1 Supplementary materials for:

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Three distinct Holocene intervals revealed in NW Madagascar: evidence from two stalagmites from two caves, and implications for ITCZ dynamics

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23 currents and convergence zones (ITCZ and ZAB, Zaire Aire Boundary). The map of the currents

- 24 was obtained from Lindesay, 1998 and Schott and McCreary (2001). The map of the ITCZ and
- 25 ZAB was adopted from Gasse (2000). The red star indicates the study location.





29 (the wettest months) in Madagascar. Red star indicates the study area. Source:

30 <u>http://iridl.ldeo.columbia.edu/</u> (accessed August 31, 2016)





34 (ANJB) and Anjokipoty (ANJK) Caves and the current extent of vegetation cover in northwestern

- 35 Madagascar.
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45 Stalagmite ANJB-2 and MAJ-5. (a) 100% calcite. (b) 50% calcite-50% aragonite. (c) 100%





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- 49 Figure S6: Photographs illustrating the mid-Holocene hiatus. Photographs showing the Type L
- 50 layer-bounding surfaces in Stalagmite ANJB-2 (a) and in Stalagmite MAJ-5 (b). Pinching of layers
- 51 toward the flank are indicated with arrows. Also, note the white and porous layer of aragonite in
- 52 Stalagmite ANB-2 that is capped with a very thin brown layer.



54 Figure S7: Stable isotope profile of δ^{18} O and δ^{13} C of the late Holocene in Madagascar from

55 Anjohibe Cave's stalagmites showing the δ^{13} C from C₃-dominated to C₄-dominated vegetation

56 (Burns et al., 2016; Voarintsoa et al., in revision). Note that the age scale is in year AD.

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Figure S8: The 8.2 ka event identified in Stalagmite ANJB-2. a) Scanned image of a portion of
Stalagmite ANJB-2 showing the 8.2 ka event and the corresponding trenches for radiometric

61 dating and X-ray diffraction analyses. b) X-ray diffraction spectra of the stalagmite layers at 195,

- 62 200, and 212 mm from the top of the stalagmite.
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65	Supplementary text 1: Possible climatic response of the latitudinal migration of the ITCZ
66	The ITCZ migrates southward in austral summer and northward in boreal summer in
67	response to seasonal insolation. Climate simulations have also reported long-term migration of
68	the ITCZ, the causes of which have been ascribed to changes in insolation and difference in
69	temperature between the Northern and the Southern Hemisphere (e.g. Chiang and Bitz, 2005;
70	Broccolli et al., 2006; Braconnot et al., 2007). The climatic responses to the ITCZ dynamics can
71	vary from region to region and from one time interval to another. Here, we attempt to provide
72	different conceptualized models to understand the climatic regime in Madagascar during the
73	early, the middle, and the late Holocene (as proposed in Sect. 5.2. of the manuscript).
74	Dry conditions in Madagascar could be conceptualized as dry years, such as modeled in
75	case 1 (Fig. S9). In that model, austral summer months receive less rainfall and austral winter
76	months could receive no rainfall. This model could be used to conceptualize the climatic
77	response in Madagascar when the ITCZ moves north, such as during the mid-Holocene.
78	Wet conditions in Madagascar could be conceptualized as wet years, such as modeled in
79	case 2 and case 3 (Fig. S9). Case 2 suggests that austral summer months receive rain as well as
80	austral winter months, thus it suggests less seasonality. Although the amount of rainfall received
81	during winter months cannot be easily estimated, this model could be used to understand the
82	climatic response in Madagascar during the early Holocene when the southern hemisphere
83	winter insolation was relatively greater than the northen Hemisphere winter insolation (Fig. 5a).
84	The ITCZ was already moved southward due to the globally cold conditions (Figs. 6a, c), and
85	because of the higher winter insolation in the SH, heating of land would bring additional
86	precipitation during austral winter. Case 3 suggests that austral summers receive more rainfall

than normal years, but austral winters stay dry or with little precipitation, thus seasonality must
have been stronger. This scenario could be used to understand the climatic response in
Madagascar during the late Holocene when the southern hemisphere summer insolation was
greater than the northern hemisphere summer insolation (Fig. 5a). Globally cooler conditions
(Figs. 6a, b) already suggest a southward migration of the ITCZ, and the greater SH summer
insolation could have intensified the monsoonal rainfall in northwestern Madagascar during
austral summers.



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95 Figure S9: Conceptualizing the different possible outcomes of the long-term latitudinal migration 96 of the ITCZ. a) Highlighting the three possible scenarios of the Holocene. b) Barplots of monthly 97 rainfall in northwestern Madagascar, using the modern data as a reference to estimating the 98 region's paleoclimate during drier and wetter conditions. c) Global rainfall maps from NASA (same 99 source as Fig. 1 in the manuscript). These maps are modern, but they are only shown here to give a better perspective of the position of Madagascar when the ITCZ is relatively north or south.

101 Tables

Dft (mm)	Sample no.	²³⁸ U (ppb)		²³² Th (ppt)		²³⁰ Th / ²³² Th (atomic x10 ⁻⁶)		d ²³⁴ U* (measured)		²³⁰ Th / ²³⁸ U (activity)		²³⁰ Th Age (yr) (uncorrected)		²³⁰ Th Age (yr) (corrected)		d ²³⁴ U _{initiai} ** (corrected)		²³⁰ Th Age (yr BP)*** (corrected)	
3	ANJ-B-2-U003	3371	±11	39850	±809	10	±0	5.2	±2.0	0.0070	±0.0001	761	±7	419	±243	5	±2	355	±243
8	ANJB-2-008	194.6	±0.3	410	±8	20	±3	3.6	±1.7	0.0026	±0.0004	284	±47	223	±64	4	±2	161	±64
25	ANJB-2-025	4646.6	±6.6	1594	±32	216	±5	3.4	±1.4	0.0045	±0.0001	489	±7	479	±10	3	±1	417	±10
47	ANJ-B-2-U047	64	±0	634	±15	31	±12	3.0	±4.3	0.0187	±0.0074	2052	±822	1762	±845	3	±4	1697	±845
53	ANJ-B-2-U053R	134	±0	2325	±47	20	±4	4.6	±2.0	0.0211	±0.0037	2313	±416	1808	±547	5	±2	1743	±547
72	ANJB-2-072	67.8	±0.1	382	±8	48	±3	8.5	±2.4	0.0163	±0.0008	1778	±93	1615	±147	9	±2	1553	±147
92	ANJ-B-2-U092	78	±0	180	±8	92	±40	5.1	±3.5	0.0129	±0.0056	1408	±610	1341	±612	5	±4	1276	±612
105	ANJ-B-2-U105	117	±0	229	±9	113	±34	13.5	±2.9	0.0134	±0.0040	1450	±432	1393	±434	14	±3	1329	±434
112	ANJ-B-2-U112	1322	±2	7456	±150	42	±2	8.5	±2.0	0.0145	±0.0005	1576	±50	1413	±125	9	±2	1348 ±12	
116	AB-1a	130.9	±0.2	530	±14	136.7	±16.7	27.0	±3.0	0.033539275	±0.00400	3620	±439	3506	±446	27.2	±3.0	3444 ±44	
118	AB-2a	2569.8	±2.9	5266	±106	580.4	±11.9	7.5	±1.6	0.072126026	±0.00032	8100	±40	8040	±58	7.6	±1.6	7978 ±5	
120	ANJ-B-2-U120	1710	±4	20753	±418	108	±2	9.1	±2.3	0.0796	±0.0004	8955	±48	8605	±252	9	±2	8541	±252
120	ANJ-B-2-U120R	2075	±3	13340	±268	197	±4	6.3	±1.5	0.0767	±0.0003	8640	±38	8454	±137	6	±2	8389	±137
130	ANJB-2-130	3042.4	±3.8	7448	±149	477	±10	6.2	±1.5	0.0709	±0.0002	7966	±26	7895	±57	6	±2	7833	±57
160	ANJB-2-160	2994.8	±4.1	2484	±50	1416	±29	4.3	±1.5	0.0712	±0.0002	8021	±30	7997	±35	4	±2	7935	±35
185	ANJ-B-2-U185	3490	±5	6040	±122	690	±14	3.6	±1.7	0.0724	±0.0003	8167	±33	8117	±48	4	±2	8053	±48
201	ANJ-B-2-U205	574	±1	1881	±38	374	±8	5.7	±1.8	0.0743	±0.0006	8367	±70	8272	±97	6	±2	8208	±97
215	ANJ-B-2-U215	3146	±4	5418	±109	713	±15	7.0	±1.5	0.0745	±0.0003	8379	±33	8329	±48	7	±2	8265	±48
251	ANJ-B-2-U251	4246	±5	7290	±147	745	±15	6.3	±1.3	0.0776	±0.0002	8750	±27	8700	±44	6	±1	8636	±44
275	ANJ-B-2-U275	6077	±9	9132	±184	861	±17	4.5	±1.5	0.0785	±0.0002	8867	±32	8823	±44	5	±2	8759	±44
280	MAJ-B-2-U280	5721	±18	5408	±110	1360	±28	2.4	±1.6	0.0780	±0.0003	8828	±36	8801	±41	2	±2	8737	±41
302	ANJ-B-2-U302	9833	±44	1617	±33	8024	±166	5.2	±1.9	0.0800	±0.0004	9045	±50	9041	±50	5	±2	8977	±50

102 Table S1: ²³⁰Th dating results for Stalagmite ANJB-2. The error is 2 σ error. Dft= distance from the top of the stalagmite.

- 103 U decay constants: $\lambda_{238} = 1.55125 \times 10^{-10}$ (Jaffey et al., 1971) and $\lambda_{234} = 2.82206 \times 10^{-6}$ (Cheng et al., 2013). Th decay constant: $\lambda_{230} = 9.1705 \times 10^{-6}$
- 104 (Cheng et al., 2013).
- $105 \qquad *\delta^{234}U = ([^{234}U/^{238}U]_{activity} 1)x1000. \\ **\delta^{234}U_{initial} \text{ was calculated based on }^{230}\text{Th age (T), i.e., } \\ \delta^{234}U_{initial} = \delta^{234}U_{measured} \times e\lambda^{234xT}.$
- 106 Corrected ²³⁰Th ages assume the initial ²³⁰Th/²³²Th atomic ratio of $4.4 \pm 2.2 \times 10^{-6}$. Those are the values for a material at secular
- 107 equilibrium, with the bulk earth 232 Th/ 238 U value of 3.8. The errors are arbitrarily assumed to be 50%.
- 108 ***B.P. stands for "Before Present" where the "Present" is defined as the year 1950 A.D.
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110 Table S2: ²³⁰Th dating results for Stalagmite MAJ-5. The error is 2 σ error. Dft= distance from the top of the stalagmite.

Dft (mm)	Sample no.	²³⁸ U (ppb)		²³² Th (ppt)		²³⁰ Th / ²³² Th (atomic x10 ⁻⁶)		d ²³⁴ U* (measured)		²³⁰ Th / ²³⁸ U (activity)		²³⁰ Th Age (yr) (uncorrected)		²³⁰ Th Age (yr) (corrected)		d ²³⁴ U _{Initial} ** (corrected)		²³⁰ Th Age (yr BP)*** (corrected)	
1	MAJ-5-U001	2734	±15	3044	±63	33	±1	-3.2	±2.7	0.0023	±0.0000	246	±5	214	±24	-3	±3	150	±24
10	MAJ-5-U010	6691	±38	22757	±474	27	±1	-3.1	±3.6	0.0056	±0.0001	609	±7	510	±71	-3	±4	446	±71
22	MAJ-5-U022	3292	±4	11633	±234	31	±1	-3.0	±1.5	0.0067	±0.0001	736	±14	633	±74	-3	±1	569	±74
41	MAJ-5-U041	1380	±3	10604	±213	32	±1	-1.1	±2.1	0.0147	±0.0001	1617	±15	1393	±159	-1	±2	1329	±159
50	MAJ-5-U050	1224	±4	4144	±84	40	±1	-2.9	±2.4	0.0082	±0.0001	898	±15	799	±71	-3	±2	735	±71
60	MAJ-5-U060	1578	±3	14591	±293	31	±1	-0.9	±2.6	0.0173	±0.0005	1901	±56	1631	±199	-1	±3	1567	±199
66	MAJ-5-U066	12609	±83	38990	±842	461	±10	-4.6	±2.9	0.0865	±0.0006	9912	±81	9821	±103	-5	±3	9757	±103
80	MAJ-5-U080	11684	±16	27838	±559	598	±12	-2.6	±1.2	0.0864	±0.0002	9882	±24	9813	±55	-3	±1	9749	±55
89	MAJ-5-U089	10930	±12	30247	±606	519	±10	-1.2	±1.3	0.0870	±0.0002	9941	±29	9860	±64	-1	±1	9796	±64

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112 U decay constants: $\lambda_{238} = 1.55125 \times 10^{-10}$ (Jaffey et al., 1971) and $\lambda_{234} = 2.82206 \times 10^{-6}$ (Cheng et al., 2013). Th decay constant: $\lambda_{230} = 9.1705 \times 10^{-6}$

113 (Cheng et al., 2013).

- 114 $\delta^{234}U = ([^{234}U/^{238}U]_{activity} 1)x1000$. ** $\delta^{234}U_{initial}$ was calculated based on ²³⁰Th age (T), i.e., $\delta^{234}U_{initial} = \delta^{234}U_{measured} \times e\lambda^{234xT}$.
- 115 Corrected 230 Th ages assume the initial 230 Th/ 232 Th atomic ratio of 4.4 ±2.2 x10⁻⁶. Those are the values for a material at secular

116 equilibrium, with the bulk earth 232 Th/ 238 U value of 3.8. The errors are arbitrarily assumed to be 50%.

- 117 ***B.P. stands for "Before Present" where the "Present" is defined as the year 1950 A.D.
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