0.1767	0.4830	0.4047	0.3449	0.1610

138

139

140

141

142

143

144

145

146

147

148

149

150

151

152

153

154

155

156

157

158

159

160

161

162

**Supplementary Table S6.** Pearson correlations (normal and detrended) between LdRS long core LDI record and different proxies for solar and volcanic forcing, North Atlantic modes, greenhouse gases, and temperatures. Normal correlations show long-term trends. Data were standardised (z-scores), normalised (squares), and a Mann-Kendall trend test was performed using PAST software (Hammer et al., 2001) in order to assess the existence of any trend over time in the data series. Afterwards, data were detrended by fitting a linear regression versus time, and a Pearson correlation was worked out with the residuals. \*: Note that inverse  $\Delta^{14}\text{C}$  and global volcanic forcing values have been used in order to show the same trends as in Fig. 4.

**Solar Proxies:**  $\Delta^{14}\text{C}$  (Reimer et al., 2013);  $\Delta\text{TSI}$ , reconstruction of the difference of the total solar irradiance from the value of the PMOD composite series during the solar cycle minimum of the year 1986 CE ( $1365.57 \text{ W m}^{-2}$ ) (Steinhilber et al., 2009); TSI, total solar irradiance (Coddington et al., 2016).

**Volcanic proxies:** Annual stratospheric volcanic sulfate aerosol injection for the past 1500 years in the North Hemisphere, and worldwide (Gao et al., 2008); global volcanic aerosol forcing ( $\text{W m}^{-2}$ ) (Sigl et al., 2015).

**North Atlantic modes:** NAO, North Atlantic Oscillation reconstruction (Trouet et al., 2009); AMO, Atlantic Multidecadal Oscillation reconstruction (Mann et al., 2009).

**Greenhouse gases:** reconstructed concentrations of atmospheric CO<sub>2</sub>, NO<sub>2</sub>, and CH<sub>4</sub> (ppm) (Schmidt et al., 2011).

**Temperatures:** Composite-plus-scaling (CPS) mean summer temperature anomaly reconstruction from tree rings records in Europe with respect to 1974-2003 CE (MSTA °C) (Luterbacher et al., 2016); Alkenone-Sea Surface Temperatures (SST °C) of the core Gol-Ho1B\_KSGC-31 (Gulf of Lion: NW Mediterranean Sea) (Sicre et al., 2016), alkenone and TEX<sub>86</sub> (from GDGTs) SST records of the cores 384B and 436B in the Alboran Sea (Nieto-Moreno et al., 2013).

Record	Lab Code	Sampling Depth (cm)	Cal yr BP	Yr CE	C <sub>28</sub> 1,13-diol	C <sub>30</sub> 1,15-diol	C <sub>30</sub> 1,13-diol	C <sub>32</sub> 1,15-diol	LDI	Observations
LdRS shc	BECS 1271	0	-58	2008	0.0449	0.0506	0.0807	0.8238	0.2871	
LdRS shc	BECS 1272	0.5	-53	2003	0.0598	0.0590	0.0772	0.8040	0.3012	
LdRS shc	BECS 1273	1	-48	1998	0.0525	0.0620	0.0792	0.8063	0.3202	
LdRS shc	BECS 1274	1.5	-43	1993	0.0539	0.0568	0.0893	0.8000	0.2840	
LdRS shc	BECS 1275	2	-38	1988	0.0721	0.0584	0.1060	0.7635	0.2470	
LdRS shc	BECS 1276	2.5	-33	1983	0.0579	0.0600	0.1032	0.7789	0.2713	
LdRS shc	BECS 1277	3	-28	1978	0.0805	0.0639	0.1043	0.7513	0.2569	
LdRS shc	BECS 1278	3.5	-23	1973	0.0676	0.0601	0.1174	0.7549	0.2452	
LdRS shc	BECS 1279	4	-18	1968	0.0789	0.0611	0.1232	0.7368	0.2320	
LdRS shc	BECS 1280	4.5	-13	1963	0.0863	0.0635	0.1313	0.7189	0.2260	
LdRS shc	BECS 1281	5	-8	1958	0.0801	0.0664	0.1308	0.7227	0.2395	
LdRS shc	BECS 1282	5.5	-3	1953	0.0819	0.0667	0.1341	0.7173	0.2359	
LdRS shc	BECS 1283	6	2	1948	0.0934	0.0564	0.1323	0.7180	0.1999	
LdRS shc	BECS 1284	6.5	9	1941	0.1281	0.0549	0.1492	0.6678	0.1654	
LdRS shc	BECS 1285	7	15	1935	0.1233	0.0563	0.1531	0.6674	0.1693	
LdRS shc	BECS 1286	7.5	22	1928	0.1067	0.0525	0.1389	0.7019	0.1762	
LdRS shc	BECS 1322	8	29	1921	0.1193	0.0485	0.1576	0.6746	0.1492	
LdRS shc	BECS 1323	8.5	35	1915	0.1326	0.0492	0.1742	0.6441	0.1382	
LdRS shc	BECS 1324	9	42	1908	0.1508	0.0489	0.1800	0.6203	0.1288	
LdRS shc	BECS 1325	9.5	49	1901	0.1345	0.0468	0.1782	0.6405	0.1302	
LdRS shc	BECS 1326	10	55	1895	0.1446	0.0475	0.1724	0.6354	0.1304	
LdRS shc	BECS 1327	10.5	62	1888	0.1315	0.0501	0.1645	0.6538	0.1448	
LdRS shc	BECS 1328	11	69	1881	0.1429	0.0508	0.1705	0.6357	0.1396	
LdRS shc	BECS 1329	11.5	75	1875	0.1306	0.0509	0.1713	0.6472	0.1443	
LdRS shc	BECS 1330	12	82	1868	0.1184	0.0524	0.1738	0.6555	0.1520	
LdRS shc	BECS 1331	12.5	89	1861	0.1171	0.0491	0.1879	0.6460	0.1386	
LdRS shc	BECS 1332	13	95	1855	0.1132	0.0515	0.1961	0.6393	0.1426	
LdRS shc	BECS 1333	13.5	102	1848	0.0918	0.0573	0.1705	0.6804	0.1792	
LdRS shc	BECS 1334	14	109	1841	0.0852	0.0607	0.1660	0.6881	0.1945	
LdRS shc	BECS 1335	14.5	115	1835	0.0741	0.0621	0.1597	0.7041	0.2098	
LdRS shc	BECS 1336	15	122	1828	0.0755	0.0625	0.1649	0.6972	0.2064	
LdRS shc	BECS 1337	15.5	129	1821	0.0746	0.0596	0.1746	0.6912	0.1930	
LdRS lgc	GMOL 1886	0	-56	2006	0.0708	0.0703	0.1041	0.7548	0.2866	
LdRS lgc	GMOL 1887	1	-41	1991	0.0630	0.0742	0.1032	0.7596	0.3087	
LdRS lgc	GMOL 1888	2	-26	1976	0.1016	0.0765	0.1399	0.6820	0.2405	
LdRS lgc	GMOL 1889	3	-11	1961	0.1264	0.0796	0.1585	0.6356	0.2183	
LdRS lgc	GMOL 1890	4	4	1946	0.2003	0.0647	0.1724	0.5625	0.1479	
LdRS lgc	GMOL 1891	5	20	1931	0.1477	0.0620	0.1338	0.6566	0.1805	
LdRS lgc	GMOL 1892	6	35	1915	0.2733	0.0516	0.1666	0.5085	0.1049	
LdRS lgc	GMOL 1893	7	50	1900	0.2283	0.0623	0.1911	0.5183	0.1294	

LdRS lgc	GMOL 1895	9	80	1870	0.1334	0.0586	0.2003	0.6076	0.1495	
LdRS lgc	GMOL 1896	10.5	106	1844	0.1093	0.0644	0.1937	0.6326	0.1752	
LdRS lgc	GMOL 1897	11.5	169	1781	0.4216	0.0551	0.1927	0.3307	0.0823	
LdRS lgc	GMOL 1898	12.5	263	1687	0.4132	0.0325	0.1851	0.3692	0.0515	
LdRS lgc	GMOL 1899	13.5	343	1607	0.3809	0.0578	0.3068	0.2545	0.0775	
LdRS lgc	GMOL 1900	14.5	437	1513	0.3456	0.0603	0.2828	0.3113	0.0876	
LdRS lgc	GMOL 1901	15.5	547	1403	0.2543	0.0923	0.2568	0.3966	0.1529	
LdRS lgc	GMOL 1902	16.5	738	1212	0.1883	0.0859	0.1677	0.5581	0.1944	
LdRS lgc	GMOL 1903	17.5	876	1074	0.2448	0.0732	0.2179	0.4641	0.1366	
LdRS lgc	GMOL 1904	18.5	1022	928	0.1784	0.1070	0.1753	0.5394	0.2322	
LdRS lgc	GMOL 1905	19.5								below quantification limits
LdRS lgc	GMOL 1906	20.5	1404	546	0.2144	0.0938	0.2602	0.4316	0.1650	
LdRS lgc	GMOL 1907	21.5	1556	394	0.2210	0.0878	0.2165	0.4746	0.1672	

163

164

165

166

167

168

169

170

171

172

173

174

175

176

177

178

179

180

181

182

183

184

185

186

187

188

189

190

191

192

193

194

195

196

**Supplementary Table S7.** Fractional abundances of C<sub>28</sub> 1,13-diol, C<sub>30</sub> 1,13-diol, C<sub>30</sub> 1,15-diol, and C<sub>32</sub> 1,15-diol in the different samples of both LdRS cores (LdRS shc and LdRS lgc) along with the obtained Long Chain Diol Index (LDI), according to Rampen et al. (2012).

### Supplementary References

Barea-Arco, J., Pérez-Martínez, C., and Morales-Baquero, R.: Evidence of a mutualistic relationship between an algal epibiont and its host, *Daphnia pulicaria*, *Limnology and Oceanography*, 46, 871-881, 2001.

Carrillo, P., Cruz-Pizarro, L., and Sánchez Castillo, P. M.: Aportación al conocimiento del ciclo biológico de *Chromulina nevadensis*, *Acta Botánica Malacitana*, 16, 19-26, 1991.

Coddington, O., Lean, J. L., Pilewskie, P., Snow, M., and Lindholm, D.: A Solar Irradiance Climate Data Record, *Bulletin of the American Meteorological Society*, 97, 1265-1282, 10.1175/bams-d-14-00265.1, 2016.

Dogniaux, R.: Prediction of Solar Radiation in Areas with a Specific Microclimate, *Prediction of Solar Radiation in Areas with a Specific Microclimate*, 108 pp., 1994.

Gao, C., Robock, A., and Ammann, C.: Volcanic forcing of climate over the past 1500 years: An improved ice core-based index for climate models, *Journal of Geophysical Research: Atmospheres*, 113, doi:10.1029/2008JD010239, 2008.

Gonzalez-Hidalgo, J. C., Peña-Angulo, D., Brunetti, M., and Cortesi, N.: MOTEDAS: a new monthly temperature database for mainland Spain and the trend in temperature (1951–2010), *International Journal of Climatology*, 35, 4444-4463, 10.1002/joc.4298, 2015.

Hammer, Ø., Harper, D. A. T., and Ryan, P. D.: PAST: Paleontological statistics software package for education and data analysis, *Palaeontologia Electronica* 4, 9, 2001.

Hansen, J., Ruedy, R., Sato, M., and Lo, K.: Global Surface Temperature Change, *Reviews of Geophysics*, 48, 10.1029/2010RG000345, 2010.

Luterbacher, J., Werner, J. P., Smerdon, J. E., Fernández-Donado, L., González-Rouco, F. J., Barriopedro, D., Ljungqvist, F. C., Büntgen, U., Zorita, E., Wagner, S., Esper, J., McCarroll, D., Toreti, A., Frank, D., Jungclauss, J. H., Barriendos, M., Bertolin, C., Bothe, O., Brázdil, R., Camuffo, D., Dobrovolný, P., Gagen, M., García-Bustamante,

197 E., Ge, Q., Gómez-Navarro, J. J., Guiot, J., Hao, Z., Hegerl, G. C., Holmgren, K.,  
198 Klimenko, V. V., Martín-Chivelet, J., Pfister, C., Roberts, N., Schindler, A., Schurer,  
199 A., Solomina, O., von Gunten, L., Wahl, E., Wanner, H., Wetter, O., Xoplaki, E.,  
200 Yuan, N., Zanchettin, D., Zhang, H., and Zerefos, C.: European summer temperatures  
201 since Roman times, *Environmental Research Letters*, 11, 024001, citeulike-article-  
202 id:14089240  
203 doi: 10.1088/1748-9326/11/2/024001, 2016.

204 Mann, M. E., Zhang, Z., Rutherford, S., Bradley, R. S., Hughes, M. K., Shindell, D.,  
205 Ammann, C., Faluvegi, G., and Ni, F.: Global Signatures and Dynamical Origins of  
206 the Little Ice Age and Medieval Climate Anomaly, *Science*, 326, 1256-1260,  
207 10.1126/science.1177303, 2009.

208 Nieto-Moreno, V., Martínez-Ruiz, F., Willmott, V., García-Orellana, J., Masqué, P., and  
209 Sinninghe Damsté, J. S.: Climate conditions in the westernmost Mediterranean over  
210 the last two millennia: An integrated biomarker approach, *Organic Geochemistry*, 55,  
211 1-10, <https://doi.org/10.1016/j.orggeochem.2012.11.001>, 2013.

212 Organization, I. C. A.: Manual of the ICAO standard atmosphere : extended to 80  
213 kilometres (262 500 feet) (Third ed.), International Civil Aviation Organization,  
214 Montreal, Quebec, 1993.

215 Reimer, P. J., Bard, E., Bayliss, A., Beck, J. W., Blackwell, P. G., Ramsey, C. B., Buck,  
216 C. E., Cheng, H., Edwards, R. L., Friedrich, M., Grootes, P. M., Guilderson, T. P.,  
217 Haflidason, H., Hajdas, I., Hatté, C., Heaton, T. J., Hoffmann, D. L., Hogg, A. G.,  
218 Hughen, K. A., Kaiser, K. F., Kromer, B., Manning, S. W., Niu, M., Reimer, R. W.,  
219 Richards, D. A., Scott, E. M., Southon, J. R., Staff, R. A., Turney, C. S. M., and van  
220 der Plicht, J.: IntCal13 and Marine13 Radiocarbon Age Calibration Curves 0–50,000  
221 Years cal BP, *Radiocarbon*, 55, 1869-1887,  
222 [https://doi.org/10.2458/azu\\_js\\_rc.55.16947](https://doi.org/10.2458/azu_js_rc.55.16947), 2013.

223 Sánchez-Castillo, P. M.: Algas de las lagunas de alta montaña de Sierra Nevada (Granada,  
224 España), *Acta Botánica Malacitana*, 13, 21 -52, 1988.

225 Sandgren, C. D.: The ecology of chrysophyte flagellates: their growth and perennation  
226 strategies as freshwater phytoplankton, in: *Growth and reproductive strategies of*  
227 *freshwater phytoplankton*, edited by: C.D., S., Cambridge University Press,  
228 Cambridge, 9-104, 1988.

229 Schmidt, G. A., Jungclauss, J. H., Ammann, C. M., Bard, E., Braconnot, P., Crowley, T.  
230 J., Delaygue, G., Joos, F., Krivova, N. A., Muscheler, R., Otto-Bliesner, B. L.,  
231 Pongratz, J., Shindell, D. T., Solanki, S. K., Steinhilber, F., and Vieira, L. E. A.:  
232 Climate forcing reconstructions for use in PMIP simulations of the last millennium  
233 (v1.0), *Geosci. Model Dev.*, 4, 33-45, 10.5194/gmd-4-33-2011, 2011.

234 Sicre, M.-A., Jalali, B., Martrat, B., Schmidt, S., Bassetti, M.-A., and Kallel, N.: Sea  
235 surface temperature variability in the North Western Mediterranean Sea (Gulf of Lion)  
236 during the Common Era, *Earth and Planetary Science Letters*, 456, 124-133,  
237 <https://doi.org/10.1016/j.epsl.2016.09.032>, 2016.

238 Sigl, M., Winstrup, M., McConnell, J. R., Welten, K. C., Plunkett, G., Ludlow, F.,  
239 Buntgen, U., Caffee, M., Chellman, N., Dahl-Jensen, D., Fischer, H., Kipfstuhl, S.,  
240 Kostick, C., Maselli, O. J., Mekhaldi, F., Mulvaney, R., Muscheler, R., Pasteris, D. R.,  
241 Pilcher, J. R., Salzer, M., Schupbach, S., Steffensen, J. P., Vinther, B. M., and  
242 Woodruff, T. E.: Timing and climate forcing of volcanic eruptions for the past 2,500  
243 years, *Nature*, 523, 543-549, 10.1038/nature14565  
244 [http://www.nature.com/nature/journal/v523/n7562/abs/nature14565.html#supplementar](http://www.nature.com/nature/journal/v523/n7562/abs/nature14565.html#supplementary-information)  
245 [y-information](http://www.nature.com/nature/journal/v523/n7562/abs/nature14565.html#supplementary-information), 2015.

246 Steinhilber, F., Beer, J., and Fröhlich, C.: Total solar irradiance during the Holocene,  
247 Geophysical Research Letters, 36, 10.1029/2009GL040142, 2009.  
248 Trouet, V., Esper, J., Graham, N. E., Baker, A., Scourse, J. D., and Frank, D. C.: Persistent  
249 Positive North Atlantic Oscillation Mode Dominated the Medieval Climate Anomaly,  
250 Science, 324, 78-80, 10.1126/science.1166349, 2009.  
251