

A refined TALDICE-1a age scale from 55 to 112 ka before present for the Talos Dome ice core based on high-resolution methane measurements

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Abstract

A precise synchronization of different climate records is indispensable for a correct dynamical interpretation of paleoclimatic data. A chronology for the TALDICE ice core from the Ross Sea sector of East Antarctica has recently been presented based on methane synchronization with Greenland and the EDC ice cores and $\delta^{18}\text{O}_{\text{ice}}$ synchronization with EDC in the bottom part (TALDICE-1). By the use of new high-resolution methane data, obtained with a continuous flow analysis technique, we present a refined age scale for the age interval from 55 – 112 ka before present where TALDICE is synchronized with EDC. New and more precise tie points reduce the uncertainties of the age scale from up to 1900 years in TALDICE-1 to below 1100 years over most of the refined interval. Thus, discussions of climate dynamics at sub-millennial time scales are now possible back to 110 ka, in particular during the inception of the last ice age. Calcium data of EDC and TALDICE are compared to show the impact of the refinement to the synchronization of the two ice cores not only for the gas but also for the ice age scale.

1 Introduction

For a good understanding of mechanisms at work in the climate system it is indispensable to know the chronology and phase relationships of climate events in the past. Precise dating of climate archives such as ice cores is therefore necessary to optimally utilize the information stored in such archives. Ice cores contain various strains of information on climate and environmental changes in the past. These comprise the water isotopic signature of the ice matrix, dissolved and particulate

33 aerosol tracers as well as the gas composition of the atmosphere all in one climate archive.
34 Accordingly, synthesizing these strains of ice core information circumvents crucial cross-dating
35 issues that affect the comparison of independent climate archives. For the comparison of different ice
36 cores absolute dating of each core is not necessary, it is sufficient to synchronize the records properly
37 by the use of a global tracer. Air trapped in polar ice cores has the unique property of containing
38 global tracers of the atmosphere, which show the same variations over time at drilling sites on both
39 hemispheres. Thus, it is possible to build relative age scales of different ice cores by synchronizing
40 the respective methane (CH₄) records (Blunier and Brook, 2001; Blunier et al., 1998; Blunier et al.,
41 2007; Chappellaz et al., 1997; EPICA, 2006). Methane is particularly well suited for such a
42 synchronization because abrupt concentration changes have been observed over large periods back to
43 800 thousand years before present (ka BP) not only at glacial-interglacial transitions but also during
44 glacial times, especially during Dansgaard-Oeschger (DO) events (Brook et al., 2000; Chappellaz et
45 al., 1997; Huber et al., 2006; Loulergue et al., 2008; Spahni et al., 2005). Most probably, present-day
46 interhemispheric mixing time of about one year (Warneck, 1988) and atmospheric lifetime in the
47 order of 10 years (Lelieveld et al., 1998) did not change substantially under the past conditions of the
48 studied period. Thus, these abrupt CH₄ concentration changes are global time markers which are well
49 archived in all polar ice cores.

50
51 The synchronization of ice cores is limited by the mixing of the air in the firn before bubble close-off
52 which causes different age distributions of the enclosed gas depending on accumulation and
53 temperature at the drilling site. This age distribution as well as the firnification process can be
54 modelled (Goujon et al., 2003; Schwander et al., 1993; Spahni et al., 2003) within its model
55 uncertainties. It has been shown that additional uncertainties of methane tie points of up to 200 years
56 can be caused by different gas enclosure characteristics at different drilling sites (Köhler et al.,
57 2011). Even larger errors may arise for very low accumulation rate sites (such as Vostok, Dome Fuji
58 or Dome C), where firnification models seem to be in contradiction with $\delta^{15}\text{N}_2$ measurements
59 (Landais et al., 2006). Another important limitation usually is the limited resolution of the methane
60 records. Records with higher resolution preserve fast concentration changes better. Therefore, tie
61 points can be defined more precisely.

62
63 The first official chronology (TALDICE-1) of the deep ice core TALDICE (TALos Dome Ice CorE)
64 at Talos Dome in the Ross Sea sector of East Antarctica (72°47' S, 159°11' E), based on an inverse
65 model (Lemieux-Dudon et al., 2010) and methane synchronization with Greenland ice cores (Blunier
66 et al., 2007) (0–50 ka BP) and the EPICA Dome C (EDC) ice core (Loulergue et al., 2008; Spahni et

67 al., 2005) (50–140 ka BP) as well as $\delta^{18}\text{O}_{\text{ice}}$ synchronization with EDC for ages older than 140 ka,
68 has recently been published by Buiron et al. (2011). The well-resolved (mean resolution of 87 years)
69 TALDICE methane record was synchronized with the Greenland record back to 50 ka BP. Relative
70 age uncertainty of TALDICE-1 remains lower than 600 years in this period (except for the Last
71 Glacial Maximum where abrupt methane variations are missing). However, for the time period from
72 50–140 ka BP where the methane synchronization was made with the EDC ice core, the age
73 uncertainty increases to 2 ka, mainly due to the coarse resolution (mean resolution of 620 years) of
74 the TALDICE methane record. Note that all ages in ka BP given in this paper are relative to 1950
75 AD.

76
77 The purpose of this paper is to apply a new continuous measurement technique for methane
78 (Schüpbach et al., 2009) and to produce a high-resolution CH_4 record for the early part of the last ice
79 age. In the new record we define 12 new age tie points which result from the high-resolution record.
80 With these additional constraints we are able to present a refined age scale (TALDICE-1a) for the
81 time period from 55–112 ka BP based on the TALDICE-1 age scale. The impact of the refinement of
82 the age scale to the synchronization of TALDICE and EDC ice cores is shown by a comparison of
83 Calcium (Ca^{2+}) records of the two cores in a selected interval. This provides an independent means
84 of verifying the quality of the revised age scale TALDICE-1a.

85
86 The paper is organised as follows. In Sect. 2 we describe the new high-resolution CH_4 data and the
87 construction of the revised age scale. Section 3 presents a discussion of the implications of the new
88 time scale, in particular on ice-based records, and conclusions are given in Sect. 4.

89

90 **2 Experimental methods and age scale construction**

91 Methane measurements on TALDICE were performed with a new on-line melting technique using a
92 Continuous Flow Analysis (CFA) system (Schüpbach et al., 2009) in the depth interval from 1187 m
93 to 1488 m. These measurements cover the section where TALDICE was synchronized with the EDC
94 ice core (1228 m to 1428 m, Buiron et al., 2011) by use of discrete methane measurements using a
95 traditional melt-refreeze extraction method (Chappellaz et al., 1997; Spahni et al., 2005). The new
96 on-line record yields a mean depth resolution of 26 cm, compared to a mean depth resolution of 1.52
97 m of the methane record used for the synchronization with the EDC record by Buiron et al. (2011).
98 Even though the precision of the on-line measurements is lower (1σ of 15 – 20 ppbv) than the one of
99 the discrete measurements (1σ of 10 ppbv) and absolute calibration is an issue, the new dataset is

100 very well suited for a refined synchronization of the TALDICE and the EDC methane records. This
101 is due to the considerably higher depth resolution and the magnitude of atmospheric CH₄ variability
102 which is in the range of 350 ppb to 750 ppb during the last glacial period. This allows for the
103 definition of more tie points with better precision. Gaps in our high-resolution CH₄ record (see Fig.
104 3) longer than 1 m were caused either by several distinct ash layers in the ice core that were not
105 measured with CFA (3 m at 86.5–88.5 ka, 2 m at 107–109 ka and 4 m at 111.5–115 ka BP) or
106 maintenance of the GC system (12 m at 61–64 ka BP) while CFA measurements were continued.

107

108 A methane record covering the Antarctic Cold Reversal (ACR) was measured with the same method
109 on TALDICE and presented in Schüpbach et al. (2009). This record features a nominal resolution of
110 3–10 years. However, no large concentration variations of methane in the air trapped in ice are
111 possible within such short time periods due to the slow bubble close-off process (Schwander et al.,
112 1993). Therefore, the data were filtered by a binomial 5-point filter to smooth out artificial variations
113 induced by the measurement uncertainty without corrupting the signal over fast concentration
114 increases or decreases. Since these high-frequency variations are a measurement artefact, the filter is
115 not applied on a constant time window but always over five consecutive data points, i.e. on a
116 constant depth interval. This same filter was applied for all the high-resolution data presented in this
117 work featuring similar depth resolution but much lower resolution in time (mean temporal resolution
118 of 103 years) than the data covering the ACR. In doing so, measurement artefacts are filtered
119 reliably, but atmospheric CH₄ variations are potentially smoothed since the mean temporal resolution
120 of the CFA-CH₄ record presented in this study is in the order of magnitude of the bubble close-off
121 time. However, the potential smoothing of atmospheric variations does not have implications on the
122 result of the synchronization of the CH₄ records of TALDICE and EDC.

123

124 The filtered high-resolution methane record was then synchronized to the EDC methane data by
125 visually matching fast transitions in the two methane records. The tie points are chosen at mid-slope
126 of the transitions at the onset of Dansgaard-Oeschger (DO) events or at the maxima of very short
127 methane peaks (e.g. the very pronounced event at 58,600 ka BP preceding DO 17 in Fig. 1). Due to
128 the high resolution of the new Talos Dome methane record the uncertainty of the visual matching
129 remains lower than 300 years (compared to 400–1500 years in Buiron et al. 2011) in the discussed
130 depth interval. This uncertainty only depends on the resolution of both records and is calculated as
131 the square root of the sum of squares of the EDC and TALDICE time resolution at the respective tie
132 point (see Table 1). Not included in this uncertainty is the additional synchronization error caused by
133 different bubble close-off characteristics of Dome C and Talos Dome. This additional error is lower

134 than the 200 years calculated by Köhler et al. (2011). This is because Köhler et al. (2011) compare
135 the EDC and NGRIP ice cores where bubble close-off characteristics are very different. The close-
136 off characteristics of TALDICE is more similar to those of EDC, resulting in smaller synchronization
137 uncertainties. In order to obtain the uncertainty of the absolute gas age of each tie point, the
138 uncertainties caused by visual matching has to be added to the inherent uncertainty of the EDC3 age
139 scale of 1–4 ka BP in the interval discussed in this work (Parrenin et al., 2007).

140

141 As the TALDICE-1 age scale both for ice and gas is based on gas tie points only (for ages younger
142 than 141 ka BP), shifting the gas tie points has direct implications on the age of the ice. The age
143 difference between the gas and the ice at the same depth (Δ age) is largely dependent on the
144 accumulation rate. Since changes in the accumulation rate in the refined age scale caused by shifting
145 tie points do not exceed 16%, i.e. stay well within the uncertainty of the accumulation rate given by
146 Buiron et al. (2011) ($\pm 20\%$), we applied the modeled Δ age of TALDICE-1 to the refined gas age
147 scale in order to derive the age of the ice at the corresponding depth.

148

149 Soluble calcium (Ca^{2+}), a tracer for mineral dust input, was analyzed on the entire TALDICE with a
150 well-established CFA system used for the determination of aerosol constituents in ice cores
151 (Kaufmann et al., 2008). In the depth interval from 1220 m to 1323 m discussed here a continuous
152 high-resolution Ca^{2+} record was obtained except for a gap of four meters (1276–1280 m), where Ca^{2+}
153 data are not available. The nominal depth resolution of the continuous Ca^{2+} record is typically 1 cm
154 (Bigler et al., 2006), for the purpose of this study the high-resolution Ca^{2+} record is down-sampled to
155 a depth resolution of 50 cm to compare with the EDC Ca^{2+} record. The mean measurement error of
156 the Ca^{2+} concentration record is estimated to be less than 10% (Röthlisberger et al., 2000).

157

158 **3 Results and discussion**

159 Figure 1 shows the Dome C methane record in the time interval from 50–86 ka BP on the EDC3 age
160 scale (Louergue et al., 2007; Louergue et al., 2008) along with the discrete methane data on the
161 TALDICE-1 age scale (Buiron et al., 2011). With the new high-resolution methane data overlaid
162 (orange line) discrepancies between the two age scales appear which could not be unambiguously
163 detected with the discrete measurements only. For example at the onset of DO 17 preceded by a
164 distinct precursor event the TALDICE-1 gas age is biased 1000 years towards older ages. Replacing
165 the tie point at 59,800 years BP with a tie point at the peak of the precursor event (58,600 years BP
166 on the EDC3 age scale) and thus shifting TALDICE-1 approx. 1000 years towards younger ages

167 while keeping the tie point at 71,200 years BP fixed (onset of DO 19) stretches the data in a way that
168 an additional tie point at the onset of DO 18 becomes apparent (see Fig. 2).

169

170 The same procedure has been applied for the entire period from 55–112 ka BP shown in Fig. 3. By
171 matching all the fast transitions at the onsets of DO 16–24, precursor events or other distinct signals
172 in the two methane records 12 new tie points were defined in the age interval from 55–112 ka BP,
173 four of the tie points proposed by Buiron et al. (2011) were adopted unchanged (see Table 1 and Fig.
174 3). Correlation coefficients between EDC CH₄ and CFA-CH₄ on the TALDICE-1 age scale and on
175 the revised TALDICE-1a age scale have been calculated by linearly interpolating the CFA-CH₄
176 record to obtain concentration values in the high-resolution record at exactly the same age as the
177 EDC data points. For the 208 data points in the investigated interval correlation on the TALDICE-1
178 age scale is $r^2 = 0.68$ compared to $r^2 = 0.81$ on the revised age scale. The TALDICE CH₄ data on the
179 whole interval of the refined age scale is shown in Fig. 3 along with the EDC CH₄ data. This new
180 TALDICE-1a age scale is not meant to replace the TALDICE-1 age scale, it is rather a refinement of
181 this age scale in the above mentioned time interval. Figure 4 (A) shows how much the gas age scale
182 has been changed by the construction of the new age scale with respect to the original age scale over
183 the entire depth interval of the age scale refinement. Largest shifts (up to 1200 years) can be
184 observed in the period from 56–70 ka BP. In the older part the changes in the age scale are lower
185 than 400 years. Due to new high-resolution methane data tie points between the TALDICE and EDC
186 CH₄ records could be significantly constrained, yielding relative age uncertainties of 150–300 years
187 compared to 400–1500 years at the tie points in the TALDICE-1 age scale. The uncertainty of the ice
188 age derived from synchronized gas records depends mainly on the Δ age uncertainties of both ice
189 cores and of the uncertainty of the CH₄ match. Thus the better constrained gas tie points also reduce
190 the uncertainty of the ice age scale, leading to relative age uncertainties between TALDICE and EDC
191 of below 1100 years in the refined interval (except for the depth interval 1267–1290 m
192 corresponding to 60–65 ka BP, where uncertainties reach up to 1500 years due to missing high-
193 resolution CH₄ data) compared to maximum uncertainties of 1900 years in the same interval with the
194 TALDICE-1 age scale. The uncertainties of the original TALDICE-1 ice age scale are compared to
195 the new uncertainties of the TALDICE-1a ice age scale in the refined period in Fig. 4 (B). The new
196 uncertainty is estimated by error propagation with unchanged Δ age uncertainties in EDC and
197 TALDICE and the reduced new uncertainty from the gas tie points. While discussions of climate
198 dynamics at sub-millennial time scales were possible back to MIS 3.3 with the TALDICE-1 age
199 scale, the refined age scale allows for such discussions back to MIS 5.3.

200

201 For the first time the precursor event of DO 21 has clearly been detected in methane in an Antarctic
202 ice core (see Fig. 1). It has been measured in high-resolution and discussed before in the GISP2 ice
203 core by Grachev et al., (2007 and 2009) and, thus, has been independently verified by our
204 measurements. Also the rapid variations of methane in the NGRIP ice core over DO 16 and 17
205 discussed in detail by Huber et al. (2006) have not been measured before in such resolution in
206 Antarctica. The existence not only of fast transitions during DO events in methane in both Antarctic
207 and Greenland ice cores, but now also the availability of precursor-like events in the methane records
208 of both hemispheres allows for a discussion of the mechanisms at work at time scales of a few
209 hundred years. However, the EDC CH₄ record does not show all the short events in methane due to
210 limited depth resolution but also due to considerable smoothing of the gas records due to low
211 accumulation and temperature. In contrast the EDML CH₄ record (Capron et al., 2010; EPICA, 2006;
212 Schilt et al., 2010), which features good depth resolution in the discussed interval, shows the distinct
213 variations over DO 15–17, which allows for even more precise synchronization with the TALDICE
214 CH₄ record. Furthermore, no additional phasing uncertainty due to the bubble close-off
215 characteristics is induced between TALDICE and EDML, since accumulation and temperature at
216 both drilling sites are very similar. In Table 1 the corresponding tie points are also proposed for the
217 EDML ice core based on synchronization with the new TALDICE CH₄ record.

218
219 To demonstrate the impact of the refinement of the TALDICE-1 age scale not only in gas records but
220 also in the surrounding ice matrix the Ca²⁺ concentration records of the TALDICE and EDC ice
221 cores are compared on a selected interval from 54–80 ka BP. This represents an independent
222 quantification of the validity of our approach and the quality of the revised age scale.

223
224 Calcium in East Antarctic ice cores mainly originates from terrestrial dust from southern South
225 America during the last glacial period (Delmonte et al., 2008; Fischer et al., 2007). Changes in the
226 flux of Ca²⁺ during this period should be synchronous across East Antarctica and can be used to
227 synchronize ice core records from this region (Mulvaney et al., 2000). Thus, we compare the Ca²⁺
228 records to demonstrate the impact of the refined age scale on TALDICE. In Fig. 5 Ca²⁺
229 concentrations from EDC and TALDICE are shown on the time interval from 54–80 ka BP on the
230 EDC3 and the original TALDICE-1 age scales (A), and the refined TALDICE-1a age scale (B),
231 respectively. In general, Ca²⁺ concentrations are approximately three times lower in TALDICE than
232 the respective concentrations in the EDC ice core (note different scales of the ordinates in Figs. 4 A
233 and B) over the discussed interval. The relative variations of the two Ca²⁺ records show high
234 correlation as expected according to Mulvaney et al. (2000). However, the variations are

235 substantially shifted in time when using the original TALDICE-1 age scale (Buiron et al., 2011) (see
236 Fig. 5 A). Especially between 60 and 70 ka BP (covering DO 17–19) where highest Ca^{2+}
237 concentrations are reached a temporal shift towards older ages in the order of 1000 years becomes
238 apparent.

239

240 When using the refined TALDICE-1a age scale instead (Fig. 5 B) the variations in the Ca^{2+} records
241 are in phase within the error limits, confirming the improved consistency of the TALDICE-1a age
242 scale and the EDC3 age scale compared to the TALDICE-1 age scale. Correlation of the TALDICE
243 and EDC Ca^{2+} records (246 data points each) in this interval is increased from $r^2 = 0.71$ using
244 TALDICE-1 to $r^2 = 0.89$ when the refined TALDICE-1a age scale is applied. Thus, not only in the
245 interval where the largest corrections in the gas age scale have been applied (around DO 17, see Figs.
246 1 and 2), but also in other sections of the refined interval a substantial improvement of the
247 synchronization in both the gas and the ice age scale has been achieved by the use of the new high-
248 resolution methane data. The first methane tie point of the refined age scale is at 55,150 years BP
249 (see Table 1) in the gas age, corresponding to an age of the surrounding ice of 56,300 years BP.
250 Thus, the ice age scale is readjusted only for ages older than 56,3 ka BP as can be seen in Fig. 5 B.
251 Δ age modeled by Buiron et al. (2011) is slightly overestimated in the age interval 55–67 ka BP,
252 whereas for older ages it seems to fit well with the refined TALDICE-1a age scale.

253

254 **4 Conclusions**

255 The refined age scale TALDICE-1a for TALDICE presented in this work complements the
256 TALDICE-1 age scale in the age interval from 55–112 ka BP. This refinement is required for
257 investigations of climate dynamics at sub-millennial time scales not only back to 50 ka BP as with
258 the TALDICE-1 age scale but back to MIS 5.3 at 110 ka BP. In particular, precise north-south
259 synchronization is essential for the study of interhemispheric connections (Raisbeck et al., 2007;
260 Stocker and Johnsen, 2003). The availability of such high-resolution CH_4 data allows for more
261 precise synchronizations with future ice cores which are also analyzed with on-line CH_4
262 measurements. For the present purpose, absolute calibration of the CFA- CH_4 measurements is not
263 necessary. This greatly enhances the value of these data. Further improvements concerning the
264 precision of the on-line measurements would then also allow for a better insight in the dynamics of
265 the methane cycle on short time scales and at low concentration variations. With additional methane
266 measurements to achieve higher resolution in the lower part of TALDICE (ages older than 130 ka
267 BP) and using Ca^{2+} for tie points in the ice matrix the synchronization of TALDICE with EDC could

268 further be improved in the future through the entire length of the ice core by using e.g. the inverse
269 model by Lemieux-Dudon et al. (2010).

270

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277

278 **References**

- 279 Bigler, M., Röthlisberger, R., Lambert, F., Stocker, T. F., and Wagenbach, D.: Aerosol deposited in
280 East Antarctica over the last glacial cycle: Detailed apportionment of continental and sea-salt
281 contributions, *J. Geophys. Res.*, 111, D08205, doi:10.1029/2005JD006469, 2006.
- 282 Blunier, T., Chappellaz, J., Schwander, J., Dallenbach, A., Stauffer, B., Stocker, T. F., Raynaud, D.,
283 Jouzel, J., Clausen, H. B., Hammer, C. U., and Johnsen, S. J.: Asynchrony of Antarctic and
284 Greenland climate change during the last glacial period, *Nature*, 394, 739-743,
285 doi:10.1038/29447, 1998.
- 286 Blunier, T., and Brook, E. J.: Timing of Millennial-Scale Climate Change in Antarctica and
287 Greenland During the Last Glacial Period, *Science*, 291, 109-112, doi:
288 10.1126/science.291.5501.109, 2001.
- 289 Blunier, T., Spahni, R., Barnola, J. M., Chappellaz, J., Louergue, L., and Schwander, J.:
290 Synchronization of ice core records via atmospheric gases, *Clim. Past*, 3, 325-330, 2007.
- 291 Brook, E. J., Harder, S., Severinghaus, J. P., Steig, E. J., and Sucher, C. M.: On the Origin and
292 Timing of Rapid Changes in Atmospheric Methane During the Last Glacial Period, *Global*
293 *Biogeochem. Cy.*, 14, 559-572, doi:10.1029/1999GB001182, 2000.
- 294 Buiron, D., Chappellaz, J., Stenni, B., Frezzotti, M., Baumgartner, M., Capron, E., Landais, A.,
295 Lemieux-Dudon, B., Masson-Delmotte, V., Montagnat, M., Parrenin, F., and Schilt, A.:
296 TALDICE-1 age scale of the Talos Dome deep ice core, East Antarctica, *Clim. Past*, 7, 1-16,
297 doi:10.5194/cp-7-1-2011, 2011.
- 298 Capron, E., Landais, A., Lemieux-Dudon, B., Schilt, A., Masson-Delmotte, V., Buiron, D.,
299 Chappellaz, J., Dahl-Jensen, D., Johnsen, S., Leuenberger, M., Louergue, L., and Oerter, H.:
300 Synchronising EDML and NorthGRIP ice cores using $\delta^{18}\text{O}$ of atmospheric oxygen ($\delta^{18}\text{O}_{\text{atm}}$) and
301 CH_4 measurements over MIS5 (80-123 kyr), *Quaternary Sci. Rev.*, 29, 222-234,
302 doi:10.1016/j.quascirev.2009.07.014, 2010.
- 303 Chappellaz, J., Brook, E., Blunier, T., and Malaizé, B.: CH_4 and $\delta^{18}\text{O}$ of O_2 records from Antarctic
304 and Greenland ice: A clue for stratigraphic disturbance in the bottom part of the Greenland Ice
305 Core Project and the Greenland Ice Sheet Project 2 ice cores, *J. Geophys. Res.*, 102, 26547-
306 26557, 1997.
- 307 Delmonte, B., Andersson, P. S., Hansson, M., Schöberg, H., Petit, J. R., Basile-Doelsch, I., and
308 Maggi, V.: Aeolian dust in East Antarctica (EPICA-Dome C and Vostok): Provenance during

309 glacial ages over the last 800 kyr, *Geophys. Res. Lett.*, 35, L07703, doi:10.1029/2008GL033382,
310 2008.

311 EPICA, c. m.: One-to-one coupling of glacial climate variability in Greenland and Antarctica,
312 *Nature*, 444, 195-198, 2006.

313 Fischer, H., Siggaard-Andersen, M.-L., Ruth, U., Röthlisberger, R., and Wolff, E.:
314 Glacial/interglacial changes in mineral dust and sea-salt records in polar ice cores: Sources,
315 transport, and deposition, *Rev. Geophys.*, 45, RG1002, doi:10.1029/2005RG000192, 2007.

316 Goujon, C., Barnola, J. M., and Ritz, C.: Modeling the densification of polar firn including heat
317 diffusion: Application to close-off characteristics and gas isotopic fractionation for Antarctica
318 and Greenland sites, *J. Geophys. Res.*, 108, 4792, doi:10.1029/2002JD003319, 2003.

319 Grachev, A. M., Brook, E. J., and Severinghaus, J. P.: Abrupt changes in atmospheric methane at the
320 MIS 5b-5a transition, *Geophys. Res. Lett.*, 34, L20703, doi:10.1029/2007GL029799, 2007.

321 Grachev, A. M., Brook, E. J., Severinghaus, J. P., and Piasias, N. G.: Relative timing and variability
322 of atmospheric methane and GISP2 oxygen isotopes between 68 and 86 ka, *Global Biogeochem.*
323 *Cy.*, 23, GB2009, doi:10.1029/2008GB003330, 2009.

324 Huber, C., Leuenberger, M., Spahni, R., Flückiger, J., Schwander, J., Stocker, T. F., Johnsen, S.,
325 Landais, A., and Jouzel, J.: Isotope calibrated Greenland temperature record over Marine Isotope
326 Stage 3 and its relation to CH₄, *Earth Planet. Sc. Lett.*, 243, 504-519, 2006.

327 Kaufmann, P. R., Federer, U., Hutterli, M. A., Bigler, M., Schüpbach, S., Ruth, U., Schmitt, J., and
328 Stocker, T. F.: An Improved Continuous Flow Analysis System for High-Resolution Field
329 Measurements on Ice Cores, *Environ. Sci. Technol.*, 42, 8044-8050, doi:10.1021/es8007722,
330 2008.

331 Köhler, P., Knorr, G., Buiron, D., Lourantou, A., and Chappellaz, J.: Abrupt rise in atmospheric CO₂
332 at the onset of the Bølling/Allerød: in-situ ice core data versus true atmospheric signals, *Clim.*
333 *Past*, 7, 473-486, doi: 10.5194/cp-7-473-2011, 2011.

334 Landais, A., Barnola, J. M., Kawamura, K., Caillon, N., Delmotte, M., Van Ommen, T., Dreyfus, G.,
335 Jouzel, J., Masson-Delmotte, V., Minster, B., Freitag, J., Leuenberger, M., Schwander, J., Huber,
336 C., Etheridge, D., and Morgan, V.: Firn-air δ¹⁵N in modern polar sites and glacial-interglacial ice:
337 a model-data mismatch during glacial periods in Antarctica?, *Quaternary Sci. Rev.*, 25, 49-62,
338 2006.

339 Lelieveld, J., Crutzen, P. J., and Dentener, F. J.: Changing concentration, lifetime and climate forcing
340 of atmospheric methane, *Tellus B*, 50, 128-150, doi: 10.1034/j.1600-0889.1998.t01-1-00002.x,
341 1998.

342 Lemieux-Dudon, B., Blayo, E., Petit, J.-R., Waelbroeck, C., Svensson, A., Ritz, C., Barnola, J.-M.,
343 Narcisi, B. M., and Parrenin, F.: Consistent dating for Antarctic and Greenland ice cores,
344 *Quaternary Sci. Rev.*, 29, 8-20, doi: 10.1016/j.quascirev.2009.11.010, 2010.

345 Loulergue, L., Parrenin, F., Blunier, T., Barnola, J. M., Spahni, R., Schilt, A., Raisbeck, G., and
346 Chappellaz, J.: New constraints on the gas age-ice age difference along the EPICA ice cores, 0-
347 50 kyr, *Clim. Past*, 3, 527-540, 2007.

348 Loulergue, L., Schilt, A., Spahni, R., Masson-Delmotte, V., Blunier, T., Lemieux, B., Barnola, J.-M.,
349 Raynaud, D., Stocker, T. F., and Chappellaz, J.: Orbital and millennial-scale features of
350 atmospheric CH₄ over the past 800,000 years, *Nature*, 453, 383-386, 2008.

351 Mulvaney, R., Röthlisberger, R., Wolff, E. W., Sommer, S., Schwander, J., Hutterli, M. A., and
352 Jouzel, J.: The transition from the Last Glacial Period in inland and near-coastal Antarctica,
353 *Geophys. Res. Lett.*, 27, 2673-2676, doi:10.1029/1999GL011254, 2000.

354 Parrenin, F., Barnola, J. M., Beer, J., Blunier, T., Castellano, E., Chappellaz, J., Dreyfus, G., Fischer,
355 H., Fujita, S., Jouzel, J., Kawamura, K., Lemieux-Dudon, B., Loulergue, L., Masson-Delmotte,
356 V., Narcisi, B., Petit, J. R., Raisbeck, G., Raynaud, D., Ruth, U., Schwander, J., Severi, M.,
357 Spahni, R., Steffensen, J. P., Svensson, A., Udisti, R., Waelbroeck, C., and Wolff, E.: The EDC3
358 chronology for the EPICA Dome C ice core, *Clim. Past*, 3, 485-497, 2007.

359 Raisbeck, G. M., Yiou, F., Jouzel, J., and Stocker, T. F.: Direct north-south synchronization of abrupt
360 climate change record in ice cores using Beryllium 10, *Clim. Past*, 3, 541-547, 2007.

361 Röthlisberger, R., Bigler, M., Hutterli, M., Sommer, S., Stauffer, B., Junghans, H. G., and
362 Wagenbach, D.: Technique for Continuous High-Resolution Analysis of Trace Substances in Firn
363 and Ice Cores, *Environ. Sci. Technol.*, 34, 338-342, doi:10.1021/es9907055, 2000.

364 Schilt, A., Baumgartner, M., Blunier, T., Schwander, J., Spahni, R., Fischer, H., and Stocker, T. F.:
365 Glacial-interglacial and millennial-scale variations in the atmospheric nitrous oxide concentration
366 during the last 800,000 years, *Quaternary Sci. Rev.*, 29, 182-192,
367 doi:10.1016/j.quascirev.2009.03.011, 2010.

368 Schüpbach, S., Federer, U., Kaufmann, P. R., Hutterli, M. A., Buiron, D., Blunier, T., Fischer, H.,
369 and Stocker, T. F.: A New Method for High-Resolution Methane Measurements on Polar Ice
370 Cores Using Continuous Flow Analysis, *Environ. Sci. Technol.*, 43, 5371-5376,
371 doi:10.1021/es9003137, 2009.

372 Schwander, J., Barnola, J. M., Andrié, C., Leuenberger, M., Ludin, A., Raynaud, D., and Stauffer,
373 B.: The Age of the Air in the Firn and the Ice at Summit, Greenland, *J. Geophys. Res.*, 98, 2831-
374 2838, 1993.

375 Spahni, R., Schwander, J., Flückiger, J., Stauffer, B., Chappellaz, J., and Raynaud, D.: The
376 attenuation of fast atmospheric CH₄ variations recorded in polar ice cores, *Geophys. Res. Lett.*,
377 30, 1571, doi:10.1029/2003GL017093, 2003.

378 Spahni, R., Chappellaz, J., Stocker, T. F., Loulergue, L., Hausammann, G., Kawamura, K.,
379 Fluckiger, J., Schwander, J., Raynaud, D., Masson-Delmotte, V., and Jouzel, J.: Atmospheric
380 Methane and Nitrous Oxide of the Late Pleistocene from Antarctic Ice Cores, *Science*, 310, 1317-
381 1321, 10.1126/science.1120132, 2005.

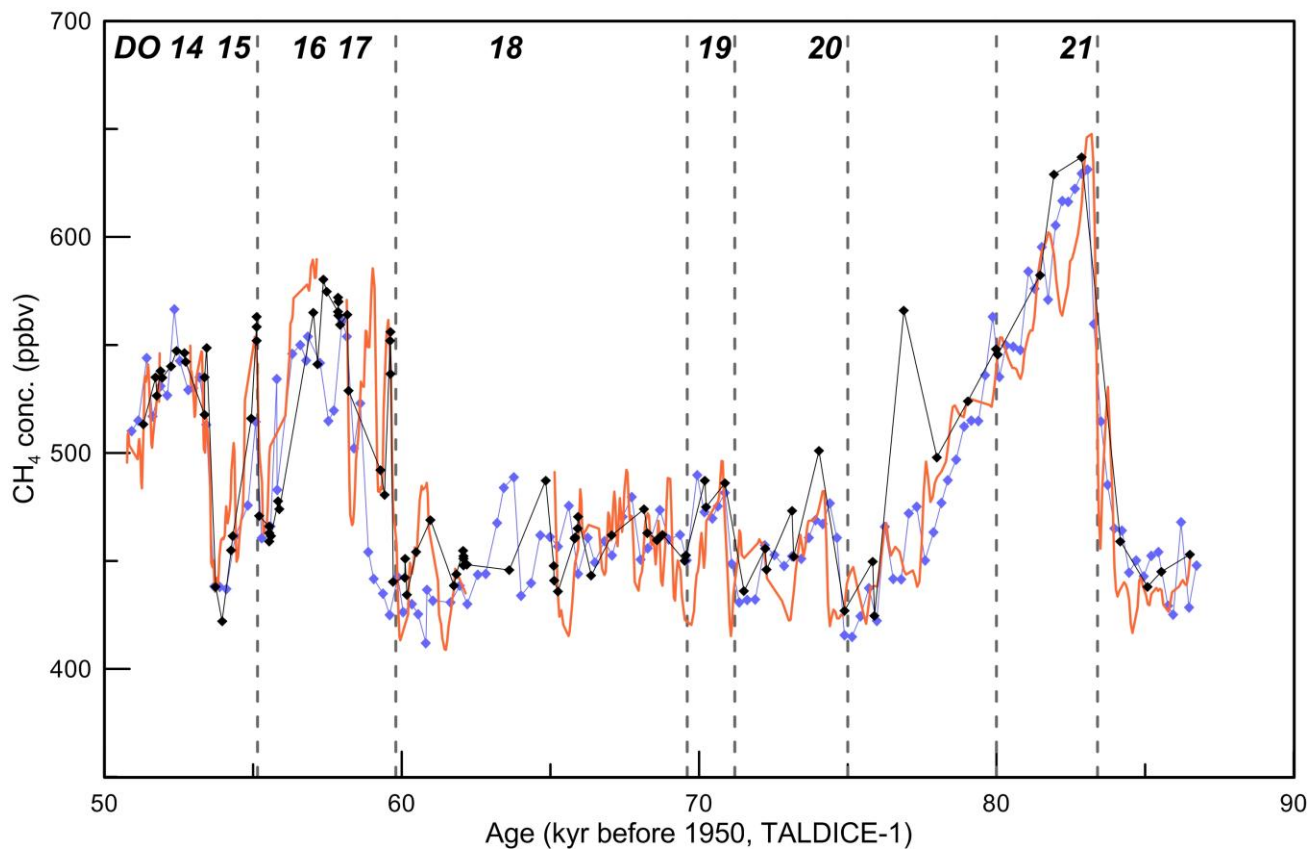
382 Stocker, T. F., and Johnsen, S. J.: A minimum thermodynamic model for the bipolar seesaw,
383 *Paleoceanography*, 18, 1087, doi:10.1029/2003PA000920, 2003.

384 Warneck, P.: *Chemistry of the Natural Atmosphere*, 2nd ed., Academic Press, 927 pp., 1988.

385
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387
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389
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391
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393
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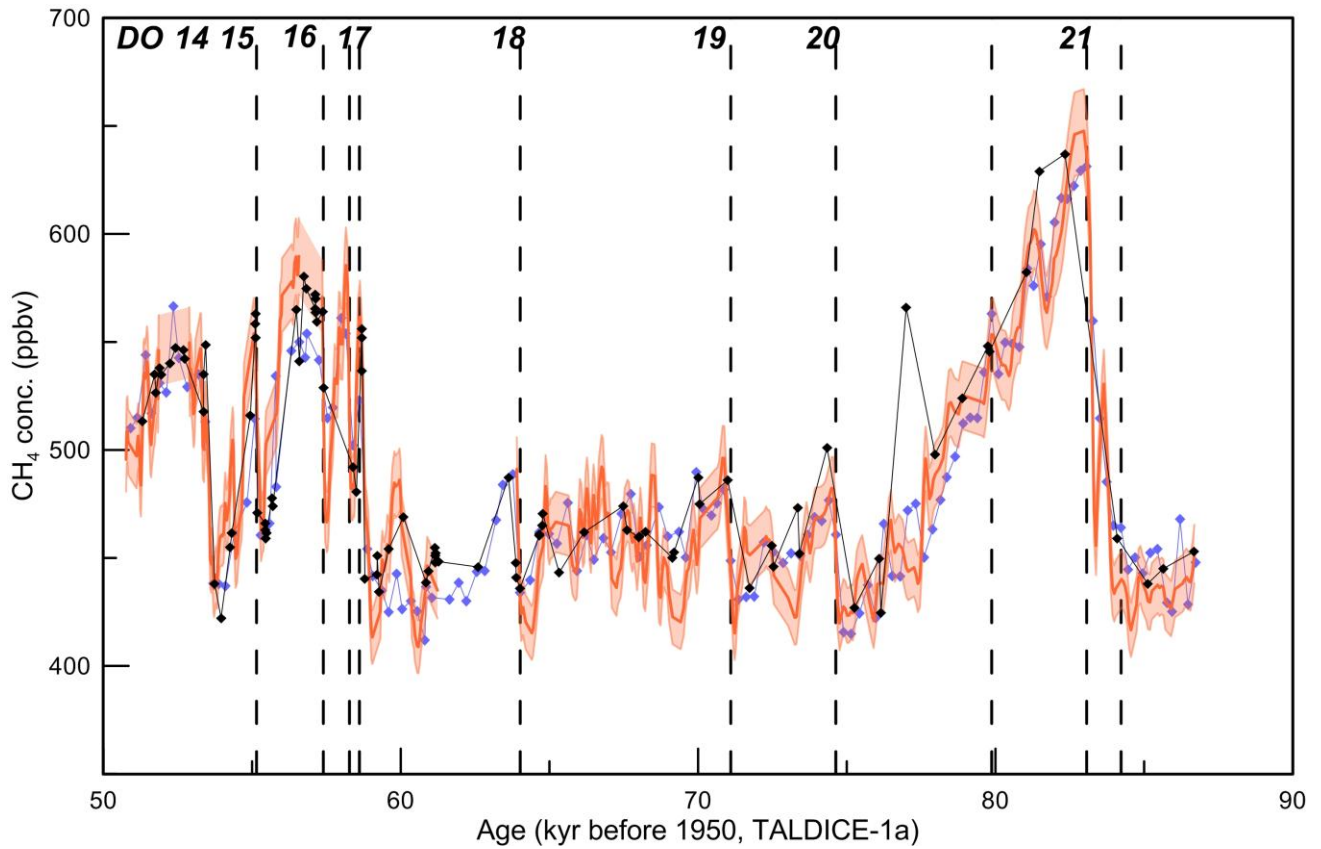
399 **Figures and Tables**

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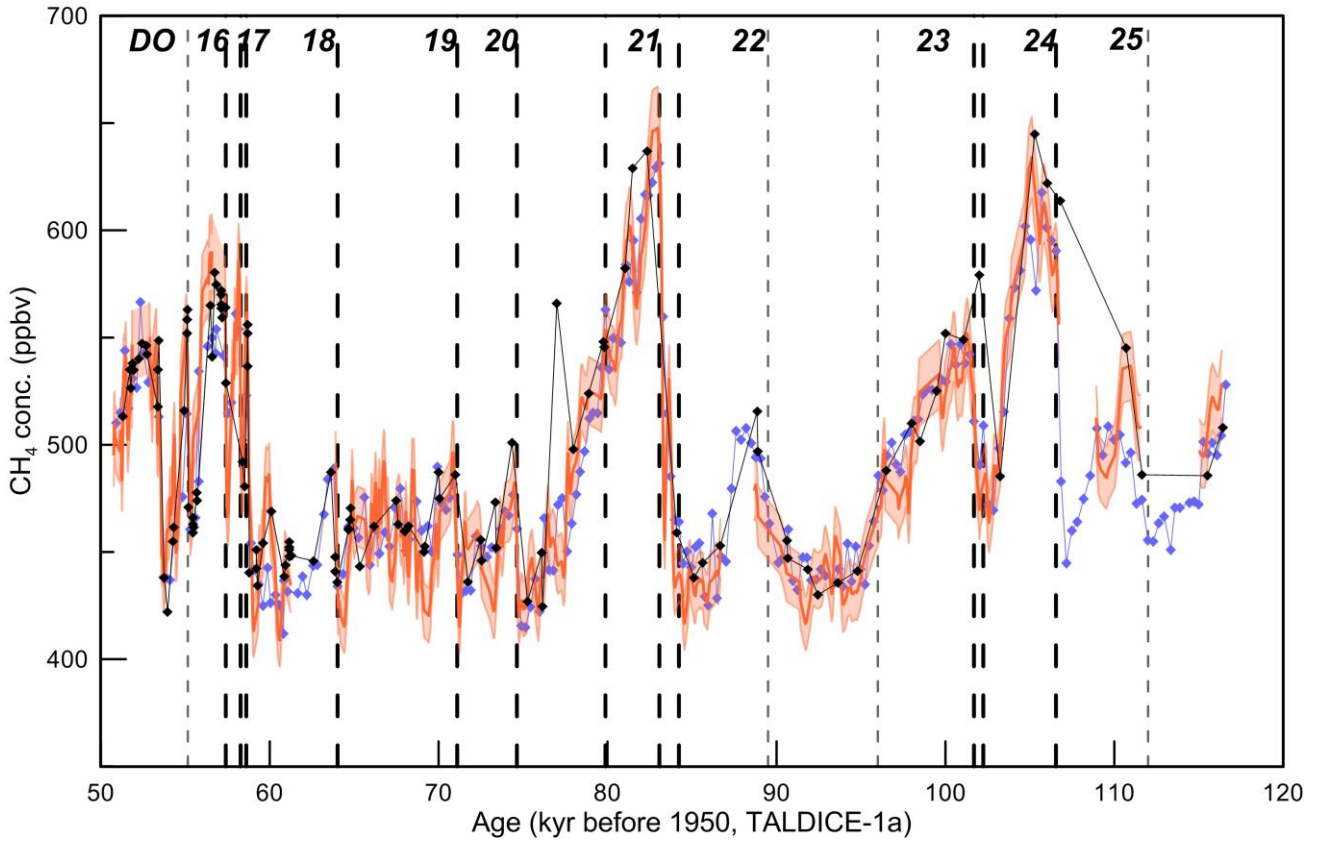
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402 **Fig. 1.** The EDC CH₄ record (blue diamonds) on the EDC3 age scale is compared to the TALDICE
403 CH₄ record (black diamonds, Buiron et al., 2011) on the TALDICE-1 age scale. The new high-
404 resolution CH₄ record (orange line) is also shown on the TALDICE-1 age scale. Dashed lines
405 indicate the tie points of the TALDICE-1 age scale used by Buiron et al. (2011). Bold italic numbers
406 indicate Dansgaard-Oeschger (DO) events.



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408 **Fig. 2.** The TALDICE CH₄ records (black diamonds: discrete data (Buiron et al., 2011); orange line:
 409 new high-resolution data with the light orange band indicating a $\pm 3\%$ error band) plotted on the
 410 refined TALDICE-1a age scale in comparison with the EDC CH₄ record (blue diamonds) on the
 411 EDC3 age scale. Dashed lines indicate new tie points of the TALDICE-1a age scale; bold italic
 412 numbers indicate Dansgaard-Oeschger (DO) events.



413

414 **Fig. 3.** The CH₄ records (EDC: blue diamonds, discrete TALDICE data: black diamonds, new high
 415 resolution TALDICE data: orange line) on the whole interval from 55–112 ka BP where the
 416 TALDICE-1 age scale has been refined. Bold dashed lines indicate the new tie point; fine dashed
 417 lines indicate tie points adopted from the TALDICE-1 age scale; bold italic numbers indicate
 418 Dansgaard-Oeschger (DO) events.

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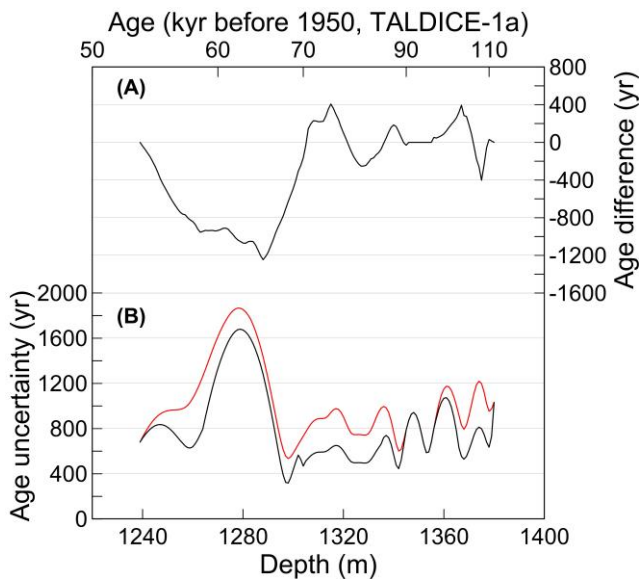
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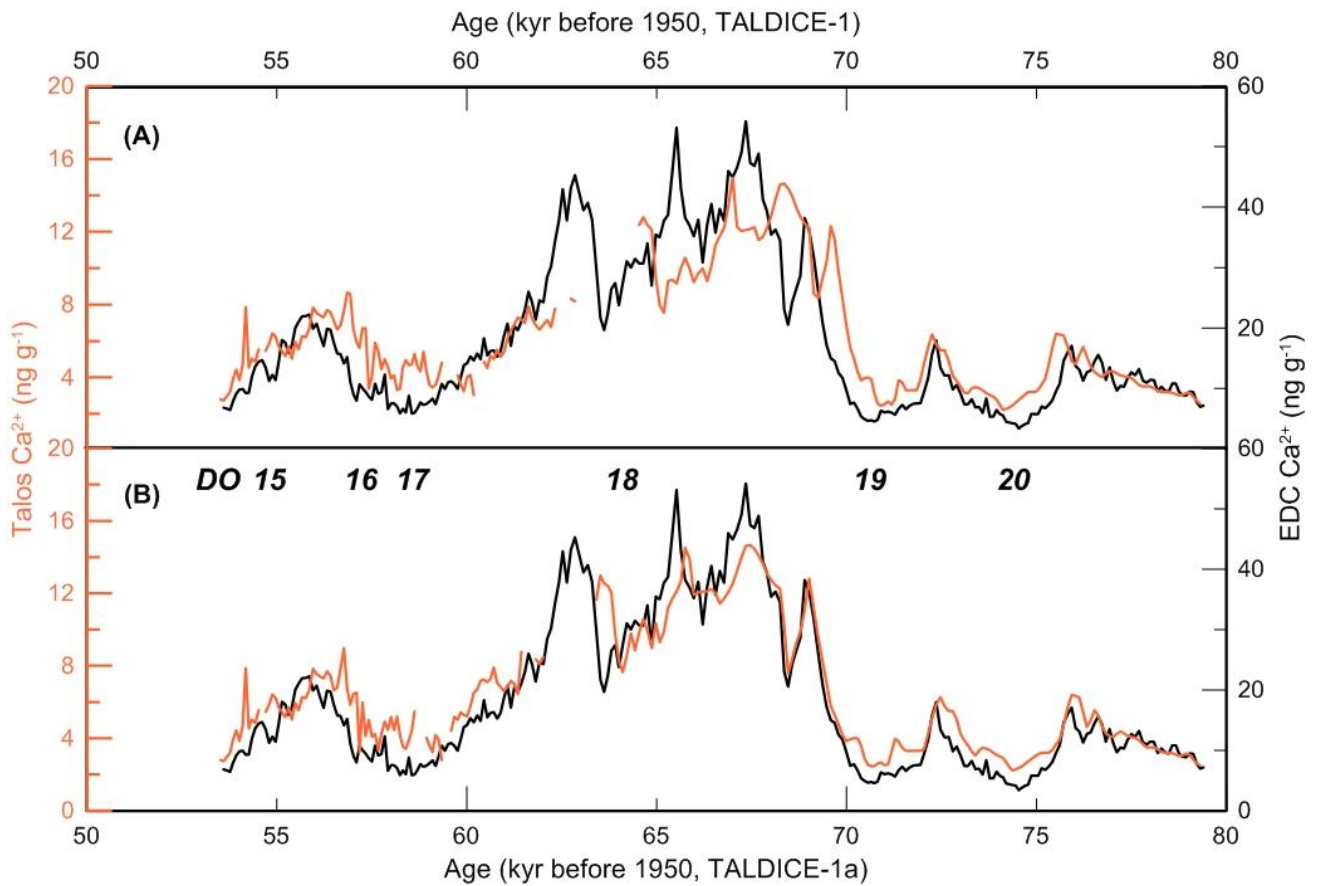
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Fig. 4. (A) The ice age differences of the refined TALDICE-1a age scale compared to TALDICE-1 (the age difference is defined as TALDICE-1a age subtracted by the TALDICE-1 age at the respective depth). (B) Age uncertainties of the original TALDICE-1 ice age scale (red line, Buiron et al., 2011) and the reduced uncertainties of the refined TALDICE-1a ice age scale (black line), respectively. The corresponding age (TALDICE-1a) is indicated on top.



439

440 **Fig. 5.** The Ca^{2+} records from EDC (black line, Bigler et al., 2006) on the EDC3 age scale and from
 441 TALDICE (orange line, new data) on the original TALDICE-1 age scale (A) and the refined
 442 TALDICE-1a age scale (B), respectively, are compared on the interval from 54–80 ka BP. EDC data
 443 are shown as 1 m averages, TALDICE data as 50 cm averages. TALDICE data are interpolated to fit
 444 the EDC data at the respective ages. Bold italic numbers indicate Dansgaard-Oeschger (DO) events.

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EDC depth (m)	TALDICE depth (m)	EDML depth (m)	Gas age EDC3 (yr BP)	Uncertainty (yr)	Comments
944.50	1239.00		55150	200	*
969.65	1255.50	1663.50	57400	220	
980.39	1260.50	1681.03	58280	160	onset DO 17
984.52	1262.67	1686.98	58610	190	precursor DO 17
1039.51	1287.75	1764.10	64020	250	onset DO 18
1105.55	1306.25	1862.50	71100	200	onset DO 19
1141.27	1314.57	1914.50	74630	200	onset DO 20
1196.27	1326.14	1978.10	79875	190	
1234.77	1332.75	2019.80	83070	200	peak DO 21
	1334	2425.20	83650		precursor DO 21
1248.52	1335.25	2031.20	84230	200	
1302.70	1345.00		89500	500	*
1369.3	1356		96000	500	*
1427.27	1367.1	2196	101690	230	onset DO 23
1432.77	1368.4	2199.32	102240	250	
1471.27	1374.75	2228.99	106550	280	onset DO 24
1515.4	1380.00		112000	1000	*

457 *: Tie point adopted from Buiron et al., (2011)

458 **Table 1.** Tie points defined in the age interval from 55–112 ka BP by synchronization of the new
459 high-resolution TALDICE CH₄ data with the EDC CH₄ record on the EDC3 age scale. The indicated
460 uncertainty is from visual matching of TALDICE and EDC only. Additionally indicated are the
461 corresponding depths of the EDML ice core for all the new tie points.