

## ***Interactive comment on “The last interglacial (MIS 5e) cycle at Little Bahama Bank: A history of climate and sea-level changes” by Anastasia Zhuravleva and Henning A. Bauch***

### **Anonymous Referee #2**

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**SUMMARY:** Zhuravleva and Bauch present a detailed consideration of the climate evolution of the Last Interglacial (LIG) for a core site on the Little Bahama Bank (LBB) using faunal assemblage and scanning XRF techniques. The high resolution faunal assemblages nicely resolve hydrographic oscillations at the site for the LIG reflecting both the insolation driven and AMOC modulated migration of the ITCZ for this region.

I would recommend the following amendments/clarifications: (a) change of title to better reflect the content of the paper; (b) removal or at the very least restructuring of the discussion of sea level. This section could be significantly trimmed and simplified (no new insights offered but a nice corroboration). Alternatively, if the authors wish to retain

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the sea-level discussion, then discussion of other sea level evidence from the region, glacio-isostatic adjustment (GIA) processes etc. are needed. (c) clearer discussion of the teleconnections between N Atlantic oceanic changes (i.e., variation in AMOC), the migration of the ITCZ and surface hydrographic change at MD99-2202.

**GENERAL COMMENTS:** The manuscript, in general, reads well. However, the structure and focus of the paper requires further thought. A clear statement of the research questions was missing and is reflected in the general tone of the introduction (and the manuscript generally).

#### **1. Title**

I found this to be somewhat misleading. The data in Zhuravleva and Bauch is not a sea-level record per se, rather a record of increased aragonite supply to the core site during interglacials, with these intervals of increased aragonite production/supply likely corresponding to < -6 m relative sea level (RSL) due to the generally shallow nature of Little Bahama Bank (i.e., you can infer periods of < -6 m relative sea level). This work nicely corroborates the Lantzsich et al., 2007 and Chabaud et al., 2016 studies but isn't a sea-level story. What is new and interesting the palaeoceanographic evolution of the Last Interglacial (LIG) at the site, and the interplay of interglacial climate (movement of the ITCZ etc.). I would suggest changing the title to better reflect this.

#### **2. Sea level**

This section requires some restructuring to help the reader. The definition of the “flooding interval” (and corresponding relative sea level, < -6 m) is key to this section of the manuscript but I struggled to clearly follow the logic of how you defined the flooding interval using your records and why a -6 m RSL for this interval was appropriate. The connection between the flooding interval and inferred RSL of < -6 m was found almost at the end of the section (line 222 to 226) when it should be at the start. All the information is there but the reader has to work hard to follow the argument.

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Perhaps something along the lines of;

modern LBB lagoon is shallow, with an average water depth < 6 m (Williams, 1985); tectonic stability of the region (refs needed);

during the Llg, increasing RSL at the site floods the generally shallow bank and increases the area for aragonite production (i.e., the carbonate shedding model, Droxler and Schlager, 1985; Schlager et al., 1994);

Conversely, during glacial intervals, the top is exposed which limits the production and export of aragonite;

As such, we define the flooding interval (and inferred <-6 m RSL) is defined by an increase in the sedimentation rate, increase in wt % aragonite, increased Sr/Ca ratio, increase % Globigerinoides/decrease in numbers of *G. menardii*.

This could then usefully be followed with your discussion of very high values of Sr/Ca due to increased saltwater (lines 192 to 211). Perhaps shade these 'problematic' Sr/Ca intervals in subsequent figures? You should also note the truncation of the Sr/Ca record in caption of Figure 4.

I would suggest confining discussion of sea level in this section to that suggested above. If you wish to make more of the sea level story, then greater consideration of other Bahamas sea-level records, as well as those from the wider area is needed. For example, the +6m notch on Little Sale Cay (LLB) (Neumann and Hearty, 1996), other geomorphological records (e.g., Hearty and Kindler, 1995; Neumann and Hearty, 1996), the elevated Last Interglacial (Llg) coral records of Chen et al 1991, Hearty et al., 2007, Thompson et al., 2011 and the regionally extensive erosional surface that is suggestive of a sea-level oscillation within the Llg (Bahamas, Florida and Yucatan; Chen et al 1991, Hearty et al., 2007, Blanchon et al., 2009; Thompson et al., 2011).

How does the timing of the highstand from the coral/other records from the Bahamas compare to the timing of the interval of enhanced aragonite production (and inferred

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sea levels < -6 m)?

How do changes in hydrography (variations in faunal assemblages) at the site compare to the timing of the Bahamas Llg highstand? The broad correspondence between climate ( $\delta^{18}O_{G.Ruber}$ ) and relative sea level (RSL) (weight % aragonite) is hinted at in lines 138 to 141 but could be developed further if you wish to keep the sea-level discussion.

Any discussion of the Llg highstand in a general sense (i.e., the eustatic record) (lines 227 to 231) and Bahamas RSL will need to consider glacio-isostatic (GIA) processes, given the intermediate location of the site on the peripheral bulge of the former Laurentide Ice Sheets. There will be a regional expression of the Llg highstand; the Bahamas would "see" a "late" Llg highstand compared to eustatic sea level (e.g., Figure 6 in Stirling et al., 1998). There seems to be good correspondence between the age of your "flooding interval" at the site (i.e., sea level < -6 m) and the predictions of RSL (Stirling et al., 1998, their Figure 6).

Given that the records presented are not strictly a sea level record, rather incidence of increased aragonite production/export, and seems to corroborate previous studies rather than adding anything new, I would confine this section to just a brief consideration of the timing of your "flooding interval".

### 3. Palaeoceanographic reconstruction

This section is much more coherent and well written. I would recommend this as the focus of the manuscript.

The discussion, while nicely documenting the site/regional changes during the Llg, was lacking in consideration of the mechanisms. This section would be strengthened by a clearer exposition of the mechanisms linking ITCZ position, insolation (precession and the migration of the ITCZ to the warming hemisphere) and AMOC (i.e., modification of the thermal condition at the surface, due to interactions with the ocean, that in

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turn act to drive atmospheric circulation). This is well documented for the last deglaciation and glacial period (and in modelling studies), where N. hemisphere extra-tropical cooling (brought about by variations in the AMOC, forced by freshwater inputs) lead to an interhemispheric thermal gradient and a southward shift in the ITCZ (e.g., review of Chiang and Friedman, 2012 or Schneider et al., 2014). This would help the reader to place the different records (Cariaco, MD99-2202 and Site 1063) within a broader climatological context. Again, all the 'threads' of the story are there, it just needs a stranger framework. For example, I found the correspondence between the % *G. ruber* and *G. sacculifer* and the XRF Mo count of the Caricao Basin striking (demonstrating the clear record of ITCZ shifts at the LBB) but the link to N. Atlantic surface density changes (AMOC slowdown with surface freshening e.g., Galaasen et al., 2014 etc.) and positive the  $\delta^{18}\text{O}_{\text{G.ruber}}$  excursion and faunal changes at MD99-2202 and southward migration of the ITCZ (Cariaco Mo, MD99-2202 decrease in % *Globigerinoides*) weak. A short introductory paragraph should fix this.

A southward shift of the ITCZ, strengthens of the trade winds increases the eolian input from the Sahara, resulting in reduced Al/Ti in Cariaco. These episodes of decreased Al/Ti ratios in Cariaco correspond to elevated salinities in the Caribbean (e.g., Yarincik et al 2000). I assume there is clear correspondence between the Cariaco Al/Ti and Mo records and hence to your % of tropical species? Do you see an increase in iron (with increased dust transport) in your record during the positive the  $\delta^{18}\text{O}_{\text{G.ruber}}$  excursion  $\sim 127$  ka? (plotting this on a log scale for the Llg might help). Dust inputs are probably better reflected in the XRF core scanning Ti record, given that your Fe inputs could change with a number of factors. It would be interesting to compare to your faunal assemblages and a calculated  $\delta^{18}\text{O}_{\text{seawater}}$  for MD99-2202. Additionally, is there any correspondence to the dated palaeosols on the Bahamas (Muhs et al., 2007)?

Given you have faunal assemblage data, could you calculate a transfer function/MAT sea surface temperature? From this you could then calculate  $\delta^{18}\text{O}_{\text{seawater}}$  at the site to think about density changes during the Llg.

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lines 372 to 381 – I found this paragraph to be highly speculative and not well supported by your data (I struggled to see the change in the sedimentological properties in your figures, even with the help of the white arrows). I would suggest removing this section.

#### TECHNICAL CORRECTIONS

1. Referencing: greater care needed with referencing. Please check manuscript. For example, the depth of submersion of the LBB (and origin of the -6 m quoted often in the paper) is Neumann and Land, 1975. Similarly the carbonate shedding model (upon which the inferred sea-level story is based) comes from Droxler and Schlager, 1985; Schlager et al., 1994

line 124 – what potential biases are you refereeing to? Please give appropriate references for these.

line 242 – what is the derivation of “minimal ice volume interval” and it's reference?

line 332 - reference required for the “full resumption of the AMOC... only by  $\sim 124$  ka”

2. Other: line 64 – capitalisation of “intertropical convergence zone”

line 133 – unit of measurement missing, add “m”.

line 330 - please clarify or add examples of the “additional forcing” or add “as discussed below”.

line 324-325 - Clarification of the age of the cooling/increase salinity event  $\sim 127$  ka is needed. The  $\sim 127$  ka age for this event is derived from your age model, whereas the U-series ages for the correlative event in core 152JPC (Bahamas, Slowey et al., 1996), dated (in duplicate) above and below the event to  $\sim 121 \pm 3$  ka and  $125.6$  ka (mean of  $127 \pm 4.8$  and  $124.1 \pm 5.1$ ) respectively. (Note, the relatively large age uncertainties associated with these U-series ages)

3. Figures: typo - Axis label in Figure 4D should read “# *G. menardii*” rather than “# *G. menradii*” I would remove the sea level records (H), or if you choose to retain these and

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your discussion of sea level, then you should remove the dashed blue line “RSL above -6 m” from the lowermost panel (H).

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