



Mid-winter (DJF) temperature reconstruction in Jerusalem since 1750 with some regional implications

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1 **Abstract**

2 This work presents a statistical reconstruction of average mid-winter (DJF)
3 temperature in Jerusalem since 1750. It is a first comprehensive attempt to reconstruct
4 the temperature in Jerusalem, as a good representation of the Eastern Mediterranean
5 (EM) climate. This representativeness is verified here. The data has been reconstructed
6 by using a statistical model based on Principal Component Regression (PCR), using
7 both instrumental data and high temporal resolution records of proxy data, including
8 tree ring chronologies from Jordan, and records of DJF precipitation and Sea Level
9 Pressure from central and Western Europe. A split validation procedure has resulted in
10 a 0.73 correlation between observed and reconstructed temperature. The warming trend
11 of last decades is well noted in the reconstruction and is in line with other studies.
12 Winters which were cold/warm were historically documented as wet/dry, respectively,
13 consistent with earlier studies pointing a strong relationship between Jerusalem
14 temperatures and precipitation. It is shown here for the first time that the 'First Aliyah'
15 (immigration) to Israel during 1882-1904 initiated during favouring climate conditions
16 (cool and wet) to establish an agricultural community in the region. These conditions
17 were found to be exceptional compared to other periods since 1750.

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19 **Keywords:** Climate reconstruction; Principal Component Analysis; climate change;
20 First Aliyah; immigration

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32 **1 Introduction**

33 Temperature and precipitation are the most important climatic elements affecting
34 humanity, economy and ecosystems. Warming and drying trends, as well as extreme
35 temperature events, including heat waves, dry spells and droughts have an essential
36 influence on human life, especially in the Eastern Mediterranean (EM) region (Saaroni
37 et al., 2003; Luterbacher et al., 2004; Touchan et al., 2005; Ziv et al., 2005; Luterbacher
38 et al., 2012; Lelieveld et al., 2012; Tanarhte et al., 2012; Saaroni et al., 2015). Several
39 studies suggested that climate change had a decisive impact on societies in the region,
40 at times even bringing about total collapse (Alpert and Neumann, 1989; Weiss and
41 Bradley, 2001; Enzel et al., 2003; Ellenblum, 2012; Torfstein et al., 2013, Kelley et al.,
42 2015). For example, recent studies suggest that Syria and the greater Fertile Crescent,
43 where agriculture had begun some 12,000 years ago, experienced the worst 3-year
44 drought in the instrumental record (Trigo et al., 2010). The drought caused massive
45 agricultural failures. The most significant consequence was the migration of as many
46 as 1.5 million people from rural farming areas to urban centers. It was further suggested
47 that the drought (2007-2010), contributed to the present conflict in Syria (Kelley et al.,
48 2015). The first study to point at global warming as a central contributor to the drying
49 of the Fertile Crescent was Kitoh et al., (2008), see also Alpert et al., (2013).

50 The EM region has received relatively little attention concerning past multi proxy
51 high resolution climate reconstructions (Finne et al., 2011; Luterbacher et al., 2012;
52 Lelieveld et al., 2012) despite the unique position of the region, as a transition zone
53 between temperate climate in the North and arid climate in the South (Bar-Matthews et
54 al., 1997; Lionello et al., 2014). The EM climate is influenced by subtropical and mid-
55 latitude circulations as well as by tropical intrusions. Fluctuations in the occurrence and
56 intensity of these circulations are known to be controlled by different large-scale
57 oscillations; This leads to complex features in the region's climatic fluctuations (Alpert
58 et al., 2006; Saaroni et al., 2010).

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60 **1.1 Recent studies of past climate in the EM**

61 The use of proxy data from natural climate archives in climate reconstruction helps
62 to expand the time scale of the study, and enables investigation of multi-decadal
63 variations (Luterbacher et al., 2012). Various publications deal with European
64 temperature reconstruction at different scales of time and space (Luterbacher et al.,
65 1999; Luterbacher et al., 2004; Casty et al., 2005a; Xoplaki et al., 2005). In Israel and



66 nearby areas a few attempts have been made at climate reconstruction with high
67 resolution data, covering the past centuries. Liphshitz and Waisel, (1967) used tree
68 ring data to develop paleo-climate records of the region. Tamari, (1976) has
69 reconstructed rainfall regimes in Northern Sinai and Southern Israel through Dendro-
70 chronological analysis since 1850. Frumkin et al., (1991), Bar-Matthews et al., (1997)
71 and Vaks et al., (2003) have shown isotopic fractionation analysis of speleothems in
72 karstic caves in Israel to infer precipitation at the caves sites. Kuperman, (2005) had
73 reconstructed flood periods during the Holocene by dating speleothems in the Northern
74 Negev, Israel. Touchan and Hughes, (1999) have reconstructed rainfall regimes in
75 Southern Jordan since 1600 using Dendro-chronological analysis. Enzel et al., (2003)
76 exposed the surface levels of the Dead Sea in a resolution of decades to centuries, using
77 it as a proxy for winter rainfall in the region. In a recent study, droughts were
78 reconstructed using Tree-ring chronologies from the Mediterranean region (Cook et al.,
79 2016). They suggested that the recent drought in the Eastern Mediterranean is the worst
80 in 900 years.

81 No comprehensive attempt has been made to reconstruct the regional climate of
82 Jerusalem, by combining and integrating different proxies, with high temporally
83 resolved records. The various proxies differ widely in their temporal resolutions. Some
84 proxy records have annual or longer time scales. Several works (Pauling et al., 2003;
85 Luterbacher et al., 2012) have reviewed the strengths and potential weaknesses of proxy
86 sources used for seasonal climate reconstructions in the Mediterranean region. They
87 concluded that an optimal combination of high quality natural and documentary proxies
88 can yield better results than using only one of the two types.

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90 **1.2 The climate of Jerusalem**

91 Jerusalem is a town located on the Judean Mountains, at an elevation of about 600-
92 800 m. It is situated on the border between temperate climate to its north and west and
93 the semi-arid and arid climate to its east, known as the Judean Desert, which is a 'rain
94 shadow' desert. Jerusalem belongs to the Mediterranean climatic region (according to
95 the Köppen classification) with an average annual precipitation of 537 mm.
96 Precipitation occurs mostly during Oct-May. January is the coldest month of the year
97 with an average temperature of 9.8°C; July and August are the hottest months with and
98 average temperature of 25°C. Temperatures vary widely from day to night (IMS 2016).



99 This study aims to reconstruct the mid-winter (DJF) temperature fluctuation in
100 Jerusalem since 1750. The methodology uses a statistical model based on Principal
101 Component Regression (PCR), combining instrumental data and different documentary
102 and natural proxies.

103

104 **2 Data**

105 We draw on homogenised mid-winter (DJF) mean temperature of Jerusalem for the
106 period of 1865-2015, as the target of our model; retrieved from the Global Historical
107 Climatology Network (GHCN-adjusted, Lawrimore et al., 2011) and verified as
108 homogenised at the Israel Meteorological Service. Our analysis also compares this
109 record with the precipitation record of Jerusalem for the same period.

110 Our preliminary data base consisted of 49 long time-series of temperature,
111 precipitation, and Sea Level Pressure (SLP) measurements taken in central and Western
112 Europe, retrieved from the GHCN adjusted station data (Lawrimore et al., 2011). Most
113 of these series start in the mid-18th century with virtually no missing values; some series
114 go further back to the end of the 17th century. The relevance of the above data for the
115 reconstruction of temperature in Jerusalem is established in section 2.3 (Figure 1).

116 In addition, ten tree-ring site chronologies originating in the Dana Reserve in Jordan
117 and in the Troodos Mountains in Cyprus (Table 1) (www.ncdc.noaa.gov) underwent
118 de-trending and standardization by ARSTAN software (Cook, 1985) in order to
119 eliminate an age trend without losing high and low frequency climatic signals.
120 Furthermore, we have used historically documented data from Jerusalem, since 1750
121 (Klein, 1985). Brazdil et al., (2005) have suggested that this kind of documentary source
122 can be used to pin point extremes.

123 The investigation of the synoptic and large-scale patterns that affect Jerusalem was
124 made possible through the use of NCEP/NCAR reanalysis archive (Kalnay et al., 1996;
125 Kistler et al., 2001), along with Kaplan's Global Sea Surface Temperature (SST)
126 reconstructions (<http://climexp.knmi.nl>). All in all we started with 59 original time-
127 series, which make the full data base. Figure 1 displays a map with the origin of each
128 data source described in this section.

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133 **3 Statistical reconstruction model**

134 Our statistical modelling approach consists of three steps:

- 135 1) Reduction of dimensionality – first choosing significant correlated
- 136 predictors with Jerusalem mean DJF temperature and then Principal Component
- 137 Analysis (PCA) on the chosen predictors.
- 138 2) Temperature reconstruction –Multiple Linear Regression (MLR).
- 139 3) Validation – Split Validation (SV).

140 PCA is a common technique used in statistical modelling, where the number of
141 predictors used in the computation is reduced for the sake of a more robust model. PCA
142 produces artificial predictors, which are uncorrelated to one another, in order to avoid
143 ambiguity in the climatic signal (Mierswa et al., 2006). In this study, significant
144 correlated predictors (95% significant level) to the mean DJF temperature in Jerusalem
145 were found. Before applying PCA we normalised the various proxies in order to avoid
146 dominance of numerically larger attributes over smaller ones (Hsu et al., 2003). Next,
147 PCA was performed, keeping 70% of the variance in the original data while retaining
148 the least number of principal components. The chosen PCs were learned by a MLR
149 model, using the period 1865-1981, for which that proxies can be related to Jerusalem
150 temperature with virtually no missing values. The model produced was applied on the
151 reconstructed period (1750-1864).

152 Validation of the reconstruction was done by Split Validation, sometimes called
153 rotation estimation, which is the statistical practice of partitioning a sample of data into
154 subsets so that analysis is initially performed on a single subset while the rest of the set
155 is retained for confirmation of the initial analysis (Michaelson, 1987). The initial subset
156 is the training set and the other is validation or testing sets. In this study, the data of
157 1865-1981 were randomly partitioned into a training set and a validation set.

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159 **4 Results**

160 **4.1 Jerusalem as a representative of Eastern Mediterranean climate**

161 Studies have shown that the Jerusalem precipitation record is a reliable proxy for
162 the hydro-climatic fluctuations in the EM (Enzel et al., 2003, Kushnir and Stein, 2010).
163 Following this, we applied two tests to evaluate how well the recorded temperature in
164 Jerusalem represents the temperature of the EM. First, we evaluated the coherence
165 among the nine longest records of instrumental surface temperature from stations
166 located in the EM; Jerusalem (Israel), Beirut (Lebanon), Amman (Jordan), Nicosia



167 (Cyprus), Port-Said (Egypt), Abbasiya (Egypt), Antalya (Turkey), Urfa (Turkey) and
168 Adana (Turkey). Using correlation coefficients between every two pairs of these
169 records, each station received a score of “representativeness”. The score of each station
170 was calculated as the ratio between the average of all correlation coefficients it has and
171 the average of the significance values for these correlations. This way, a higher average
172 correlation or a lower average significance both contribute to a higher overall score.
173 Jerusalem turned out to be the most representative of them, followed by Nicosia and
174 Amman (Table 2).

175 The second criterion to evaluate the representativeness of Jerusalem was displaying
176 a comparison of the spatial correlation between the average temperature in Jerusalem
177 and GHCN surface temperature fields in other places in the region and in Europe for
178 DJF, during the period of 1900-1990 (Figure 2). The resulting high significant
179 correlations with other locations in the EM, lead us again to conclude that Jerusalem is
180 a suitable representative for the climate of the region. Accordingly, the reconstruction
181 of temperature variations in Jerusalem is an important first step towards a better
182 understanding of climatic variability in the EM.

183 To justify our use of both temperature, precipitation and SLP data in Central Europe
184 as good predictors for the temperature in Jerusalem, we analysed the relationship
185 between these variables during the 20th century, using correlation maps, to identify tele-
186 connection patterns. Figure 2 displays correlation maps between mid-winter (DJF)
187 temperature in Jerusalem and temperature, precipitation and SLP in Europe for the same
188 period. The results indicate a strong relationship between the values in the two regions:
189 a positive correlation pattern between precipitation in south Europe and mean
190 temperature in Jerusalem, and a negative one between temperature/SLP in Europe and
191 mean temperature in Jerusalem.

192 These findings suggest that connection exists between mid-winter (DJF) temperature
193 in Jerusalem and large-scale circulation patterns. Figure 3 presents teleconnection
194 patterns between Jerusalem mid-winter temperature and SST's. A negative relationship
195 ($R < -0.3$) between SST in the North Sea and a positive ($R > 0.4$) with the tropical Atlantic
196 and Jerusalem temperature is observed. Figure 4 presents teleconnection patterns
197 between Jerusalem mid-winter temperature and 300mb geopotential heights. A strong
198 negative relationship is observed over central and western Europe. These findings along
199 with the negative SLP relationship with Jerusalem mid-winter temperature (Figure 2)
200 reinforce the statistical relationship between mid-winter precipitation in Europe and



201 temperature in Jerusalem, with a dynamical one. European high pressure causes
202 northerly winds in Israel which bring cold air from western Russia to the region, in a
203 process governed by the phase configuration of Rossby waves (Ziv et al., 2006).

204

205 **4.2 The Statistical reconstruction model and validation results**

206 Table 3 shows the final predictors found to be significantly (95% level) correlated
207 to Jerusalem mid- winter (DJF) temperature. 70% of variance in the data was kept by
208 the first **five Principle Components (PC)**, while the first PC explained **41%** of the
209 variance. The final regression equation for Jerusalem mid-winter (DJF) temperature is
210 displayed: $0.25*PC1 + 0.16*PC2 + 0.06*PC3 - 0.11 * PC4 + 0.18*PC5 + 8.45$

211 The split validation has resulted in relatively good performance; with 0.73
212 correlation coefficient (Table 4 displays performance indicators for the reconstruction
213 model). Figure 5 displays the ability of the statistical model to reconstruct the average
214 winter temperature of Jerusalem. It is shown that the model can simulate relatively well
215 the mid-winter (DJF) temperature fluctuations in Jerusalem.

216

217 **4.3 Reconstruction results**

218 Final results for the reconstruction of normalised mean Jerusalem mid-winter (DJF)
219 temperatures are displayed in Figure 6. Furthermore, precipitation fluctuations during
220 the instrumental period are super-imposed on the temperature fluctuations. Although a
221 precipitation reconstruction is not the subject of this paper, the variations seen in Fig. 6
222 indicate the strong negative correlation between temperature and precipitation in the
223 EM in agreement with Striem, (1979), who found that a decrease in 1°C is equivalent
224 to an increase of 100mm in seasonal precipitation in Jerusalem. To verify this, we
225 compared our reconstructed temperature series with independent historical documents
226 of Klein (1985). The results are in agreement with Striem, (1979); i.e., the winters of
227 1763/4 1771/2, 1772/3, 1784/5, 1827/8, 1828/9, 1844/5, 1845/6 which were winters
228 with higher than average reconstructed temperature were historically documented as
229 dry winters (Klein, 1985), while the winters 1796/7, 1797/8, 1806/7, 1835/6, 1849/50,
230 1861/2 with lower than average reconstructed temperature were documented as wet
231 (Klein, 1985). These findings also strengthen the accuracy of our reconstructed series.

232 It is worth referring to the period 1882-1904 known as the 'First Aliyah' (FA) or the
233 'Agricultural Aliyah' period. This term is used to describe a large immigration of Jewish
234 Zionists to Israel, ~25,000-35,000, who managed, for the first time, to establish an



235 agricultural community. It is suggested here that this period had favouring climate
236 conditions for dry farming agriculture in the relatively hot and dry conditions of Israel,
237 i.e., below normal temperatures and above normal precipitation. Before the FA, the
238 rainy seasons of 1873/4 and 1877/8 had exceptional high precipitation amounts
239 (~1000mm/year in Jerusalem) as recorded in observations, on the same order of
240 magnitude as 1991/2, which was the most exceptional rainy year on record, related to
241 Mt. Pinatubo eruption (Bookman et al., 2014). Furthermore, 1879/80, 1881/2, 1882/3
242 and 1883/4 were historically documented as extremely wet years (Klein, 1985). The
243 temperature reconstruction presented here suggests that a partially relief in the hardness,
244 of high temperatures and lack of precipitation, for agriculture probably helped to the
245 success of this agricultural immigration. These conditions have not been seen since
246 1750. The warming trend of recent years is unprecedented in the last 264 years and is
247 in line with other studies.

248

249 **5 Summary**

250 The study reconstructed the mid-winter temperatures in Jerusalem since 1750, using
251 data from multiple climate proxies. The hypothesis that Jerusalem is a good
252 representative of the EM climatic fluctuations is proved, based on comparison of the
253 representativeness scores of the nine longest records of instrumental surface
254 temperature in the EM. Also, the high spatial correlations between the temperature in
255 Jerusalem at the 20th century and the surface temperature fields in the EM and in
256 Europe, shown through correlation maps, strengthen this conclusion. The suitability of
257 European station records to serve as highly resolved predictors for the average
258 temperature in Jerusalem is well demonstrated, suggesting they can be used to
259 reconstruct the larger EM climate. Historical documents used pointed at periods of
260 extreme temperatures and therefore, validate the reconstructed values.

261 The warming trend of recent decades is unprecedented since 1750, in agreement with
262 global and regional studies (IPCC 2013). It is suggested here that the period of the First
263 Aliyah (FA) to Israel coincided with favouring climate conditions for dry farming
264 agriculture in the relatively hot and dry conditions of Israel. These conditions have not
265 been seen since 1750, and probably helped in its success.

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279

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413 **Figure Captions**

414 **Fig. 1** Map displaying sites of climate proxies and measurement records used in this
415 work: instrumental records of precipitation temperature and SLP (blue icons), tree rings
416 (green icons), and Jerusalem mean temperature target record (in a black question mark).
417 Additional records used, which do not originate in a single site include, 300 hPa geo-
418 potential heights from NCEP/NCAR Reanalysis, Kaplan global SST reconstructions
419 (<http://climexp.knmi.nl>), and historic documents (Klein, 1985). These records are
420 represented on the map as red labels.

421 **Fig. 2** Correlations of winter (DJF) temperature in Jerusalem ($P < 5\%$) with surface
422 temperature (left column), precipitation (middle column) and SLP fields (right column).
423 The correlations were calculated for the period of 1900-1990 and for three sub periods
424 of 30 years each. Warm colours correspond to positive correlation and cold colours to
425 negative correlation.

426 **Fig. 3** Correlation maps of winter (DJF) temperature in Jerusalem with Kaplan global
427 seasonal SST reconstruction (1856-2003) (<http://climexp.knmi.nl>). Warm colours
428 indicate positive correlations and cold colours indicate negative correlations. Coloured
429 regions are found to be significant ($P < 5\%$).

430 **Fig. 4** Correlation map of winter (DJF) temperature in Jerusalem with NCEP/NCAR
431 300-hPa geo-potential height (1948-2014). Warm colours indicate positive correlations
432 and cold colours indicate negative correlations. Coloured regions are found to be
433 significant ($P < 5\%$).



434 **Fig. 5** The randomly selected validation set of points retained from the 1865-1981
435 period. Observed vs. Reconstructed mean mid-winter (DJF) temperature of Jerusalem
436 marked in black and red lines, respectively.

437 **Fig. 6** Normalized reconstructed Jerusalem mean winter (DJF) Temperature between
438 1750 and 2014 (black line), and winter precipitation between 1865 and 1995 (red line).
439 Five year moving average lines are marked in thick black and red, respectively. The
440 reconstructed and observed periods are divided by a black dashed line in the year 1865.
441 The First Aliyah period (1882-1904) is marked by a grey rectangle.

442 **Table 1** Tree ring site chronologies information from Cyprus and Jordan
443 www.ncdc.noaa.gov.

444 **Table 2:** Scores of “representativeness” for the nine stations with the longest records
445 of instrumental surface temperature in the Eastern Mediterranean. The score shows how
446 well a station represents all the other stations. Jerusalem received the highest score. The
447 score is calculated as the ratio between the average of the correlations with the other
448 stations and the average of the respective significance scores. The station with the
449 highest score is the best representative of the temperature fluctuations in the region.

450 **Table 3** Predictors found to be significantly (95%) correlated to Jerusalem DJF
451 temperature. The correlations between predictors and Jerusalem mid-winter
452 temperature are indicated in the right column.

453 **Table 4** Split validation results for the statistical reconstruction model.

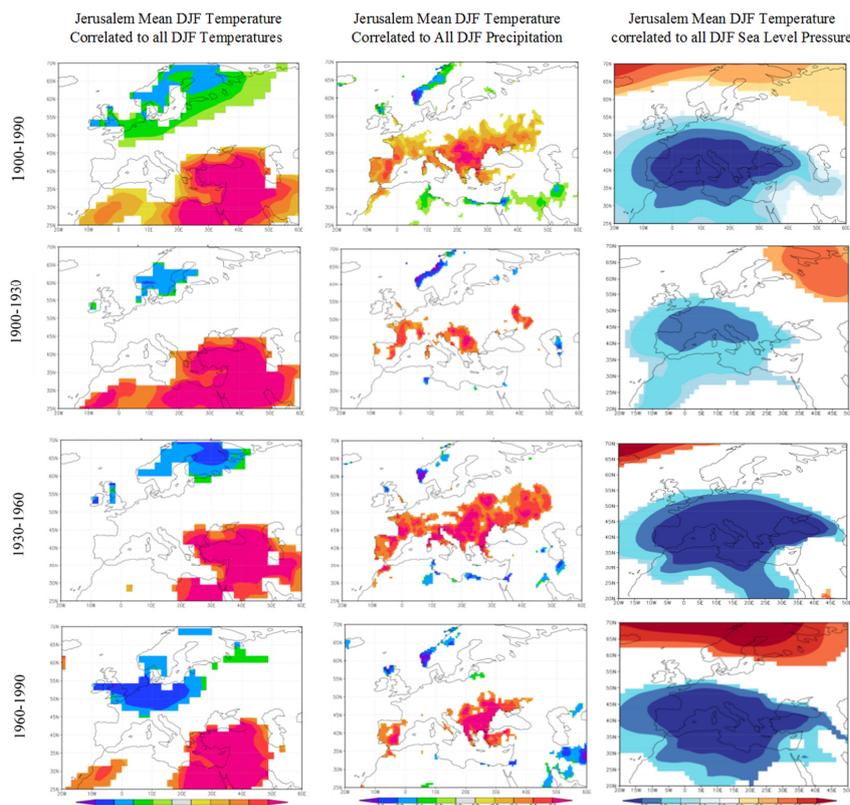
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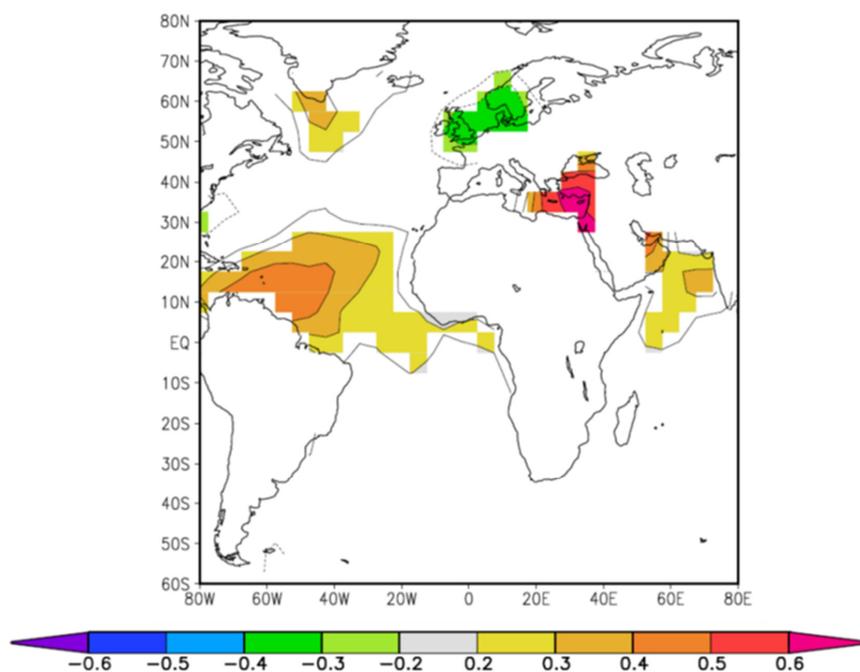
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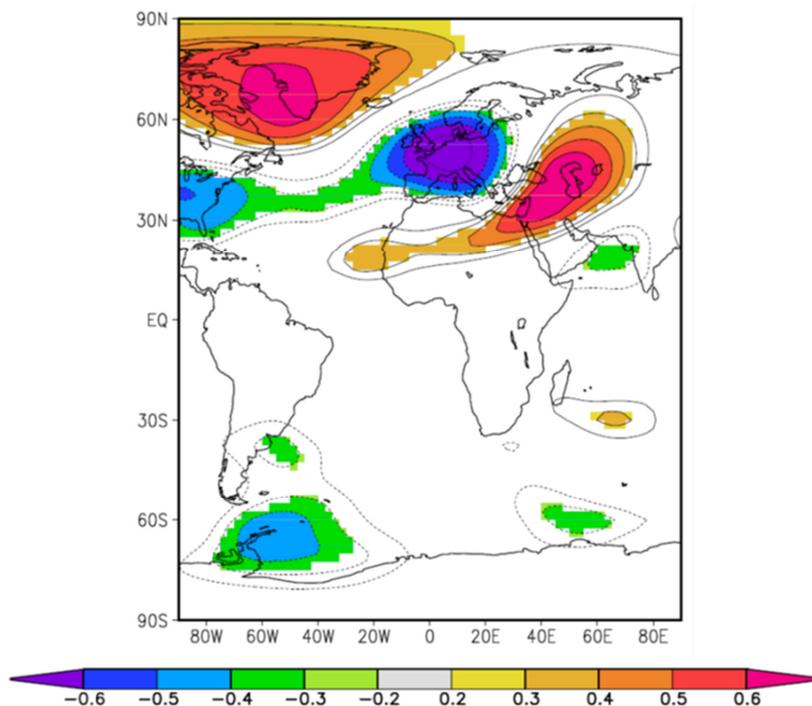
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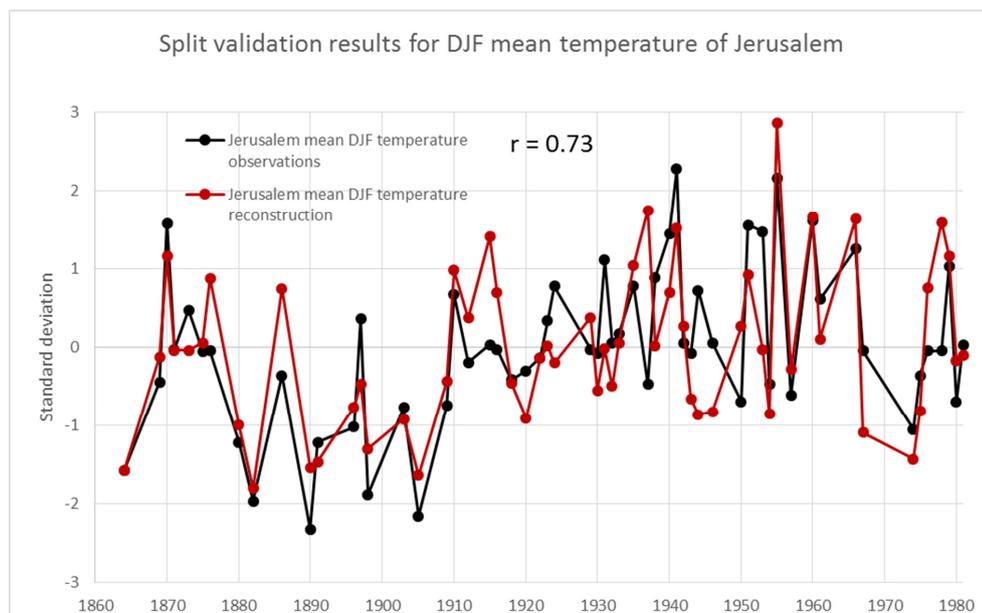
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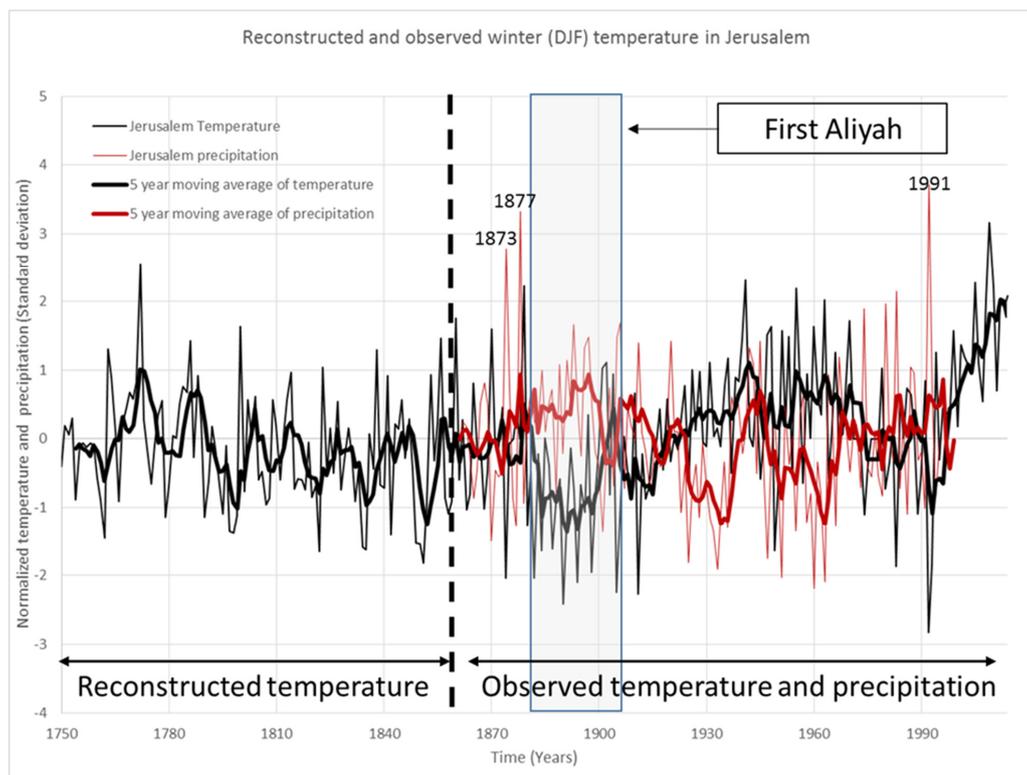
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Country	Site name	Site elevation	Grid point	Number of trees in chronology	Time	Tree type
Cyprus	Ceadar Valley Troocken	1280m	34.59N 32.41S	21	1675-1981	Cyprian Cedar
Cyprus	Ceadar Valley Mittel	1380m	34.59N 32.41S	20	1806-1981	Cyprian Cedar
Cyprus	Plano Platres Feucht	1600m	34.54N 32.54S	22	1742-1981	Austrian Pine, Black Pine
Cyprus	Plano Platres Troocken	1620m	34.54N 32.54S	20	1703-1981	Calbria Pine, Brutia Pine
Cyprus	Trodoos		34.56N 32.52S	32	1628-1981	Austrian Pine, Black Pine
Cyprus	Trodoos Feucht	1900m	34.56N 32.52S	24	1608-1981	Austrian Pine, Black Pine
Cyprus	Trodoos Mittel	1820m	34.56N 32.52S	21	1616-1981	Austrian Pine, Black Pine
Cyprus	Trodoos Mountains	1600m	34.55N 32.53S	13	1594-1981	Austrian Pine
Cyprus	Stavzos Poskas	1050m	35.01N 32.38S	37	1739-2002	Calabrian Pine
Jordan	Dana Resrve and Tor al Iraq	1250m	30.38N 35.30S	17	1469-1995	Phoenician Juniper

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539 **Table 1** Tree ring site chronologies information from Cyprus and Jordan540 (www.ncdc.noaa.gov).

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Station	Representative Score for Mean Temperature Records (Mean correlations/Mean Significance)	Mean correlations	Mean Significance
Jerusalem	381.36	0.56	0.001
Nicosia	322.13	0.62	0.002
Amman	102.1	0.69	0.007
Urfa	8.78	0.65	0.074
Adana	7.78	0.56	0.072
Port-Said	6.12	0.51	0.084
Antalya	4.14	0.51	0.122
Abbaisia	4.01	0.66	0.165
Beyrouth	3.09	0.47	0.153

547

548 **Table 2:** Scores of “representativeness” for the nine stations with the longest records
 549 of instrumental surface temperature in the Eastern Mediterranean. The score shows how
 550 well a station represents all the other stations. Jerusalem received the highest score. The
 551 score is calculated as the ratio between the average of the correlations with the other
 552 stations and the average of the respective significance scores. The station with the
 553 highest score is the best representative of the temperature fluctuations in the region.

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Predictors/Correlation	Jerusalem DJF temperature
Gibraltar SLP	0.51
Dana Reserve Tree ring chronology (Jordan)	0.35
Florence DJF precipitation	0.42
Karlsruhe DJF precipitation	0.35
Larochelle DJF precipitation	0.33
Padua DJF precipitation	0.34
Pouilly DJF precipitation	0.34
Frankfurt DJF precipitation	0.3
Paris DJF precipitation	0.37
Strasbourg DJF precipitation	0.33
Lisbon DJF precipitation	0.45
Udine Rivolto DJF precipitation	0.5

555

556 **Table 3** Predictors found to be significantly (95%) correlated to Jerusalem DJF
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 558 temperature are indicated in the right column.

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Correlation	0.73
Root Mean Squared Error	0.95
Absolute Error	0.76
Relative Error	9%
Prediction Average	8.59

560

561 **Table 4** Split validation results for the statistical reconstruction model.