

The knowledge of the natural climate variability has been deepened, and the availability of longer time series of relevant variables contributes to a more accurate understanding of the anthropogenic effects on climate.

Different authors have synthesised the types of sources used to reconstruct past climates from documentary evidence (diaries, memoir books, daily weather reports, ship logbooks, administrative and ecclesiastic archives, among others), from instrumental data and from natural proxies, such as tree-rings, corals, speleothems, boreholes, ice-cores (Brázdil et al., 2005, 2010a). The lack of sufficient documentary information has hampered studies dealing with the reconstruction of the climate of Portugal during the 17th and 18th centuries (e.g. Luterbacher et al., 2006; Camuffo et al., 2010 for a review). Therefore the relevance of all the 18th century meteorological observations is clear. In this paper instrumental data obtained by Portuguese and foreign observers for Continental Portugal, Madeira and Rio de Janeiro (former capital of the Portuguese colony of Brazil) will be dealt with.

The Enlightenment period was favourable to the development of scientific initiatives; the first observations begun in Europe, in the late 16th and in the 17th centuries (Table 1), that is some time after the invention and development of the meteorological instruments. In the beginning of the 18th century, meteorology had much in common with both natural philosophy and applied mathematics; it overlapped these scientific branches and contributed to their development and to the progressive quantification of “natural philosophy”. The meteorologists made essential contributions to the design of physical instruments, including the so-called meteorological instruments used to observe weather.

Several European scientific academies coordinated the first observers’ work and encouraged systematic measurement of different weather elements. Unlike other European countries, meteorological observations at the end of the 18th century in Portugal were not performed at standard hours, neither with well-calibrated instruments (Taborda et al., 2004; Trigo et al., 2009). In fact, at this stage such observations were already standard procedure in more developed regions of Europe, such as Scandinavia

3401

(Bergstrom and Moberg, 2002), Northern Italy (Camuffo, 2002a,b; Maugeri et al., 2002) and France and UK (Jones et al., 1999; Slonosky et al., 2001; Slonosky, 2002).

However, climate analysis was not a common procedure. “Climatology [...] existed in neither deed nor word in the early 18th century” (Feldman, 1990). The term would be used more widely only in the early 19th century, with the meaning of compiling meteorological data, and performing temporal and spatial analysis. Synoptical interpretation and simple forecast would be coined later, with increasing use throughout the 20th century.

Most of the early instrumental records cover short periods, although a few meteorological sites are still active. A detailed description may be found in Kington (1988), Camuffo (2002a) and Brazdil et al. (2005, 2010b). There was a great progress in meteorological observations until the 1790s. Around the 1790s several political events and war incidents, including the French revolution in 1789, and the following Napoleonic wars disrupted a great number of networks. In any case at the end of the 18th century these activities had made an impact on the Iberian scientists, thanks to the relative numerous contacts with foreign scientific academies, and the support of the state in the enlightened spirit of the time (Barriendos et al., 2000). The scientists’ main areas of expertise were grouped in accordance with their foreign counterparts (Medicine, Physics, Mathematics, and Astronomy).

The main objectives of the present work are: (1) to disclose the labour of the first meteorologists working in Portugal (both Portuguese and foreigners), (2) to synthesise the existent early Portuguese instrumental meteorological records of the 18th century including data and metadata and to show their potentials for historical climatic studies and, (3) to demonstrate how the relatively short datasets obtained can be used to expand our knowledge on specific climatic events (e.g. the Laki eruption in 1783/1784) or long-term variability (e.g. NAO and precipitation).

In the following section, an overview of meteorological observations in Europe during the 18th century is carried out and it is intended to introduce the meteorologists working in Portugal and overseas within the context of the scientific community

3402

passes through a window where it pours the water into a lead cylindrical pot which has the same diameter as the mouth of the funnel. A body floats in this pot and a stick in the centre passes through the cover of the pot. It marks the elevation of the water on a ruler graduated according to the divisions of the *French king foot* (Velho, 1797a, p.450). Velho indicates very clearly that the thermometer is constantly exposed outside. The position of the instruments outside was advised at the end of the 18th century to make the data more easily comparable and more representative of the area (Kington, 1988). In the following years Velho changed for another barometer that he also describes in detail (Velho, 1797b, p.475). As no other indication concerning instrument type is included, we may infer that no changes occurred in instrument type after 1784.

Data from Lamego, Lisbon (Lisboa) and Mafra provide sufficient information to characterize the rainfall of the 1780s and to confirm Kington's statement (1988, p.2) that the 1780s contain several positive and negative rainfall extremes. Like in other Mediterranean regions (Font Tullot, 1988), an extremely rainy period occurred between 1784 and 1789, following and followed by dry periods, as can be observed in Fig. 4, although no conclusions may be drawn relatively to absolute rainfall differences (owing to discrepancies in measurement methodologies and instruments). Data from Lisbon and Lamego, as well as documentary data confirm the occurrence of a drought period in Portugal in 1779, and the corresponding tragic influences on agriculture and water availability (Taborda et al., 2004). This can be related to the persistence of a blocking anticyclone on Central Europe, as confirmed by Luterbacher's sea level pressure (SLP) field reconstructions (Luterbacher et al., 2002a). In this regard, the anticyclone reconstructed for March 1779 would generate an Easterly flux over Iberia (Fig. 5 left).

The three following years were also rather dry (638.4 mm in 1781 and 554.9 mm in 1782 in Lisbon) with the exception of April 1782 when 162.4 mm were measured by Pretorius. Frequent low pressure centres located SW from the British Isles, as one can infer from SLP reconstructions (Fig. 5 right, Luterbacher et al., 2002a) would have then caused heavy precipitation in Lisbon.

3413

1783 seems to be a turning point, as rainfall increases in Lisbon and Lamego, but in 1784 the increment continues (Fig. 4). Pretorius (1785, p.271) writes: "The year of 1784 is the wettest on record in Lisbon since 1777 [...]". In Lisbon, 906.8 mm of rain were measured in 1784 (162.4 mm more than in the preceding year) and 1221.4 mm were recorded in Mafra (480 mm more than in 1783). Documentary sources mention that *pro-serenitate* rogations (Alcoforado et al., 2000) took place in 1784: "If this country has been exempted of intense cold, (...) as is occurring in the rest of Europe, rain has been so continuously strong that the Cardinal Patriarch of Lisbon ordered public rogations in every church, to obtain serenity of the sky" (Newspaper *Gazeta de Lisboa*, no. 11, 16 March 1784). December 1784 is the wettest month of the Mafra series (312.5 mm) and was also very rainy in Lisbon (239.1 mm) (Fig. 6). Precipitation intensity was the highest at the end of the month (55.1 mm on the 29th, Fig. 7). Taborda et al. (2004) shows that the interpretation is in agreement with the daily synoptic charts reconstruction for Europe by Kington (1988), presenting an extensive low pressure centre influencing the Iberian Peninsula during the last week of December 1784 (Fig. 8).

Brázdil et al. (2010a) refer the devastating floods that took place in Central Europe in winter and spring 1783–1784. Most of them were due to intense snow melt and heavy rainfall. However, the floods in Southern France in December 1783 were mostly caused by sustained rains (Brázdil et al., 2010a). According to our results, December 1783 was also wet in Lisbon and in Mafra, but the precipitation was lower than in December 1784 (Figs. 6 and 7).

The following years of 1785–1787 can be considered as the continuation of the rainy period (Figs. 4 and 6). Although 1785 is slightly drier than 1784, total rainfall increased once more in the following years. In Mafra, the 1786 annual value attains 1429.7 mm. From Schulze writings one may infer that precipitation was also in excess of 1000 mm in Lisbon in 1786 (Schulze, 1790, fl.3 and Taborda et al., 2004). Although one should be careful while comparing with current values, those are certainly very high precipitation annual totals: Mafra/Tapada precipitation more recent average refers to the period

3414

1941–1960 and accounts 751.6 mm. In Lisboa/Geofísico, the main Lisbon station, the rainfall average for the period 1941–1970 reaches 714.4 mm. Moreover, there are further reports of *pro serenitate* rogations in March 1786 in Lisbon (Newspaper Gazeta de Lisboa, no. 12, 24 March). After 1789, precipitation decreases once more and in 1791 there are documents referring to the lack of water for agriculture in Southern Portugal (Taborda et al., 2004).

Although Velho infers wind direction with a crude method (only based on the movement of the clouds and smokes), it is interesting to combine his precipitation and wind data (Fig. 9): the highest amounts of rain occur with compass wind direction predominantly from W, SW and S winds, as Veiga had also noted for Lamego and which is similar to modern climatological analysis (Trigo and Dacamara, 2000; Fragoso and Tildes, 2008).

3.6 Lopes (Oporto, 1792)

José Bento Lopes (Lopes, hereafter) was a physician, graduated by the University of Coimbra, concerned about public health issues. His readings were performed in 1792 and took place twice a day; in the morning, between 7 and 8 h, and in the afternoon, between 15 and 16 h. Data on air pressure are expressed in English inches while the thermometer had a Fahrenheit scale. Lopes used a hygrometer with 60 subdivisions (30 to measure the degree of dryness and 30 to measure the degree of moisture). The meteorological and medical data were published in volume I of a medical journal (Anno Medico, 1796).

3.7 Dorta (Rio de Janeiro, 1781–1788)

Bento Sanches Dorta (BSD hereafter, Coimbra, Portugal, 1739 – São Paulo, Brazil, 1794) was an astronomer and geographer sent to Brazil by the King of Portugal in charge of a geographical mission, with the main task of determining the actual limits

3415

of the Portuguese colony (Carvalho, 1982). He was a member of the LRAS having published numerous papers in its *Memoirs* (Dorta, 1797, 1799a,b; 1812a,b).

The sheer quantity and apparent quality of his wide range scientific observations have been used recently in different publications by some of us. Thus Vaquero et al. (2005) have showed the added value of the sunspot observation undertaken by BSD during the solar eclipse that took place on 9 February 1785. Based on the outstanding number of geomagnetic declination measurements observations performed by BSD, Vaquero and Trigo (2005b) analyzed the instruments and measure methods. Unusual observations of low latitude auroras australis obtained by BSD between 1781 and 1788 were linked with the geomagnetic declination values carefully registered in his papers and the solar storms that occurred (Vaquero and Trigo, 2006). Furthermore, the observations of BSD and the out of the ordinary weather conditions, such as abnormal mist (haze) and dry fog that happened in the years following the Laki eruption are dealt with in Trigo et al. (2010). Only very recently have we looked at the meteorological observations of BSD to assess the late 18th century climate of Rio de Janeiro (Farrona et al., 2011).

In the initial collection of instruments (thermometer, pluviometer, and evaporation pan) there was not a barometer that only arrived from Europe in 1784. The instruments were located “in the camera of my room”, located 50 spans and 4 inches above the sea level. This camera had three opened windows toward the southwest (Dorta, 1797b, p.346) and was situated in the castle hill of the city.

BSD measured other important weather parameters such as rainfall and evaporation and computed the monthly mean value of all these variables. Finally, he registered the wind direction in an 8-directional compass from which he derived the monthly mean wind direction in the morning and in the afternoon (Dorta, 1812a, p.76).

BSD was a keen observer performing seven times per day, from 06:00 to 18:00 LT every two hours during the period 1781–1783, raising to eight times per day after the arrival of the barometer in August 1784 until December 1788 (Dorta, 1799a, p.346). Based on these sub-daily observations he computed monthly averages namely, mean

3416

temperature and pressure in the morning, at midday and in the afternoon. These values were calculated as the arithmetic mean of all the observations made. In the diary he also wrote the maximum and minimum temperature, however he admitted he did not know if these values were correctly measured because, as mentioned above, he only checked the thermometer at some pre-defined hours throughout the day (Dorta, 1812a, p.74).

Additionally to these quantitative observations BSD was interested to gather relevant qualitative information describing the state of the sky. As he said, “The four qualities of the day which shape the years, namely: clear, variable, cloudy, overcast; there should add 365” (Dorta, 1812b, p.115). He computed the monthly mean of number of days of thunder, rain, fog, aurora australis, zodiacal light, and clear, variable and cloudy days. In fact, at the end of each year BSD would summarize all weather patterns observed during that year, where he described the extreme and/or interesting events and he specified the date on which they happened.

3.8 Murdock (Madeira, 1793–1802)

William Gourlay, Fellow of the Royal College of Physicians of Edinburgh and physician of the British Factory at Madeira, published in 1811 a treatise on the Natural History of Madeira containing monthly accounts of weather from 1793 to 1802. For each month, data on monthly temperature and pressure (maximum, minimum and mean) and a description of weather are available (Gourlay, 1811, 39–66 pp.).

According to Barral (1854), the observer was James Murdock and the place of observation was named “Sitio do Vale” [Place of the valley] (400 feet above the sea level). Barral (1854) indicates that there are no metadata for this observer and he classified this record as “doubly”. While temperature time series present a reliable seasonal range and inter-annual variability, the same cannot be said in relation to the pressure data. In fact, pressure does not present the usual inter-annual variability and the number of consecutive similar values is very high.

3417

Figure 10 shows the temperature record after some minor typographical errors corrected. The extreme monthly temperature recorded were 28.9 °C (September 1802) and 10.6 °C (January 1802). The yearly extreme temperature recorded were 19.1 °C and 19.0 °C (years 1798 and 1799, respectively) and 17.5 °C (1800). Murdock includes in his observations interesting comments about the state of weather and its relation with agriculture, especially on the plantations at different altitudes.

4 Some examples

4.1 Laki eruption impact in 1783

The year 1783 deserves particular attention due to “foggy weather” and the “almost permanent fog during day and night”, which dominated the state of weather for a considerable number of days during June and July. According to Pretorius (1785, 270 and 272 pp.), the extract of meteorological observations on the 1783 and 1784 of this event gives the following testimony: “But what makes this year [1783] more notable among many in the past was the summer haze. Between June 22 and July 6 we had 14 days with a permanent haze throughout the day and night, and shortly thereafter this was succeeded by the same phenomena during 8 days, from July 12 to 20” and it should be noted that the same kind of fog and haze dominated over our boreal hemisphere.

It is now widely accepted among the scientific community that the main culprit for this unusual weather was the Icelandic Laki eruption, that took place between June 1783 and February 1784. The amount of gases and aerosols released by Laki were so large that provoked widespread impacts in Europe (Thordarson and Self, 2003). These impacts include increase mortality in the UK (Witham and Oppenheimer, 2005) and Continental Europe (Grattan et al., 2005), but also the devastation of pastures and livestock in Iceland (Steingrimsson, 1998), and to a less extent in the British Isles, France, and Benelux. The large-scale circulation pattern prevalent at the time advected the plume of smoke and haze towards Northern Europe (Stothers, 1999; Démarée et al.,

3418

1998) but affecting the entire European continent from Lisbon to Moscow (Thordarson and Self, 2003).

The year of 1783 is relatively well covered by Portuguese observers being two based in Continental Portugal (Pretorius in Lisbon and Velho in Mafra, Taborda et al., 2004) but also with important complementary observations being made in Rio de Janeiro by BSD (Fig. 11).

In fact similarly to Pretorius, Velho also recorded foggy and dry fog (haze) during periods that the German refers, in particular between 26 June and 2 July and between 13 and 20 July. However, Velho makes further references to fog and misty sky during the months that span between June and August (Taborda et al., 2004). During these three months, 67.2% of allusions to fog and cloudy sky are associated with winds from the north and northwest winds and 81.3% to winds from the north, northwest and west (Taborda et al., 2004).

Perhaps more unexpected is the recent description of possible association between unusual dry fogs over Rio de Janeiro and the Laki eruption, although further work must be done to evaluate the robustness of such link (Trigo et al., 2010). The progress of the number of foggy days per month witnessed by BSD between 1781 and 1788 can be seen in Fig. 11.

The summer of 1783 for both sites with observations in Portugal (Lisbon and Mafra) were characterized by relatively mild (even slightly colder) conditions, possibly due to aerosol concentration in the atmosphere that would decrease summer maximum temperatures; on the contrary most of Western Europe experienced particularly hot weather. As an example, we are presenting average summer temperatures in Mafra: in 1783 and 1784 they are circa 2°C inferior to those of the two following years (Table 2). Thus, in Central England, the summer of 1783 and, in particular the month of July, is considered one of the hottest the last three centuries (Kington, 1980, 1988), as well as in Southern Moravia (Brazdil et al., 2003) and in Germany (Jacobeit et al., 1997). The most probable physical mechanism responsible for such high temperatures is the existence of persistent blocking patterns that dominated the atmospheric circulation for

3419

several months and are known to induce warm and dry conditions in Western Europe (Trigo et al., 2004). In fact, the early summer months of 1783 were dominated by a strong anticyclone over Europe, that promoted the subsiding of air including the dry sulphurous fog towards the surface (see Figs. 6 and 7 in Thordarson and Self, 2003).

5 4.2 Correlation between the precipitation and the NAO index between 1780 and 1793

The North Atlantic Oscillation (NAO) is relatively well known since the early works of Walker (1924) and represents the most important large-scale pattern of circulation in the Northern Hemisphere (Wallace and Gutzler, 1981). In a nutshell, the NAO mode corresponds to a large-scale meridional alternation of atmospheric mass between the polar low pressure system closes to Iceland and the subtropical anticyclone near the Azores (Hurrell, 1995; Trigo et al., 2002). In the last two decades, a growing number of research papers have examined the relationship between winter precipitation in Europe and the NAO index and, especially, for the Iberian Peninsula (Hurrell, 1995; Trigo et al., 2004; Gallego et al., 2005). In particular it has been shown that extreme events such as the outstanding drought of 2005 (Garcia-Herrera et al., 2007) and record winter precipitation of 2009–2010 (Vicente-Serrano et al., 2011) are closely related to prolonged episodes of positive and negative values of the NAO index, respectively. Besides the long-term precipitation dataset available for Lisbon since 1865 we now have two additional small datasets of rainfall data for the Lisbon region; Lisbon and Mafra (see Sects. 3.3 and 3.5) that we can use to confirm the relationship between winter rainfall in Lisbon, during the 1780s, and the existing reconstructions of the NAO index (Cook et al., 2002; Luterbacher et al., 2002b).

Historical precipitation data recovered shows higher values compared with the values of the more recent period (1865–2010). For example, the accumulated values of winter precipitation (DJFM) in Lisbon during the winters of 1783–1784 and 1784–1785 recorded by Pretorius are 555 and 582 mm, respectively. These values correspond to the 92 and 89 percentile values of winter precipitation in the official series of Lisbon

3420

during the period 1865–2010. Moreover, the NAO index reconstructions of Luterbacher et al. (2002b) and Cook et al. (2002) show negative or near zero values in the years that have historical records of precipitation in Lisbon and Mafra. Therefore, both data sets are compatible.

5 The relationship between the average monthly precipitation for winter (DJFM) and the winter NAO index (DJFM) is shown in Fig. 12. The black dots represent the values of the instrumental period (1865–2000). We used the historical rainfall series of Lisbon employed in recent assessment of drought and wet events over Western Iberia (Garcia-Herrera et al., 2007; Vicente-Serrano et al., 2011) and the NAO index values
10 provided by Jones et al. (1999). Squares and circles represent the historical values of precipitation of Mafra (1783–1787) and Lisbon (1783–1785, 1789 and 1793), respectively while colour indicates the origin of the reconstructed NAO index time series: red (Cook et al., 2002) and blue (Luterbacher et al., 2002b).

In spite of the small set of precipitation data available for Lisbon and Mafra in the late
15 18th century, it shows a similar behaviour to that provided by the long-term dataset for the modern period (1865–2010). In particular we can note that the decade of 1780s was wetter than the long-term average in the Lisbon region and that the corresponding reconstructed NAO values were predominately negative.

We would like to stress the values observed for the winter of 1783–1784. According
20 to several authors this winter was wetter than the average over large sectors of Europe (e.g. D'Arrigo et al., 2011) and may be influenced by a number of different physical mechanisms. In fact while some authors stress the influence of external driving mechanisms such as the recent eruption of Laki in June 1783 (Thordarson and Self, 2003), or the deep minimum of 11-yr solar cycle around 1784 (Vaquero, 2004), others point
25 to large variability of internal modes such as a combined negative phase of the North Atlantic Oscillation and an El Niño-Southern Oscillation (ENSO) warm events.

3421

5 Conclusions

European wide temperature and precipitation reconstructions clearly point to the fact that more information is still needed for a better understanding of climate change in Europe. Therefore, any new data from Portugal may help bridging existing spatial and
5 temporal gaps in coverage of Southwest Europe.

The early instrumental data from Portugal and Brazil that were presented in this paper are the first short series available in a period for which climate reconstruction has been mostly based on documentary data (Taborda et al., 2004; Trigo et al., 2009). Systematic observations of different weather elements were encouraged by scientific
10 academies and data quality is sometimes very good. Given the novel and innovative nature of the information, some scientists (particularly Velho and Pretorius) were concerned with legitimizing their results through the descriptions concerning instrument exposure, type and location, as well as instrument site. Pretorius and Velho were among the first in Iberia to take their measurements outside (Barriendos et al., 2000).
15 However, most data values may not be directly compared to current ones in absolute terms due to the following facts: (i) instruments were not always properly calibrated; (ii) scales were not normalized; (iii) exact station location and site are often unknown; (iv) instruments (e.g. thermometers) were not shielded (temperatures are often higher than nowadays); or (v) no metadata is available. Although units used by the different
20 observers (or the same observer) are quite different, it has been possible to convert them into SI units.

Similarly to other European countries early Portuguese meteorologists had to face serious difficulties to carry on regular and reliable observations: (i) the instruments were not easy to get; (ii) as they were not shielded, they were frequently damaged by
25 storms; (iii) there were no explicit and general measuring rules; (iv) most of the observers were not trained (lack of precision in readings, lack of discipline in observation time).

3422

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3425

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3426

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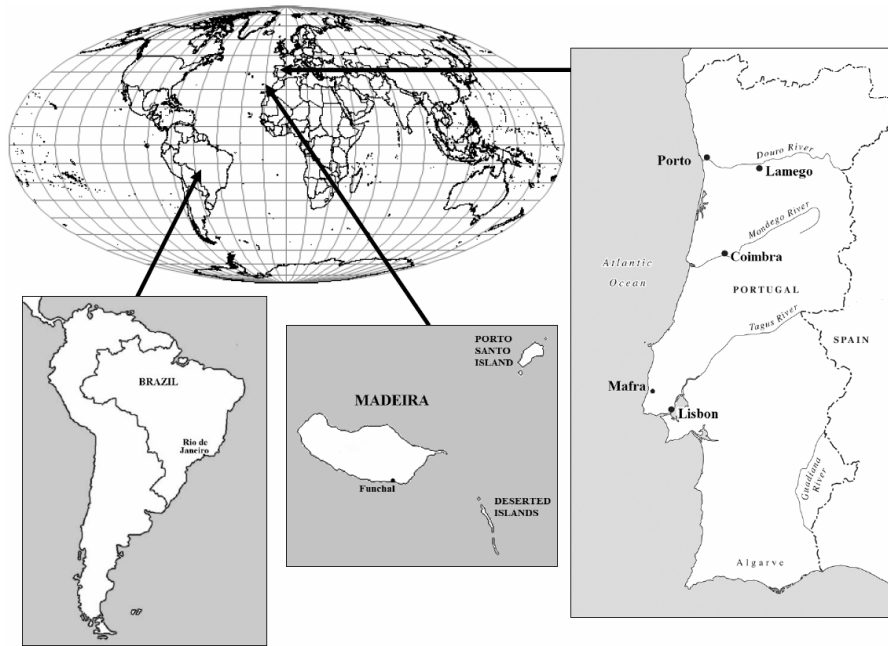


Fig. 1. Location maps.

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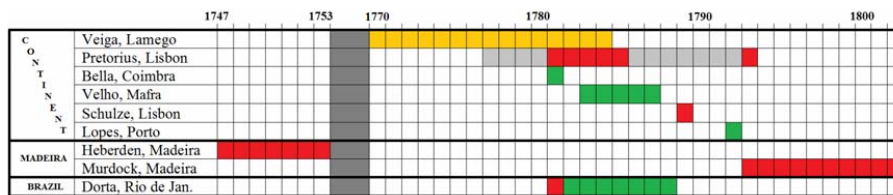


Fig. 2. Meteorological data availability in Portugal and overseas during the 18th century (yellow: annual data; red: monthly data; green: daily data; grey: lost data).

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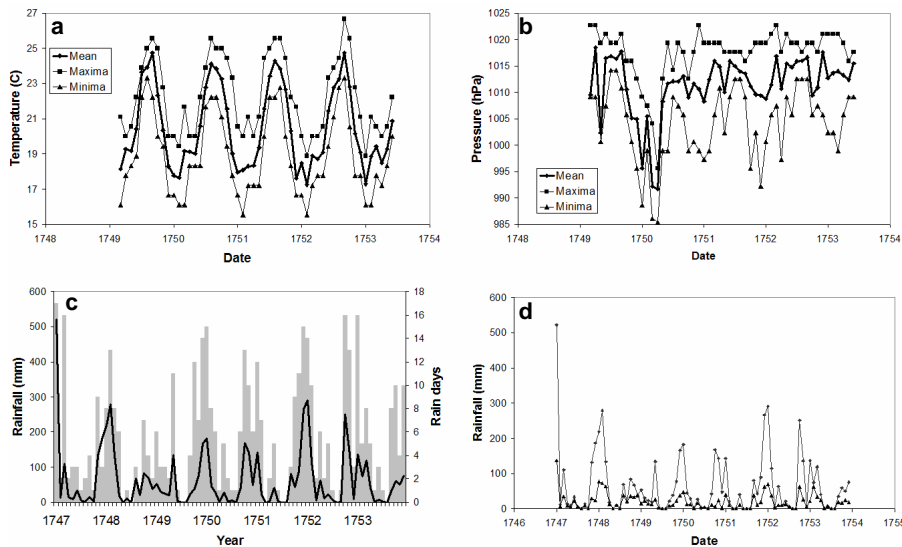


Fig. 3. Meteorological parameters recorded in Funchal by Heberden from March 1749 to June 1753 (temperature and pressure) and from January 1747 to December 1753 (rainfall). **(a)** Monthly temperature values **(b)** Monthly pressure values. **(c)** Monthly rainfall (dark line) and number of rain days (grey bars) **(d)** Monthly rainfall (grey line) and largest daily amount of water, which fell in any day of the respective month (dark line).

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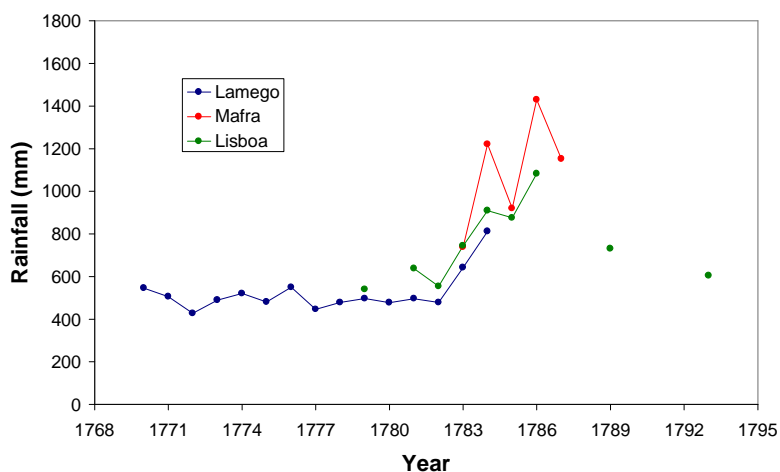


Fig. 4. Precipitation in three meteorological Portuguese sites in the 18th century (Taborda et al., 2004).

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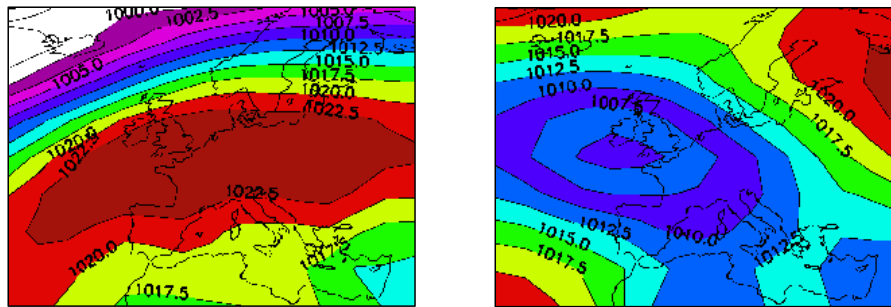


Fig. 5. Reconstruction of monthly sea level pressure fields in Europe in March 1779 (left) and in April 1782 (right) (Luterbacher et al., 2002a).

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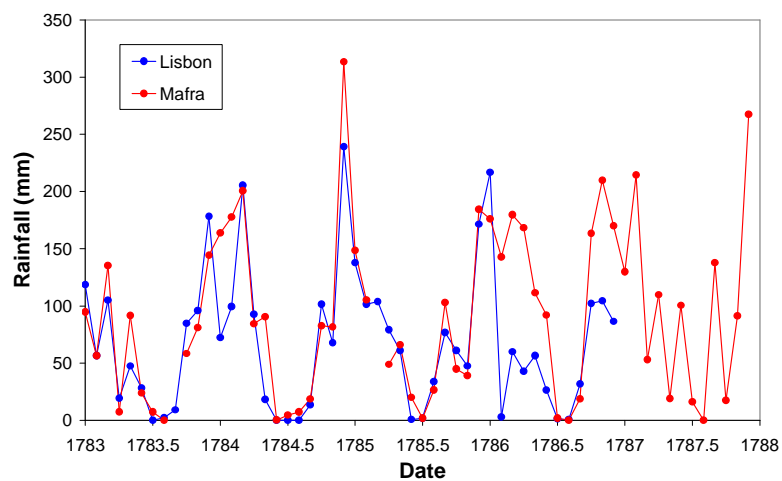


Fig. 6. Monthly rainfall in Lisbon and Mafra from 1783 to 1789.

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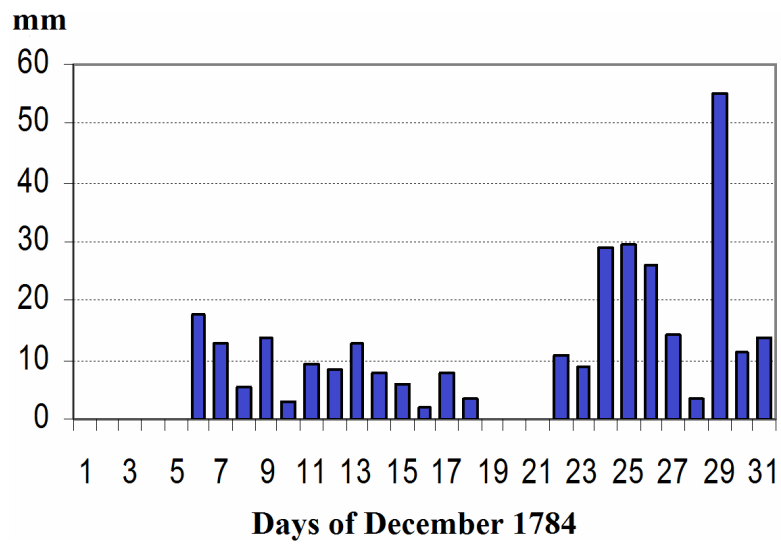


Fig. 7. Daily precipitation in Mafra in December 1784.

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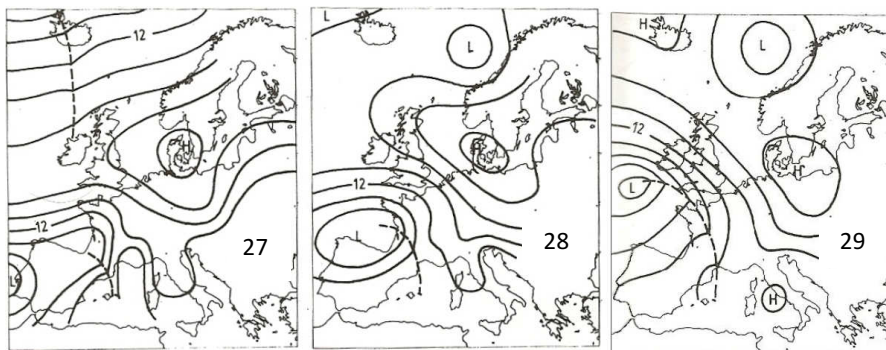


Fig. 8. Reconstruction of daily sea level pressure fields in Europe from 27 to 29 December 1784 (Kington, 1988).

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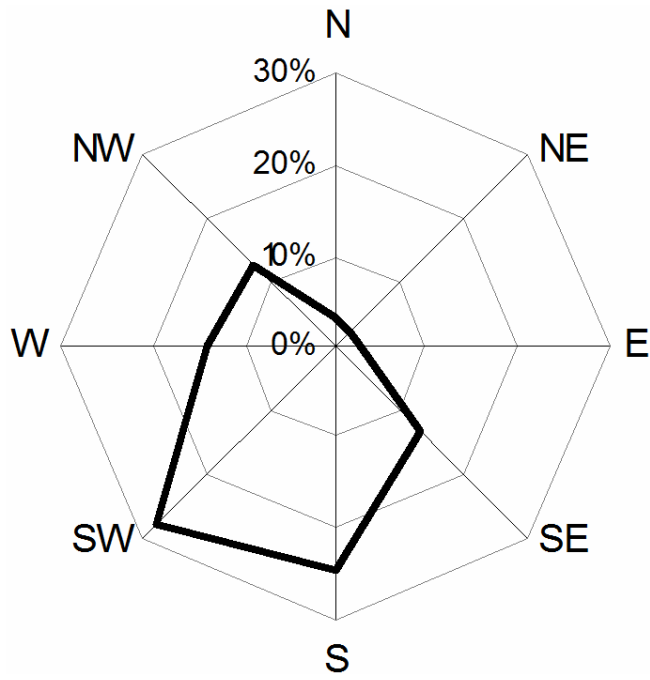


Fig. 9. Percentage of rainfall in Mafra (1783–1787) by wind direction, according with Velhos data (Taborda et al., 2004).

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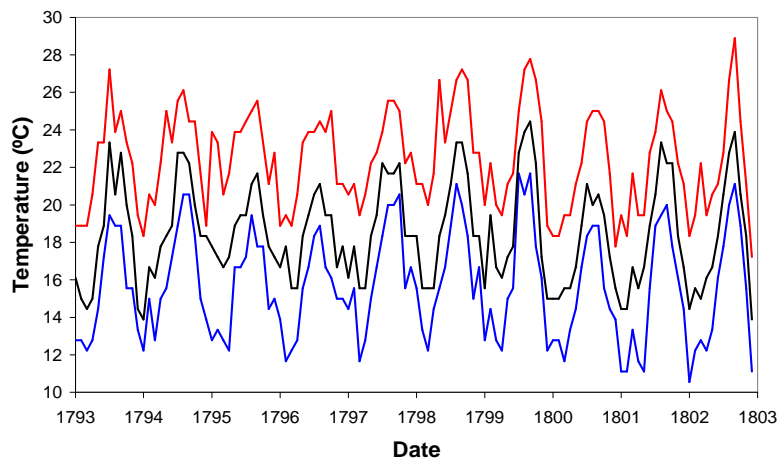


Fig. 10. Maxima (red), minima (blue) and mean (black) temperature at Madeira recorded by Murdock from 1793 to 1802.

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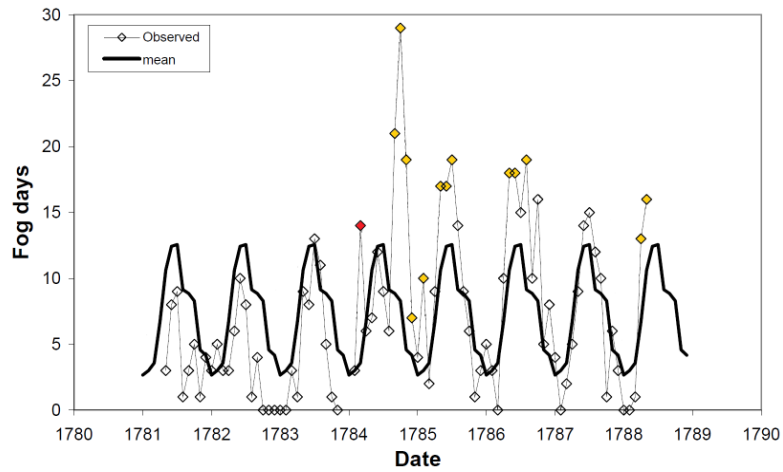


Fig. 11. Monthly values of number of fog days (nfd) recorded by BSD between 1781 and 1788 compared with the monthly climatology (bold line). The outstanding values that lie above the corresponding 1s (2s) values were filled with yellow (red) color.

3447

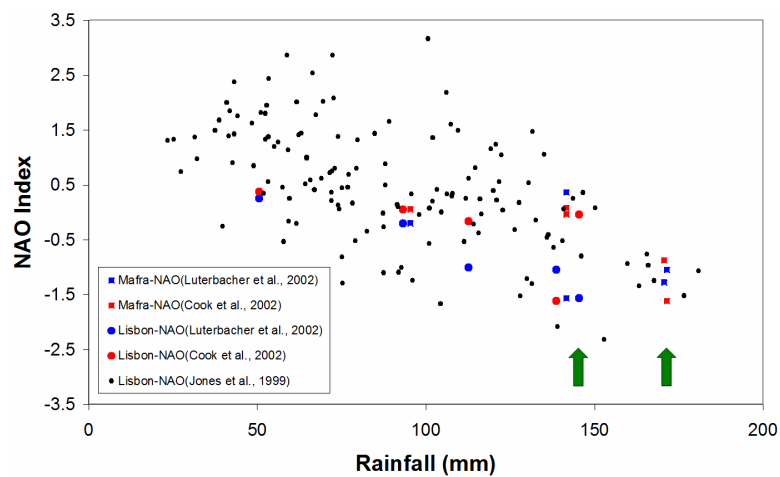


Fig. 12. Relationship between the average monthly precipitation for winter (DJFM) and the winter NAO index (DJFM). Arrows show the values for the 1783–1784 winter (see text for details).

3448