Middle and Late Pleistocene climate and continentality inferred from ice wedges at Batagay megaslump in the Northern Hemisphere's most continental region, Yana Highlands, interior Yakutia

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Abstract

Ice wedges in the Yana Highlands of interior Yakutia—the most continental region of the Northern Hemisphere—were investigated to elucidate the winter climate and continentality during the Middle to Late Pleistocene. The Batagay megaslump exposes ice wedges and composite wedges that were sampled from three cryostratigraphic units: the lower sand unit of Marine Isotope Stage (MIS) 6 age, the upper Ice Complex (Yedoma) and the upper sand unit (both MIS3 to MIS2). A terrace of the nearby Adycha River provides a Late Holocene (MIS1) ice wedge that serves as a modern endmember for analysis. Stable-isotope values of ice wedges in the MIS3 upper Yedoma Ice Complex at Batagay are more depleted (mean $\delta^{18}O$ about –35‰) than those from 17 other ice-wedge sites across coastal and central Yakutia. This observation points to lower winter temperatures and, therefore, higher continentality in the Yana Highlands during MIS3. Likewise, more depleted isotope values compared to other sites in Yakutia are found in Holocene wedge ice (mean $\delta^{18}O$ about –29‰). Ice-wedge isotopic signatures of the MIS6 lower sand unit (mean $\delta^{18}O$ about –33‰) and of the MIS3-2 upper sand unit (mean $\delta^{18}O$ from about -33 to -30‰) are less distinctive regionally and preserve traces of fast formation in rapidly accumulating sand sheets and of post-depositional fractionation.

1 Introduction

Interior Yakutia is currently the most continental region of the Northern Hemisphere. At Verkhoyansk, the pole of cold of the Northern Hemisphere, mean monthly air temperature ranges exceed 60°C and absolute temperatures ranges can exceed 100°C

(Lydolph, 1985). Mean annual precipitation is only 155 mm (Lydolph, 1985). Scarce stable isotope data indicate that modern winter precipitation in this region has the most depleted isotopic composition in Siberia (Kurita et al., 2004; Kurita et al., 2005) according to the "continental effect" in stable isotope fractionation, i.e. the farthest distance from the source region (Dansgaard, 1964). However, warm summers enrich stable isotopes in summer precipitation at interior sites compared to coastal sites in the North, which also reflects the strong continentality of interior Yakutia. During Late Pleistocene cold stages, when lower sea level exposed larger areas of continental shelves, interior Yakutia probably experienced even stronger continentality, with further decreased annual precipitation and larger annual temperature amplitudes. Such increased amplitudes must have resulted from either lower winter temperatures or higher summer temperatures or a combination of both.

Although biological proxies such as pollen or plant macro remains can be used to reconstruct summer climate, winter conditions in this region may be inferred only from ice wedges, which are widespread in Middle and Late Pleistocene permafrost deposits. Ice wedges provide winter temperature information due to their specific seasonality, with frost cracking in winter and crack infilling by snowmelt in spring (Opel et al., 2018). They integrate the stable isotope composition of cold season precipitation (Meyer et al., 2015), which in high northern latitudes is mainly a function of temperature (Dansgaard, 1964). However, a knowledge gap exists about the winter palaeoclimate of interior Yakutia, because ice-wedge studies in the Yana Highlands—between the Verkhoyansky and Chersky ranges—began only recently (Vasil'chuk et al., 2017). In contrast, other east Siberian sites have been extensively studied over recent decades, including the coastal lowlands to the north (e.g. Meyer et al., 2002a; Meyer et al., 2002b; Opel et al., 2017b), the Kolyma region to the east (e.g. Vasil'chuk et al., 2001; Vasil'chuk, 2013; Vasil'chuk and Vasil'chuk, 2014) as well as central Yakutia to the south (e.g. Popp et al., 2006) (Figure 1). Here we report cryostratigraphic field observations of ice wedges and composite wedges made during a reconnaissance expedition from 27 July to 5 August 2017 together with stable-isotope data from the Batagay megaslump and Adycha River floodplain (Figure 1). New radiocarbon ages from organic remains in ice wedges and host sediments from the former help constrain the late Quaternary chronology. The observations and ages elucidate the history of wedge development, past winter temperatures and continentality since MIS 6.

2 Study site characteristics, stratigraphy and chronology

Neither study site has been glaciated during at least the last 50,000 years, and the Batagay megaslump provides access to permafrost formations since the Middle Pleistocene (Ashastina et al., 2017; Murton et al., 2017). The present mean annual air temperature at Batagay is -15.4°C (mean of the coldest month: -47.7°C, mean of the warmest month: 15.4°C) and the mean annual precipitation is 170 to 200 mm (Ivanova, 2003). Prevailing wind directions are southwest in winter (January) and north or northeast in summer (July) (Murton et al., 2017). The mean annual temperature of the permafrost is -5.5 to -8.0°C and the active-layer depth between 0.2-0.4 m beneath forest and moss covers and 0.4 to 1.2 m beneath open sites (Ivanova, 2003).

The **Batagay megaslump** (67.58°N, 134.77°E), close to the city of Batagay on the right bank of the Yana, is the world's largest known thaw slump (Figure 2) (Ashastina et al., 2017; Murton et al., 2017). It is located on a hillslope in the taiga

landscape and exposes a sequence of permafrost deposits at least 60 m thick, with underlying bedrock cropping out in places (Kunitsky et al., 2013). The slump has formed during recent decades, with increasing headwall retreat rates of up to 30 m per year. In 2016 it reached a width of 840 m and a total area of >70 ha (Günther et al., 2016).

The slump exposes Holocene and Pleistocene permafrost formations ranging in age from MIS 6 (or older) to MIS 1 (Ashastina et al., 2017; Murton et al., 2017; Ashastina et al., 2018). Above bedrock and a basal diamicton, four major cryostratigraphic units contain ice wedges (Figure 3). The lowest unit is a pebbly dark sand 3-7 m thick, with ice wedges at least 2-3 m high and 1 m wide truncated by a thaw unconformity. This **lower Ice Complex** has neither been dated nor sampled previously for wedge ice.

Above it is the **lower sand** unit, reaching 20 to 30 m in thickness and composed of yellowish pore-ice cemented (fine) sand with grey horizontal bands. It contains tall, narrow syngenetic ice wedges up to 0.5 m wide. The middle part of the unit was dated to 142.8±25.3 and >123.2 ka by optically-stimulated luminescence (OSL) and to 210.0±23.0 ka by infrared stimulated luminescence (IRSL) (Ashastina et al., 2017), pointing to its deposition during MIS6 or even MIS7. A thin organic-rich layer of silty sand with abundant woody debris (in lenses up to 3 m thick) overlies a distinct erosional surface near the top of the lower sand unit. A wood fragment yielded a conventional radiocarbon age of 49.32±3.15 ^{14}C ka BP, outside of the range of calibration (Murton et al., 2017), though palaeoecological analysis points to a Last Interglacial age for the organic material (Ashastina et al., 2018).

The overlying **upper Ice Complex** is 20-40 m thick and dominated by huge syngenetic ice wedges up to few metres wide and at least several metres high within silty and sandy deposits. Finite radiocarbon ages from 49.0±2.0 ^{14}C ka BP to 12.66±0.05 ^{14}C ka BP (Ashastina et al., 2017) indicate sedimentation gaps or erosive events and reveal a MIS3 to late MIS2 age for the upper part of the upper Ice Complex. This unit has numerous chronostratigraphic analogies in the coastal lowlands of northern Yakutia, i.e. the Yedoma Ice Complex (Schirrmeister et al., 2011b; Murton et al., 2015). The upper Ice Complex forms the highest Pleistocene stratigraphic unit in the upper central part (upslope) of the slump headwall, whereas downslope towards the slump mouth it is overlain and partially grades into the **upper sand unit** (Figure 3).

The upper sand is up to about 20 m thick and consists of pore-ice cemented brown to grey sand with narrow (≤ 0.5 m wide) syngenetic ice wedges and composite wedges. Radiocarbon ages between 36.30±0.70 ^{14}C ka BP and 26.18±0.22 ^{14}C ka BP (Ashastina et al., 2017; Murton et al., 2017) indicate deposition during MIS3 and MIS2, similar to the upper Ice Complex, but in a more complex spatiotemporal pattern, likely due to varying palaeotopography and sediment supply.

A near-surface layer 1-1.5 m thick of brown sand and modern soil covers the sequence. This layer was dated to 0.295±0.03 ka ^{14}C BP (Ashastina et al., 2017). Distinct (thermo-)erosional contacts can be found between the lower Ice Complex and the lower sand, the lower sand and the upper Ice Complex and below the near-surface layer.

According to Murton et al. (2017) exposed floodplains of proximal rivers such as the Batagay and Yana, within 2 and 10 km, respectively, of the slump are the assumed major source of the sediments exposed in the headwall, which implies upslope directed transport by wind. Periglacial, proluvial, or nival processes on nearby hillslopes may also have contributed to sediment supply (Ashastina et al., 2017).

The second study site, chosen as a Late Holocene reference, is the actively eroding Holocene to recent bank of a small cut-off channel of the **Adycha River**, about 40 km east of the Batagay megaslump (Figure 1). The Adycha site is located about 3 km upstream of the 60-65 m high and 1.2 km long river bluff at Ulakhan Sullar (Kaplina et al., 1983; Sher et al., 2011; Germonpre et al., 2017). Organic-rich silty sands of alluvial origin are exposed in the Adycha riverbank section.

3 Material and Methods

3.1 Fieldwork

Cryostratigraphic observations of ice wedges, composite wedges, unit contacts and sediments at the Batagay megaslump and a terrace of the Adycha River provided a framework with which we selected samples. A total of six ice wedges and composite wedges from the lower sand, the upper Ice Complex and the upper sand were described in detail and sampled at two sections of the slump (Figure 2, Figure 4, Figure 5). The lower Ice Complex was inaccessible due to dangerous outcrop conditions, and no wedges were observed in the near-surface layer. At the Adycha River, we sampled a single ice wedge. Ice samples about 4 cm wide, mostly in horizontal profiles, were obtained by chain saw. Details about the studied ice and composite wedges are given in Table S1. The ice samples were melted on site in sample bags, and the meltwater filled up 30 ml PE bottles that were then then tightly closed and stored cool until stable-isotope analysis.

In addition, water from small streams draining the slump was sampled in several parts of the slump floor and the main outflow stream (n=7) as well as from summer precipitation (n=4). Supernatant water from selected Batagay sediment samples (n=9) was taken for stable-isotope analysis of intrasedimental pore and segregated ice.

3.2 Stable-isotope analysis

The stable oxygen ($\delta^{18}O$) and hydrogen (δD) isotope ratios of all samples were determined at the Stable Isotope Laboratory of the Alfred Wegener Institute Helmholtz Centre for Polar and Marine Research in Potsdam, using a Finnigan MAT Delta S mass spectrometer with an analytical precision of better than ±0.1‰ for $\delta^{18}O$ and ±0.8‰ for δD (Meyer et al., 2000). The isotopic composition is expressed in δ per mille values (‰) relative to the V-SMOW standard. The deuterium excess d was calculated as $d = \delta D - 8 \times \delta^{18}O$ (Dansgaard, 1964). In some cases, lateral ice-wedge samples with isotopic values distinctly different from the wedge centres and similar to host sediments due to exchange processes between wedge ice and host sediments were excluded from further analysis.

3.3 Radiocarbon dating

Organic material from ice wedges and host sediments was radiocarbon dated at the MICADAS ^{14}C dating facility of the Alfred Wegener Institute Helmholtz Centre for Polar and Marine Research in Bremerhaven. The organic material was subjected to chemical leaching and cleaning following the ABA (acid-base-acid) procedure. Material was first submerged in 1 M HCl (for 30 min at 60 °C) to remove carbonate contamination. Following rinsing with Milli-Q water, samples were submerged in 1 M

NaOH (for 30 min at 60 °C) to leach out humic acids. The base leaching was repeated for a minimum of 7 times or until no coloring of the solution was evident. Afterwards the material was subjected to a final acid treatment (1 M HCl; 30 min at 60 °C), rinsed to neutral with Milli-Q and dried for a minimum of 12 h at 60 °C. Milli-Q used for cleaning and acid/base solutions was pre-cleaned by liquid-liquid extraction with dichloromethane to remove remaining organic contamination.

5   Samples were packed in tin capsules and combusted individually using an Elementar vario ISOTOPE EA (Elemental Analyzer). If samples were smaller than approximately 200 µgC, the CO_2 produced by combustion with the EA was directly injected into the hybrid ion source of the Ionplus MICADAS using the gas interface system GIS (Fahrni et al., 2013). For larger samples, the CO_2 produced was graphitized using the Ionplus AGE3 system (Automated Graphitization System; (Wacker et al., 2010c). The radiocarbon content of the samples was determined alongside reference standards (oxalic acid;

10   NIST 4990c) and blanks (phthalic anhydride; Sigma-Aldrich 320065) using the Ionplus MICADAS dating system (Synal et al., 2007; Wacker et al., 2010a), and blank correction and standard normalization were performed using the BATS software (Wacker et al., 2010b). Results are reported as $F^{14}C$ (Reimer et al., 2004) and conventional radiocarbon (^{14}C) ages in year BP. Several analyses yielded lower ^{14}C intensities than the blanks included in the respective sequence. In these cases, results are reported as $F^{14}C_{sample} < F^{14}C_{blanks}$. Conventional ^{14}C ages were calibrated using Oxcal 4.3 (Bronk Ramsey, 2009) based on the

15   IntCAL13 dataset (Reimer et al., 2013).

4 Results

4.1 Batagay megaslump

4.1.1 Field observations of ice wedges

The present study confirms cryostratigraphic observations about ice wedges and composite wedges from earlier studies

20   (Ashastina et al., 2017; Murton et al., 2017) and provides new observations about wedge ice relationships across contacts between the cryostratigraphic units, which are outlines below.

In the lower Ice Complex, ice wedges are truncated by a thaw unconformity, though the toes of some narrow syngenetic composite wedges in the overlying lower sand unit extend down across the contact (Figure S1). The basal contact of the lower Ice Complex was not observed.

25   In the lower and upper sand units, narrow syngenetic wedges vary in apparent width from a few centimetres to about 0.5 m. Width commonly varies irregularly with height along individual wedges, sometimes gradually, sometimes abruptly. Wedge height varies from a few metres to at least 12 m. Wedges tend to be oriented approximately at right angle to colour bands in the sand, such that wedges in the upper sand unit—whose bands dip downslope, parallel to the ground surface—are characteristically subvertical, with their tops inclined downslope (Figure S2). The wedge infills grade between end members

30   of ice-wedge ice and icy sand wedges, with slightly sandy ice-wedge ice (i.e. composite wedges) the most common type. A single example of clear ice sharply overlying the shoulder of a composite wedge in the lower sand unit (B17-IW1, Figure 4)

contained brown plant remains and round, few-mm diameter organic bodies identified as hare droppings. This ice lacked the foliation characteristic of wedge ice and is interpreted as pool ice.

Near the base of the upper Ice Complex, syngenetic ice wedges, a few metres wide tend to narrow downwards in the lower several metres of the unit and terminate with irregular or flattish or U- to V-shaped toes over a vertical distance of about 1–3

5   m (Figure S2). At the northeast part of the slump headwall, however, the base of the upper Ice Complex is a sharp contact that cuts down into and truncates bands in the underlying lower sand unit over a vertical distance of up to at least several metres (Figure 3). As a result, the upper ice complex thickens substantially downslope.

The top of the upper ice complex displays a variety of wedge relationships with the overlying upper sand unit (Figure S3). Some wide syngenetic wedges terminate along planar to gently undulating contacts that are horizontal to gently dipping and

10 have apparent widths of about 1–3 m; some taper upward into narrow wedges characteristic of the upper sand unit; some taper irregularly upward, marked by shoulders up to about 0.5 m wide; and some end upwards with a narrow offshoot. These changes occur over a vertical distance of about 1–3 m.

4.1.2 Lower sand unit (MIS7-6)

A small truncated ice wedge B17-IW1 (0.5 m wide, 1.1 m high) from the lower sand unit was sampled in section 1 (Figure 4)

15 near the slump bottom, about 50 m below the ground surface (about 5 m below the altitude level dated by luminescence to 142.8 ± 25.3 and >123.2 ka by OSL and to 210.0 ± 23.0 ka by IRSL (Ashastina et al., 2017). It shows a mean stable-isotope composition of –33.1‰ for $\delta^{18}O$, –256.7‰ for δD, and 8.0‰ for d excess (n=6), with very little variability (Table 1, Figure 6). The regression line in a $\delta^{18}O–\delta D$ bi-plot is $\delta D = 12.78\ \delta^{18}O + 166.23$ ($R^2 = 0.80$).

The isotopic composition of the intrasedimental ice (n=4) shows similar mean δ values (–32.6‰ for $\delta^{18}O$, –244.3‰ for δD),

20 but higher d excess values (14.7‰). The co-isotopic regression line is $\delta D = 20.06\ \delta^{18}O + 405.23$ ($R^2 = 0.77$).

Radiocarbon dating of twigs, Cyperaceae stems and roots from a sediment sample 0.5 m above the truncated ice wedge yielded nonfinite ages of >53,400 and >37,500 ^{14}C yr BP (Table 2). A hare dropping from the ice wedge was dated to 15,792±358 ^{14}C yr BP, which can be only explained by redistribution or contamination.

4.1.3 Upper Ice Complex (MIS3-2)

25 Two large syngenetic ice wedges from the overlying ice-rich upper Ice Complex in section 2 (Figure 5) in the southern part of the slump show more depleted isotope values as compared to the lower sand unit (Table 1, Figure 6). Ice wedge B17-IW6 (about 0.5 m wide, sampled about 26 m below surface) shows a mean isotopic composition of –35.1‰ for $\delta^{18}O$, –269.4‰ for δD, and 11.0‰ for d excess (n=4). The co-isotopic regression is $\delta D = 12.25\ \delta^{18}O + 159.79$ ($R^2 = 0.99$). The mean stable isotope values of ice wedge B17-IW5 (about 1.6 m wide and sampled about 20 m below surface) are –34.9‰ for $\delta^{18}O$, –271.0‰ for

30 δD, and 8.1‰ for d excess (n=12). The regression line in a $\delta^{18}O–\delta D$ bi-plot is $\delta D = 7.41\ \delta^{18}O - 12.36$ ($R^2 = 0.99$). The intrasedimental ice data (n=2) show distinctly enriched mean values (–26.2‰ for $\delta^{18}O$, –195.0‰ for δD, 14.5‰ for d excess).

Unidentified plant and insect remains from ice wedge B17-IW6 as well as Cyperaceae remains and roots from host sediment at the lowest studied level in section 2, about 26 m below surface, both revealed nonfinite ages (>37,500 ^{14}C yr BP; Table 2). At the level of about 20 m bs rootlets from the host sediment yielded ages of >37,500 and 47,550±677 ^{14}C yr BP, whereas unidentified plant remains from within ice wedge B17-IW5 were dated to 24,858±536 ^{14}C yr BP.

4.1.4 Upper sand unit (MIS3-2)

Two narrow sand-ice wedges (composite wedges) and one ice wedge were sampled from the ice-poor upper sand unit in section 2 (Figure 5) at the southern part of the slump. Composite wedge B17-IW2 (0.25 m wide) was sampled 2.6 and 2.8 m bs, and composite wedge B17-IW3 (0.2 m wide) was sampled 2.0 m bs. The stable-isotope composition of the composite wedges from the upper sand unit differs from those of both underlying units, i.e. the lower sand unit and the upper Ice Complex (Table 1, Figure 6). The mean values are –32.8‰ and –31.9‰ for $\delta^{18}O$, –247.5‰ and –240.5‰ for δD, and 15.1‰ and 14.7‰ for d excess (n=4 and n=2), respectively, and plot well above the Global Meteoric Water Line (GMWL, $\delta D = 8\ \delta^{18}O + 10$). The corresponding regressions are $\delta D = 12.51\ \delta^{18}O + 163.29$ ($R^2 = 0.92$) and $\delta D = 14.47\ \delta^{18}O + 220.91$, respectively.

Ice wedge B17-IW4 (about 0.4 m wide and containing a few thin sand veins) was sampled in the upper sand 2.0 m bs and exhibits a more enriched stable-isotope composition (mean values –29.9‰ for $\delta^{18}O$, –227.8‰ for δD, and 11.6‰ for d excess (n=6). The co-isotopic regression is $\delta D = 15.03\ \delta^{18}O + 222.01$ ($R^2 = 0.91$) (Table 1, Figure 6).

The intrasedimental ice (n=3) shows similar mean δ values (–30.6‰ for $\delta^{18}O$, –220.0‰ for δD) but higher d excess (24.4‰) values compared to the composite wedges and the ice wedges of the upper sand. The regression line in a $\delta^{18}O$–δD bi-plot is $\delta D = 10.61\ \delta^{18}O + 104.19$ ($R^2 = 0.99$).

Two radiocarbon ages (analysed as gas) of small samples of *Empetrum nigra* leaf, twigs, and Cyperaceae remains from host sediment about 11 m bs are nonfinite, whereas a graphite target yielded an age of 38,348±236 ^{14}C yr BP.

4.1.5 Slump outflow

The samples from the outflow stream draining the slump plot well above the GMWL (Table 1, Figure 6). Their average stable-isotope values are –31.7‰ for $\delta^{18}O$, –238.6‰ for δD, and 15.2‰ for d excess ($\delta D = 8.84\ \delta^{18}O + 41.83$ ($R^2 = 0.97$, n=7).

4.1.6 Summer precipitation

The samples of summer precipitation from both sites show much more enriched stable-isotope compositions than the wedges, as expected (Table 1). The samples taken at the Batagay megaslump plot on a co-isotopic regression line of $\delta D = 7.93\ \delta^{18}O + 1.65$ ($R^2 = 1$). Mean values are –14.3‰ for $\delta^{18}O$, –112.0‰ for δD, and 2.6‰ for d excess (n=4).

4.2 Adycha River

Ice wedge A17-IW3 was about 0.3 m wide and sampled about 1.4 m above river level and about 2 m below the ground surface. Its mean stable isotope values are –29.0‰ for $\delta^{18}O$, –226.3‰ for δD, and 5.5‰ for d excess (n=10, Table 1, Figure 6). The regression line in a $\delta^{18}O$–δD bi-plot is $\delta D = 7.01 \, \delta^{18}O - 23.25$ ($R^2 = 0.99$).

5 Rootlets and larch needles from the host sediment were dated to 213±109 ^{14}C yr BP and confirm the Late Holocene age of the alluvial sediments and the ice wedge.

5 Discussion

5.1 Chronostratigraphy of the Batagay megaslump

To date, the exceptionally long record of permafrost history under highly continental climate conditions as preserved in the
10 Batagay megaslump is still not entirely established in terms of its geochronology. Previous studies by Ashastina et al. (2017) and Murton et al. (2017) provide the general differentiation into four main units overlain by Late Holocene cover deposits, which are the lower Ice Complex (pre-MIS6), the lower sand unit (MIS7-6), the upper Ice Complex (MIS3-2) and the upper sand unit (MIS3-2). Thaw unconformities as observed in sampled sections of this study (Figure 4, Figure 5), but also in previous studies question the completeness of the archive. Such (thermo-)erosional features are often observed in late
15 Quaternary permafrost chronologies (e.g. Wetterich et al., 2009) and most likely caused by intense permafrost degradation during warm stages such as the Last Interglacial with widespread thermokarst (e.g. Kienast et al., 2011), but might also relate to palaeotopography.

The Batagay megaslump exposes the single stratigraphic units in different contacts. Its complex structure, the rapid erosion and the huge dimensions, however, hold potential to identify additional records and units in future studies. With the present
20 study and additional radiocarbon dating, further confidence of the existing stratigraphy is reached although the wedge-ice derived age information asks for careful interpretation. Due to the formation mechanism and downward directed growth of wedge ice, organic remains from inside the ice are commonly younger than those of the host deposits at the same altitude (sampling depth). The risk to date redeposited organic material, that entered into or later froze onto the surface of ice recently exposed in a slump, is always given, in particular in erosional features such as a slump with gullies. It can be seen in the **lower**
25 **sand unit** (Figure 4) where an age of about 16 ^{14}C ka BP was obtained from the ice wedge, while radiocarbon dates of the host deposit are infinite and luminescence ages reach 142.8±25.3 ka and >123.2 ka by OSL as well as 210.0±23.0 ka by IRSL (Ashastina et al., 2017).

A distinct layer of wood and other plant remains up to 1.5 m thick and traceable along the exposure is situated above the lower sand unit and the upper Ice Complex. Radiocarbon dating of this layer revealed an infinite radiocarbon age of >44 ^{14}C ka BP
30 (Ashastina et al., 2017) as well as an out of calibration age of 49.32±3.15 ^{14}C ka BP (Murton et al., 2017). The palaeobotanic

proxy data indicate warm climate. Hence the unit is aligned to the Last Interglacial (Ashastina et al., 2018) that caused particular thaw of the underlying lower sand unit and formed a discordance in the sequences.

The expected difference between ice-hosted and sediment-hosted ages of the same sample depth is clearly seen in the **upper Ice Complex**, where organic remains from the ice wedge are dated to about 25 ^{14}C ka BP and the host deposits yield nonfinite

5 to about 48 ^{14}C ka BP ages (Figure 5). However, even though the age offset is very large and a contamination of the dated sample in the field cannot be excluded, such information may be reliable and lies within the previously dated overall accumulation period of the upper Ice Complex between about 49 and 13 ^{14}C ka BP (Ashastina et al., 2017; Murton et al., 2017).

The new radiocarbon ages from the **upper sand unit** largely confirm the assumed accumulation period of this unit from about

10 36 to 26 ^{14}C ka BP (Ashastina et al., 2017; Murton et al., 2017), but predate it by about 2000 years, at least with a date of about 38 ^{14}C ka BP and nonfinite ages of >37.5 ^{14}C ka BP from the host deposits. We note that the upper sand has been studied at two different sites of the slump in different altitude levels (north-northeast of the central headwall (Ashastina et al., 2017; Murton et al., 2017) and together with the upper Ice Complex more downslope in the southern part of the slump, (Figure 2), which complicates the correlation of dating results from different sampling sites. Furthermore, close to our section 2, the upper

15 sand could not clearly be differentiated from the upper Ice Complex during fieldwork in 2014 (Ashastina et al., 2017). Hence, the more or less simultaneous accumulation of both the upper Ice Complex and the upper sand unit during MIS3-2 but the differing spatial distribution in the slump highlights the importance of paleo-topography in a hillslope setting and varying sediment sources and velocities of deposition and require future sampling and analysis at higher vertical resolution.

Interestingly, no indications for an upper sand layer have been found in the topmost part of the modern headwall. Additionally,

20 there are no significant Holocene deposits exposed at all in the upper part of the slump. However, carcasses of *Equus lenensis* and baby *Bison priscus* (dated to 4.45±0.35 ka ^{14}C BP and 8.215 +0.45/-0.40 ka ^{14}C BP) were found at the exit of the slump (Murton et al., 2017). These observations may be related to (a) the palaeo-topographic situation and transport and deposition regimes of windblown dust during MIS3-1, likely decreasing upslope, (b) widespread permafrost degradation since the Late Glacial Early Holocene transition, and/or (c) changed hydrological and vegetation patterns in the Holocene preventing

25 transport and/or deposition of windblown dust.

The **lower Ice Complex** is not dated so far. Given the dating results of the overlying lower sand unit it may correspond to the Yukagir Ice Complex from Bol'shoy Lyakhovsky Island, which has been dated by radioisotope disequilibria (^{230}Th/U) in peat to 178±14, 221±27 ka (Wetterich et al., in revision) and to 201±3 ka (Schirrmeister et al., 2002). Hence this unit most likely has survived two interglacials.

30 **5.2 Cryostratigraphy at the Batagay megaslump – implications for ice-wedge growth regime**

The shape and composition of the wedges in the Batagay megaslump correspond to the assumed genesis of the respective units and may also provide palaeoclimate information (Table 3). Both the lower and upper Ice Complex units are characterized by ice wedges up to several m wide, whereas both the lower and upper sand units contain tall narrow wedges. The former indicates

more stable surface conditions with lower accumulation rates and sufficient melt water supply during the formation of both Ice Complexes that allowed more horizontal growth of ice wedges. The tall narrow wedges of both sand units, in contrast, point to rapid upslope directed aeolian deposition of sediments and, therefore, a predominant vertical ice-wedge growth in persistent polygonal patterns. The downslope inclination of narrow syngenetic wedges in the upper sand unit indicates upward growth subvertically at right angle to the aggrading depositional surface, a hillslope. The tall and narrow wedges in the sand units are similar in size and shape to chimney-like sand wedges in thick aeolian sand sheets in the Tuktoyaktuk Coastlands of western Arctic Canada, attributed to rapid aggradation of aeolian sand and rapid wedge growth (Murton and Bateman, 2007). Ice wedges in the lower sand unit point to moister conditions than the composite wedges in the upper sand unit, which imply dry conditions with only little melt water supply. However, as most ice wedges exhibit a significant sediment content, a steady supply of windblown sediment and a rather thin snow cover during ice-wedge formation is likely.

At the top of the upper Ice Complex, ice wedges tend to narrow and partly transform into narrow composite wedges of the upper sand unit (Figure S3). This indicates a rather gradual change of the deposition regime towards higher sedimentation rates and possibly lower melt water supply due to drier conditions. Interestingly the polygonal pattern has not been affected by this change, as indicated by the consistent frost-cracking positions. The upward transition of wedges from the upper ice complex to upper sand unit, however, was interrupted episodically by thaw, producing a number of thaw unconformities at different depths. In contrast, the erosional event truncating the lower Ice Complex seems to have changed the polygonal pattern in which the lower sand unit has been deposited.

A major erosional surface attributed to gullying by water flowing down a palaeo-hillslope is inferred from the large concave-up lower contact of the upper ice complex in the northeast part of the headwall (Figure 3). Numerous gullies on the present hillslope near the megaslump are indicated on satellite images between 1968 and 2010 reported by Kunitsky et al. (2013), which suggests that gullies are characteristic landforms of such terrain under present environmental conditions. Water is supplied to such gullies by snowmelt in spring and rainfall and melt of ground ice in summer. Erosion of the underlying lower sand unit provided substantial accommodation space for development of an unusually thick upper ice complex with wide ice wedges above it. Stratigraphically, this erosional surface is at a similar level as the upper woody trash layer that is thought to be of last interglacial age (Ashastina et al., 2017; Ashastina et al., 2018), which in turn suggests that the erosion also probably took place during warm (interglacial) conditions.

5.3 Yana Highlands' ice-wedge stable isotopes – regional palaeoclimate implications

Unfortunately, no comprehensive modern precipitation stable-isotope data are available for the Yana Highlands. The nearest sites with available data are Zhigansk about 500 km to the west (Kurita, 2011), Tiksi about 500 km to the northwest (Kloss, 2008) and Yakutsk about 650 km to the south (Kloss, 2008; Papina et al., 2017). These sites show similar isotopic characteristics for the cold season (October to March). Mean $\delta^{18}O$ values decrease along the north-south transect from about -29‰ in Tiksi to -31‰ in Yakutsk, d excess varies strongly without a clear geographical pattern and shows values between about 8 and 15‰. In general, the mean winter isotopic data plot around the GMWL. However, these sites are characterized by

less cold winters compared to the Yana Highlands. Furthermore, they are located east and south, respectively, of the Verkhoyansky Range (Figure 1) that constitutes an orographic barrier for atmospheric moisture transport from the westerlies and thereby contributes to the extreme cold and dry winters of the Yana Highlands. Hence, more depleted $\delta^{18}O$ values can be expected for winter precipitation.

5   The Yana Highlands' ice-wedge stable isotope data form two major clusters. The first cluster plots below but mainly parallel to the GMWL (Figure 6) and comprises ice wedges B17-IW1 (lower sand unit), B17-IW5 (upper Ice Complex) and A17-IW3 (recent Adycha riverbank). The second cluster plots clearly above the GMWL and comprises all data from the upper sand unit (composite wedges B17-IW2, B17-IW3, ice wedge B17-IW4). The sediment-rich ice wedge B17-IW6 (upper Ice Complex) plots in-between the clusters. Furthermore, all except one of the intrasedimental-ice samples as well as all slump outflow

10 samples—interpreted as integrated signal of the entire exposed Batagay sequence—plot well above the GMWL. In summary, the higher the sediment content of our samples, the higher are the d excess values and, hence, the offset above the GMWL. Considering the cryostratigraphic interpretation of composite-wedge and ice-wedge formation, the high d excess values in the wedges from the upper sand unit might also be related to a generally dry environment during sediment deposition and wedge formation, whereas the lower d excess values of the other cluster might point to more abundant melt water supply during ice-

15 wedge formation.

Generally, the high d excess values indicate a higher kinetic fractionation at the moisture source, e.g. due to initial evaporation under low relative humidity and/or high sea-surface temperature conditions. As the cold season in Eastern Siberia is characterized by higher d excess values than the warm season, one might conclude that cold season precipitation does not only feed the ice and composite wedges but also contributes significantly to the formation of intrasedimental ice. However,

20 secondary fractionation during snow cover evolution, snow melt, refreezing and freeze-thaw cycles in the active layer may also affect d excess in wedge and intrasedimental ice.

The regional palaeoclimatic implications for the Yana Highlands drawn from our ice-wedge $\delta^{18}O$ data are summarised in Table 3. Extremely low winter temperatures ($\delta^{18}O$ values around -35‰) are recorded by the ice wedges from the Upper Ice Complex, which has formed during the MIS3 interstadial (Kargin interstadial in Siberian stratigraphy). Similar values including d excess

25 have been reported by Vasil'chuk et al. (2017) for three undated ice wedges from the upper Ice Complex sampled close to the headwall of the Batagay megaslump (Figure 3). Very low winter temperatures during the formation of the lower sand unit during MIS6 (Taz stadial) are also indicated by $\delta^{18}O$ values around -33‰. However, given the small size of the corresponding ice wedge, the old age and similar isotope values of intrasedimental ice in the host deposits some postdepositional alteration of the originally recorded signal (Opel et al., 2018) cannot be excluded. Hence, these values should be treated with caution.

30 Very low to low winter temperatures can be inferred for composite wedges ($\delta^{18}O$ values around -33‰ and -32‰) and an ice wedge ($\delta^{18}O$ values about -30‰) from the upper sand unit, which is attributed to the transition from MIS3 to MIS 2. Again, similar values have been reported by Vasil'chuk et al. (2017) for an undated narrow ice wedge from the upper sand unit sampled in another part of the Batagay megaslump (Figure 3). However, the palaeoclimatic significance of the rather small composite wedges with few samples has to be treated with caution. Ice wedge B17-IW4 might even be an epigenetic ice wedge of

Holocene age, given the quite similar δ^{18}O values (about -29‰) of the definitively Holocene ice wedge from the Adycha river bank (A17-IW3), which indicate moderate to low winter temperatures in the last centuries to millennia. As there is usually an offset between the stable-isotope composition of modern snow cover and wedge ice (Grinter et al., 2018; Opel et al., 2018), these values suggest that the winter precipitation in the Yana Highlands is isotopically more depleted than that at Tiksi, Zhigansk and Yakutsk.

No palaeoclimatic information based on ice-wedge isotopes is available yet from Batagay for the MIS4 (Zyryan) and MIS2 (Sartan) stadials. However, at Bol'shoy Lyakhovsky Island in northeastern Siberia ice wedges from both periods show δ^{18}O values about 6‰ lower than those of MIS3 Ice Complex ice wedges (Wetterich et al., 2011; Opel et al., 2017b; Wetterich et al., in revision). Hence, even more depleted values of about -40‰ could be expected in LGM ice wedges at the Batagay megaslump.

In general, the inferred lower winter temperatures meet the expectations of an increased continentality during the Late Pleistocene cold stages as well as partly drier conditions indicated by the formation of composite wedges. Palaeoecological analysis (plant macro remains, pollen and beetles) of Batagay megaslump deposits (Ashastina et al., 2018) indicate generally much drier conditions compared to recent times (Table 3), in particular during the cold stages. Tree and beetle indicators imply warm summers for most units, whereas macro remains indicate cold to very cold winters for the upper Ice Complex. These patterns are generally in line with our ice-wedge interpretation and support the hypothesized higher continentality.

5.4 Large-scale implications for palaeoclimate and past continentality

The spatial significance of our data can be considered in terms of the upper Ice Complex of the Batagay megaslump, formed in the MIS3 Kargin interstadial (about 50 to 30 ka ago) as well as the Holocene.

5.4.1 MIS3

To assess the climate and continentality across much of western Beringia during the MIS3 Kargin interstadial we compare our Yana Highlands ice-wedge stable-isotope data to a dataset of 17 other ice-wedge sites (Figure 1, Figure 7). Most of the wedges are from the Yedoma Ice Complex (Schirrmeister et al., 2011b; Schirrmeister et al., 2013), which is widely distributed in east and central Siberia. We selected one ice wedge per study site in this analysis and only those that provide both δ^{18}O and *d* excess data (Table S2). Additionally, all except two of the considered wedges or their host sediments have been directly dated by radiocarbon methods to between about 50 and 30 ka ago. We appreciate that an ice wedge may not contain a full record of this time period and that the climate was not uniform throughout this period; in particular, a warmer period around 40 ka BP is well known from several palaeoecological proxies for some of the study sites (Wetterich et al., 2014).

The Batagay upper Ice Complex ice wedge B17-IW5 (and the sediment-rich ice wedge B17-IW6, which has not been included in this analysis as well as the data of Vasil'chuk et al. (2017)) show δ^{18}O values of around -35‰, which is 6 to 2‰ lower than ice-wedge δ^{18}O data from today's coastal lowlands of northern Yakutia, the Kolyma region and central Yakutia (Figure 8). As hypothesized, the relatively depleted δ^{18}O values from Batagay clearly point to lower MIS3 winter temperatures in the Yana

Highlands than in the other regions, probably due to the more continental location framed by the glaciated Verkhoyansky and Chersky ranges. Even the ice-wedge $\delta^{18}O$ data from highly continental sites in central Yakutia do not reach such low delta values as observed in the Batagay ice wedges. Interestingly, the lowest MIS3 $\delta^{18}O$ values beside Batagay originate from the northeastern-most site, Novaya Sibir Island. If an Atlantic moisture source for MIS3 wedge ice is assumed, Novaya Sibir

5 would represent the farthest location from the source region and therefore show much depleted $\delta^{18}O$ values, almost reaching those of Batagay.

The peculiarity of the Yana Highlands is also apparent in the d excess data of the MIS3 ice wedges. The Batagay ice wedges exhibit the highest d values, between 8 and 11‰. Similar values have been found in ice wedges from central Yakutia, whereas today's coastal lowlands of northern Yakutia, the Laptev Sea islands and the Kolyma region sites show values between 2 and

10 7‰ (Figure 9). Again, from all northern sites, Novaya Sibir Island shows the value closest to those of the Yana Highlands and central Yakutia.

Little is known about d variation in northeast Siberian ice wedges in general, and particularly within specific time periods. Previous studies have interpreted changes in the d values between the Late Pleistocene (low values) and Holocene (higher values) as changes in the moisture sources and transport pathways (Meyer et al., 2002a; Meyer et al., 2002b). Higher d values

15 in central Yakutia and the Verkhoyansk Mountain forelands have also been found in modern winter snow along a west-to-east transect over eastern Siberia from Yakutsk to Magadan (Kurita et al., 2005). This spatial pattern has been attributed to the influence of different moisture sources, with greater kinetic effects during initial evaporation, as resulting from a lower relative humidity (Kurita et al., 2005). Analogously, during MIS3, the Yana Highlands and central Yakutia may have received precipitation from other moisture sources and/or transport pathways compared to more northern sites. Due to their more

20 southern location, the Yana Highlands might have received a contribution of moisture from a Mediterranean source. Such moisture usually has higher d excess values compared to North Atlantic moisture (Gat et al., 2003) and may have reached this region via a different southern moisture transport pathway.

5.4.2 Holocene

The mean $\delta^{18}O$ values of the Holocene ice wedge sampled at the Adycha River also point to lower winter temperatures—due

25 to greater continentality—than other Holocene ice wedges in the data set. The Adycha samples (mean $\delta^{18}O$ -29‰) are more depleted than Holocene wedges from northeast Siberian coastal sites, which show typically mean $\delta^{18}O$ values between -27 and -23‰ (e.g. Meyer et al., 2002b; Meyer et al., 2015; Opel et al., 2017a; Wetterich et al., 2018). Holocene ice wedges from the Bykovsky Peninsula (Meyer et al., 2002a) show exceptional values (mean $\delta^{18}O$ of 28 to -26‰), which do not fit into the regional pattern and may relate to local climate or environmental effects. In contrast, Holocene ice wedges from central Yakutia

30 exhibit mean $\delta^{18}O$ values of -28 to -26‰ (Popp et al., 2006), which are closer to those from the Yana Highlands and also reflect the greater continentality and lower winter temperatures than the Holocene coastal sites.

The d values from Holocene ice wedges in Siberia are highly variable. Mean d values at the coastal sites are higher (compared to MIS3) with values between 4 and 12‰ (e.g. Opel et al., 2011; Meyer et al., 2002b; Wetterich et al., 2008; Wetterich et al.,

2018). In contrast, in the Yana Highlands the mean d value of 6‰ is lower compared to the MIS3, whereas d values in central Yakutia (between 6 and 11‰) are rather similar to MIS3 values (Popp et al., 2006). The exceptional ice wedges from the Bykovsky Peninsula (Meyer et al., 2002a) even reach d values of 16‰. In summary, Holocene ice wedge d excess values do not show a clear spatial pattern that is readily interpreted in terms of moisture sources or continentality.

6 Conclusions and outlook

Our stable-isotope data clearly show that winter temperatures in the Yana Highlands during MIS 3, represented by the upper Ice Complex (Yedoma) at Batagay, and the Holocene were significantly lower than at other ice-wedge study sites in coastal and central Yakutia. This indicates the persistence of enhanced continentality of the Yana Highlands region during at least part of the Late Pleistocene in western Beringia and during the Holocene. The stable-isotope data from narrow composite wedges of the upper sand unit (MIS3-2) and an old ice wedge from the lower sand unit (MIS6) are less indicative and require additional studies.

High-resolution systematic sampling and dating now needs to be carried out for all cryostratigraphic units of the Batagay permafrost sequence to validate our findings of increased continentality during MIS3 and the Holocene, to improve the temporal resolution of the Batagay ice-wedge record, and to elucidate the palaeoclimatic history from other time slices. Of particular interest are the lower and uppermost parts of the upper Ice Complex, likely representing MIS4 and MIS2 stadials, as well as the lower Ice Complex, possibly representing late MIS7 or early MIS6. To establish reliable ice-wedge chronologies, radiocarbon dating should include macro remains, dissolved organic carbon and CO_2 from gas bubbles (Lachniet et al., 2012). Additionally, previous luminescence dating results (Ashastina et al., 2017) need to be validated. Finally, the Batagay megaslump may be an appropriate study site to validate and further develop ice-wedge dating approaches using [36]Cl/Cl (Blinov et al., 2009) as well as uranium isotope methods (Ewing et al., 2015).

Author contributions

Thomas Opel initiated and designed the present study and wrote the paper with contributions by the other co-authors. Thomas Opel, Julian Murton, and Kseniia Ashastina sampled and described ground ice and host sediments, supported by Petr Danilov and Vasily Boeskorov. Hanno Meyer carried out stable-isotope analyses and supported interpretation. Hendrik Grotheer and Gesine Mollenhauer conducted the radiocarbon dating. Frank Günther provided GIS analysis and maps. Sebastian Wetterich and Lutz Schirrmeister collected ice-wedge data for comparison and supported data analysis and interpretation. All co-authors contributed to the final discussion of the results and interpretations.

Competing interests

The authors declare that they have no conflict of interest.

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Data availability

The ice-wedge and new ^{14}C data will be made available at https://www.pangaea.de after acceptance of this paper.

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Table 1. Stable isotope (δ^{18}O, δD and d excess) minimum, mean and maximum values, standard deviations as well as slopes and intercept in the δ^{18}O-δD diagram for all ice and composite wedges, intrasedimental ice, outflow water and rain water, respectively.

Ice wedge/ Type of water	Width (m)	Samples (n)	δ^{18}O min (‰)	δ^{18}O mean (‰)	δ^{18}O max (‰)	δ^{18}O sd (‰)	δD min (‰)	δD mean (‰)	δD max (‰)	δD sd (‰)	d min (‰)	d mean (‰)	d max (‰)	d sd (‰)	Slope	Intercept	R^2
Batagay megaslump																	
Upper sand unit																	
B17-IW2 (CW)	0.25	4	-33.04	-32.83	-32.47	0.25	-250.2	-247.5	-242.8	3.3	13.7	15.1	16.9	1.5	12.51	163.29	0.92
B17-IW3 (CW)	0.2	2	-32.14	-31.87	-31.60	0.38	-244.2	-240.3	-236.4	5.5	12.9	14.7	16.4	2.4	14.47	220.91	1.00
B17-IW4	0.4	6	-30.10	-29.93	-29.71	0.13	-229.6	-227.8	-223.9	2.2	10.0	11.6	13.8	1.3	15.03	222.01	0.82
Intrased. ice	n/a	3	-32.61	-30.55	-29.21	1.82	-242.2	-220.0	-207.3	19.3	18.7	24.4	28.2	5.1	10.61	104.19	0.99
Upper Ice Complex																	
B17-IW5	1.6	12	-36.17	-34.89	-33.50	0.86	-280.6	-271.0	-260.8	6.4	6.8	8.1	10.3	0.9	7.41	-12.36	0.99
B17-IW6	0.5	4	-35.52	-35.05	-34.27	0.57	-274.9	-269.4	-259.7	7.0	8.5	11.0	14.4	2.5	12.25	159.79	0.99
Intrased. ice	n/a	2	-27.63	-26.19	-24.75	2.04	-198.5	-195.0	-191.5	5.0	6.5	14.5	22.5	11.3	2.44	-131.03	1.00
Lower sand unit																	
B17-IW1	0.5	6	-33.18	-33.09	-32.91	0.10	-258.3	-256.7	-254.8	1.5	6.9	8.0	9.4	0.8	12.78	166.23	0.80
Intrased. ice	n/a	4	-32.56	-32.38	-32.24	0.14	-247.4	-244.3	-241.6	3.2	12.4	14.7	16.8	2.3	20.06	405.23	0.77
Other																	
Outflow water	n/a	7	-32.29	-31.73	-31.35	0.34	-244.0	-238.6	-235.4	3.1	14.3	15.2	16.1	0.6	8.84	41.83	0.97
Rain water	n/a	4	-16.82	-14.33	-10.17	2.92	-131.7	112.0	-79.0	23.1	2.3	2.6	2.9	0.2	7.93	1.65	1.00
Adycha River																	
Holocene river bank Adycha River																	
A17-IW3	0.3	10	-30.40	-28.98	-28.15	0.74	-236.6	-226.3	-220.5	5.2	4.4	5.5	7.3	0.9	7.01	-23.25	0.99

Table 2. Radiocarbon ages of organic remains in ice wedges (sample ID includes IW) and host sediments of the Batagay megaslump and at the Adycha River. The samples were radiocarbon dated as gas targets (Lab ID ends with 1.1) and graphite targets (Lab ID ends with 2.1).

Sample ID	Depth (m bs)	Lab ID	$F^{14}C$	Radiocarbon age (yr BP)	Calibrated age 95.4% (cal yr b2k)	Dated material	Remarks
Batagay megaslump							
Upper sand unit							
B17-S2-AMS4-1	11	1689.1.1	<0.0094	>37,500	n/a	*Empetrum nigra* leaf, twigs, Cyperaceae remains (sediment)	
		1689.2.1	0.0084±0.0002	38,348±236	42,863-42,140		
B17-S2-AMS4-2	11	1690.1.1	<0.0094	>37,500		Cyperaceae stems and roots (sediment)	
Upper Ice Complex							
B17-IW5-02	20.3	1684.1.1	0.0453±0.0030	24,858±536	30,355-27,926	unidentified plant remains (ice wedge)	
B17-IW6-04	25.8	1686.1.1	<0.0094	>37,500	n/a	unidentified plant and insect remain (ice wedge)	
B17-S2-AMS5	20.3	1691.1.1	<0.0094	>37,500	n/a	leaf fragment, roots (sediment)	
		1691.2.1	0.0027±0.0002	47,550±677	>50,033		
B17-S2-AMS6	25.8	1692.1.1	<0.0094	>37,500	n/a	Cyperaceae remains, roots (sediment)	
Lower sand unit							
B17-IW1-04	~50	1683.1.1	0.1400±0.0062	15,792±358	20,044-18,402	hare dropping (ice wedge)	Likely redistributed
B17-S1-AMS2	~49.5	1688.1.1	<0.0094	>37,500	n/a	twigs, Cyperaceae stems and roots (sediment)	
		1688.2.1	<0.0013	>53,400	n/a		
Adycha River							
Holocene river bank Adycha River							
A17-K-03	2	1687.1.1	0.9738±0.0132	213±109	<514	rootlets and larix needle (sediment)	

Table 3. Summary table on past climate in the Yana Highlands. Palaeoecological interpretations are from (Ashastina et al., 2018) and based on fieldwork in 2014.

Unit Ice wedge Chronology	Cryostratigraphy	Interpretation	IW isotopes	Interpretation	Palaeoecological interpretation
Adycha, river bank Ice wedge A17-IW3 Late Holocene to recent	Narrow ice wedge, little sediment	Moist (high melt water supply)	δ^{18}O: -29‰ d excess: 7‰	Cold winters rather moist winters	-
Batagay, upper sand unit Composite wedges B17-IW2/3 MIS3/2	Narrow composite wedges	Dry (little snow melt water) and/or high sediment supply	δ^{18}O: -33‰, -32‰ d excess: 15‰	Very cold to cold winters dry winters	Dry summers, T_{july} >12°C (trees)
Batagay, upper sand unit Ice wedge B17-IW4 MIS3/2	Narrow ice wedge, little to medium sediment content	Rather moist (higher melt water supply compared to composite wedges)	δ^{18}O: -30‰ d excess: 12‰	Cold winters rather dry winters (Epigenetic Holocene IW?)	Dry summers, T_{july} >12°C (trees)
Batagay, upper Ice Complex Ice wedge B17-IW5 MIS3	Wide ice wedge, little to medium sediment content	(rather) moist (high melt water supply)	δ^{18}O: -35‰ d excess: 8‰	Extremely cold winters rather moist winters	Dry and (very) warm summers, very cold and dry winters (beetles)
Batagay, upper Ice Complex Ice wedge B17-IW6 MIS3	Medium ice wedge, high sediment content	Rather dry (drier than IW5, moister than IWs 2/3)	δ^{18}O: -35‰ d excess: 11‰	Extremely cold winters rather dry winters	Dry summers, cold winters
Batagay, lower sand unit Ice wedge B17-IW1 MIS6	Narrow ice wedge, little sediment content	Moist, high melt water supply	δ^{18}O: -33‰ d excess: 8‰	Very cold winters rather moist winters	Dry summers, T_{july} >12°C (trees)
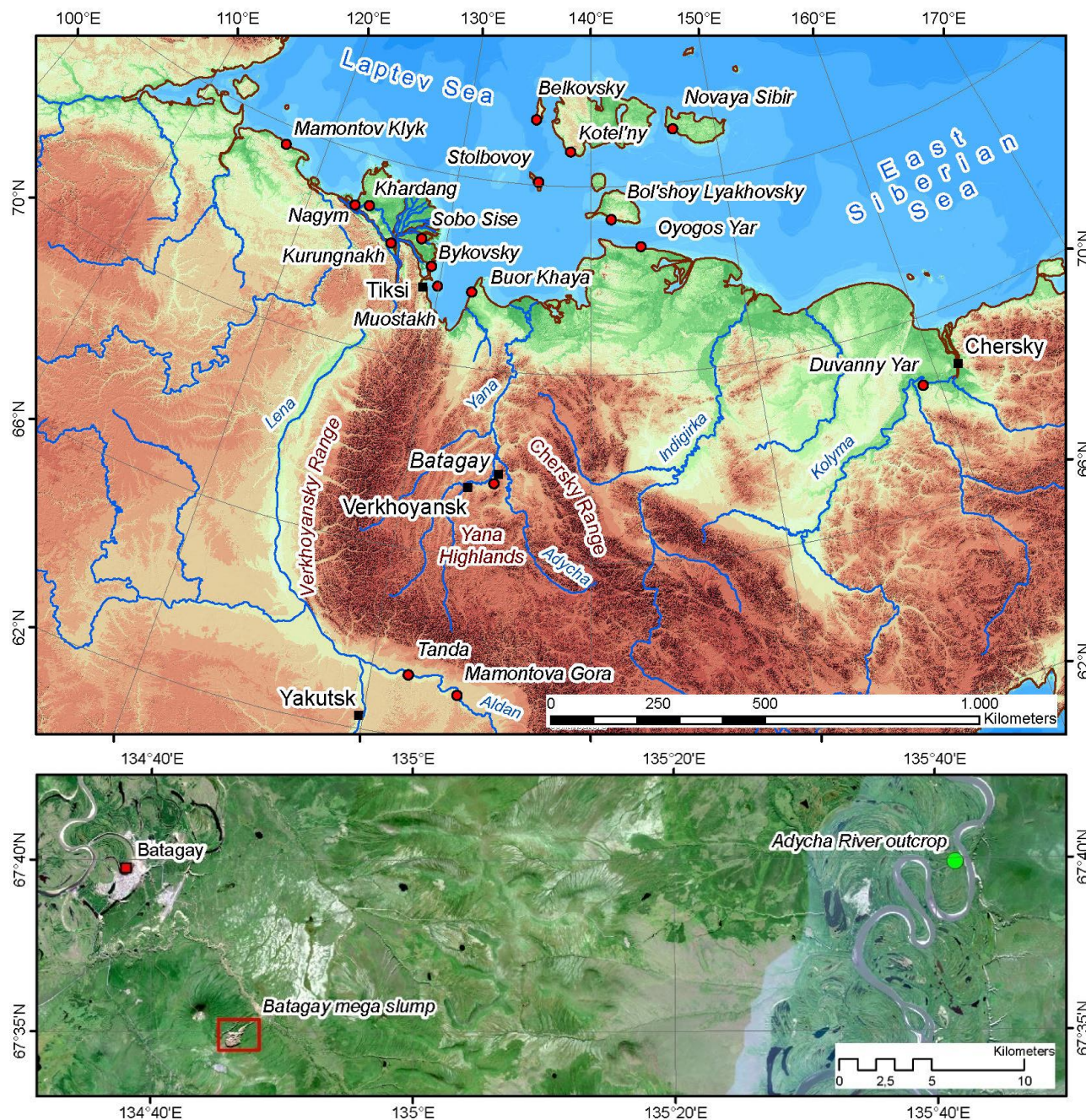
Figure 1. (top) Study region in northeastern Siberia showing all ice-wedge study sites mentioned. (bottom) Study sites in the Yana Highlands (Background image: Sentinel-2 mosaic, August 7, August 21, 2017).
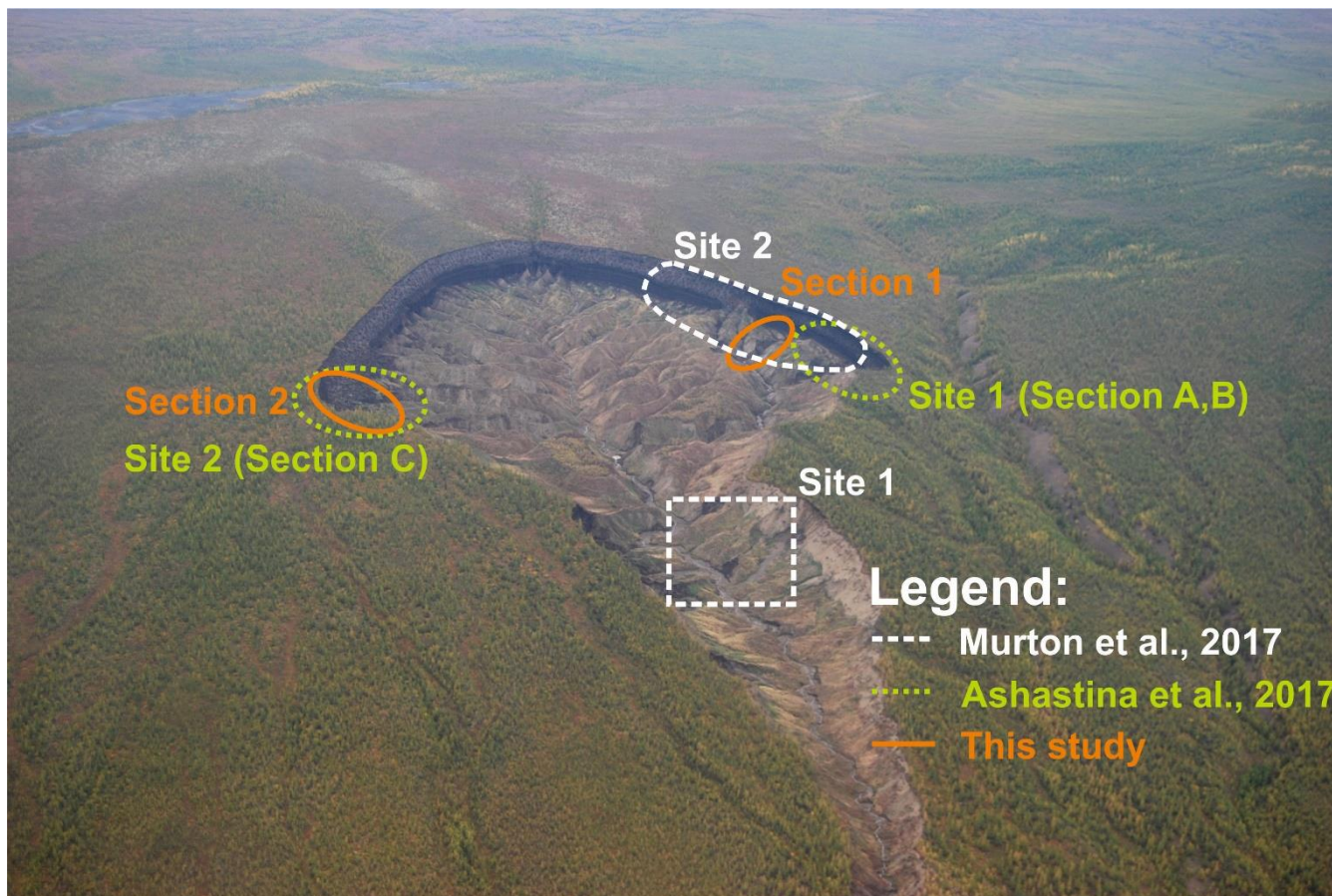
Figure 2. Overview photograph of the Batagay megaslump showing the study sites of (Ashastina et al., 2017; Murton et al., 2017) and this study. Photograph taken by Alexander Gabyshev in 2015.

Figure 3. Panoramic photograph of the Batagay megaslump with main stratigraphic units and approximate sampling locations of ice wedges (blue arrows). Furthermore, red circles and numbers indicate the approximate sampling locations of ice wedges used by (Vasil'chuk et al., 2017).
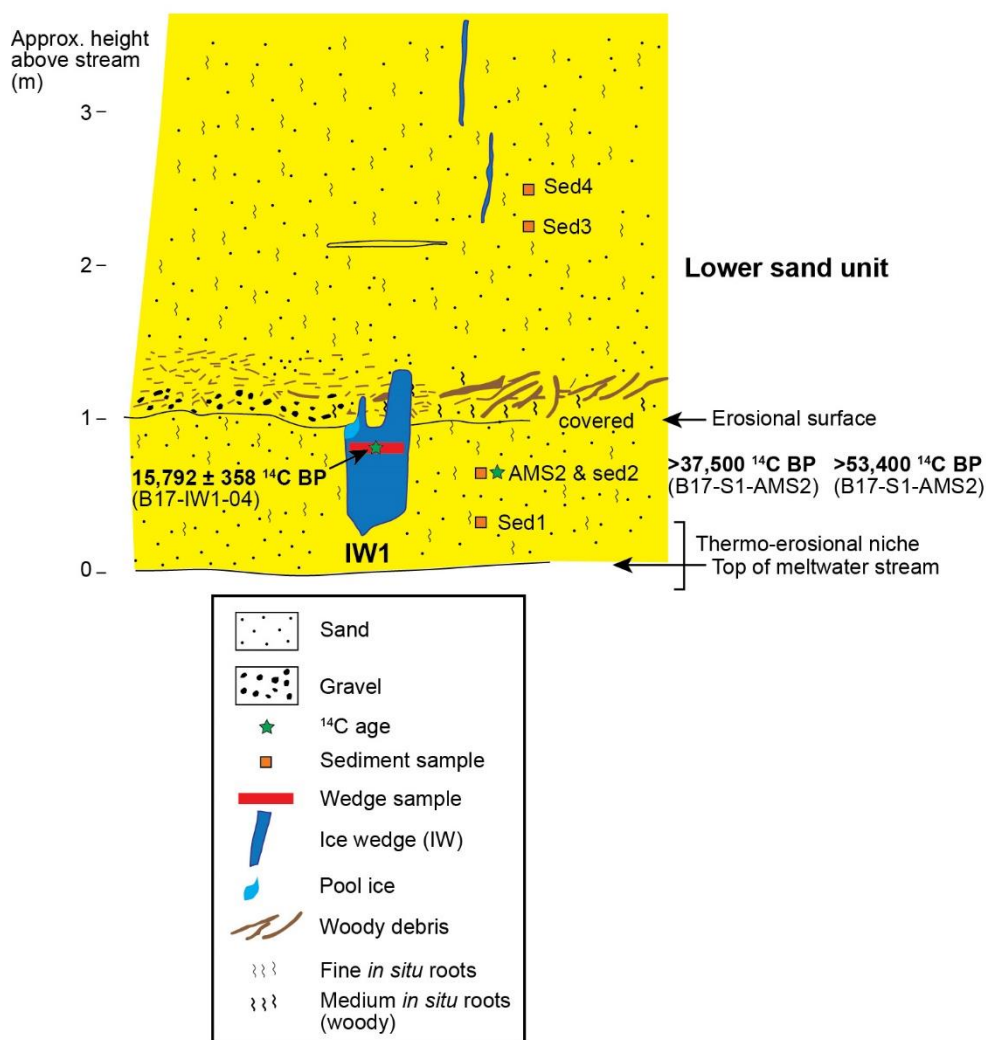
Figure 4. Schematic cryostratigraphy of Section 1, Batagay megaslump.
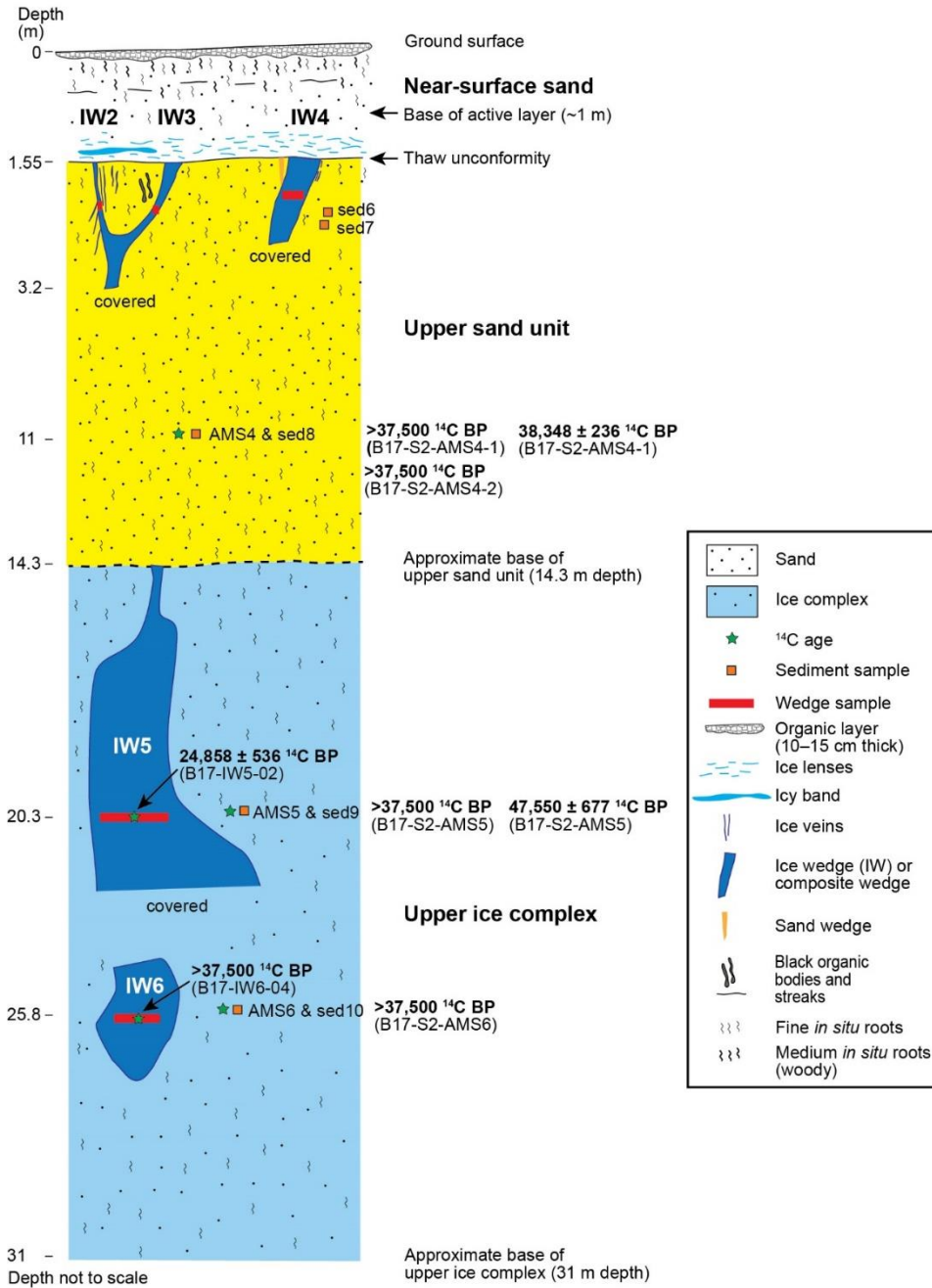
Figure 5. Schematic cryostratigraphy of Section 2, Batagay megaslump.
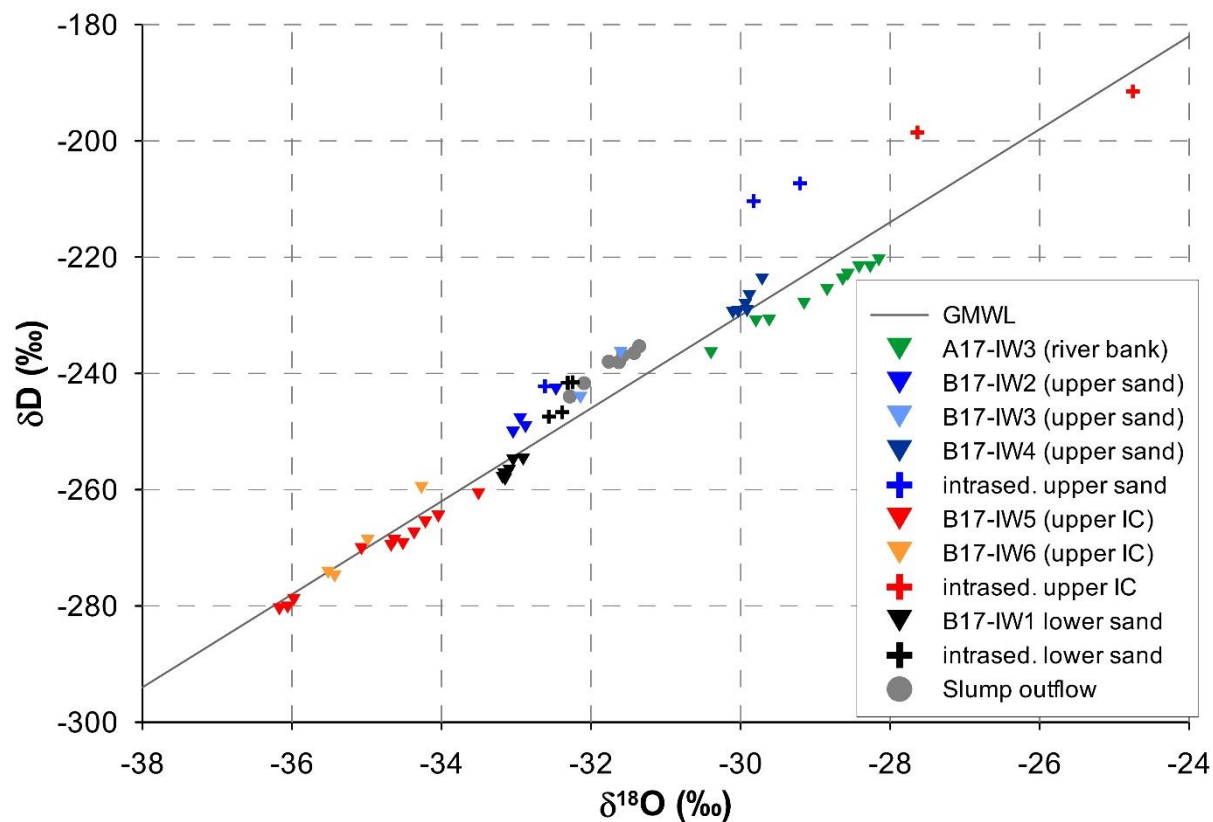
Figure 6. Co-isotopic δ¹⁸O-δD diagram of ice wedges and intrasedimental ice sampled at the Batagay megaslump and the Adycha River. Furthermore, the data of the Batagay megaslump outflow samples are shown which display an integrated signal of the entire Batagay permafrost sequence and modern precipitation.
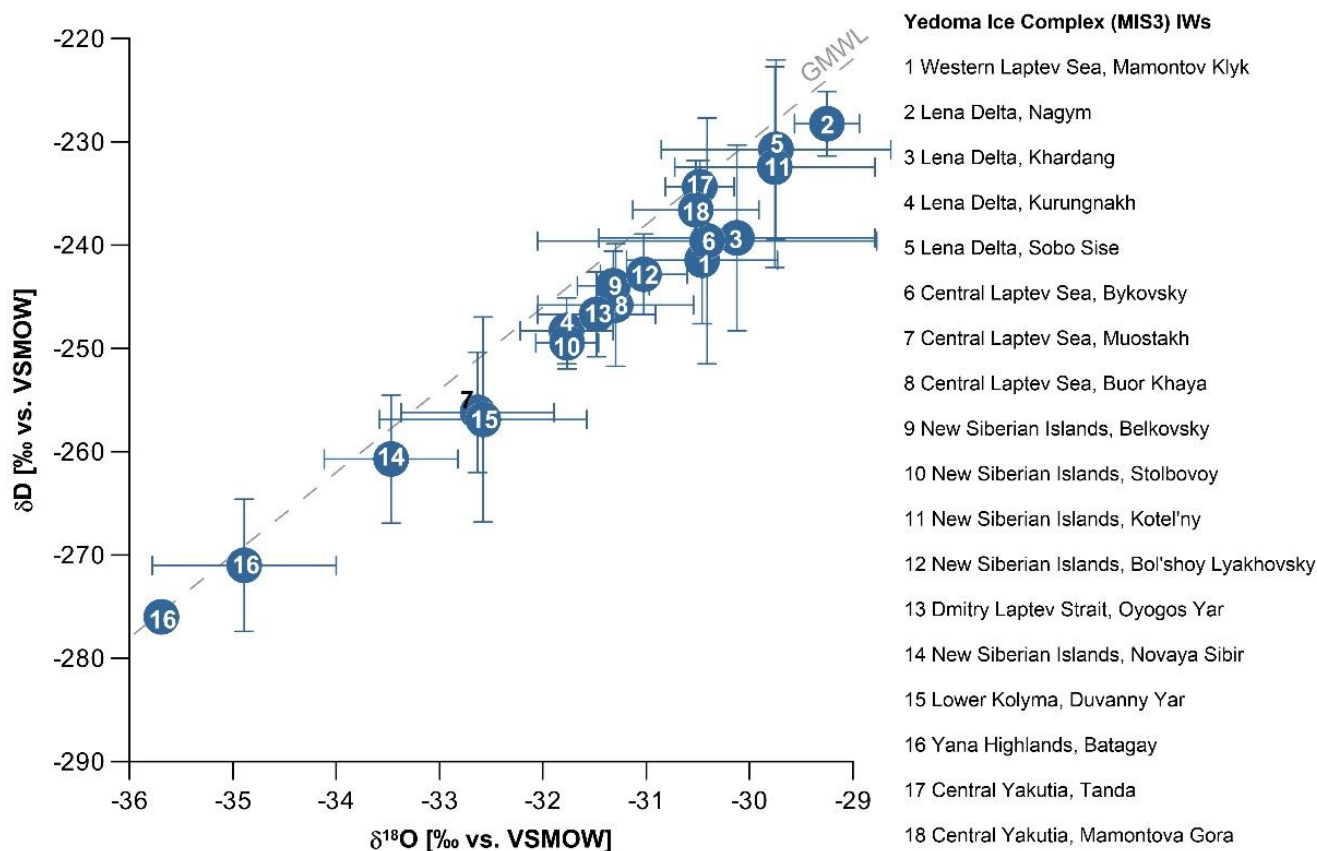
Figure 7. Comparison of stable isotope data from Siberian ice wedges attributed to the Kargin MIS3 interstadial (about 50 to 30 ka), comprising data from the Western Laptev Sea (Magens, 2005), the Lena River Delta (Schirrmeister et al., 2003; Schirrmeister et al., 2011a; Wetterich et al., 2008; Opel, unpublished), the Central Laptev Sea (Meyer et al., 2002a; Meyer/Opel, unpublished; Schirrmeister et al., 2017), the New Siberian Islands and the Dmitry Laptev Strait (Schirrmeister, unpublished; Wetterich et al., 2014; Opel et al., 2017b), the Kolyma Lowland (Strauss, 2010), the Yana Highlands (this study; Vasil'chuk et al., 2017), and Central Yakutia (Schirrmeister, unpublished; Popp et al., 2006). Further information is given in Table S2.
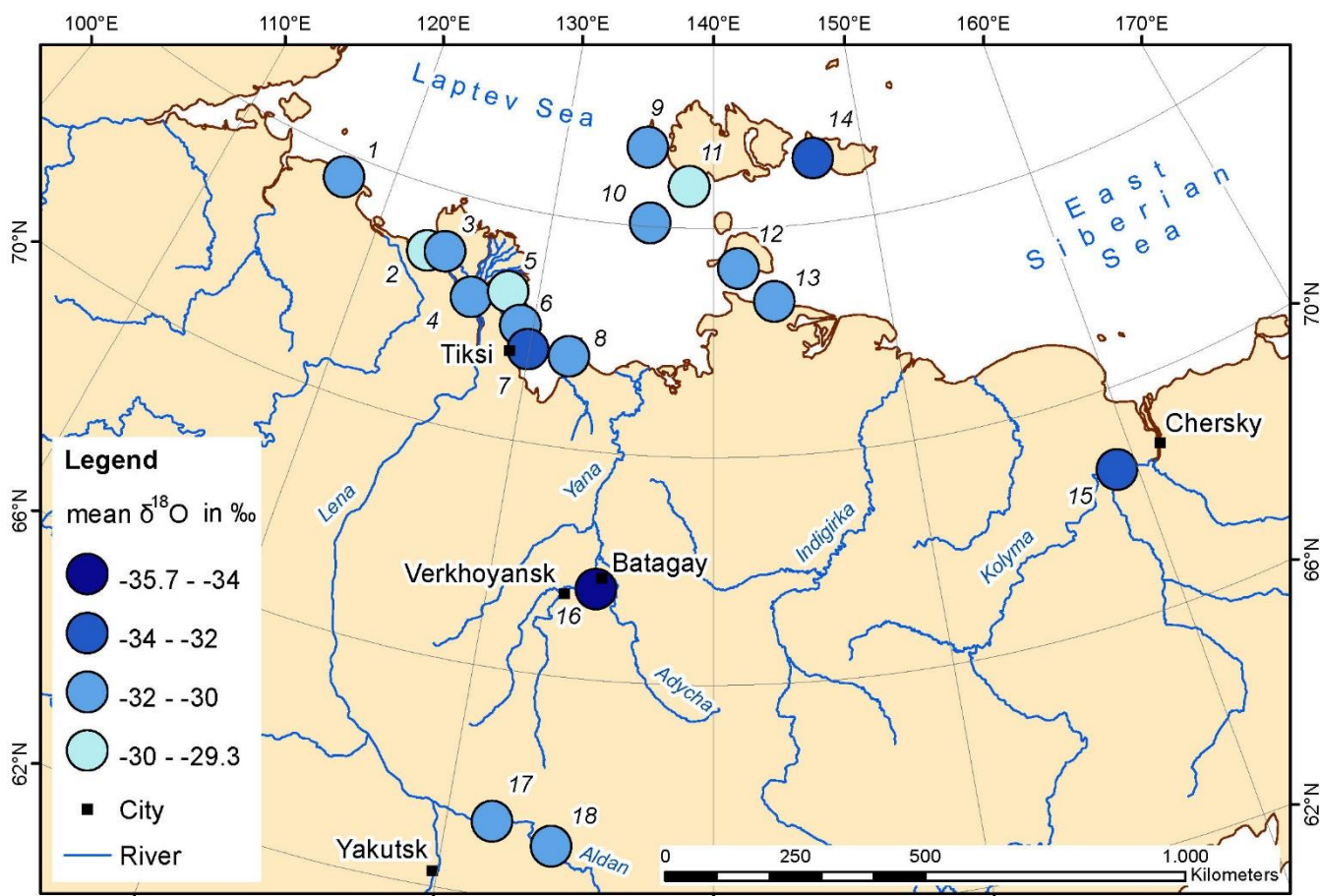
Figure 8. Map of mean δ¹⁸O data from Siberian ice wedges attributed to the Kargin MIS3 interstadial (about 50 to 30 ka), Site IDs are given in Figure 7 and further details in Table S2.
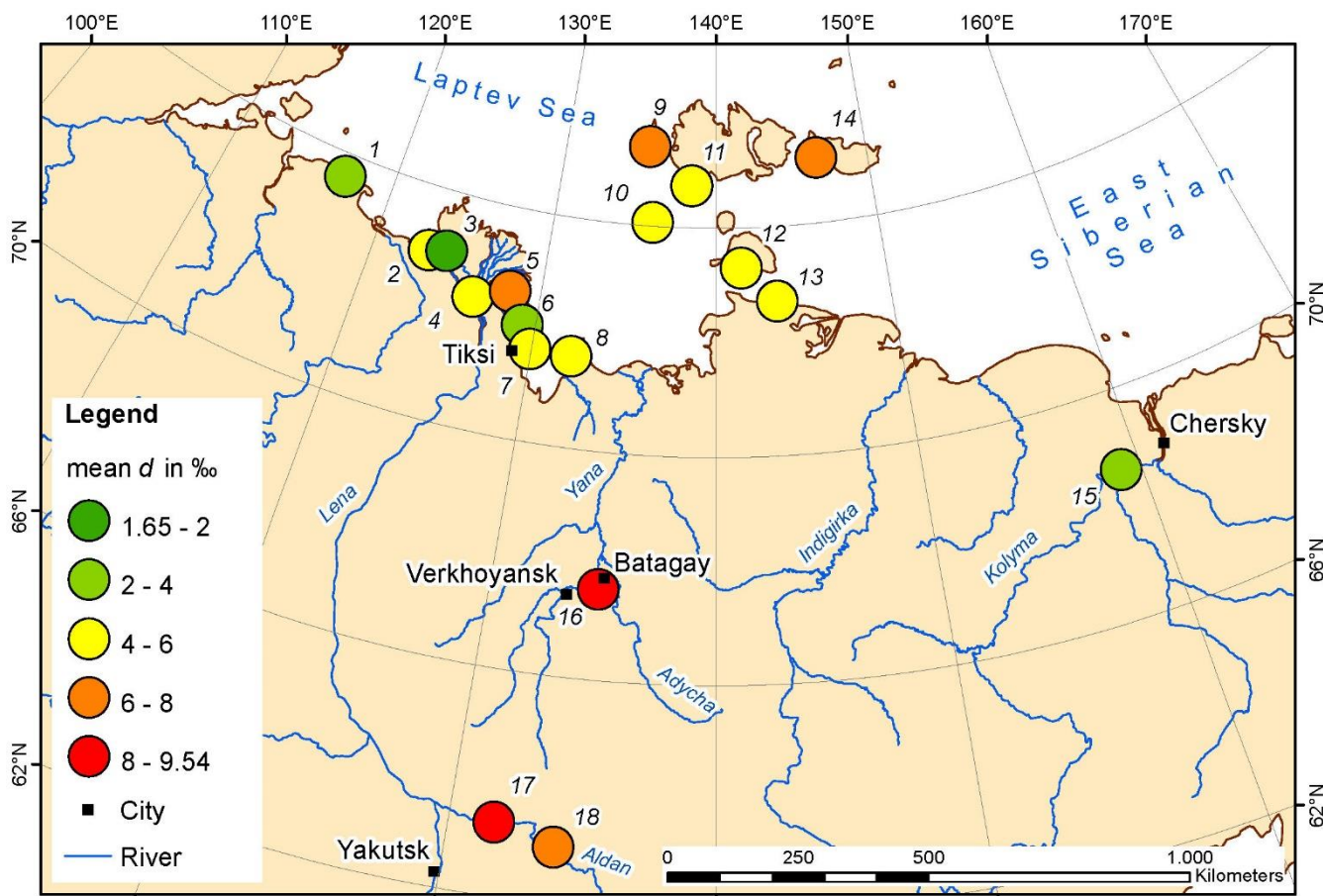
Figure 9. Map of mean *d* data from Siberian ice wedges attributed to the Kargin MIS3 interstadial (about 50 to 30 ka), Site IDs are given in Figure 7 and further details in Table S2.