### UiO Centre for Ecological and Evolutionary Synthesis University of Oslo

The Editor Climate of the Past

Re: Cp-2016-60, 2016

Date: 5 October 2016, Oslo

### **Dear Editor**

Thank you for the positive and constructive referee reports on our paper on Jens Esmark's early temperature observations from the Norwegian capital. We hereby submit a revised version of our paper where due consideration has been made to the comments of the referees. More generally we have altered the structure of the Results chapter to make it more logical and readable, and changed several figures for the same reason.

Our answers to particular points raised by the referees are attached below, also to issues rise in last correspondence with Editor. We thus hope that the paper is substantially improved and will now be suitable for publication in Climate of the Past.

Sincerely yours

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# UiO 🖁

### To editorial remarks 28. Sept. 1016

You have opted not to perform relative testing with Stockholm and Copenhagen, but you now show the Stockholm series. However, visually the correlation seems very good. Does this support your argument that changes in climate patterns could matter here? Please comment on the correlations.

--The correlation for annual values between the homogenized Stockholm series and the homogenized Esmark's series is 0.74, whereas correlation in the modern station networks of Norway is greater than 0.95. Without our corrections of the Esmark's series the correlation to Stockholm would have been less. The test does not tell us which station (or maybe both) of them that might have a possible inhomogeneity. We agree that the chance for a large inhomogeneity of the Stockholm series is less than for the Esmark's series because the Stockholm series is already tested against the nearby Uppsala series. However, the arguments against relative testing are strong. 1) The distance is 350 km between the stations so there is a risk for different climate development during a shorter time interval. 2) Neighbouring stations are only in a narrow sector. Therefore the situation described in 1) could easily occur 3) It would be a pity for climate research of the past if we miss important climate variability in the homogenisation process. There are certainly pro and contras in this issue, but we find the risk too high for "smoothing" away real variability

Section 4.4. using the term "shift" might be misleading here as it is a trend.

--We used a shift test so the only outcome of the test is "shift". See also new comment on Fig 6 below.

Tables: add unit of the shift, perhaps mark those that were corrected in the end.

--Units are added.

Fig. 6: Explain the line.

--Additional text is added, see 4.4

All figures: Remove bounding box.

--Done, and all figures are transferred to pdf-format

# *Conclusions: A quick reader might conclude that the corrections can be applied to the daily scale. Please clarify, as the reviewers also repeatedly were asking you to consider the daily time scale.*

--The decision of using monthly and seasonal data for the homogenization process rather than daily or sub daily values was taken before we started. It seemed too ambiguous homogenizing Esmark's observations on daily data mainly for two reasons, i.e. lacking reference stations and uncertain observation times. Also, our main interest is to study the long term variability in climate, and later, establish an Oslo series starting with Esmark's observations. See also our answer to R3. If we have understood R3 correctly he thinks that





using daily or in particular sub daily data should help us to better assess the observation times. However, the frequency distribution for neighbouring hours does not change much from hour to hour, so in this respect daily data will be of little help. What have helped us, is to study the difference of the mean values between temperatures at the midday observations (where temperature does not vary much from hour to hour) and temperatures at the morning and evening (where temperature vary much from hour to hour). The stability of this difference is our key method of this work. May be one should have used a robust mean, for example the median, rather than the arithmetic mean, but for the Esmark's observations this would not have given other results according to the quality control.

### Ad ref. Przybylak:

- The original fig. 4 has been deleted as superfluous, and figures after no. 3 renumbered in accordance with this.

### Ad anonymous ref. 1:

- The many suggestions for improvements of language and clarifications have all been adopted.

Reference series are not used for the homogeneity testing. The authors justify this choice with the unavailability of contemporary temperature observations near Oslo (Lines 355-356). However, I expect temperature series such as Bergen, Stockholm, Uppsala, and Copenhagen (all available in public datasets such as GHCN-M: https://www.ncdc.noaa.gov/ghcnm/) to be correlated enough to be used as reference series, at least at annual resolution. The absolute tests carried out by the authors give valuable information but could potentially overlook inhomogeneities due for example to changes in the instrumentation, therefore I think that the choice of not using reference series should be reconsidered and an additional relative homogeneity test should be applied after monthly means are calculated (Sect. 4.5).

We think it is important not to use reference stations too far away, in particular when they do not represent the whole circle around the station under testing. In this way wrong conclusions might be drawn as spatial temperature differences could be interpreted as inhomogeneities. For the Esmark series there exist contemporary stations mainly to the east (Stockholm/Uppsala) and south (Copenhagen). From Bergen original data are lacking, and daily data exist only in these intervals: 1818.01-1818.05, 1823.08-1824.07, 1824.08-1824.09, mostly printed in newspapers, and mean values, of questionable quality. We will not use Bergen as reference for the homogeneity testing as it could "contaminate" the results rather than improving them.

In the discussion part we have already compared Esmark's observations with other series from Oslo. Now this section is extended by binging in the Stockholm/Uppsala and the Copenhagen series. However, in the result part we exclude relative testing over those long distances (350 km and 450 km for Stockholm and Copenhagen respectively), also because (as mentioned above) the narrow sector available.

Taking into account some further work (see text) and also your consideration we have dropped the correction for overheating of the midday observation. We have also changed formula for calculating monthly mean temperature. Now, we use Mohn's formula, which is more robust when there is not complete knowledge of the observation times.

When correcting for the inhomogeneity in the evening observation, 1816-1821, a trivial sign error entered the original table submitted. This has been corrected.

Line 355: "For much of Esmark's period of observation there was no other nearby station in operation so internal testing was the only possibility" I disagree on this. Monthly mean temperature anomalies at stations 400 km apart are usually still strongly correlated (r>0.8) (e.g., Auer et al., 2007). It should be possible to use data from Sweden, Denmark for reference and integrate the internal testing results with a relative test.

Same point as above. If Esmark's series was given the same pattern of variability as the stations far away, important spatial climate variability could be hidden. However, we now open for using the available stations Stockholm/Uppsala and Copenhagen for evaluation of the final result, but detecting inhomogeneities by internal testing is our focus. It should be kept in mind that we do not know the exact observation times for the whole period of Esmark's observations. Here internal testing is the only possible way for detecting changes over time.

Line 367: "But this break in homogeneity was much less than that of the morning observation." Less what?

A reference to Table 2 is added.

Line 449: What about the other seasons? According to your formula, you adjust the minimum temperature of 28 February 1838 by 2°C, while the minimum temperature of the following day is not corrected at all! I think here a correction function should be estimated for each month of the year (with some smoothing to better represent the annual cycle if the correction parameters are too noisy).

We do not disagree, but prefer to keep this correction as simple as possible. The aim for the article is to homogenise a series of monthly mean temperatures only. Also, the correction is applied to



minimum temperature only, so the amount of the correction on the monthly means will be much less than 2°C, cf. Fig. 7.

You also ignore other significant breakpoints without explaining why (e.g., 1835 for III vs II in spring in Table 2).

New text has been added.

Line 453: I am not convinced by the evidences for an overheating of the midday observations. You compare with a station with arguably a very different microclimate (different elevation, distant from the sea, etc.). I think that your conclusions should be more conservative, considering the limited information you have on the thermometer exposition and the surrounding environment. You could say that a correction might be necessary for some applications (e.g., analysis of extremes), but for the analysis carried out in this manuscript I don't really see the point of applying such a correction.

There are certainly good arguments for correcting the midday observations, but as you say we have limited knowledge of the microclimate at the station area. We have now made a new comparison with the station Oslo II, where thermometers were well protected by the Astronomical Institute building, and this supports no correction, Fig. 9.

*Line 544: "Also the midday observation is warmer by Hansteen than by Esmark. This is harder to understand." Isn't this because of the overheating correction that you applied to Esmark's data?* 

No, the comparison was done before Esmark's data were corrected (However, now they are not corrected)

*Line 567: Here also the differences with the Astronomical Observatory in summer (Fig. 12) are in large part created by the overheating correction.* 

Yes, an important reason for dropping the correction of the midday observation, see above.

Line 591: The role of Tambora on the climate anomalies in Europe and North America is still debated. Besides, the "paradigm" of the Year Without a Summer is related not only to temperature, but also (and probably more) to precipitation and cloud cover anomalies.

We now note these points in the Discussion

Lines 605-606: Isn't it somewhat surprising that in Bergen, only 200 km or so from Oslo, 1816

was one of warmest years? Is this consistent with the instrumental temperature series of Bergen? Can you comment on the uncertainty of the reconstructions for individual years?

Bergen is situated in quite another climate region than Oslo – oceanic vs. semi-continental.. The distance between the cities is about 300 km. See new comment in the text and reference to literature.

Line 608-616: You cannot reach conclusions on "weather patterns, excessive rain, frost, snowfall" in the summer of 1816 just by analysing the seasonal mean temperature.

Our conclusion concerns temperature only. However, it is important not to forget that Europe is larger than southern and western Europe. In the media we often see the summer of 1816 reckoned as the coldest one in our newest history. The present article contributes to a more nuanced view.

You should rather answer the question: How consistent are Esmark's observations with the results shown by Luterbacher and Pfister (actually, the temperature reconstruction that they use is from Casty et al. (2005))? It would be interesting if you could add the series of the nearest gridpoint of that reconstruction in Fig. 10 and comment on the differences.

Good point. We have done this. See new figure and text.

Re CONCLUSION: This section is incomplete and too synthetic, it should be much improved.

See new text

Table 8: It would be practical to have an additional column with the reference for each reconstruction.

Done

Relevant new references have been included.

### Ad anonymous ref. 2.

While the authors have made extensive efforts to account for data quality and to homogenize the readings for long term climatic analysis in the face of sparse metadata, I am particularly uneasy about the lack of information concerning the observation times, and it is my opinion that further analyses may help reduce this uncertainty. In particular, the authors could make use of frequency analysis as exemplified by the work of Bergström and Moberg (2002) and Slonosky (2014) to compare Esmark's daily morning, afternoon and evening observations to the nearly 25





years of modern hourly observations mentioned in Table 1 and possibly obtain an approximate idea of the times of observation. It may be necessary to sub-divide the historical record for suspected changes in observation time derived from the SNHT analysis and to consider the possibility of observation times, especially in the morning, changing with the season, if this is supported by other metadata (e.g the statement of observation times quoted on lines 188-189). If probable times of observation can be established, the entire analysis will stand on much firmer ground. As it stands, there are many adjustments made on a statistical basis which add to the uncertainty of the final values of the observations, particularly given the differences seen when compared with other nearby observations.

We think that the difference between a relatively stable midday temperature and the temperature at times of the day when temperature are changing most rapidly is the most efficient tool for detecting changes in observation times. On the other hand the shape of the frequency distribution does not change much from one hour to the next, so it will not help us much. The main problems of Bergstrøm and Moberg (2002) and for Slonosky (2014) are not observation times, but the environment of the thermometers. We agree that for those purposes (and in particular for the data quality control) frequency distribution analysis is very well suited. See also our comments to reviewer No. 1.

The accounting of the adjustments due to inhomogeneities detected by the SNHT and other intraseries comparisons is extremely thorough and to be commended, but as is presented leaves the reader confused. A plethora of monthly adjustments is proposed in Tables 2-5, but it is not clear which adjustments were finally applied to which observations, the sequence of the adjustments nor whether the adjustments were applied to the daily data or to the monthly means. If daily, there will be artificial jumps between the end of one month and the beginning of the next - see Vincent et al (2002). In general, more use might be made of the advantages gained by having daily, rather than monthly, observations to analyze; much work had been done in the field of historical climatology in the past decade or two on analyzing daily observations directly.

The Results chapter has been restructured to improve readability. We think that before starting the corrections of the observations it is important to have detected the inhomogeneities. This is what we have done, cf. Table 2 and Table 3. We agree that to interpret the tables might be challenging, but we want to give the reader the opportunity to be able to better judge our work. You claim that "it is not clear which adjustments were finally applied to which observations". We think we have now made this easier to follow, as we have first a detection part (4.1) and then the corrections are discussed for each of the three shifts (4.2, 4.3 and 4.4). We also made it clear that the homogeneity testing was done with monthly and seasonal values, so we have no aim of adjusting daily values. From this article the outcome will not be a series of homogenized daily values. However, we agree that it is important to specify what data and how the results will be stored in the database at the Norwegian Meteorological Institute. See new text.





The fairly large differences shown between these data and other nearby stations, less than 1 km away, also give reasons for concern about the final quality of the data. Comparisons with other series, such as Uppsala -Bergström and Moberg (2002) - and Stockholm –Moberg and Berström (1997), although a considerable distance away, may still give valuable indicators as to the character of each month and help decide which series in the comparisons are the more reliable.

Yes, we have now proceeded further with this comparison, see new text.

Finally, all the data, including the raw data, should be placed in an online archive.

Yes, they will. This is easy in Norway, where the entire network of stations is freely available for everybody.

Introduction, line 45 and thereafter: "protocol" usually refers to a method; it would be less confusing if the authors could use a word like "logbook" or "weather registers" if they mean the actual physical records of Esmark's weather observations.

As Webster's Dictionary defines a protocol as 'an original draft, minute or record' and as neither editor nor any of the other referees have had objections to our use of 'protocol' here, we think the term is appropriate as used.

Line 350: The authors should take note of Gauvin's 2012 article on the Réaumur thermometer: The authors should be aware from Gauvin's work that theoretical adjustment of 1.25 for Réaumur to Celsius may not be accurate. This could help explain some of the large differences seen when comparing Esmark's values to the nearby observations in section 5.

The problems related to the 'Reaumur scale' vs. other temperature scales were already thoroughly discussed in Middleton's book *A History of the Thermometer and Its Uses in Meteorology* (1966), where the conclusion is that by the 1780s a 'Reaumur scale' and 'Reaumur thermometer' had stabilized with thermometer manufacturers that was in fact rather different from the several scales and designs proposed by Reaumur, who died in 1757. For instance alcoholic solutions had been substituted by mercury, and the zero point defined by the melting rather than the freezing of ice etc. etc. Gauvin's paper mainly concerns the confusion surrounding this scale before this period, and is not particularly relevant (and largely recapitulates Middleton's work). The R-scales used by Esmark were certainly of the late 18<sup>th</sup> century kind, where notably the work of Swiss savant Jean-Andre De Luc had cleared up most of the confusion.



It is today a standard convention to convert R-values to C-values by the formula R x 1.25 = C, and although some R-thermometers may not exactly exhibit this calibration with the C, such nonlinear deviations would be much too small to explain the large value differences in Sect. 5.

Section 4.1: The SHNT results seem somewhat ambiguous. Can the SNHT be run on all the 7665 days of observations, rather than dividing up into months and seasons? This might give a clearer indication of the actual break date. If this is too large a number for computational purposes, the series could be tested on running sub-portions (i.e first six months, move forward three months and test next six month period, and so on). Testing on other variables such as pressure might also give a potential indication of a change in the positioning of the instruments. It may help to further divide section 4.1 into subsections dealing with all of the adjustments to each of the three observation times separately and consecutively.

Barometric data have yet to be digitized, and as stated in the metadata we have no reason to believe that barometers and thermometers were situated in exactly the same place.

Line 380: A synopsis of the shifts and dates for each of the observation series would make these clearer. What were the final adjustments made to each series? A table summarizing the actual adjustments applied and the order in which they were applied would be helpful.

See new text.

Line 381: The authors appear to be postulating a replacement of an hourly observation in the morning with a minimum thermometer. Hourly temperature observations and minimum temperature observations are not the same entity. If the authors think that a minimum thermometer was in use, a new series labelled "minimum temperature" should be analyzed. Rather than an inhomogeneity, this is a new variable.

We think our arguments for the minimum thermometer are very strong. We put the minimum temperature to the 'morning' observation only because listed by Esmark as such. Of course information will be to find in the metadata file of the MET-Norway.

Line 390-2/871: This reasoning needs to be better explained, especially given the actual observation times are unknown. What does the description in the title of Table 4 "minimum temperature at 0800 UTC" mean?

We have amended the text to provide a better explanation.





Line 405: More specifics are needed to explain this conclusion: 26% of interpreted "minimum" values being higher than the evening temperature is a high proportion. This unusual temperature trend is a situation which could occur with the passage of frontal systems overnight. What is the proportion of such unusual diurnal temperature trends in the modern record?

In a modern record the minimum temperature comes from the same sensor as the ordinary temperature so this percentage will be 0%. On manual stations it is difficult to say because the meteorological institutes correct there minimum thermometer with a so called "formal test". In practice this means that the minimum temperature will be corrected. In the case of Esmark one should expect more violation of this formal test as he seems not to have used a screen for the thermometers. For modern manual stations both thermometers are located in a common Stevenson screen.

Table 4/ Line 872: What is the authors' interpretation of the negative summer differences for 1816-1828 and 1822-1828, compared to the modern differences? How are these differences changed with the selection of different observation times in the modern period (e.g. 0700, 0600, and sunrise?). Why is the period 1816-1821 corrected but not 1822-1828? Are these results from before or after the application of the adjustment of the 1821 inhomogeneity?

Our first hypothesis was that changed observation time could explain the inhomogeneities, but this is hardly probable, see discussion. The answer to the second question is yes. See new text.

*Line 442: How does this weakened diurnal temperature wave affect the reasoning section 4.1 concerning the minimum thermometer?* 

The effect is zero as the problem with the minimum thermometer only concerned the winter. See new text.

Page 14, lines 447-451, Figure 5: Adjusting from one postulated unknown time to a second unknown time is a procedure beset with uncertainty, particularly as the linear trend does not appear to apply as well in the middle of the period, 1833-1836, when the points would give a much less steep slope. Have the authors explored regressions and residuals for other, finer time resolutions than the three-month period shown in Fig 5? What is the value of the sum of squares error? If better estimates of actual times of observation can't be made, some portion of the data may just have to be classified as unusable.

The same procedure is used as for the other inhomogeneities, detection and then correction. For correction linear regression analysis is used with regression lines shown as equation (2). The standard error of the estimates is  $0.15^{\circ}$ C.





Section 4.5 Again, how and why are these adjustment values derived? This is not clear. See the last answer and text.

Line 457, Figure 8: These adjustments should be presented in a Table separate from the Figure.

We think tables and regression equation is sufficient in addition to Fig. 7 that gives an overview.

Line 482: Section 4.6 should be in the discussion section, while section 5.1 would perhaps be better placed as a summary in section 4. The comparisons with other observers and discussion of the thermometer error would be better placed in a data quality and comparison section, with the climatic discussion in a separate section.

As these errors are not used in the analysis, they have been moved to Appendix B.

Line 541: Again, if we don't know the observation times, it's impossible to attribute the difference between the observers to a specific cause such as instrument location.

For Hansteen we know the observation times. (for Esmark we only know the exact observation time in 1833).

Line 550: is 2100 UTC after sunset in summer?

Yes

Figure 12: This would seem to suggest that the unadjusted values for Esmark are closer to the Observatory than the adjusted values.

Yes, this is one of the reasons for not having applied correction of Esmark's midday observation in this new version.

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2	Jens Esmark's Christiania (Oslo) meteorological observations	
3	1816-1838: The first long term continuous temperature record	
4	from the Norwegian capital homogenized and analysed	Formatted: Underline
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6		
7	Geir Hestmark <sup>1</sup> and Øyvind Nordli <sup>2</sup>	Formatted: Font: Bold
8		
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12	Research and Development Department, Division for Model and Climate Analysis,	
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14		
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16		
17	Abstract	
18	In 2010 we rediscovered the complete set of meteorological observation protocols	
19	made by professor Jens Esmark (1762-1839) during his years of residence in the	
20	Norwegian capital of Oslo (then Christiania). From 1 January 1816 to 25 January	
21	1839 Esmark at his house in Øvre Voldgate in the morning, early afternoon and	
22	late evening recorded air temperature with state of the art thermometers. He also	
23	noted air pressure, cloud cover, precipitation and wind directions, and	
24	experimented with rain gauges and hygrometers. From 1818 to the end of 1838 he	
25	twice a month provided weather tables to the official newspaper Den norske	
26	Rigstidende, and thus acquired a semi-official status as the first Norwegian state	
27	meteorologist. This paper evaluates the quality of Esmark's temperature	
28	observations, presents new metadata, new homogenization and analysis of monthly	
29	means. Three significant shifts in the measurement series were detected, and	
30	suitable corrections are proposed. The air temperature in Oslo during this period is	
31	shown to exhibit a slow rise from 1816 towards 1825, followed by a slighter fall	
32	again towards 1838.	
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36	1 Introduction	
37	The current concern with climate change has increased the interest in early	
38	meteorological observation series and evaluation of their quality ( e.g. Bergström	
39	& Moberg, 2002; Auer et al., 2007). In a recent paper we analysed the temperature	
40	record for the Norwegian capital made 1837-2012 by the astronomical	
41	Observatory at the University of Oslo and the Norwegian Meteorological Institute	
42	(MET Norway) (Nordli et al., 2015). Previous to 1837 long term observations of	
43	the Oslo weather were known to have been made by Jens Esmark (1762-1839),	
44	professor of mining sciences at the University of Oslo (then Christiania). A first	
45	reanalysis of Esmark's observations was made by meteorologist B. J. Birkeland	
46	(Birkeland, 1925). Our rediscovery in 2010 of Esmark's original meteorological	
47	observation protocols has provided an opportunity to digitize, homogenize and	
48	analyze his data with modern methods.	
49	Esmark is today mostly remembered for his pioneer ascents of many of	
50	Norway's highest peaks (Esmark 1802, 1812; Hestmark 2009), his discovery of	
51	Ice Ages, and his astronomical explanation of such dramatic climate change as	
52	caused by variations in the eccentricity of the orbit of the Earth, a hypothesis now	Deleted: is
53	recognized as a precursor of the theories of James Croll and Milutin Milankovich	
54	(Esmark, 1824, 1826; Andersen, 1992; Worsley, 2006; Rudwick, 2008; Berger,	
55	2012; Krüger, 2013). In his own lifetime he was primarily known as a <u>skilful</u>	Deleted: skillful
56	mineralogist and geologist. Throughout his life Esmark maintained a passion for	
57	meteorological observation with instruments he crafted himself in accordance with	
58	the highest contemporary standards. His main inspiration for this activity were his	
59	teachers at Copenhagen University, which he attended 1784-89; first among them	
60	the Astronomer Royal, professor Thomas Bugge (1740-1815), who in his	
61	observatory tower Rundetarn in the middle of Copenhagen made daily	
62	measurements of the weather (Willaume-Jantzen 1896). Esmark also befriended	
63	Bugge's instrument maker, the Swede Johan(nes) Ahl (1729-1795) (Esmark, 1825;	
64	Anonymous 1839). In addition Esmark followed the lectures of Christian Gottlieb	
65	Kratzenstein (1723-1795), professor of medicine and experimental physics, a	
66	'hands on' practical man who enjoyed crafting instruments and all sorts of	
67	mechanical machines (Kratzenstain 1791, Snorrason, 1974, Splinter, 2007). From	
68	1789 to 1791 Esmark studied mining sciences at the Norwegian silver town of	
69	Kongsberg, and after further studies in Freiberg, Saxony and Schemnitz, in today's	Deleted: , Austria-Hungary Formatted: Footer, Right: 0,63 cm

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73	Slovakia, he in 1798 moved back to Kongsberg to take up a position as Assessor in	
74	the central mining administration (Overbergamtet) of the dual kingdom Denmark-	
75	Norway. At Kongsberg he also lectured in mineralogy, geology and experimental	
76	physics at the Royal Norwegian Mining Seminar, acting as its temporary Inspector	Deleted: (Headmaster)
77	from 1799, and permanent Inspector 1802-1815. From 1 January 1799 he three	
78	times a day recorded observations of the Kongsberg weather - air pressure on	
79	mercury barometers (in inches and lines), and air temperature in degrees of	
80	Reaumur; documented in a series of small notebooks running continuously with	
81	some lacunae until 16 September 1810, and rediscovered by the authors in 2010	
82	(Esmark 1799-1810). When Esmark in 1815 moved to the Norwegian capital	
83	Christiania (now Oslo) to become the first professor in the mining sciences at the	
84	University he continued this habit. At least from January 1816 up to and until the	
85	day before his death on 26 January 1839 he recorded air temperature and	
86	barometric pressure three times a day. The complete set of his 23 Christiania	
87	observation protocols, long believed lost, was rediscovered in 2010 by the authors,	
88	and is now safely deposited in the Norwegian National Archive (Riksarkivet)	Deleted: archives
89	(Esmark 1816-1838). They provide a unique and detailed picture of the weather in	
90	Oslo in the early 19 <sup>th</sup> century. From January 1818 to December 1838 tables of	
91	Esmark's observations were published every fortnight in the official newspaper	
92	Den norske Rigstidende (cf. Appendix A), and he thus acquired a semi-official	
93	position as Norway's first state meteorologist. Based on a number of previously	
94	unpublished documents (cited as Document 1 etc, with archival location in	
95	Reference list) we here present new metadata for Esmark's meteorological	
96	observations from Christiania, and homogenize, analyse and evaluate his original	Deleted: reanalyse
97	temperature data with modern statistical tools to characterize the temperature	Deleted: weather
98	variations in the Norwegian capital in this period,	Deleted:
99		

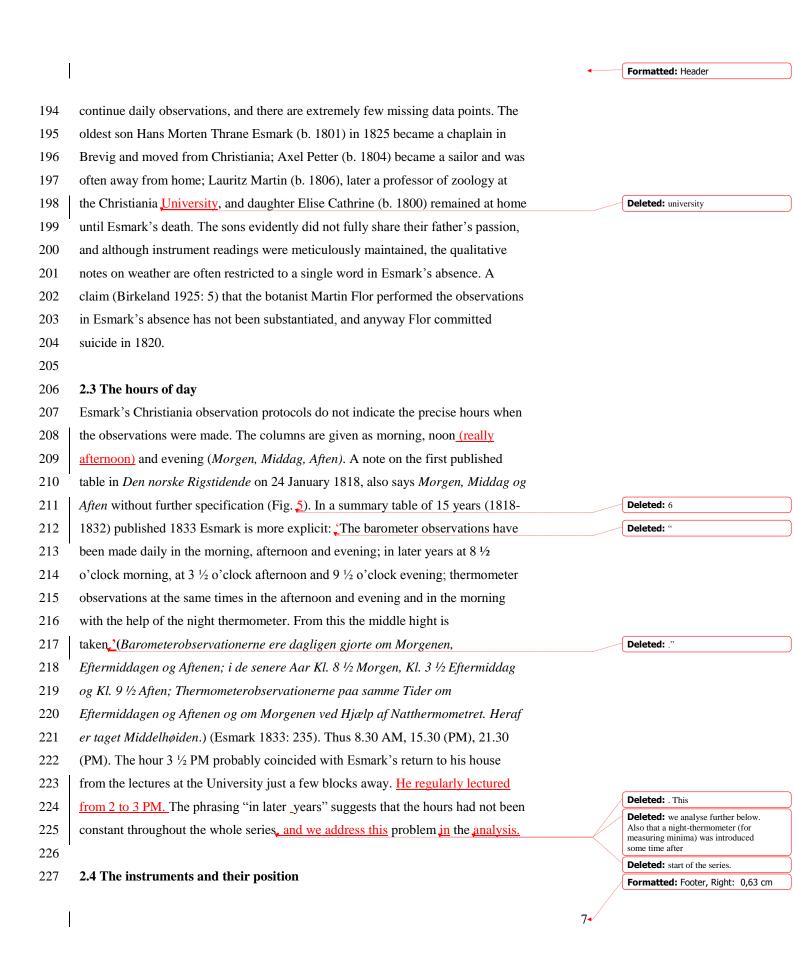
### 100 2 Metadata

101	2.1 The location - No. 308, Vestre Rode - Øvre Vollgate 7.	
102	Esmark's observations were made at his home (cf. Esmark <u>1823</u> : <i>De ere tagne i</i>	Deleted: 1823b
103	min Bopel), and there is no evidence indicating that he changed the location. On 19	
104	August 1815 Esmark was registred as owner of property No. 308 in Vestre Rode	
105	(i.e. Western Quarter), one of the four old quarters of Christiania town (Document	
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112	1). It was a modest one-and-half storey house built late in the 18 <sup>th</sup> century with an	
113	adjoining garden. Esmark's continued residence at this address until his death is	Deleted: a
114	documented in annual censuses and tax protocols (Document 2). Property No. 308	Deleted: & 3
115	was situated on the north-western side of the street Øvre Vollgate (Øvre	
116	Woldgaden), laid out literally on what used to be the outermost western rampart	
117	(voll) of nearby Akershus Castle and Fortress (Fig. 1). It was a natural rock	
118	promontory above a meadow to the west where the poor fishing village Pipervigen	
119	would develop later in the 19 <sup>th</sup> century, today the site of Oslo Town Hall. In 1815	
120	Øvre Vollgate constituted the south-western limit of Christiania, a town with only	
121	about 15000 citizens (Myhre 1990). Until 1814 the main administration centre of	
122	the dual kingdom was in Copenhagen, but with Christiania in that year acquiring	
123	the new parliament _and government after the separation of Norway from	
124	Denmark, the town expanded rapidly. When street numbers were introduced,	
125	Esmark's property was numbered Øvre Vollgate No. 7. The present Øvre Vollgate	Deleted: Vollgt
126	7 – an office highrise – comprises previous numbers $\emptyset$ vre Vollgate 3, 5 and 7.	
127	Esmark's property No. 308 and all neighbouring properties were measured	
128	and mapped for the new matriculation of Christiania in the summer of 1830, and	
129	thus we have very precise data on his house and the surrounding properties at the	
130	relevant time (Document $\underline{3}$ ). The whole property roughly constituted an elongated	Deleted: 4
131	rectangle, approximately 14 m x 60 m (Fig. 2). The unit used in these	
132	measurements was the 'Norwegian alen' (Norsk alen), determined by law in 1824	
133	to be 62.75 cm. It was divided into two feet, each divided into 12 inches, each	
134	divided into 12 lines. No. 308 was measured to 2026 square alen, of which the	
135	house (including a yard) was 733 1/2 and the garden 1292 1/2 square alen (1 square	
136	alen = $0.3937 \text{ m}^2$ ). Thus the whole property was ca. 800 m <sup>2</sup> , and the house	
137	(including yard) ca. 290 m <sup>2</sup> . The house had a 22 alen 6 inch (ca. 14 m) long	
138	façade towards the street Øvre Voldgate, constituting the south eastern border of	
139	the property, with windows, doors, and a gate leading in to the back yard (Fig. 3).	
140	Øvre Vollgate street runs from SW to NE at an angle of roughly 32° NE (400	
141	degrees). At the back the house surrounded a small yard, with a narrow passage	
142	opening out to the garden in the NW. As it would have been hazardous to place the	
143	meteorological instruments on the street-side of the house, where passers-by could	
144	tinker with them, it is almost certain that they were placed in Esmark's back yard,	<b>Deleted:</b> have tinkered
145	a <u>wellguarded</u> space. When the house was finally demolished in 1938, it was in	Deleted: well guarded
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152	such <u>a</u> bad condition that the Oslo city health authorities demanded the whole	
153	property to be sprayed with hydrocyanic acid and that none of the fungus-infected	
154	material be used for construction elsewhere (Document <u>4</u> ).	Deleted: 5
155	Esmark's garden on the NW side of the house and court yard was a	
156	continuous slope, dropping ten alen (6,25 m) down along 66 alen length towards	
157	Pipervika. Here it was most probably limited by a fence towards the Præste Gade	
158	street which later changed name to today's Rosenkrantz street. In 1841, a couple of	Deleted: todays
159	years after Esmark's death, most of this garden was indeed sectioned out and sold	Deleted: gate
160	to form the new property Rosenkrantz gate 26. In Esmark's time, however, the	
161	promontory remained an open garden space. His neighbours on both sides (No.	
162	307 and No. 309) had the same arrangement of house and garden, with facades to	
163	Øvre Vollgate and gardens sloping down on the back to Præstegaden (Document	
164	5). To the north of the lowermost part of Esmark's property was an open space	Deleted: 6
165	called Jomfru Wold's Løkke (No. 368). South of this lower part of the garden was	
166	the street Pipervigbakken, leading down from Rådhusgaten street passing by the	
167	outer ramparts of Akershus fortress and Castle. The sea with Pipervika bay	
168	(Piperviks Bugten) was less than 200 m south of Esmark's garden. His garden was	
169	not an entirely constant environment. In 1823 for instance, he received several	
170	fruit trees from a Danish friend which he planted in the garden (Document <u>6</u> ).	Deleted: 7
171	It was a modest residence for a professor, situated in a comparatively poor	
172	part of town, with mainly craftsmen, tradesmen and artisans in the neighbourhood	
173	(Myhre 1990: 40). Here Esmark, a widower since 1811, moved in with his three	
174	sons Hans Morten, Petter and Lauritz, a maid and a manservant (Document _2).	Deleted: & 3
175	His daughter Elise resided with her grandparents in Copenhagen, but later returned	
176	to Norway to take up residence in No. 308.	
177		
178	2.2 The observers	
179	The great majority of the Christiania observations were made and noted down by	
180	Esmark himself who has an easily recognizable handwriting. His position as	
181	professor in the mining sciences did however sometimes cause him to leave town	
182	on short or long field excursions, some lasting several months. He was away from	
183	Christiania on long voyages in 1818 (Hallingdal), 1819 (Kristiansand), 1822	Deleted: );
184	(Bergen), 1823 (round-trip south Norway), 1826 (Setesdalen), 1827 (Trondhjem)	
185	and 1829 (Copenhagen). In his absence his sons seem to have been instructed to	Deleted: absense Formatted: Footer, Right: 0,63 cm



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January 1818, Esmark provides a few details of his measurements: "The	
observations are made 34 Rhinelandic feet [i.e. 10.68 m] above the sea, and are the	
middle value of observations made morning, noon and evening. The barometer	
heights are corrected as they would have been if the barometer was subject to a	
temperature of 0°. The thermometer hangs freely against north, (Observationerne	Deleted: ."
ere anstillede 34 Rhinlandske Fod over Havet, og ere Middeltallet af	
Observationer, anstillede Morgen, Middag og Aften. Barometerhøiderne ere	
corrigerede saaledes, som de skulle være, dersom Barometret havde været udsat	
for 0° Temperatur. Thermometret hænger frit imod Nord.) (Fig. 5). Esmark also	Deleted: 6
notes that The barometer height is reduced to $0^{\circ}$ R. If one wants it reduced to sea	Deleted: for these (average?) data
level, one must add a line or 1/12 of an inch to its height, so that the barometer	Deleted: "
height at sea level becomes 28.1,20 in French measure (Barometerhøiden er	Deleted: ."
reduceret til 0° R. Vil Man have den reduceret til Havets Overflade, maa Man til	
den anførte Høide lægge en Linie eller 1/12 Deel af en Tomme, saa at	
Barometerhøiden ved Havets Overflade bliver 28.1,20 i Fransk Maal.) (Esmark	Formatted: English (U.S.)
1833: 235).	
Thermometers. Esmark all his life used the Reaumur scale; R. The precision of his	Deleted: "
Reaumur thermometer was 1/2 of a degree. On a table of averages for the years	Deleted: ".
1816-1822 Esmark notes: The thermometer observations are made in shadow in	Deleted: "
free air with a Reaumur thermometer, which boiling point is determined at 28	
inches 2 lines (French measure) barometric height, '(Thermometerobservationerne	Deleted: ."("
ere gjorte i Skyggen i fri Luft med et Reaumurs Thermometer, hvis Kogepunkt er	
bestemt ved 28 Tommers 2 Liniers (fransk Maal) Barometerhöide <u>.) (Esmark</u>	
•	
1823). In Esmark's observation protocol for the year 1816 some instrumental	
	Moved (insertion) [1]
1823). In Esmark's observation protocol for the year 1816 some instrumental	Moved (insertion) [1]
1823). In Esmark's observation protocol for the year 1816 some instrumental corrections are given for what is claimed to be Esmark's thermometer, They are	Moved (insertion) [1]
1823). In Esmark's observation protocol for the year 1816 some instrumental corrections are given for what is claimed to be Esmark's thermometer, They are not written by Esmark himself, most probably they are notes written by Birkeland,	Moved (insertion) [1] Deleted: . ") (Esmark 1823).
1823). In Esmark's observation protocol for the year 1816 some instrumental corrections are given for what is claimed to be Esmark's thermometer, They are not written by Esmark himself, most probably they are notes written by Birkeland, who says he has them after Hansteen 1821-23, but it is not certain that they belong	
1823). In Esmark's observation protocol for the year 1816 some instrumental corrections are given for what is claimed to be Esmark's thermometer, They are not written by Esmark himself, most probably they are notes written by Birkeland, who says he has them after Hansteen 1821-23, but it is not certain that they belong to the thermometer used by Esmark. The corrections are listed in Appendix B but	

271 simple barometer, the tube of which is 2 1/2 line in diameter and which capsul is 40

In a note to his first table presented in the journal Den norske Rigstidende, on 24

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Formatted: Header 283 lines in diameter, and calibrated after a siphon barometer.' (Barometret er et **Deleted:** hevertharometer 284 simpelt Barometer, hvis Rør er 2 <sup>1</sup>/<sub>2</sub> Linie i Diameter og hvis Capsel er 40 Linier i Diameter, samt justeret efter et Hævertbarometer.) 285 Moved (insertion) [2] 286 287 2.5 The protocols and data recorded 288 Esmark's Christiania protocols are handmade, folded sheets of white paper cut up 289 and sewn in with a thin grey cardboard cover, one protocol for each year, 23 Moved up [2]: ¶ 2.5 The protocols and data recorded Esmark's Christiania protocols are 290 protocols in all (Esmark 1816-1838). Esmark interfoliated the official printed handmade, folded sheets of white paper cut up and sewn in with a thin grey 291 Almanach for Christiania. This had for each month 16 days on each page, and thus cardboard cover, one protocol for each vear 292 Esmark wrote down his data for 15 or 16 days on the first page of a month and the Deleted: (Fig. 4), 293 remaining days from 17 to 28, 29, 30 or 31 on the next page (Fig. 4). The protocols Deleted: 5 294 start on 1 January 1816 and end 31 December 1838, only 26 days before his death; 295 altogether 8401 days of continuous measurements. There are only a few small 296 lacunae. Photographs of all the protocols are available at MET Norway (Klimadata 297 samba server, HistKlim skanna dokument), and digitized values, converted from Deleted: might 298 <sup>o</sup><u>R to <sup>o</sup>C, can</u> be downloaded from MET Norway's home page: <u>http://www.met.no</u>. 299 Esmark & sons continued observations in January 1839 until the day before his death 26 January, but these observations are only known through the newspaper 300 301 Morgenbladet, which had published Esmark's daily measurements since 1834. Formatted: Font: Not Bold, English (U.S.) 302 Three times a day Esmark recorded temperature to a half degree, and air 303 pressure in inches and lines (Fig. 4). In the right hand margin he noted the weather Deleted: with one or two decimals Deleted: 5 304 (Veirliget) with qualitative terms; see also Esmark (1833). He used a fairly limited 305 number of categories: Precipitation: lidt Regn (a little rain); Fiin Regn (drizzle); Deleted: drissle 306 Regn (rain); Regn Bygger/Bÿgger (showers); Regn af og til (Rain now and then); 307 megen Regn (much rain); Sne (snow); Sne Flokker (snowflakes); Sne Bygger Deleted: snow 308 (snow showers). Cloud cover: Klart (clear), enkelte Skyer (a few clouds); tynde 309 *Skyer* (thin clouds); *skyet* (<u>cloudy</u>); *skyer i Horizonten* (clouds in the horizon); Deleted: cludy 310 disig (haze); Taage (fog). The most common category was tykt (thick) which 311 means a grey day with haze, often with precipitation. Wind: Wind direction was usually recorded only once a day, in the afternoon, with categories N, S, V and O, 312 Deleted: at midday 313 and combinations, e.g. N. O. (nord ost/north easterly). Other: Torden (thunder); 314 Nordlys (northern lights); Flekker i Solen (sunspots); one or two circles around the 315 sun; *Høyt vand* (high sea level). In June 1818 Esmark introduced a new parameter: 316 precipitation, measured with a rain gauge, and in the June summary, he could Formatted: Footer, Right: 0,63 cm

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334 a	announce: In this month there has, according to the rain gauge, fallen rain to a	Deleted: "
	height, which, if it had been standing, had constituted a height of 1 inch and 9 and	
336 7	7/12 line. The rain gauge is situated 15 feet above sea level. The low altitude of	Deleted: ."
	the rain gauge suggests that it was placed at the lower part of the slope in his	Formatted: Not Superscript/ Subscript
338 g	garden. In October 1820 he presented the readers of <i>Rigstidende</i> to his new design	
339 f	for a hygrometer – an instrument to measure air humidity (Esmark, 1820). It was	
340 r	modified from a model developed by John Livingstone, a M.D. from Canton,	Deleted: and
341 (	China, published in the <i>Edinburgh Philosophical Journal</i> in 1819 (Livingstone	
342 1	1819). The general idea was to put a moisture absorbing/releasing chemical	
343 s	substance (Livingstone used pure sulphuric acid, which was also used to produce	
344 i	ce) on one side of a balance, balanced against a weight on the other side. The	
	palance was placed under a glass jar open in the bottom to let air freely flow in and	
346 o	but, and to protect it from precipitation. Esmark made two new hygrometers	
347 a	according to this model. Anyone who desires to see these hygrometers, can see	Deleted: "
348 t	hem at my <u>house' (</u> Enhver, som har Lyst dertil, kan see disse Hygrometere hos	Deleted: house" ("
349 n	nig.)(Esmark, 1820) He had tested them for several months, and thought they	Deleted: . ")(
350 c	could be used by farmers to predict weather change as <u>a</u> substitute for barometers.	
351 H	He did not, however, use the hygrometer data for his meteorological tables. For the	
352 y	year 1821 he presented more regular monthly data on precipitation in inches –	
353 f	from 1 May through October – apparently the months without frost.	
354		
355 2	2.6 The published tables	
356 \$	Starting on Saturday 24 January 1818, with a table presenting weather data for the	
357 f	first half of the month, the semi-official daily Den norske Rigstidende published	
358 H	Esmark's meteorological observations, which thus acquired an official air. (Fig. 5).	Deleted: 6
359 I	It became a regular series, published twice a month – one table for the first half of	
360 t	the month, one for the second half $-a$ total of 24 tables each year, all with the	
361 s	same title "Meteorologiske lagttagelser i Christiania [year], anstillede af Prof.	
362 H	Esmark." (Meteorological observations in Christiania [year], made by Prof.	
363 H	Esmark) etc, This series running from 1 January 1818 to 15 December 1838 is	Deleted:
364 a	absent from all previously published bibliographies of Esmark's works, but in fact	
365 r	runs to no less than 503 published tables (!) (Appendix A). They present 7665 days	
366 o	of continuous observations. In addition comes the two full years of 1816 and 1817,	
367 o	only published summarily by Esmark (1823) but with complete record preserved	Formatted: Footer, Right: 0,63 cm
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276	in the original material. The whole even 1919 was summed on an 9 January 1910	
376 377	in the original protocols. The whole year 1818 was summed up on 8 January 1819	
	with means etc., and here Esmark also compared the Christiania data to those	
378	obtained by Wargentin in Stockholm, by Bugge in Copenhagen, and (no	
379	observator given) in St. Petersburg, Russia. It was not a weather forecast but	
380	rather a weather 'backlog', and this may have dimmed their public interest	
381	somewhat. The data given in these published tables differ from the raw data of the	
382	protocols by being daily averages. For each day he gave the barometric pressure	
383	and temperature, averaged from observations made in the morning, <u>afternoon</u> , and	Deleted: at noon Deleted: in the
384	evening (at first without further precision of hour). To calculate these averages he	
385	apparently used the formula:	
386	$T_m = \frac{1}{4} \left( T_I + 2T_{II} + T_{III} \right) \tag{1}$	
387	where $T_m$ is Esmark's daily <u>temperature</u> , and $T_I$ , $T_{II}$ , and $T_{III}$ are the	Deleted: "mean"
388	observed temperature morning, afternoon and evening, respectively. To the tables	Deleted: noon
389	for the second half of each month, he also appended a note with the mean	
390	barometric pressure and temperature for the entire month, and indicated which	
391	days had the maximum and minimum air pressure and temperature. The mean	
392	temperature was given to $1/100^{\text{th}}$ degree (a spurious precision). The series	
393	continued in 1820, now also with the daily wind direction. Esmark evidently	
394	trusted only himself to calculate the means and set up the tables, and thus the	Deleted: averages
395	readers of Rigstidende sometimes had to wait for months to read the weather for	
396	the last fortnight, when he was off on some excursion. From 1834 Esmark's	Deleted:
397	observations were also published in the Christiania newspaper Morgenbladet every	
398	day, with two days delay, i.e. observations for the 1 <sup>st</sup> day of the month were	
399	published on the 3 <sup>rd</sup> etc. This was initiated after Christiania doctors suspected a	
400	connection between the weather and the cholera epidemics which struck Norway	
401	from 1833 and forward,	Deleted:
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404	3 Methods	
405		
406	3.1 Homogeneity testing	

407 A homogenous climatic time series shows variations in climate without being disturbed by

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408 other factors involved, like changes in the environment, observational procedures or

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416	instrument calibration. For the study of climate variations the use of homogenous series is of	
417	paramount importance, otherwise the climate analysis might be wrong (e.g. Auer et al., 2007;	
418	Moberg and Alexandersson, 1997; Tuomenvirta, 2001). For testing the homogeneity of	
419	Esmark's temperature series we selected the Standard Normal Homogeneity Test (SNHT)	
420	with significance level = $0.05$ , which has been widely used for testing of both precipitation	Deleted: that
421	series and temperature series (Alexandersson, 1986; Alexandersson and Moberg, 1997;	
422	Ducré-Robitaille et al., 2003). The first version of the test (Alexandersson, 1986) had one step	
423	change as the only possibility, whereas in the version of 1997 both double shifts and a trend	
424	were possible outcomes of the test. In any year the significance of a potential break is	
425	examined. The testing followed the principle of comparing a candidate series (the series under	
426	testing) against a reference series. The reference might be series from one or more	
427	neighbouring stations. A candidate series might also be observations at one particular time of	
428	the day, which are compared with other observation times for the same station. In the latter	
429	case we call it "internal testing". <u>Contemporary</u> neighbouring <u>series overlapping Esmark's</u>	<b>Deleted:</b> Without contemporary
430	observations are too short to be used in the homogeneity testing. The nearest stations that	Formatted: English (U.S.)
431	could have been used are Stockholm/Uppsala about 350 km from Oslo. The problem with	
431	using series so far away is that spatial temperature variations could be interpreted as	
		Deletedu is the only needikility. If as
433	inhomogeneities. Therefore our chosen method is internal testing. Later measurement series	<b>Deleted:</b> is the only possibility. If no significant break occurs the series is considered homogenous. Esmark's station
434	from observation stations in the Oslo area may however be of some use in some analyses, and	at Øvre Vollgate 7 as well as other
435	these are listed with Esmark's in Table 1, with their national station number (identifier) and	<b>Deleted:</b> used in this article are given
436	name. While the official names of the stations refer to their sites we will in the text for	Deleteu.,
437	convenience often refer to the names of the observers, i.e. the column 'additional information'	
438	in Table 1. Before the analysis started all observations in degrees of Reaumur were converted	Deleted: were calculated from degree
439	into degrees of Celsius by multiplying by the factor 1.25.	<b>Deleted:</b> to degree <b>Deleted:</b> Esmark's Reaumur readings
440		
441	4 Results	
442		Deleted: Homogeneity testing¶ For much
443	4.1 Detection of inhomogeneities	Formatted: Font: Bold
444	First we will use SNHT for detection of the inhomogeneities and thereafter treat each	Deleted: Esmark's period
445	inhomogeneity in more detail, and come up with corrections. The testing was performed both	<b>Deleted:</b> observation there was no other nearby station
446	for seasonal (Table 2) and monthly (Table 3) resolutions where observations taken in the	<b>Deleted:</b> operation so internal testing was the only possibility.
447	morning (I), <u>midday</u> (II) and evening (III) were compared with each other. By comparing	Deleted: see
448	several test results it was possible to decide at which observation time a shift (inhomogeneity)	Deleted: see
449	occurred. Most striking are the huge shifts detected in spring, summer and autumn when the	Deleted: noon

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471	morning observation was involved. The most probable year for the shift was 1827; in	
472	particular this was true for the single shift test. Here we apply the common convention to	
473	define the shift year as the last year before the shift. We have to conclude that the morning	
474	observation is inhomogeneous. Further investigation of the daily observations (not shown)	Deleted: A further
475	suggested that the change took place in the month of March 1828.	Deleted: within
	When the evening observation was tested against the midday observation a shift	Deleted. within
476		
477	seemed to occur in 1820 or 1821, most probably in 1821. But this break in homogeneity was	
478	much less than that of the morning observation <u>, cf. Table 2.</u> The shift seems to be absent or	Deleted: .
479	very weak during winter so exact dating was impossible. For convenience the end of 1821	
480	was adopted as the time of the inhomogeneity.	Deleted: year Deleted:
481	Tests including the midday observation revealed no additional shifts than those	
482	already detected. The occurrence of the shifts in the tests I vs II and III vs II seemed to reflect	
483	shifts either in the morning or in the evening observations. For the winter season a shift in the	
484	last part of the series was detected, possible shift years were 1832, 1833 or 1834.	
485	The large shift in the morning observation could have masked possible smaller shifts in the	
486	series on both sides of this shift. Therefore the single shift SNHT was applied on two different	
487	parts of Esmark's series: 1816.01-1828.02 and 1828.03-1838.12, parts 2 and 3 in Table 2.	
488	However, no further shifts in the series were detected	<b>Deleted:</b> The shifts detected in part 1 in the evening observations of 1821 and in the
488 489	However, no further shifts in the series were detected.	the evening observations of 1821 and in the morning observation in the 1830s for the
	However, no further shifts in the series were detected. Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in	the evening observations of 1821 and in the
489	•	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed.
489 490	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed.
489 490 491	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in 1827 (probably 1828.02) for the morning observation and one during winter with possible	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed.
489 490 491 492	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in 1827 (probably 1828.02) for the morning observation and one during winter with possible	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed.
489 490 491 492 493	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in 1827 (probably 1828.02) for the morning observation and one during winter with possible shift years 1832, 1833 or 1834. We now proceed to propose corrections.	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed.
489 490 491 492 493 494	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in 1827 (probably 1828.02) for the morning observation and one during winter with possible shift years 1832, 1833 or 1834. We now proceed to propose corrections.	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed.
489 490 491 492 493 494 495	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in 1827 (probably 1828.02) for the morning observation and one during winter with possible shift years 1832, 1833 or 1834. We now proceed to propose corrections. <b>4.2 Correcting the shift in 1821.12 in the evening observation</b> This inhomogeneity was corrected by using the midday observation that came out of the	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed. <b>Deleted:</b> The reliability
489 490 491 492 493 494 495 496	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in 1827 (probably 1828.02) for the morning observation and one during winter with possible shift years 1832, 1833 or 1834. We now proceed to propose corrections. <b>4.2 Correcting the shift in 1821.12 in the evening observation</b> This inhomogeneity was corrected by using the midday observation that came out of the SNHT as homogenous. The monthly mean difference between the midday observation and the	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed. <b>Deleted:</b> The reliability <b>Deleted:</b> results was further tested on
489 490 491 492 493 494 495 496 497	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in1827 (probably 1828.02) for the morning observation and one during winter with possibleshift years 1832, 1833 or 1834. We now proceed to propose corrections. <b>4.2 Correcting the shift in 1821.12 in the evening observation</b> This inhomogeneity was corrected by using the midday observation that came out of theSNHT as homogenous. The monthly mean difference between the midday observation and theevening observation on each side of the shift was calculated. Then the evening observation	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed. <b>Deleted:</b> The reliability <b>Deleted:</b> results was further tested on
489 490 491 492 493 494 495 496 497 498	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in 1827 (probably 1828.02) for the morning observation and one during winter with possible shift years 1832, 1833 or 1834. We now proceed to propose corrections. <b>4.2 Correcting the shift in 1821.12 in the evening observation</b> This inhomogeneity was corrected by using the midday observation that came out of the SNHT as homogenous. The monthly mean difference between the midday observation and the evening observation on each side of the shift was calculated. Then the evening observation was corrected by adding monthly correction terms so that this mean difference was constant	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed. Deleted: The reliability Deleted: results was further tested on Deleted: resolution and Deleted: resolution and
489 490 491 492 493 494 495 496 497 498 499	<ul> <li>Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in 1827 (probably 1828.02) for the morning observation and one during winter with possible shift years 1832, 1833 or 1834. We now proceed to propose corrections.</li> <li><b>4.2 Correcting the shift in 1821.12 in the evening observation</b></li> <li>This inhomogeneity was corrected by using the midday observation that came out of the SNHT as homogenous. The monthly mean difference between the midday observation and the evening observation on each side of the shift was calculated. Then the evening observation was corrected by adding monthly correction terms so that this mean difference was constant on each side of the shift. It is most common to correct the early part of the series so this was</li> </ul>	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed. <b>Deleted:</b> The reliability <b>Deleted:</b> results was further tested on <b>Deleted:</b> resolution and
489 490 491 492 493 494 495 496 497 498 499 500	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in 1827 (probably 1828.02) for the morning observation and one during winter with possible shift years 1832, 1833 or 1834. We now proceed to propose corrections. <b>4.2 Correcting the shift in 1821.12 in the evening observation</b> This inhomogeneity was corrected by using the midday observation that came out of the SNHT as homogenous. The monthly mean difference between the midday observation and the evening observation on each side of the shift was calculated. Then the evening observation was corrected by adding monthly correction terms so that this mean difference was constant on each side of the shift. It is most common to correct the early part of the series so this was done also here. Therefore the period 1816.01-1821.12 was corrected, whereas the rest of the	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed. Deleted: The reliability Deleted: results was further tested on Deleted: resolution and Deleted: resolution and
489 490 491 492 493 494 495 496 497 498 499 500 501	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in 1827 (probably 1828.02) for the morning observation and one during winter with possible shift years 1832, 1833 or 1834. We now proceed to propose corrections. <b>4.2 Correcting the shift in 1821.12 in the evening observation</b> This inhomogeneity was corrected by using the midday observation that came out of the SNHT as homogenous. The monthly mean difference between the midday observation and the evening observation on each side of the shift was calculated. Then the evening observation was corrected by adding monthly correction terms so that this mean difference was constant on each side of the shift. It is most common to correct the early part of the series so this was done also here. Therefore the period 1816.01-1821.12 was corrected, whereas the rest of the	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed. Deleted: The reliability Deleted: results was further tested on Deleted: resolution and Deleted: resolution and
489 490 491 492 493 494 495 495 496 497 498 499 500 501 502	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in 1827 (probably 1828.02) for the morning observation and one during winter with possible shift years 1832, 1833 or 1834. We now proceed to propose corrections. <b>4.2 Correcting the shift in 1821.12 in the evening observation</b> This inhomogeneity was corrected by using the midday observation that came out of the SNHT as homogenous. The monthly mean difference between the midday observation and the evening observation on each side of the shift was calculated. Then the evening observation was corrected by adding monthly correction terms so that this mean difference was constant on each side of the shift. It is most common to correct the early part of the series so this was done also here. Therefore the period 1816.01-1821.12 was corrected, whereas the rest of the series was not. The corrections are given in Table 4.	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed. Deleted: The reliability Deleted: results was further tested on Deleted: resolution and Deleted: resolution and
<ul> <li>489</li> <li>490</li> <li>491</li> <li>492</li> <li>493</li> <li>494</li> <li>495</li> <li>496</li> <li>497</li> <li>498</li> <li>499</li> <li>500</li> <li>501</li> <li>502</li> <li>503</li> </ul>	Thus there are three shifts that seem reliable, one in 1821 for the evening observation, one in 1827 (probably 1828.02) for the morning observation and one during winter with possible shift years 1832, 1833 or 1834. We now proceed to propose corrections. <b>4.2 Correcting the shift in 1821.12 in the evening observation</b> This inhomogeneity was corrected by using the midday observation that came out of the SNHT as homogenous. The monthly mean difference between the midday observation and the evening observation on each side of the shift was calculated. Then the evening observation was corrected by adding monthly correction terms so that this mean difference was constant on each side of the shift. It is most common to correct the early part of the series so this was done also here. Therefore the period 1816.01-1821.12 was corrected, whereas the rest of the series was not. The corrections are given in Table 4. The corrections are largest in the months where the daily temperature wave is largest, so one	the evening observations of 1821 and in the morning observation in the 1830s for the winter season were confirmed. Deleted: The reliability Deleted: results was further tested on Deleted: resolution and Deleted: resolution and

Formatted: Header 519 speaking we know Esmark's observation times only in 1833, so this hypothesis is not in 520 contradiction to metadata. But observation times cannot be the only reason for the shift, 521 because it appeared also in midwinter when the daily temperature wave is weak. Moreover, 522 the amounts of the corrections are so large that only observation times near midnight would 523 compensate for the low values of the evening observation. Observation times that late seem 524 unlikely. There is some indication that a changed environment could have played a role for 525 this inhomogeneity as Esmark in 1823 planted fruit trees in his garden, cf. Metadata. A one year mismatch of the shift detected by the SNHT is not uncommon. 526 527 4.3 Correcting the shift in 1828.02 in the morning observation 528 529 Esmark (1833) relates that he uses "a night thermometer" for the morning observation. Our Deleted: tells 530 hypothesis is that in Esmark's terminology the "night thermometer" was a minimum Deleted: means ' 531 thermometer, That means that he at some point started to note the night minimum temperature Deleted: ", and that the introduction of the minimum thermometer is the reason for in the column for the morning temperature, rather than the actual morning temperature when 532 the shift in March 1828. 533 he read the barometer. This hypothesis was tested by studying the difference between 534 Esmark's evening observation and the morning observation the following day for the three 535 homogenous intervals, Table 5, (the winter inhomogeneity in the 1830s was ignored). For Deleted: (see Deleted: 4) comparison we used the hourly observations (1993.09-2015.09) at the modern station Oslo -536 Deleted: this was also done for 537 Blindern (18700 Oslo), where the difference between the observation at 21 UTC and the Deleted: 538 minimum temperature for the following night is presented in row 4 in Table 5. The interval 539 for the night minimum was from 21 to 08 UTC, i.e. the same observation times as Esmark 540 used at least for his barometric observations in 1833. 541 542 In the earliest time interval (row 1) the differences in Esmark's observations are very much Deleted: were 543 smaller than those from Blindern, so it is impossible that Esmark in this early interval could 544 have recorded the nightly minimum temperature in the column for the morning observation. Deleted: noted 545 In the next interval (row 2) the differences are somewhat larger, but far too small compared to 546 Blindern so the same conclusion has to be drawn: no minimum thermometer was in use. 547 However, in the third interval (row 3) the differences are nearly the same as those for 548 Blindern. Even the monthly variations throughout the year correlate well. We conclude that Deleted: are realistic. 549 Esmark for the 'morning observation' used a minimum thermometer in the period 1828.03-Deleted: observation 550 1838.12. Before that he observed temperature in the morning with an ordinary thermometer, Deleted:

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565	Minimum thermometers were certainly available by 1828. Already in 1790 a spirit		
566	thermometer with a glass index, very much like those used up to this day at manual stations,	/	Formatted: Indent: First line: 0 cm
567	was described to the Royal Society in Edinburgh (Middleton, 1966: 152).		Deleted: notes
	was described to the Royal Society in Edinburgh (Middleton, 1900, 192).	$\ $	Deleted: logical Deleted: 2 The shift in 1821¶
568 569 570 571	If the minimum thermometer was set at the evening observation, the <u>values</u> in the column for morning observation should always be equal or lower than the evening temperature the previous day. In December this is not true for 26% of the observations and in June for 6%.		An inhomogeneity in the evening observation was detected by the homogeneity testing. It was adjusted for by the mean difference between the midday observation and the evening observation on each side of the shift, cf. Methods. The adjustments terms are presented in Table
572	These figures reduce to 6% and 2% in December and June respectively for violations no more		Deleted: .3 A
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573	than 1°C. In practice different exposure of the two thermometers may violate this test, and one		Formatted: English (U.S.)
574	should also take into account the possibility of instrumental errors in Esmark's thermometers.		Formatted: English (U.S.)
575	We may conclude that the percentage of violation is not large enough to contradict our		Formatted: Line spacing: single
576	conclusion that a night minimum thermometer was in use. The normal procedure for		<b>Formatted:</b> Expanded by 0,35 pt <b>Deleted:</b> in winter,
577	meteorological institutes when minimum thermometers are introduced is to change the		Formatted: Expanded by 0,45 pt
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578	formula for monthly mean calculation. Therefore the morning temperature will not be		Formatted: Expanded by 0,4 pt
579	corrected. Homogeneity in the monthly means will be obtained by changing formula for		Formatted: Expanded by 0,35 pt
580	monthly mean calculation, see section 4.5.	/	Formatted: Expanded by 0,45 pt
581			Deleted: as well as the
	4.4 Compating the skift in the 1920s		Formatted: Expanded by 0,45 pt
582 583	4.4. Correcting the shift in the 1830s         A significant inhomogeneity in winter for the morning observation (in this period		Formatted: Expanded by 0,45 pt
			Formatted: Expanded by 0,4 pt
584	identified as minimum temperature), was detected by the SNHT double shift, Table 2 part		Formatted: Character scale: 102%
585	<u>1 I vs II, and also by the single shift test when the time window was 1828.03-1838.12</u>	_	Formatted: Expanded by 0,4 pt
586	Table 2, part 3. Formally a significant shift in spring was also detected, Table 2 III vs	$\langle \rangle$	Formatted: Expanded by 0,5 pt
587	II, but with only three years on one side of the shift its significance was considered	$\left( \right) \right)$	Formatted: Expanded by 0,45 pt
588	doubtful. The shift in winter was firstly examined by plotting the morning temperature	())	<b>Deleted:</b> for the test
			Formatted: Expanded by 0,5 pt
589	against the evening temperature, which revealed that there was not an abrupt shift in the		Deleted: . The shift has the character of
590	difference, but rather a steady state 1829-1931 followed by a trend. The graphical plotting		Deleted: continuous
591	was followed by applying the Multiple Linear Regression procedure (MLR) also known as the		Formatted: Expanded by 0,45 pt Formatted: English (U.S.)
592	Vincent test (Vincent, 1998). The significant inhomogeneity was confirmed and also the		<b>Deleted:</b> (Fig. 7). The difference betw(
		$\langle \cdot \rangle$	Formatted: English (U.S.)
593	change point year of 1831, The trend line was found by least scare regression analysis, Fig 6.		<b>Deleted:</b> the morning observation
594	An explanation <u>for the trend might</u> be a change in the observation times. According to Esmark	$\langle \rangle$	Formatted: English (U.S.)
595	(1833) his observation times were, see Metadata.	$\sim$	Deleted: to 1838, whereas it was
596	• Morning: 08:30 ChT = 08:43 CET = 7:43 UTC	$\langle \rangle \rangle$	Formatted: English (U.S.)
			Deleted: may
597	• Midday, (afternoon): 15:30 ChT = 15:43 CET = 14:43 UTC		Formatted: English (U.S.)
500	- E		Deleted:

Evening: 21:30 ChT = 21:43 CET = 20:43 UTC •

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649	ChT = Christiania time i.e. local time for Christiania (Oslo), CET = Central European		
650	Time, UTC = Universal Time Coordinated.		
651	These observation times were for the barometric pressure, but <u>in the afternoon</u> and evening		Deleted: at midday
652	the thermometer was read at the same time as the barometer, but Esmark does not explicitly		Deleted: were
653	say that the morning thermometer was read at the same time as the barometer. He also use the		
654	term "in the <u>later</u> " years so we do not know from which year these observation times were		Deleted: latest
655	introduced or if he continued to use them also in the following years 1834-1838.		
656	Our hypothesis is that Esmark has had another observation time for the temperature		
657	observations in the morning than for the pressure observations. Pressure could be observed		Deleted: was Deleted: at
658	inside the house, but for the temperature observations he possibly had to leave the house for	/	Deleted: time
659	his garden. Esmark might originally have observed temperature and pressure at the same time		Deleted: often
660	also in the morning, but with the introduction of the minimum thermometer he could have		Deleted: might
661	thought that the observation time for the morning temperature was not important. In spring,		Deleted: gone outside for carrying out Deleted: observation
662	summer and autumn he obviously was right in his thinking as minimum temperature occurs		Deleted: This might explain
663	earlier than the morning observation (8:30 ChT), but in winter the minimum temperature		Deleted: trend shift
			Deleted: observation Deleted: adjusted
664	often occurs later in the day as the systematic daily temperature wave is weak. This can		Deleted: 7
665	explain the changing difference during winter and the stable differences during the other		Deleted: adjustments were
666	seasons. As Esmark grew older and more frail he may have got up in the morning later and	_]/////	Formatted: Font: Bold, Not Italic
667	later. Progressive illness and susceptibility to cold in his later years (Anonymous 1839) could	]]////	<b>Deleted:</b> 4.4 Overheating of the midday observation¶
668	have made it less convenient to leave the house for the garden in the morning, Following this	J///	The midday observation turned out to be homogenous, but it might have been
669	hypothesis the minimum temperature was <u>corrected</u> , $\Delta T$ , by use of formula (2) for the winter	]	overheated by insufficient radiation protection in Esmark's yard. This was tested by comparison with the Oslo –
670	season in accordance with the regression line shown in Fig. $6$ , where a = year (period 1832-		Blindern station that is well protected by a Stevenson screen. Difference between the
671	1838). No correction was undertaken for the period 1829-1831.		midday observation and the evening observation reveals a characteristic pattern
672			(Fig. 8). Whereas the differences were almost equal in the months September –
673	$\Delta T = 0.2861 \cdot a - 523.85 \tag{2}$		March, the differences in the Esmark series were larger than the differences in the
674			Blindern series for the months April – August. They were particularly large in
675	4.5 Homogenisation of the monthly mean temperature.		MJA where the sun is highest on the sky and the radiation reaches its annual
676	Esmark observed only three times a day, so it is far from obvious how monthly mean	/	maximum. Therefore our interpretation is that Esmark's thermometer was overheated at the midday observation by (reflected)
677	temperature should be calculated without bias. This problem confronts meteorological		short wave radiation in the period April – August, but not for the rest of the year.
678	institutes worldwide so formulas for such calculations have been developed (see Appendix $\underline{C}$ ).		Based on the differences between the two curves the adjustments of the midday
679	The formulas contain specific constants valid for each month and site. Strictly speaking the	$\mathcal{A}$	observation are also given (lower panel in Fig. 8).¶
680	constants were unknown for Esmark's observation site at Øvre Vollgate, but <u>are</u> well known		¶ Deleted: the calculation are
			Deleted: B
681	for the station 18700 Oslo – Blindern, <u>situated</u> 3.4 km to the north of Esmark's site.		Deleted: lying
682	Fortunately there are indications that the constants for Blindern could be used also for Øvre		Formatted: Footer, Right: 0,63 cm

#### 731 Vollgate (see Appendix $\underline{C}$ ). Given the constants the calculation of homogenous monthly mean Deleted: 2 732 temperature was trivial when the homogenised version of the observations at fixed hours was Deleted: adjustments 733 used. We found that the <u>corrections</u> for seasonal means vary from $0.0^{\circ}C$ to $+0.4^{\circ}C$ , the annual De corrections from 0.0°C to +0.3°C. How the corrections changed throughout the period of 734 De De 735 observation are shown in Fig. 7. For the period 1822.12-1831.12 no corrections were applied. De ne 736 4.6 The Christiania (Oslo) climate in Esmark's period of observation, 1816-1838 se De 737 Esmark's observations exhibit a long-term variation pattern characterised by lower values in ar 0. 738 the start and in the end of the period, whereas the middle of the period was somewhat warmer, De Fo 739 cf. Fig. 8. This is true not only for the annual means, but also for all seasons of the year except De 740 for winter. For individual years 1822 is warmest except in summer and autumn. The coldest De year is 1838 followed by the years 1816, 1829 and 1820, 741 De De 742 ye sea 743 The year 1816 is of particular interest as it has gone into history as "the year without an M 744 summer,", with an average decrease in global temperatures often ascribed to volcanic activity, Fo 745 resulting in a food shortage many places in the Northern Hemisphere. However, Esmark's De De observations show that this summer (JJA) was not extraordinary in Oslo, as the following 746 De 747 summer of 1817 and 1821, were approximately 1°C colder. The spring temperature in 1816 is of De however the coldest one in the series. The three first years of Esmark's series must have been 748 De 749 very unfavourable for agriculture due to low temperature. In the grain growing months De De 750 (AMJJA) the mean temperature was about 10°C for the three consecutive years 1816, 1817 zei 751 and 1818, i.e. the lowest temperatures in Esmark series of observation. De D

#### 752 **5** Discussion

753		Deleted: in
754	5.1 Overheating of the midday observation?	<b>Deleted:</b> From 1816 to the mid-1820s the annual Christiania temperature as recorded
755	The midday observation turned out to be homogenous, but it may have been overheated by	by Esmark rose by approximately 1.5°C, then subsequenly slowly fell by almost 1°C towards 1840 (Fig. 10). This general
756	insufficient radiation protection in Esmark's yard or simply the confined space allowing less	pattern is consistent with that found for the same time interval in the Swedish capital
757	air flow (wind). This was tested by comparison with the Oslo – Blindern station (18700),	Stockholm (compare with Fig. 5 in Moberg et al., 2002). ¶
758	which is well protected by a Stevenson screen. Differences between the midday observation	S.1 Adjusting for inhomogeneities An important inhomogeneity was detected
759	and the evening observation exhibit characteristic variations throughout the year, not only for	in Esmark's data at the end of 1822 in the evening observation, and was adjusted for.
760	Blindern, but also for the Esmark series and the Oslo II series (Astronomical Observatory,	Alternatively the inhomogeneity could be considered only as a change of
761	18651), cf. station list Table 1 and Fig. 9. Whereas the differences between the Blindern series	observational time, and not adjusted for by the testing. The series of mean temperatures could then have been kept homogenous [
762	and Esmark's series were relatively small in the months August – April, they are much larger	Formatted: Footer, Right: 0,63 cm

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ormatted: Indent: First line: 0 cm
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Adjusting for inhomogeneities¶ important inhomogeneity was detected Esmark's data at the end of 1822 in the ening observation, and was adjusted for. ternatively the inhomogeneity could be nsidered only as a change of

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871	in the months May – July, when the sun is highest on the sky and the radiation reaches its
872	annual maximum. Therefore one possible interpretation is that Esmark's thermometer was
873	overheated at the midday observation in midsummer, MJJ, by (reflected) short wave radiation.
874	However, when compared to the diurnal pattern at the Oslo II station (Astronomical
875	Observatory), it is seen that the curve representing Esmark's observations quite closely
876	follows the Oslo II curve, also in midsummer, Fig 9. At the Astronomical observatory there
877	were three thermometers on different walls - N, E and W. (Nordli et al. 2015). At least one of
878	these thermometers was in shadow and therefore available for use at every observation time.
879	This is our main reason for not correcting for a possible overheating of Esmark's midday
880	observation, see also the following 5.2 and 5.3. The deviation of the Blindern station may be
881	due to this site being more exposed to wind chill and its situation significantly higher above
882	sea level than Esmark's house and the Astronomical Observatory, cf. Table 1.
883	

884 The meteorological observations at the Astronomical Observatory started in April 1837 885 (Nordli et al., 2015), so this series overlaps Esmark's series by 21 months. The difference of 886 their uncorrected monthly means is shown in Fig. 10. It is evident that for all seasons but 887 winter Esmark's temperatures are somewhat lower than those from the Observatory. Esmark 888 died on 26 January 1839 (see Metadata), so possibly the quality of the latest months of his 889 series might be questioned. However, we cannot see any decline in quality directly from his 890 observation protocols. This is relevant also for the discussion of a possible correction of 891 Esmark's midday observation due to overheating. If Esmark's midday observation had been 892 corrected the discrepancy between Esmark's series and Observatory series would have been 893 larger.

#### 895 **5.2** Comparison with <u>Hansteen's</u> observations at the street Pilestredet in Oslo

896	During the period 1822.11-1827.02 the Christiania professor, Christopher Hansteen carried
897	out observations at his home in Pilestredet at the corner of Keysersgate, at the center of town
898	(Hansteen 1823, 1824, 1828; Birkeland, 1926: 12), cf Table 1 for some further information.
899	The distance from Esmark's site was only about 600 m. Hansteen's observation times varied
900	much but for each month he gives the observation times together with the data (Hansteen,
901	1824). The distribution of the observation times in UTC is as follows: morning 06 <sup>h</sup> 4%, 07 <sup>h</sup>
902	44%, 08 <sup>h</sup> 52%; midday 13 <sup>h</sup> 20%, 14 <sup>h</sup> 78%, 15 <sup>h</sup> 2%; evening 21 <sup>h</sup> 6%, 22 <sup>h</sup> 88%, 23 <sup>h</sup> 6%.
903	Hansteen's observations were <u>corrected</u> to Esmark's observation times, approximately 08, 15
904	and 21 UTC by use of the mean daily temperature wave at Blindern so that Esmark's

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922	observations could be compared with the <u>corrected</u> ones of Hansteen, Fig 11. It is <u>seen</u> that				
923	Hansteen's morning observation is much warmer than that of Esmark except during winter.				
924	Most likely the thermometers of Hansteen had been overheated as the the thermometers at the				
925	southern and northern side of the house (Birkeland, 1925: 12). Then it must have been				
926	difficult to find shadow in the morning. Also the midday observation is warmer at Hansteen's				
927	site than by Esmark. This is probably due to the fact that Hansteen's garden was protected by				
928	the surrounding houses and gardens of the town which reduced wind, while Esmark's garden				
929	was directly exposed to the winds from the adjacent bay. The evening temperatures at				
930	Hansteens house, however, agrees well with those from Esmark during summer unlike for the				
931	two other observation times. The evening observations occurred after sunset at both sites,				
932	whereas the two other observations occurred after sunrise.				
933					
934	Unlike the situation during summer, Hansteen's temperatures are lower than those of Esmark				
935	in the period November – March (Fig. 11). In many weather situations the air loses energy by				
936	long wave radiation because the short wave radiation is too small to compensate for the loss.				
937	The result is that the coldest air is found at the lowest places in the local terrain, not				
938	necessarily at the lowest sites above sea level. Esmark's house lies high in the local terrain at				
939	the edge of a slope down to Pipervika cf. Metadata, whereas Hansteen's house lies low in the				
940	local terrain at a floor of a small valley. The difference in winter temperature is therefore				
941	possibly an effect of topography.				
942					
943	5.3 Comparison with Stockholm and Copenhagen				
944	The Stockholm and Copenhagen series were not used as reference stations for the				
945	homogeneity testing. Their distances from Oslo were considered to be too long, 350 km and				
946	450 km respectively. However, comparison with the Stockholm Observatory and Copenhagen				
947	old Botanical Garden (Closter et al. 2006) with Esmark's observations may provide some				
948	indications of the quality of the homogenisation, see Fig 12. Compared to Esmark Stockholm				
949	seems to be relatively warmer in the first four years, 1816-19, than the rest of the series.				
950	Without correction for the years 1816-21 the differences would have been even larger.				
951	Therefore comparison with Stockholm supports the correction of the series. Probably there				
952	might be another shift in the series in 1819. Some support for this is seen in the homogeneity				
953	testing cf. Table 2, part 2. However, the reason might also be spatial temperature differences				
954	between Stockholm and Oslo, the long distance between the stations taken into account. And,				
955	in spite of homogenisation there might also be small inhomogeneities in the Stockholm series.				

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**Deleted:** by Hansteen than by Esmark. This is harder to understand. If Birkeland's account of the thermometer at the north wall of the house is correct the house is expected to give sufficient protection of that thermometer (Nordli et al., 2015), but as nothing is known about the environment other factors might have been involved.¶ The evening temperature, however, is much in agreement with that of

**Deleted:** This supports the suggestion that the differences at the morning and midday observations are due to radiation errors.

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#### Deleted: At The Astronomical Observatory in Oslo meteorological observations started in April 1837 that lasted almost for one hundred years (Nordli et al., 2015), so this series overlaps Esmark's series by 21 months. For comparison of the two series we have used unadjusted observations from the observatory, whereas both adjusted and unadjusted Esmark observations are used (

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**Deleted:** 12). It is evident that for all seasons but winter Esmark's temperatures are lower than those from the Observatory. Esmark died on 26 January 1839 (see Metadata), so probably the quality of the latest months of his series may be questioned. However, we cannot see any declined quality directly from is observation protocols, but it is possible that the last two years of his observations are not representative for Esmark's observational practice. Moreover, the overlapping period is very short; only two years for most of the months, and only one year for the months January to March. It is therefore possible that the present comparison is not valid for Esmark's entire period of observation.

5.3 The accuracy of the thermometers .... Moved up [1]: . They are not written by Esmark himself, most probably they are notes written by Birkeland, who says he ....

**Moved down [5]:** It should also be kept in mind that Esmark used another thermometer, i.e. a minimum thermomet

**Deleted:** The corrections are very small for the frequent winter temperatures, but as high as 0.5°C for frequent summer

**Deleted:** However, Esmark was a skilled instrument builder, so it is not likely that

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1065	Comparison between Copenhagen and Oslo give no reason for expecting any shift in the
1066	series, but four years is missing from the Botanical Garden series
1067	
1068	5.4 The summer of 1816 in Christiania (Oslo)
1069	Several volcanic eruptions affected global climate in the first years of Esmark's period of
1070	observation, the Tambora eruption in Asia in 1815 being the largest in terms of sulphur mass
1071	ejected and general impact (Stothers 1984, Oppenheimer, 2003). It has given rise to the
1072	paradigm for 1816: "the year without a summer". Esmark's observations show, however, that
1073	the summer of 1816, though cold, was not extraordinary cold in Oslo. And in Stockholm,
1074	("Bolin Centre Database,") that summer was rather warm, No 17 of the 23 summers from
1075	1816-1838, <u>ranked</u> from low to high (Table <u>6</u> ). May, however, was very cold in both cities,
1076	and July, quite warm in both cities, but in June and August Oslo was much colder relative to
1077	the mean value than Stockholm.
1078	
1079	Esmark's observations may also be compared to other independent reconstructions of
1080	temperature in Norway in the period 1816-1838, (Table 7). One reconstruction for FMA for
1081	Austlandet, South Eastern Norway, is based upon ice loss mainly from Lake Randsfjorden
1082	(Nordli et al., 2007), Four reconstructions are based upon the first date of grain harvest;
1083	Austlandet (Nordli, 2001a), Vestlandet (Bergen), Western Norway, (Nordli et al., 2003),
1084	Lesja (Nordli, 2001b) and Trøndelag, Mid Norway (Nordli, 2004), The grain harvest date is a
1085	proxy for AMJJA temperature in the southern lowland areas, whereas in the mountain valleys
1086	(Lesja) and northern areas (Trøndelag) it is a proxy for MJJA temperatures. We also included
1087	a gridded multi proxy series for the nearest grid point to Oslo (Luterbacher et al. 2004). The
1088	three reconstructions for Austlandet all have the spring-summer of 1816 as the coldest one in
1089	the period, whereas in the Esmark series it is listed as No. 3. The reconstructions for the two
1090	other temperature regions, Vestlandet and Trøndelag, show a very different picture with
1091	relatively warm 1816 summers like the summer in Stockholm based on instrumental
1092	observations. Vestlandet and Trøndelag belong to other climate regions than Austlandet
1093	(Hanssen-Bauer and Førland, 2000), so for a specific summer it might reflect real temperature
1094	differences. The very low temperature for spring in 1816 seems to have had a strong influence
1095	on agriculture so the harvest had been delayed in south eastern Norway. This is reflected in
1096	the AMJJA temperature reconstruction. In Fig. 13 proxy and instrumental summer
1097	temperatures (JJA) are shown for the whole period of Esmark's observations. The proxy data

**Deleted:** temperature proxies for the seasons April-August and May-August (Table 8). The three reconstructions within the county of South-Eastern Norway are all in agreement with Esmark's observations that the summer of 1816 was among the coldest in the grain growing seasons, whereas the reconstructions for the two other counties, Western and Mid Norway, show relatively warm summers, even more so than those in Stockholm.¶

Anomalies of surface temperature and precipitation for the summer months of 1816 has been reconstructed (Luterbacher and Pfister, 2015). They show a positive gradient from a cold core of air lying over France towards Eastern and Northern Europe, so the paradigm of the severe summer of 1816 has to be modified when it comes to Scandinavia and Eastern Europe. It looks like this is easy to forget, e.g. "...weather patterns were disrupted

worldwide for months, allowing for excessive rain, frost, and snowfall through much of the Northeastern U.S. and Europe in the summer of 1816"(Klingaman and Klingaman, 2014). It is therefore important that the temperature gradient is recognised. The results in Table 8 are a part of the pattern showing the spatial variability in Europe that summer.¶

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1152	summers 1816-18 were quite cold, not warm like those in Stockholm. The summer of 1819,			
1153	however, was warm in Oslo (and also in Stockholm) but not in the reconstruction. It is also			
1154	evident that the variability in the reconstructed series is too small.			
1155				
1156	The summer temperatures of 1816 have recently been analysed by Luterbacher and Pfister			
1157	(2015). Their study shows a positive gradient from a cold core of air lying over France with a			
1158	positive temperature gradient towards Eastern and Northern Europe, so the paradigm of the			
1159	severe summer of 1816 has to be modified when it comes to Scandinavia and Eastern Europe			
1160	to take into account significant geographical variation. The authors state that "in eastern			
1161	Europe, western Russia and parts of eastern Scandinavia, summer temperatures were normal			
1162	or slightly warmer than average".			

#### 1163 6 Conclusions

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1164	Homogeneity testing (SNHT) of Esmark's temperature observations 1816-1838 in Christiania		<b>Deleted:</b> are almost complete for the years
1165	(Oslo) demonstrated three significant shifts, and we propose corrections for these. First there		Deleted: . Homogeneity testing revealed
1166	is a shift in the evening observation in 1821-22. Before the shift the evening observation was		Deleted: at
1167	corrected by about +1.3° for the summer months, but only by about +0.5°C in winter.		Deleted: end of 1822.
1168	A very large shift in the morning temperature was detected in 1827-28. From Esmark himself		Deleted: March 1828
1169	we know that he used a "night thermometer" in 1833, identified as minimum thermometer.		Deleted: noted nightly
1170	This change of instrumentation explains the lower values for the morning observation. During		<b>Deleted:</b> temperature instead his previous notation
1171	the years 1831 to 1838 the nightly minimum temperature decreased steadily in the winter		Deleted: morning temperature.
1172			Deleted: increased almost
	season, i.e. it was inhomogenous. The reason seems to be later and later reading of the		
1173	minimum temperature in the morning. The seasonal corrections of the series are less than		
1174	0.5°C, and for annual means less than 0.4°C. In the time interval 1822-1831 no corrections are		
1175	applied. The homogenized temperature series 1816-1838 exhibit low temperature at both		Deleted: showed
1176	ends, with higher temperature in the middle, i.e. in the 1820s. The starting year, 1816, is of		Deleted: in
1177	particular interest as it has been referred to as 'the year without a summer'. That summer in	$\overline{}$	Deleted: of the series
		$\sim$	Deleted: s.
1178	Oslo was cold, but not extraordinary cold, as it was only the <u>fifth</u> coldest in the period of		Deleted: summer. The
1179	observation. However, March and May that year were the coldest ones in the period of		Deleted: third
1180	Esmark's data, and 1816 and 1838 had the lowest annual means. The first three years of	$\overline{\mathbb{N}}$	<b>Deleted:</b> the annual mean of 1816 and also the months
1101			Deleted: that
1181	Esmark's observation, 1816-1818, were particularly cold in the grain growing season, April-		Deleted: .
1182	August, and lends support to the historians' view that these were years of hardship and		
1183	famine.		

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1508	Vincent, L. A.: A technique for the identification of inhomogeneities in Canadian temperature		
1509	series. J. Clim. 11: 1094–1104, doi: 10.1175/1520-		
1510	<u>0442(1998)011&lt;1094:ATFTIO&gt;2.0.CO;2, 1998.</u>		
1511	Willaume-Jantzen, V.: Meteorologiske Observationer i Kjøbenhavn. Résumé des		Formatted: English (U.S.)
1512	Observations Météorologiques de Copenhague. Det Danske Meteorologiske Institut.		Deleted: . 1896.
			Formatted: English (U.S.)
1513	Kjøbenhavn, i commission hos universitets-boghandler G.E.C. Gad. <u>1896</u> .		Formatted: Danish
1514	Worsley, P.: Jens Esmark, Vassryggen and early glacial theory in Britain. Mercian Geologist		Deleted: . 2006.
1515	16: 161-172 <u>, 2006</u> ,		Formatted: English (U.S.)
	10. 101 172, 2000,	$\langle \rangle$	Formatted: English (U.S.)
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1527	APPENDIX A. ESMARK'S METEOROLOGICAL TABLES IN	
1528	DEN NORSKE RIGSTIDENDE.	
1529		
1530	Esmark, J. 1818/19. Meteorologiske Iagttagelser i Christiania 1818, anstillede af	
1531	Prof. Esmark. Den Norske Rigstidende 1818, No. 7 (24 January); No. 10 (4	
1532	February); No. 14 (18 February); No. 18 (4 March); No. 23 (21 March), No.	
1533	28 (8 April), No. 32 (22 April); No. 37 (9 May); No. 40 (20 May), No. 45 (6	
1534	June), No. 49 (20 June), No. 54 (8 July); No. 59 (25 July); No. 63 (8	
1535	August); No. 67 (21 August); No. 71 (5 September); No. 83, (17 October);	
1536	No. 84 (21 October), No. 86 (28 October); No. 88 (4 November); No. 95 (28	
1537	November); No. 98 (9 December); No. 102 (23 December); No. 3 (8 January	
1538	1819).	
1539	Esmark, J. 1819/20. Meterologiske Iagttagelser i Christiania 1819, anstillede af	
1540	Prof. Esmark. Den Norske Rigstidende No. 6 (19 January); No. 11 (5	
1541	February); No. 16 (23 February); No. 19 (5 March); No. 24 (23 March); No.	
1542	26 (6 April); No. 33 (23 April); No. 36 (4 May); No. 41 (21 May); No. 48	
1543	(15 June); No. 49 (18 June); No. 54 (6 July); No. 62 (3 August); No. 65 (13	
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	Prof. Esmark. Den Norske Rigstidende, No. 7 (25 January); No. 11 (8
	February), No. 14 (18 February); No. 18 (3 March); No. 24 (24 March) ; No.
	28 (7 April); No. 32 (21 April); No. 37 (9 May); No. 41 (23 May); No. 47
	(13 June); No. 50 (23 June); No. 54 (7 July); No. 58 (21 July); No. 63 (8
	August); No. 68 (25 August); No. 72 (8 September); No. 77 (26 September);
	No. 81 (10 October); No. 85 (24 October); No. 88 (3 November); No. 94 (24
	November); No. 98 (8 December); No. 103 (26 December); No. 3 (9 January
	1821).
Esma	rk, J. 1821/22. Meteorologiske Iagttagelser i Christiania 1821, anstillede af
	Professor Esmark. Den Norske Rigstidende, No. 7 (23 January), står bare
	snee, men ikke mengde, ; No. 11 (6 February); No. 16 (23 February); No. 21
	(13 March); No. 23 (20 March); No. 29 (10 April); No. 33 (24 April), No. 38
	(11 May); No. 41 (22 May); No. 45 (5 June); No. 52 (29 June); No. 55 (10
	July); No. 58 (20 July); No. 63 (6 August); No. 68 (24 August); No. 72 (7
	September); No. 76 (21 September); No. 80 (5 October); No. 85 (22
	October); No. 89 (5 November); No. 93 (19 November)(nytt moderne
	plusstegn); No. 98 (7 December); No. 102 (21 December); No. 2 (7 January
	1822).
Esma	rk, Jens 1822/23. Meteorologiske Iagttagelser i Christiania 1822, anstillede
	ved Professor Esmark. Den Norske Rigstidende, No. 5 (18 January); No. 10

- 1981 (4 February); No. 15 (22 February); No. 18 (4 March); No. 23 (22 March);
- 1982 No. 28 (8 April); No. 32 (22 April); No. 36 (6 May); No. 42 (27 May); No.
- 1983 45 (7 June) not nedbørmåling; No. 50 (24 June); No. 81 (11 October); No. 82

August); No. 67 (20 August); No. 78 (28 September); No. 79 (1 October)

November); No. 99 (10 December); No. 103 (24 December); No. 2 (7

Esmark, J. 1820/21. Meteorologiske Iagttagelser i Christiania 1820, anstillede af

No. 82 (12 October); No. 84 (19 October); No. 89 (5 November); No. 95 (26

- 1984 (14 October); No. 83 (18 October); No. 84 (21 October); No. 87 (1
- 1985 November); No. 89 (8November); No. 90 (11 November); No. 92 (18
- 1986 November); No. 94 (25 November); No. 96 (2 December); No. 98 (9
- 1987 December); No. 102 (23 December); No. 2 (6 January 1823).
- 1988 Esmark, J. 1823/24. Meteorologiske Iagttagelser i Christiania 1823, anstillede ved
- 1989 Professor Esmark. Den Norske Rigstidende No. 7 (24 January); No. 11 (7

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1990	February) ; No. 15 (21 February); No. 20 (10 March); No. 24 (24 March);
1991	No. 27 (4 April); No. 31 (18 April); No. 36 (5 May); No. 40 (19 May); No.
1992	46 (9 June); No. 49 (20 June); No. 75 (19 September); No. 76 (22
1993	September); No. 77 (26 September); No. 78 (29 September); No. 79 (3
1994	October); No. 81 (10 October); No. 82 (13 October); No. 84 (20 October);
1995	No. 88 (3 November); No. 93 (21 November); No. 98 (8 December); No. 102
1996	(22 December); No. 2 (5 January 1824).
1997	Esmark, J. 1824/25. Meteorologiske lagttagelser i Christiania 1824, anstillede ved
1998	Professor Esmark. Den Norske Rigstidende No. 6 (19 January); No. 11 (5
1999	February); No. 15 (19 February); No. 20 (8 March); No. 24 (22 March); No.
2000	29 (8 April); No. 33 (22 April); No. 37 (6 May); No. 42 (24 May); No. 45 (3
2001	June); No. 50 (21 June); No. 54 (5 July); No. 59 (22 July); No. 64 (9
2002	August); No. 68 (23 August); No. 74 (13 September); No. 77 (23
2003	September); No. 80 (4 October); No. 86 (25 Oktober); No. 89 (4 November);
2004	No. 96 (29 November); No. 98 (6 December); No. 103 (23 December); No. 2
2005	(6 Januar 1825).
2006	Esmark, J. 1825/26. Meteorologiske Iagttagelser i Christiania 1825, anstillede ved
2007	Professor Esmark. Den Norske Rigstidende No. 7 (24 January); No. 11 (7.
2008	February), No. 15 (21 February); No. 18 (3. March); No. 24 (24 March); No.
2009	29 (11 April); No. 33 (25 April); No. 36 (5 May); No. 40 (19 May); No. 45
2010	(6 June); No. 49 (20 June); No. 53 (4 July); No. 70 (1 September); No. 71 (5
2011	September); No. 73 (12 September); No. 74 (15. September); No. 76 (22
2012	September); No. 79 (3 October), No. 85 (24 October); No. 89 (7 November);
2013	No. 93 (21 November); No. 97 (5 December); No. 102 (22 December); No. 2
2014	(5 January 1826).
2015	Esmark, J. 1826/27. Meteorologiske lagttagelser i Christiania 1826, anstillede ved
2016	Professor Esmark. Den Norske Rigstidende No.8 (26 January); No. 12 (9
2017	February); No. 17 (27 February); No. 19 (6 March); No.23 (20 March); No.
2018	28 (6 April); No. 33 (24 April); No. 36 (4 May); No. 43 (29 May); No. 45 (5
2019	June); No. 50 (22 June); No. 55 (10 July): No.58 (20 July); No. 62 (3
2020	August); No. 67 (21 August); No. 72 (7 September); No. 77 (25 September);
2021	No. 80 (5 Oktober); No. 84 (19 October); No. 88 (2 November); No. 93 (20
2022	November); No. 97 (4 December); No. 102 (21 December); No. 2 (4 January
2023	1827).

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2024	Esmark, J. 1827/28. Meteorologiske lagttagelser i Christiania 1827, anstillede ved
2025	Professor Esmark. Den Norske Rigstidende, No. 7 (22 January); No. 11 (5
2026	February); No. 16 (22 February); No. 19 (5 March); No. 24 (22 March); No.
2027	28 (5 April); No. 32 (19 April); No. 37 (7 May); No. 43 (28 May); No. 48
2028	(14 June); No. 50 (21 June); No. 54 (5 July); No. 58 (19 July); No. 79 (1
2029	October); No. 80 (4 October); No. 81 (8 October); No. 82 (11 October); No.
2030	83 (15 October); No. 84 (18 October); No. 89 (5 November); No. 94 (22
2031	November); No. 97 (3 December); 102 (20 December); No. 2 (7 January
2032	1828) - also sums up last ten years, compares with Stockholm, the coldest
2033	years have been 1819 and 1820, the mildest 1822 and 1826.
2034	Esmark, J. 1828/29. Meteorologiske lagttagelser i Christiania 1828, anstillede ved
2035	Professor Esmark. Den Norske Rigstidende, No. 6 (21 January); No. 10 (4
2036	February); No. 15 (21 February); No. 18 (3 March); No. 24 (24 March); No.
2037	27 (3 April - mange solpletter); No. 32 (21 April); No. 36 (5 May); No. 40
2038	(19 May); No. 45 (5 June); No. 49 (19 June); No. 53 (3 July); No. 59 (24
2039	July); No. 63 (7 August); No. 78 (29 September); No. 79 (2 October); No. 81
2040	(9 October); No. 84 (20 October); No. 88 (3 November); No. 94 (24
2041	November); No. 98 (8 December); No. 102 (22 December); No.2 (5 January
2042	1829).
2043	Esmark, J. 1829/30. Meteorologiske lagttagelser i Christiania 1829, anstillede ved
2044	Professor Esmark. Den Norske Rigstidende, No. 8 (26 January); No. 11 (5
2045	February); No. 15 (19 February); No. 19 (5 March – den strengeste vinter på
2046	mange år); No. 24 (23 March); No. 27 (2 April); No. 33 (23 April); No. 37 (7
2047	May); No. 42 (25 May); No. 46 (8 June); No. 50 (22 June); No. 54 (6 July);
2048	No. 78 (28 September); No. 79 (30 September); No. 80 (5 October); No. 81
2049	(8 October); No. 85 (22 October); No. 87 (29 October); No. 89 (5
2050	November); No. 90 (9 November); No. 94 (23 November); No. 99 (10
2051	December); No. 103 (24 December); No. 2 (7 January 1830).
2052	Esmark, J. 1830/31. Meteorologiske lagttagelser i Christiania 1830, anstillede ved
2053	Professor Esmark. Den Norske Rigstidende, No. 7 (25 January); No. 11 (8
2054	February); No. 14 (18 February); No. 18 (4 March); No. 22 (18 March); No.
2055	27 (5 April); No. 31 (19 April); No. 36 (6 May); No. 40 (19 May); No. 46 (9
2056	June); No. 50 (23 June); No. 53 (5 July); No. 57 (19 July); No. 63 (9
2057	August); No. 70 (1 September); No. 73 (13 September); No. 78 (29

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2058	Septmerber); No. 81 (11 October); No. 84 (21 October); No. 91 (15
2059	November); No. 95 (29 November); 98 (9 December); No. 102 (23
2060	December); No. 3 (10 January 1831).
2061	Esmark, J. 1831/32. Meteorologiske lagttagelser i Christiania 1831, anstillede ved
2062	Professor Esmark. Den Norske Rigstidende, No. 10 (3 February); No. 11 (7
2063	February); No. 17 (28 February); No. 20 (10 March); No. 25 (28 March); No.
2064	28 (7 April); No. 33 (25 April); No. 39 (12 May); No. 43 (22 May); No. 52
2065	(12 June); No. 57 (23 June); No. 63 (7 July); No. 70 (24 July); No. 75 (4
2066	August); No. 85 (28 August); No. 88 (4 September); No. 97 (25 September);
2067	No. 102 (10 October); No. 110 (3 November); No. 112 (10 November); No.
2068	118 (1 December); No. 119 (4 December); No. 1 (1 January 1832) ; No. 2 (5
2069	January 1832).
2070	Esmark, J. 1832/33. Meteorologiske lagttagelser i Christiania 1832, anstillede ved
2071	Professor Esmark. Den Norske Rigstidende, No.10 (2 February); No. 11 (5
2072	February); No. 19 (4 March); No. 20 (8 March); No. 26 (26 March); No. 30
2073	(12 April); No. 33 (22 April); No. 37 (6 May); No. 43 (20 May); No. 52 (10
2074	Juni); No. 57 (21 Juni); No. 63 (5 July); No. 70 (22 July); No. 78 (9 August);
2075	No. 86 (28 August – usedvanlig kold sommer); No. 92 (11 September); No.
2076	98 (25 September); No. 103 (7 October); No. 108 (25 October); No. 111 (4
2077	November); No. 117 (25 November); No. 122 (13 december); No. 127 (30
2078	December); No. 4 (13 Januery 1833).
2079	Esmark, J. 1833/34. Meteorologiske lagttagelser i Christiania 1833, anstillede ved
2080	Professor Esmark. Den Norske Rigstidende, No.10 (3 February); No. 12 (10
2081	February); No. 18 (3 March); No. 24 (24 March); No. 25 (28 March); No. 30
2082	(14 April); No. 35 (2 May); No. 37 (9 May); No. 44 (26 May); No. 50 (9
2083	June); No. 58 (27 June); No. 63 (9 July); No. 77 (11 August); No. 80 (18
2084	August); No. 86 (1 September); No. 91 (12 September); No. 97 (26
2085	September); No. 103 (13 October); No. 105 (20 October); No. 110 (7
2086	November); No. 115 (24 November); No.120 (12 December); No. 123 (22
2087	December); No. 2 (5 January 1834).
2088	Esmark, J. 1834/35. Meteorologiske lagttagelser i Christiania 1834, anstillede ved
2089	Professor Esmark. Den Norske Rigstidende, No. 7 (23 Januery); No. 10 (2
2090	February); No. 16 (23 February); No. 18 (2 March); No. 24 (23 March); No.
2091	27 (3 April); No. 32 (20 April); No. 37 (4 May); No. 43 (18 May); No. 53

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2092	(10 June); No. 60 (26 June); No. 68 (15 July)(regnet som falt på en
2093	kvadratfods flate utgjorde 4 rhinlandskae tommer eller 576 kubikktommer);
2094	No. 71 (22 July); No. 79 (10 August), No. 83 (19 August); No. 90 (7
2095	September); No. 96 (21 September); No. 102 (5 October); No. 107 (23
2096	October); No. 111 (6 November); No. 117 (27 November); No. 119 (4
2097	December); No. 126 (28 December); No. 2 (8 January 1835).
2098	Esmark, J. 1835/36. Meteorologiske Iagttagelser i Christiania 1835, anstillede ved
2099	Professor Esmark. Den Norske Rigstidende, No. 10 (1 February); No. 12 (8
2100	February); No.15 (19 February); No. 20 (8 March); No. 24 (22 March); No.
2101	28 (5 April); No. 34 (26 April); No. 40 (10 May); No. 50 (2 June); No. 54
2102	(11 June); No. 58 (21 June); No. 65 (7 July); No. 72 (23 July); No. 79 (9
2103	August); No. 88 (30 August); No. 91 (6 September); No. 99 (24 September);
2104	No. 105 (11 October); No. 107 (18 October); No. 112 (5 November); No.
2105	118 (26 November); No. 120 (3 December); No. 126 (24 December); No. 3
2106	(10 January 1836).
2107	Esmark, J. 1836/37. Meteorologiske Iagttagelser i Christiania 1836, anstillede ved
2108	Professor Esmark. Den Norske Rigstidende, No. 7 (24 January); No. 15 (21
2109	February); No. 17 (28 February); No. 19 (6 March); No. 23 (20 March); No.
2110	27 (3 April); No. 32 (21 April); No. 38 (5 May); No. 45 (22 May); No. 50 (2
2111	June); No. 59 (23 June); No. 66 (10 July); No. 70 (19 July); No. 78 (7
2112	August); No. 85 (23 August?) ; No. 92 (8 September); No. 98 (22
2113	September); No. 105 (9 October); No. 111 (30 October); No. 112 (3
2114	November); No. 119 (27 November); No. 125 (18 December); No. 126 (22
2115	December); No. 3 (5 January 1837).
2116	Esmark, J. 1837/38. Meteorologiske Iagttagelser i Christiania 1837, anstillede ved
2117	Professor Esmark. Den Norske Rigstidende, No. 10 (22 January); No. 17 (7
2118	February); No. 22 (19 February); No. 22 (2 March); No. 34 (19 March); No
2119	41 (4 April); No. 48 (20 April); No. 53 (2 May); No. 61 (21 May); No. 67 (4
2120	June); No. 74 (20 June); No. 82 (9 July); No. 86 (18 July); No. 93 (3
2121	August); No. 100 (20 August); No. 106 (3 September); No. 113 (19
2122	September); No. 120 (5 October); No. 126 (19 October); No. 132 (2
2123	November); No. 139 (19 November); No. 145 (3 December); No. 152 (19
2124	December); No. 2 (4 January 1838).

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2125	Esmark, J. 1838. Meteorologiske Iagttagelser i Christiania 1838, anstillede ved
2126	Professor Esmark. Den Norske Rigstidende, No. 10 (18 January); No. 19 (3
2127	February); No. 29 (20 February); No. 36 (4 March); No. 45 (20 March); No.
2128	53 (3 April); No. 62 (19 April); No. 70 (3 May); No. 79 (19 May); No. 87 (2
2129	June); No. 98 (19 June); No. 108 (4 Junly); No. 117 (19 July); No. 127 (2
2130	August); No. 137 (19 August); No. 148 (6 September); No. 156 (20
2131	September); No. 164 (4 October); No. 173 (20 October); No. 181 (3
2132	November); No. 190 (18 November); No. 199 (4 December); No. 207 (18
2133	December).
2134	

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2135	Appendix B	<u>. Correc</u>	tions of l	Esmark'	<u>s thermo</u>	meter?				•	$\sim$	Formatted: Font: Bold
2136	The correcti	ons are v	ery smal	l for the f	frequent v	winter ter	nperatur	es, but as	high as 0.:	<u>5°C for</u>		Formatted: Normal
2137	frequent sur	nmer tem	peratures	. Due to t	the uncer	<u>tainty wi</u>	th the ide	entificatio	n of Esma	<u>ırk's</u>		
2138	thermometer	we have	not appl	ied these	correctio	ons to his	observat	tions, It sh	ould also	<u>be kept</u>		Moved (insertion) [5]
2139	in mind that	<u>Esmark ι</u>	ised anot	her therm	nometer,	i.e. a min	imum th	ermomete	r for the p	period		
2140	<u>1828.03-183</u>	8.12, wh	ich might	t also hav	<u>e instrun</u>	nental con	rrections	. <u>Howeve</u>	r, he was a	a skilled		
2141	instrument b	uilder, so	it is not	likely that	at he used	l thermor	neter wit	h larger c	orrections	that		
2142	those in Tabl	<u>le B1.</u>										
2143												
2144	Table B1In	strument o	correction	(Corr) for	thermome	eter readir	ngs (Tem	<u>o.). The the</u>	ermometer	<u>may</u>		Moved (insertion) [31]
2145	<u>have been us</u>	ed by Esn	nark, 1816	<u>6-1838.</u>								Moved (insertion) [32]
	<u>Temp. (°C)</u>	<u>25.00</u>	<u>18.75</u>	<u>12.50</u>	<u>6.25</u>	<u>0.00</u>	<u>-6.25</u>	<u>-12.50</u>	<u>-18.75</u>	<u>-25.00</u>		
	<u>Corr. (°C)</u>	+0.50	<u>+0.50</u>	+0.38	+0.38	<u>+0.13</u>	<u>+0.13</u>	<u>+0.13</u>	<u>+0.13</u>	+0.63		
2146												
2147												
2148	Appendix (	2										
2149	MET Norwa		tes mont	hlv mean	temperat	tures for	manual s	tations by	Mohn's (	also		Deleted: Føyn's
2150	called the C-	•		•	-			•	•			
2150	al., 2015), sc		-	•				•				
2152	mean temper									•	_	Deleted: Føyn's
2153	formula, Tab									-		Deleted: A1
2154	1011111111, 140											
2151	Table <mark>C1</mark> . For	mulas for	calculatio	on of mont	hlv mean <sup>.</sup>	temperatu	ıre T wh	ere Too Tu	and Tation	re		Deleted: A1
2156	monthly mear											Deleted. Al
2157	temperature,					-	-		-	5		
2158		5									/	Deleted: Føyn's
	1	<u>Mohn's</u> for	mula	$T = T_c +$	С		$T_{\rm T}$	$T_{18} + T_{15} + T_{15}$	21			Deleted: $T = T_g + k_g (T_{15} - T_g)$
				¥			$I_c = -$	$\frac{1}{3} + T_{15} + T_{15}}{3}$				Field Code Changed
	ŀ	Köppen's f	formula	$T = T_c -$	$k(T_f - T_f)$	')	7	$T_{15} + T_{21}$				Deleted: $T_g = \frac{T_{08} + T_{21}}{2}$
				z1		n /	$T_f = -$	$\frac{T_{15} + T_{21}}{2}$			//	۷
2159												Field Code Changed
2160	A "true" mor	nthly mea	an tempe	rature, T,	may be o	calculated	d by the a	arithmetic	mean of l	hourly		Deleted: $T = T_f - k(T_f - T_n)$
2161	observation	-	-		-		-			5	,	Field Code Changed Deleted: kg
		-						•		lac Ecr		Deleted: Føyn's
2162	constants, $\underline{C}$	anu k <sub>f</sub> , ai	le easily	calculated	u by rear		<u>nonn s</u> a	nu Koppe	n s tormu	1as. roi		Formatted: Footer, Right: 0,63 cm
										36•		

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2173	Esmark's series from Øvre Vollgate the constants were unknown. It was assumed that the		
2174	constants from Blindern could be used also for Øvre Vollgate. An indication of the robustness		
2175	of this assumption was tested by comparison with a short series of hourly observations from		
2176	the station 18815 Oslo – Bygdøy, 15 m a.s.l. The test procedure started with calculation of the	/	Deleted: Føyn's
2177	constants for the Blindern series based on the period 2012.12-2015.09. These constants were		Deleted: 18 Deleted: March
2178	then used for the calculation of mean monthly temperatures for Bygdøy for the same period,		Deleted: 05
2179	which were compared with the "true" monthly means, i.e. those calculated by the hourly		<b>Deleted:</b> These differences are so small that the lack of exact knowledge of the
2180	observations. For <u>Mohn's</u> formula the deviation from the true means varied from -0.06°C in		constants does add practically no uncertainty to the monthly temperatures.
2181	December to $+0,31$ °C in <u>September</u> that gave $+0,10$ °C for the whole year. For seven of the	//	Deleted: ¶ ¶
2182	months the deviation from the true value was less than ±0.1°C. Corresponding figures for		¶ ¶
2183	Köppen's formula were -0.06°C in July, +0.16°C in September and +0.01°C for the whole		1
2184	year,	/	1 ¶ ■
2185	-		¶ ¶



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2206										
2207	<b>Tables</b>								_	Moved down [37]: ¶
2208	<b></b>								M	¶ Fig.
2200	Table 1 Esm	ark's station at		7 as well as oth	er obse	rvation	etations used i	in this article:		Moved (insertion) [23]
2209				ne, period of ob						<b>Deleted:</b> 1. Map of Christiania (now Oslo) 1811 with the location of
2210				start year marks						Esmark's house in Øvre Vollgt. 7 marked.¶
2212	above sea lev			<u>art jour marie</u>				<u></u>	autoconectors:	
2212									CARCEASCONSICCAS	
	No. and name	<u>e</u>	Period (from-	to; year, month	, day I	H <u>s</u> (m)	Additional info	ormation		Jonif: A.
	18651 Oslo II		1837.04.02-1	933.12.31		25	Astronomical	Observatory	-	JK 2
	<u>18654 Oslo -</u>	Øvre Vollgate	<u>1816.01.01-1</u>	838.12.31	-	<u>11</u>	Esmark's obs	ervations		A CALES
-	<u> 18655 Oslo -</u>	Pilestredet	<u>1822.10.19-1</u>	827.02.28	:	<u>16</u>	Hansteen's ol	bservations		
	18700 Oslo -	Blindern	<u>*1993.01.05 t</u>	o present	9	<u>94</u>	Main building	, MET Norway		jaal
	<u>18815 Oslo -</u>	Bygdøy	<u>*2012.01.01 t</u>	o present	-	<u>15</u>	Mainly rural s	tation		A Carlo and a cardia
2214	▼								1	199118
2215										Ver.
2216				n of observatior						Formatted: Font: Arial, 10 pt
2217 2218				<u>ear of each par</u> e homogenous						Moved down [38]: ¶ Fig.
2219				nomogeneity (N						Deleted: 2. Matriculation and survey
2220	<b>v</b>									1830 of Esmark's property No. 308, Øvre Voldgate 7, in Oslo Byarkiv (City
-		<u>)1-1838.12: Th</u>			-					archives). ¶
	SNHT tests	Obs. times	<u>Winter</u>	<u>Spring</u>	Summ		Autumn	<u>Year</u>		Formatted: Font: Arial, 10 pt, English
-	Single shift	<u>l vs II</u>	<u>1833; -1.1</u>	<u>1827; -2.1</u>	<u>1827;</u>	_	<u>1824; -1.4</u>	<u>1827; -1.8</u>		(U.S.) Moved (insertion) [24]
	Single shift	<u>l vs III</u>	<u>1832; -1.5</u>	<u>1826; -2.8</u>	<u>1827;</u>		<u>1827; -1.7</u>	<u>1827; -2.4</u>	`	Moved (insertion) [24]
	Single shift	<u>III vs II</u>	<u>1821; 0.7</u>	<u>1820; 1.5</u>	<u>1821;</u>		<u>1821; 0.6</u>	<u>1821; 0.9</u>		
	Double shift	<u>l vs ll</u>	<u>1826; 1834</u>	<u>1818; 1827</u>	<u>1817;</u>		<u>1824; 1829</u>	<u>1823; 1827</u>		
	Double shift	<u>l vs III</u>	<u>1819; 1832</u>	<u>1820; 1826</u>	<u>1818;</u>		<u>1823; 1829</u>	<u>1818; 1827</u>		
	Double shift	<u>III vs II</u>	<u>1821; 1832</u>	<u>1819; 1835</u>	<u>1821;</u>	1835	<u>1817; 1834</u>	<u>1821; 1835</u>		
	Part 2, 1816.0		14/2 /					N .		
	<u>SNHT-tests</u>		<u>Winter</u>	Spring	Summ		Autumn	Year		
	Single shift	<u>  / </u>	<u>1826; 0.8</u>	<u>1818; 0.7</u>	<u>1817;</u>		<u>1824; 1.0</u>	<u>1823; 0.5</u>		
	<u>Single shift</u>	<u>1 /III</u>	<u>1818; -1.0</u>	<u>1820; -1.7</u>	<u>1818;</u>		<u>1821; -0.9</u>	<u>1818; -1.3</u>		Moved (insertion) [26]
	Single shift	<u>    /   </u>	<u>1821; -0.6</u>	<u>1819; -1.4</u>	<u>1821;</u>	<u>-1.2</u>	<u>1817; -0.8</u>	<u>1821; -0.8</u>		
	Part 3, 1828.0		14/:	O	0.		A t			Moved (insertion) [27]
	<u>SNHT-tests</u>	Obs. times	Winter	Spring	Summ		Autumn	<u>Year</u>		
	Single shift	<u>1/11</u>	<u>1834; -1.0</u>	<u>1834; 0.4</u>	<u>1830;</u>		<u>1829; -0.4</u>	<u>1830; -0.5</u>		
÷	<u>Single shift</u>	<u>1/III</u>	<u>1832; -1.3</u>	<u>1836; -0.6</u>	<u>1836;</u>		<u>1829; -0.9</u>	<u>1836; -0.8</u>		
2221	Single shift	<u>     /   </u>	<u>1833; 0.4</u>	<u>1835; 0.8</u>	<u>1835;</u>	<u>0.9</u>	<u>1834; 0.6</u>	<u>1835; 0.7</u>		
2221										Formatted: Footer, Right: 0,63 cm
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### 2246 2247

Table 3. The same as Table 1, but the single shift test used on monthly resolution. In the 1<sup>st</sup> and 3<sup>rd</sup>

Jun

1827

<u>-3.4</u>

<u>1826</u>

1.3

Jul

1827

<u>-3.5</u>

<u>1821</u>

1.3

Aug

1827

<u>-2.9</u>

<u>1821</u>

1.3

Sep

1825

<u>-1.9</u>

<u>1821</u>

<u>0.8</u>

Oct

1827

<u>-1.1</u>

1820

<u>0.9</u>

Nov

1824

<u>-1.5</u>

<u>1834</u>

<u>0.6</u>

Dec

1833

<u>-1.2</u>

1820

0.7

rows the years of the shifts are shown, and in the 2<sup>nd</sup> and 4<sup>th</sup> rows the adjustments. Period of

2248 observation 1816.01-1838.12.

▼	<u>Jan</u>	<u>Feb</u>	Mar	<u>Apr</u>	May
<u>I/II</u>	<u>1834</u>	<u>1826</u>	<u>1826</u>	<u>1830</u>	<u>1827</u>
	<u>-1.2</u>	<u>-1.4</u>	<u>-1.0</u>	<u>-2.2</u>	<u>-3.3</u>
<u>   /  </u>	<u>1828</u>	<u>1832</u>	<u>1820</u>	<u>1819</u>	<u>1819</u>
	<u>0.6</u>	<u>0.7</u>	<u>1.1</u>	<u>1.7</u>	<u>1.8</u>

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Table 4. Corrections (°C) of the evening observation during the period 1816.01-1821.12

Jan	<u>Feb</u>	<u>Mar</u>	<u>Apr</u>	May	<u>Jun</u>	<u>Jul</u>	Aug	<u>Sep</u>	<u>Oct</u>	Nov	Dec	
<u>0.6</u>	<u>0.6</u>	<u>1.0</u>	<u>1.2</u>	<u>1.3</u>	<u>1.2</u>	<u>1.3</u>	<u>1.3</u>	<u>0.9</u>	<u>0.8</u>	<u>0.3</u>	<u>0.5</u>	

2254

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2253 Table 5. Difference, Diff (°C), of median temperature between Esmark's evening observations and the

observations the following morning. For comparison the differences between the observation at 21

2255 UTC and the minimum temperature the following night are shown for the modern station Oslo -2256 Blindern. The night is defined by the interval 21 - 08 UTC. STD (°C) = standard deviation for the

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differences.

▼		<u>Jan</u>	<u>Feb</u>	Mar	<u>Apr</u>	May	<u>Jun</u>	<u>Jul</u>	Aug	<u>Sep</u>	<u>Oct</u>	Nov	Dec
<u>Esmark</u>	Diff	0.0	0.0	0.0	-0.7	-1.8	-1.6	-1.3	-1.2	0.0	0.5	0.0	0.0
<u>1816.01-1821.12</u>	<u>STD</u>	<u>3.4</u>	<u>2.6</u>	<u>2.4</u>	<u>2.1</u>	<u>2.4</u>	<u>2.3</u>	<u>2.6</u>	<u>2.1</u>	<u>2.1</u>	<u>2.0</u>	<u>2.6</u>	<u>2.2</u>
<u>Esmark</u>	Diff	<u>0.9</u>	<u>0.7</u>	<u>1.2</u>	<u>0.6</u>	0.6	<u>-0.7</u>	<u>-0.6</u>	<u>0.0</u>	<u>1.2</u>	<u>0.6</u>	<u>0.8</u>	<u>0.6</u>
<u>1822.01-1828.02</u>	<u>STD</u>	<u>3.1</u>	<u>2.5</u>	<u>2.3</u>	<u>1.8</u>	<u>2.2</u>	<u>2.4</u>	<u>2.2</u>	<u>2.1</u>	<u>2.9</u>	<u>2.5</u>	<u>2.5</u>	<u>2.4</u>
<u>Esmark</u>	Diff	1.3	<u>1.5</u>	<u>1.9</u>	<u>2.2</u>	<u>3.1</u>	<u>3.1</u>	<u>3.1</u>	<u>3.1</u>	<u>2.5</u>	<u>1.9</u>	<u>1.6</u>	1.3
<u>1828.03-1838.12</u>	<u>STD</u>	<u>2.6</u>	<u>2.3</u>	<u>2.5</u>	<u>1.8</u>	<u>2.1</u>	<u>2.2</u>	<u>2.4</u>	<u>2.3</u>	<u>2.2</u>	<u>2.1</u>	<u>1.9</u>	<u>2.7</u>
Blindern	Diff	<u>1.0</u>	<u>1.5</u>	<u>2.3</u>	<u>2.6</u>	<u>3.2</u>	<u>3.0</u>	<u>2.7</u>	<u>2.4</u>	<u>2.0</u>	<u>1.5</u>	<u>1.0</u>	1.0
<u>1993.09-2015.09</u>	<u>STD</u>	<u>1.7</u>	<u>1.8</u>	<u>1.8</u>	<u>1.7</u>	<u>1.8</u>	<u>1.8</u>	<u>1.7</u>	<u>1.6</u>	<u>1.6</u>	<u>1.6</u>	<u>1.5</u>	<u>1.6</u>

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)	Table 6. The	e ran	k of	mea	an te	mpe	ratur	e in 1	816 f	or mo	onth	s and	seaso	ns du	ring t	he ye	ears 1	<u>816-1838</u>		-[1	Moved (insertion) [3	33]
L	for Oslo (Es	mark	's ol	bser	vatio	ons).	For	comp	ariso	n also	o St	ockho	m is ii	nclude	ed. <mark>T</mark> ł	ne rai	nk rui	ns from low				
2	to high value	es, s	o tha	at the	e lov	vest	temp	o. is ra	anked	no.1	e									-[1	Moved (insertion) [3	34]
1		J	E	Μ	Α	M	J	Ţ	A	<u>s</u>	C	<u>N</u> D	Yr	Wi	Sp	Su	Au	]				
	Oslo	14	6	1	<u>5</u>	1	7	13				<u>8 11</u>	2		1	5	<u>2</u>					
	Stockholm	14	3	6	9	1		18				8 12		6	4	17	3				Moved (insertion) [3	351
ļ												-		-						Ċ		
																				C		
	Table 7. The																			-[!	Moved (insertion) [3	86]
	observation	<u>s), ar</u>	nd fo	or cli	mate	e rec	onsti	ructio	ns fro	m pr	сху	data a	t diffe	rent p	laces	in N	orway	<u>/. For</u>				
	<u>comparison</u>	also	Sto	ckhc	olm i	s inc	lude	<u>d. Th</u>	e rank	<u>runs</u>	s fro	m low	to hig	<u>h valı</u>	ies, s	o tha	t the	lowest				
	temp. is ran	ked 1	I. Th	ne gr	rid p	oint	<u>(59.7</u>	′5⁰N,	10.75	∘E) d	iffe	only s	lightly	from	Esm	ark's	hous	<u>e (59.91ºN</u>	1			
	<u>10.74∘E).</u>	<u>10.74°E).</u>																				
	Place, Cour	nty					F	eb-	Apr-	Ма	v-	Jun-	Refe	erence	s							
								Apr	Aug	Au		Aug										
	Oslo, South	-east	ern	Non	wav		-	2	3	3		5	Fsm	ark's	ohse	rvatio	nns					
	Austlandet,								<u> </u>	-	_	2	-									
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ļ	Austlandet,	Sout	h Ea	aster	<u>'n N</u> o	orwa	У		<u>1</u>					Norc	lli 200	<u>1a</u>						
	Lesja, South	h-eas	tern	Nor	rway	_				1	-			Norc	<u>lli 200</u>	<u>1b</u>						
Ì	Bergen, We	stern	No	rway	L				<u>18</u>				<u>N</u>	lordli	et al.	2003						
ij	Trøndelag,	Mid N	lorw	vay						1	<u>8</u>			Nor	dli 20	04						
ļ	Stockholm,	Swed	den					<u>3</u>	<u>10</u>	<u>g</u>	1	<u>17</u>	Boli	n Cer	tre D	ataba	ase					
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<ul> <li>Figure texts</li> <li>Fig. 1. Map of Christiania (now Oslo) 1811 with the location (red star) of Esmark's house in Ovre Vollgt. 7 marked.</li> <li>Fig. 2. Matriculation and survey 1830 of Esmark's property No. 308, Ovre Voldgate 7, in Oslo Byarkiv (Oslo City Archive). Arrow indicates N. Garden to the left, house surrounding backyard to the right.</li> <li>Fig. 3. Street view of Esmark's house in Ovre Voldgate 7. Photograph from around 1900.</li> <li>Oslo Bymuseum, No. OB.F00897. <u>High buildings on each side built late 19<sup>th</sup> century.</u></li> <li>Fig. 4. The January page from Esmark's meteorological observation protocol from 1823, the year he discovered ice ages. Now deposited at Riksarkivet (National archives). Oslo. S-1570. Det norske meteorologiske institutt. F/Fa. Materiale etter professorer. L0002.</li> <li>Fig. 5. Esmark's first published Christiania weather table, from <i>Den norske Rigstidende</i>, 24 January 1818. <u>Maltese crosses are intended as + signs</u>.</li> <li>Fig. 6. The temperature difference (*C) between Esmark's evening observation and the morning observation the following day for the winter season (Dec-Feb) in the period 1831-1838.</li> <li>Fig. 7. Corrections added to Esmark's series for each season during his period of observation 1816-1838.</li> <li>Fig. 9. Temperature differences (*C) between the observations at 15 UTC and at 21 UTC for the following stations: Oslo - 18 Indern for the poriod 1993.01-2015.09. Esmark 1816.01-1838.12. (The corrections, of the evening observators, Table 4, are added to the data for the period 1816.01-1838.12. (The corrections of the evenice) observators, Table 4, are added to the data for the period 1837.04-1867.12.</li> <li>Fig. 10. Differences in mean monthly temperature between Esmark's observations at Ovre Vollgate and those at the Astronomical Observatory (Esmark minus Observatory) during the period 1837.04-1838.12. Themperatures are not corrected.</li> <li>Fig. 11. Differences hetween Esmark's observatory during the period 1837.04-1838.12. Themperatu</li></ul>		
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<ul> <li>1838.12. (The <u>corrections</u> of the evening observations, Table <u>4</u>, are added to the data for the period 1816.01-1821.12 before the calculation of the differences) and Oslo II (Astronomical Observatory) 1837.04-1867.12.</li> <li>Fig. <u>10</u>. Differences in mean monthly temperature between Esmark's observations at Øvre Vollgate and those at the Astronomical Observatory (Esmark minus Observatory) during the period 1837.04-1838.12. <u>Temperatures are not corrected</u>.</li> <li>Fig. <u>11</u>, Difference between Esmark's observations at Øvre Vollgate and Hansteen's observations at Pilestredet (Esmark minus Hansteen) during the period 1822.11-1827.02 at</li> </ul>		
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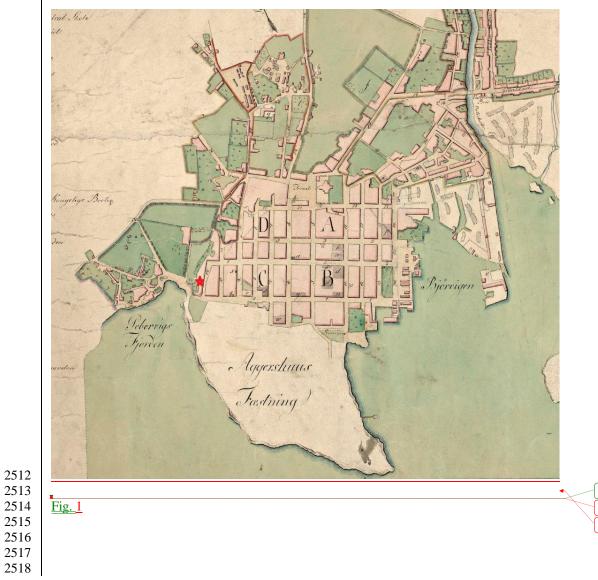
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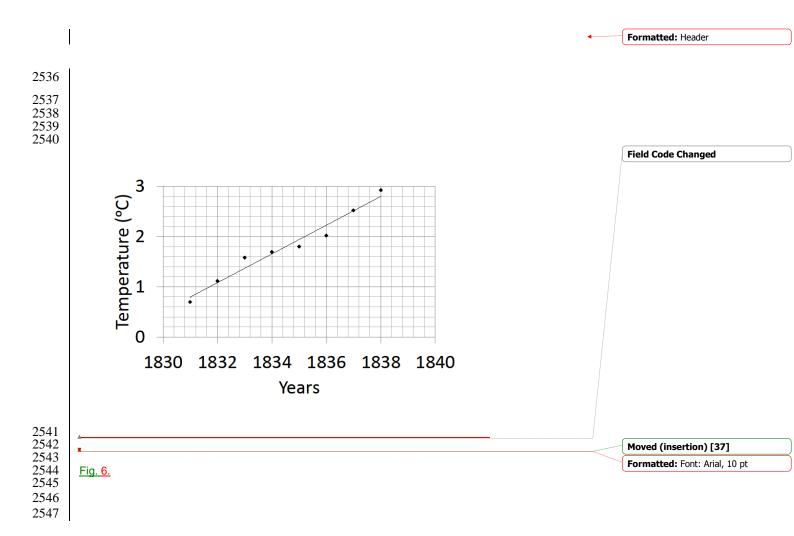
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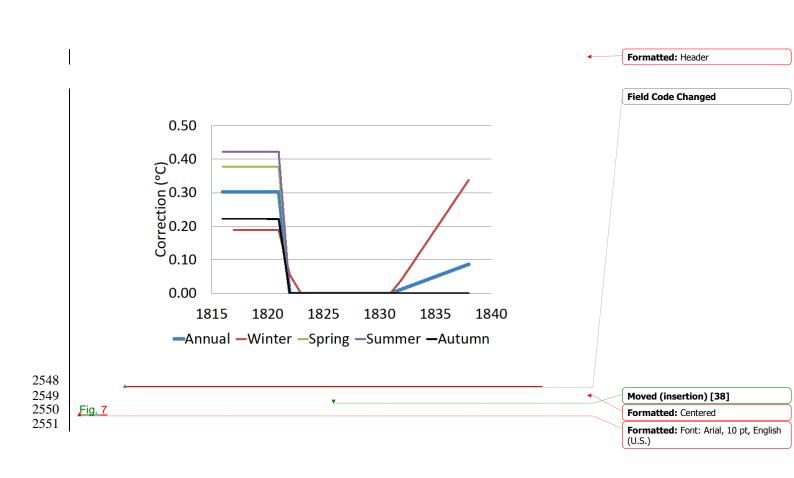
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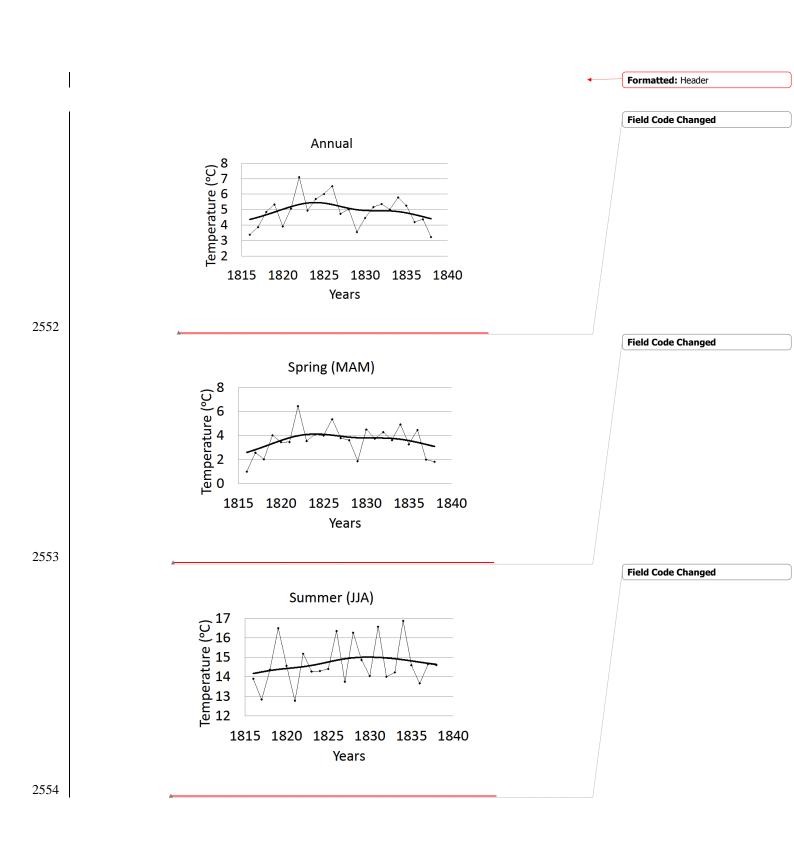
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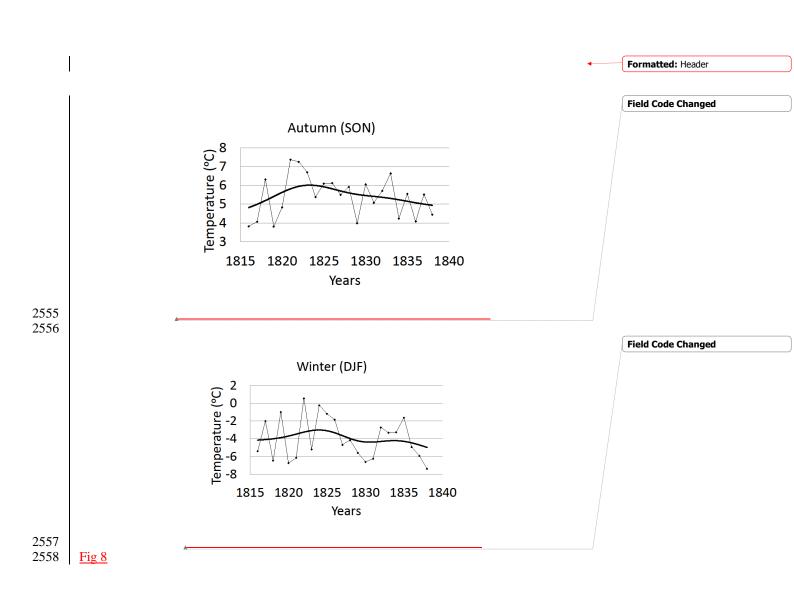


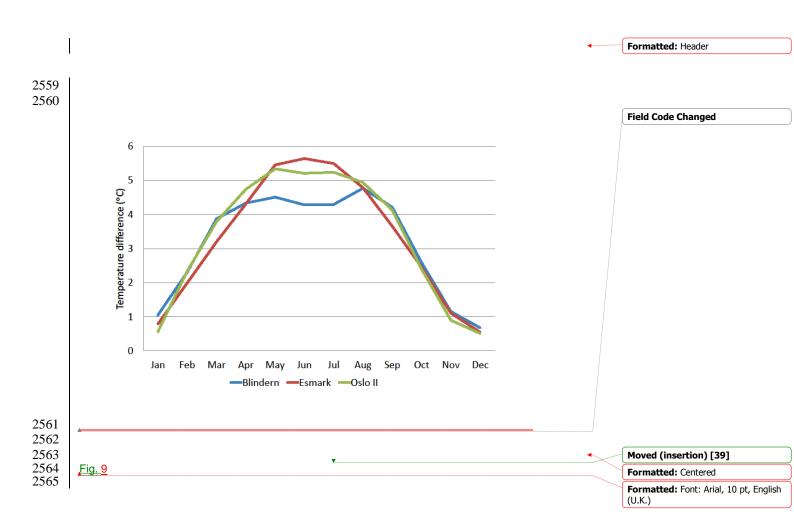
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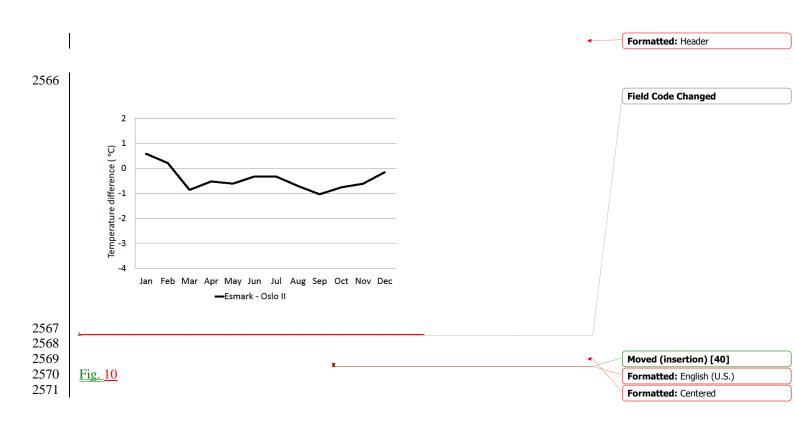


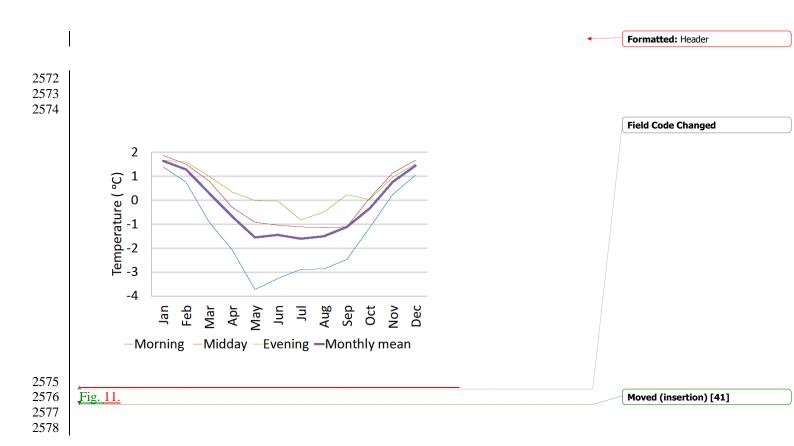
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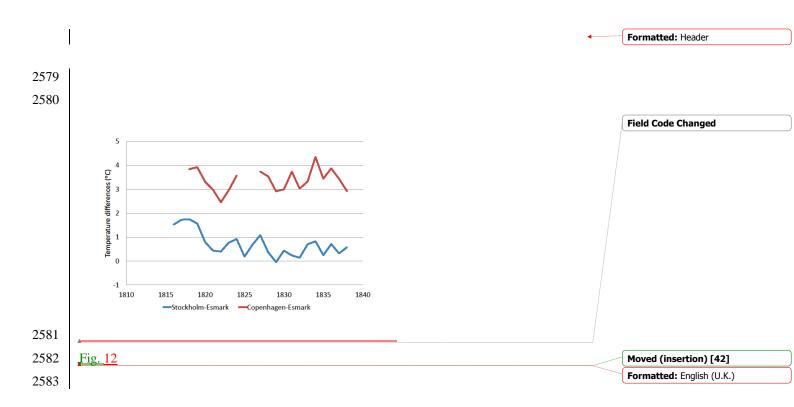
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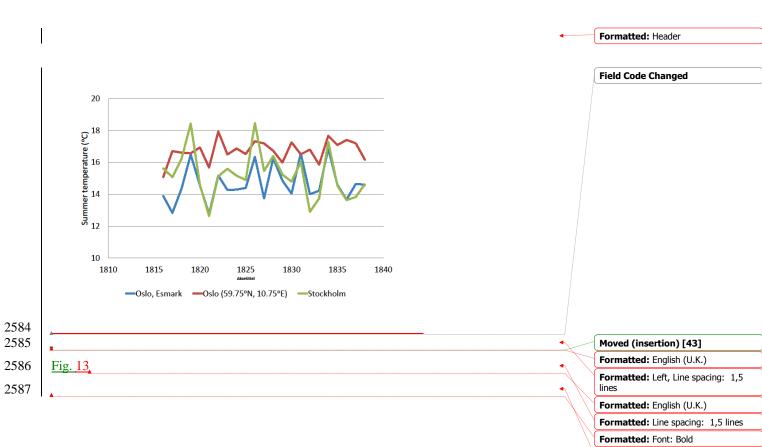












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