

Interactive comment on “The response of the Peruvian Upwelling Ecosystem to centennial-scale global change during the last two millennia” by R. Salvattecí et al.

R. Salvattecí et al.

renatosalvattecí@gmail.com

Received and published: 16 January 2014

General comments of the anonymous referee #1

In general I found that the interpretations in section 5.2 were rather far reaching considering the difficult nature of developing the time scale and interpreting each proxy. For example, parts of the DACP and MCA are missing, limiting general statements about the uniform behavior of climate during warm versus cold phases. As the authors point out, there are also discrepancies between the export production proxies, and I mention some potential problems in interpreting TOC below. This being said, everything is presented in a logical framework, the proxies are directly comparable, and the reader has

C3135

all the information needed to judge the strength of the hypotheses.

Reply: In this version of the manuscript we took into account the general comments of the anonymous referee #1 and the interpretations in section 5.2 are now more cautious.

Specific comments

1. I do not fully agree with the generic interpretation of TOC as an export production proxy. TOC percentage is necessarily related to export production, redox conditions, and the delivery of terrestrial organic carbon. It may also be biased by selective winnowing from ocean currents. The table in SM7 demonstrates that within reason, TOC is as closely correlated with V and Re (redox indicators) as it is other export proxies such as Ni, Cu, and Cd. Furthermore, the elevated TOC% during the CWP may simply be a function of enhanced preservation in recently deposited sediments. I thus encourage the authors to be more tentative in their treatment of TOC as an export proxy. For example, interpretations of export production in Fig. 4 are based partly on TOC, and it may not be a valid proxy in this sense.

Reply: In the revised manuscript we took into account the reviewer's comment. Interpretations of export production are now based on the standardized average of Ni and Cu (Fig. 4). These two metals, as explained in the text, reflect the original presence of organic matter even if it is partially lost after deposition (Tribovillard et al., 2006).

2. The interpretations in terms of ITCZ variability may be a bit tenuous. Is the ITCZ clearly defined on the Peruvian coast, or is precipitation variability tied more closely to ENSO events? The Pisco record is in anti-phase with the Haug et al. (2001) record during the MCA, as would be expected for a straightforward ITCZ relationship. But then why are these records in phase during the early part of the record? More discussion is required to explain this change in character and the implications for the ITCZ.

Reply: At interannual timescales, precipitation variability on the northern Peruvian coast is associated with ENSO events, but only very extremes El Niño events (e.g.

C3136

1998) can trigger precipitation and floods further south. Precipitation variability at seasonal timescales is associated with meridional displacements of the ITCZ and the SPSH. A weakening of the SPSH and the southward displacement of the ITCZ during austral summer induce the development of a rainy season (Garreaud et al. 2009, Lavado et al. 2012). In the revised manuscript we further discuss our terrigenous input record and other continental records (Reuter et al., 2009; Bird et al., 2011) with regards to the ITCZ displacements and SPSH expansion/contraction. The Pisco record is in anti-phase with the Haug et al. (2001) record from the MCA towards the present as expected for a straightforward ITCZ relationship, but there is an apparent in phase relationship from 300 to 800 AD. In this period the magnitude of precipitation change is higher compared to the precipitation changes during the LIA, but the timing and direction of the precipitation changes are associated with ITCZ displacements. The Haug et al. (2001) record shows a decreasing trend from ~400 to ~750 AD indicating a southward migration of the ITCZ while the Pisco record starts an increasing trend in the same time interval.

3. Are the RWP and DACP well developed in other parts of the tropics, or is this a novel record of their expression?

Reply: To our knowledge the present manuscript presents the most reliable oceanic record for the last 2,000 years in the Eastern Tropical Pacific (ETP). There are very few studies showing the precipitation changes in the ETP to the RWP and DACP such as for example, the grain size record of El Junco Lake in Galapagos (Conroy et al., 2008). This record suggests large changes in precipitation during the last 2 millennia. Increased precipitation is inferred in Galapagos during part of the DACP and the LIA, and lower precipitation during the RWP, part of the DACP and the MCA. These results are in part coherent with the results presented in our manuscript.

4. The authors cite erosion by ocean currents as a potential problem early in the manuscript (bottom of p. 5487) but then do not discuss it later. Could bottom currents have caused bulk or selective sediment redistribution, for example between the marine

C3137

and terrestrial proxies in the authors' records?

Reply: The intensification of the Poleward undercurrent has been made responsible for hiatuses in the sediment record at longer timescales (Reinhardt et al. 2002), but little is known about the effects of this current on sedimentation processes at shorter timescales. It is known that strong bottom currents can cause selective sediment redistribution, if currents are stronger, there would be a higher chance to re-suspend light-density organic matter, and therefore to leave behind high density components (i.e. detritics). A more thorough analysis should be done to understand the effect of change in this current like a ratio between a heavy and a light element or changes in grain size. However the strong correlation of the terrestrial proxies (Al and Ti) with $\delta^{15}\text{N}$ (which should not be affected by selective sediment redistribution) suggests that the bottom currents didn't cause large changes in sediment redistribution and that the proxies are truly reflecting biogeochemical changes.

5. Are there geochemical results from the slumped sections that were removed from the stratigraphic sequence? If so, are the data reasonably homogenous?

Reply: We developed on the slumped sections of core B-14 the same suite of proxies that are shown in the manuscript. The homogeneous sections show lower variability in comparison to the contiguous laminated sequences. Additionally we develop higher resolution records on several cores as for example gray level measurements and XRF measurements. The results of these analyses effectively indicate that the slumped/homogeneous sections show lower variability than the laminated sections of the core.

6. On p. 5498 (lines 7-8), the authors suggest possible control of export production by changes in OMZ intensity (paraphrased). How would this work? Wouldn't the OMZ normally respond to changes in organic carbon export? Could something else be contributing to changes in the OMZ, perhaps involving changing current regimes? In any case, more explanation is needed on this point.

C3138

Reply: A vertical expansion of the OMZ produced by changes in ocean circulation (i.e. reduced ventilation) or global warming (Stramma et al., 2008) could reduce the vertical oxygenated habitat for some organisms. This condition is adverse for a large group of macro organisms that cannot survive or avoid hypoxic zones. Then, a shallower oxycline could reduce the habitat for several pelagic organisms, diminishing their abundance, and modifying the flux of organic matter to the sediments. In the geologic past, reductions of ocean oxygen were responsible for massive extinctions (e.g. Wignall and Twitchett, 1995). In the corrected version of the manuscript we removed this idea that is beyond the scope of our work.

7. The strong correlations between the Pisco and Cascayunga Cave records are very interesting. Perhaps the authors could further interpret their records specifically in relation to the findings of Reuter et al. (2009).

Reply: In the revised manuscript we discuss in more detail our findings with those obtained from Cascayunga Cave (Reuter et al., 2009) and Pumacocha Lake (Bird et al., 2011).

Technical corrections

The minor language errors were corrected.

On at least one of the records figures, it would be nice to see the stratigraphic positions of the radiocarbon dates, and the slumped and laminated sections.

Reply: Figure 2 was modified according to the reviewer's suggestions

Literature cited

Bird, B. W., Abbott, M. B., Vuille, M., Rodbell, D. T., Stansell, N. D., and Rosenmeier, M. F.: A 2,300-year-long annually resolved record of the South American summer monsoon from the Peruvian Andes, *PNAS*, 108(21), 8583-8588, 2011.

Conroy, J. L., J. T. Overpeck, J. E. Cole, T. M. Shanahan, and M. Steinitz-Kannan.

C3139

Holocene changes in eastern tropical Pacific climate inferred from a Galapagos lake sediment record. *Quaternary Science Reviews* 27(11-12):1166-1180, 2008.

Garreaud, R. D., Vuille, M., Compagnucci, R., and Marengo, J. Present-day South American climate. *Palaeogeography, Palaeoclimatology, Palaeoecology* 281(3-4):180-195, 2009.

Gutierrez, D., A. Sifeddine, D. B. Field, L. Ortlieb, G. Vargas, F. Chavez, F. Velazco, V. Ferreira, P. Tapia, R. Salvattecchi, H. Boucher, M. C. Morales, J. Valdes, J. L. Reyss, A. Campusano, M. Boussafir, M. Mandeng-Yogo, M. Garcia, and T. Baumgartner. Rapid reorganization in ocean biogeochemistry off Peru towards the end of the Little Ice Age. *Biogeosciences* 6:835-848, 2009.

Haug, G. H., Hughen, K. A., Sigman, D. M., Peterson, L. C., and Rohl, U.: Southward migration of the Intertropical Convergence Zone through the Holocene, *Science*, 293, 1304-1308, 2001.

Lavado Casimiro, W. S., Ronchail, J., Labat, D., Espinoza, J. C., and Guyot, J. L. Basin-scale analysis of rainfall and runoff in Peru (1969-2004): Pacific, Titicaca and Amazonas drainages. *Hydrological Sciences Journal* 57(4):1-18, 2012.

Reinhardt, L., H.-R. Kudrass, A. Lückge, M. Wiedicke, J. Wunderlich, and G. Wendt. High-resolution sediment echosounding off Peru: Late Quaternary depositional sequences and sedimentary structures of a current-dominated shelf *Marine Geophysical Researches* 23(4):335-351, 2002.

Reuter, J., Stott, L., Khider, D., Sinha, A., Cheng, H., and Edwards, R. L.: A new perspective on the hydroclimate variability in northern South America during the Little Ice Age, *Geophys. Res. Lett.*, 36, L21706, 2009.

Tribouillard, N., Algeo, T. J., Lyons, T., and Riboulleau, A.: Trace metals as paleoredox and paleoproductivity proxies: An update, *Chem. Geol.*, 232, 12-32, 2006.

Wignall, P. B. and Twitchett, R. J. Oceanic Anoxia and the End Permian Mass Extinc-

C3140

tion. *Science* 272:1155-1158, 1996.

Interactive comment on *Clim. Past Discuss.*, 9, 5479, 2013.

C3141