

Climate of the Past Discussions is the access reviewed discussion forum of *Climate of the Past*

The EDC3 chronology for the EPICA Dome C ice core

F. Parrenin¹, J.-M. Barnola¹, J. Beer², T. Blunier³, E. Castellano⁴, J. Chappellaz¹,
G. Dreyfus⁵, H. Fischer⁶, S. Fujita⁷, J. Jouzel⁵, K. Kawamura⁸,
B. Lemieux-Dudon¹, L. Loulergue¹, V. Masson-Delmotte⁵, B. Narcisi⁹,
J.-R. Petit¹, G. Raisbeck¹⁰, D. Raynaud¹, U. Ruth⁶, J. Schwander³, M. Severi⁴,
R. Spahni³, J. P. Steffensen¹¹, A. Svensson¹¹, R. Udisti⁴, C. Waelbroeck¹, and
E. Wolff¹²

¹Laboratoire de Glaciologie et Géophysique de l'Environnement, CNRS and Joseph Fourier University, Grenoble, France

²Department of Surface Waters, EAWAG, Dübendorf, Switzerland

³Climate and Environmental Physics, Physics Institute, University of Bern, Bern, Switzerland

⁴Department of Chemistry, University of Florence, Florence, Italy

⁵Laboratoire des Sciences du Climat et de l'Environnement, IPSL/CEA/CNRS/UVSQ, Gif-Sur-Yvette, France

⁶Alfred-Wegener-Institute for Polar and Marine Research, Bremerhaven, Germany

⁷National Institute of Polar Research, Research Organization of Information and Systems (ROIS), Tokyo, Japan

⁸Center for Atmospheric and Oceanic Studies Graduate School of Science, Tohoku University, Sendai, Japan

The EDC3 chronology

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

⁹ENEA, C. R. Casaccia, Roma, Italy

¹⁰CSNSM/IN2P3/CNRS, Orsay, France

¹¹Niels Bohr Institute, University of Copenhagen, Copenhagen, Denmark

¹²British Antarctic Survey, Cambridge, UK

Received: 20 February 2007 – Accepted: 28 February 2007 – Published: 12 March 2007

Correspondence to: F. Parrenin (parrenin@ujf-grenoble.fr)

CPD

3, 575–606, 2007

The EDC3 chronology

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

EGU

Abstract

The EPICA (European Project for Ice Coring in Antarctica) Dome C drilling in East Antarctica has now been completed to a depth of 3260 m, at only a few meters above bedrock. Here we present the new official EDC3 chronology, which is based on the use of 1) a snow accumulation and mechanical flow model, and 2) a set of independent age markers along the core. These are obtained by pattern matching of recorded parameters to either absolutely dated paleoclimatic records, or to insolation variations. We show that this new time scale is in excellent agreement with the Dome Fuji and Vostok ice core time scales back to 100 kyr within 1 kyr. Discrepancies larger than 3 kyr arise during MIS 5.4, 5.5 and 6, which points to anomalies in either snow accumulation or mechanical flow during these time periods. We estimate that EDC3 gives accurate event durations within 20% (2σ) back to MIS11 and accurate absolute ages with a maximum uncertainty of 6 kyr back to 800 kyr.

1 Introduction

The EPICA project has provided two records in East Antarctica, one at Dome C (EDC, EPICA community members, 2004), and one in the Dronning Maud Land area (EDML, EPICA community members, 2006). The completion of the Dome C core was delayed when the first drilling became stuck at 788 m in 1999. This shorter EDC96 core provided 45 kyr of paleoclimatic reconstructions (e.g., Jouzel et al., 2001; Monnin et al., 2001). The next EDC99 drilling was voluntarily stopped at a depth of 3260 m, about 15 m above bedrock, above which seismic soundings suggest the presence of melt water. EDC provides the longest in time ice core record available so far, with so far ~740 kyr records of Antarctic temperature (EPICA community members, 2004) and chemical impurities in Antarctica (Wolff et al., 2006), and ~650 kyr records of atmospheric composition (Siegenthaler et al., 2005; Spahni et al., 2005). All these records are currently being extended to 800 kyr.

CPD

3, 575–606, 2007

The EDC3 chronology

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

EGU

**The EDC3
chronology**

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

An accurate age scale is the basis for the interpretation of paleoclimatic records. We distinguish different types of accuracies. First, age scales need to be accurate in terms of absolute ages: we want the estimated age at a certain depth to be as close as possible to the real age (with an accuracy expressed in yr). This absolute accuracy is crucial when examining the phasing of two absolutely dated paleoclimatic records, and with insolation variations calculated by modelling of planet movements in the past (Laskar, 1990). For example, the insolation/climate phase relationship has been studied for terminations I and II thanks to accurate absolute age scales (Jouzel et al., 1995; Henderson and Slowey, 2000). Second, sequences of events can be analysed in detail without absolutely perfect age scales, provided that the studied records are stratigraphically linked. Here a relative age scale, (with an accuracy expressed in years) suffices. For example, the phasing between Antarctic temperature and CO₂ variations during the last deglaciation has been obtained from the Dome C core by estimating the ice/gas bubbles age difference (Monnin et al., 2001). Other examples include the phasing between Greenland and Antarctic temperature during the last glacial period obtained by a synchronisation of those records with the CH₄ atmospheric composition, which varies in phase at both poles (Blunier et al., 1998; EPICA community members, 2006). Third and finally, the last important accuracy is in the duration of climatic events (expressed in per cent). Indeed, this duration is characteristic of the climatic mechanisms involved, and will impact the frequency analysis of the records. We can cite as an example the duration and pacing of the so-called Dansgaard-Oeschger (D-O) events during the last glacial period which has been extensively studied (e.g., Schulz, 2002).

In the lack of radiochronologic method, numerous methods have been developed to date ice cores. They fall into 4 categories: (1) counting of layers representing a known time interval, e.g. annual layers, (2) ice flow modelling, (3) wiggle matching on other precisely dated time series, in particular insolation variations, and (4) use of climate independent age markers, like volcanic eruptions.

All these methods have advantages and drawbacks. Layer counting (Andersen et

**The EDC3
chronology**

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

al., 2007) and ice flow modelling (Parrenin et al., 2007¹) are accurate in terms of event durations because they are based on the evaluation of the annual layer thickness. On the other hand, errors cumulate and the precision on absolute ages decreases with depth. The new layer counted chronology for Greenland (GICC05, Vinther et al., 2006; Rasmussen et al., 2006; Andersen et al., 2006; Svensson et al., 2006) uses an improved multi-parameter counting approach, and currently extends back to around 42 kyr BP with a maximum counting error of 4 to 7% during the last glacial period. Unfortunately, layer counting is not feasible in central Antarctica where annual cycles are barely distinguishable (Ekaykin et al., 2002).

Comparison of paleoclimatic records to insolation variations (so-called orbital tuning methods) are generally applicable to a whole ice core, as long as the stratigraphy is preserved (e.g., Martinson et al., 1987; Dreyfus et al., 2007²). On the other hand: (1) the precision in terms of event durations is poor, (2) the precision in terms of absolute ages is limited by the hypothesis of a constant phasing between the climatic record used for the orbital tuning procedure and the insolation variations (and, by definition, does not allow one to infer this phasing). The advantage is that the achieved precision is not decreasing with depth. As a consequence, it is currently the most precise method to date the bottom of deep ice cores. Recently, the search for local insolation proxies in ice cores as, e.g. the O₂/N₂ ratio (Bender et al., 2002; Kawamura et al., 2007³) or the air content record (Raynaud et al., 2007⁴) has opened new prospects to remove this

¹Parrenin, F., Dreyfus, G., Durand, G. et al.: Ice flow modelling at EPICA Dome C and Dome Fuji, East Antarctica, submitted, 2007.

²Dreyfus, G., Parrenin, F., Lemieux-Dudon, B., et al.: Anomalous flow below 2700 m in the EPICA Dome C ice core detected using $\delta^{18}\text{O}$ of atmospheric oxygen measurements, submitted, 2007.

³Kawamura, K., Parrenin, F., Uemura, R., et al.: Northern Hemisphere forcing of climatic cycles over the past 360 000 years implied by absolute dating of Antarctic ice cores, in preparation, 2007.

⁴Raynaud, D., Lipenkov, V., Lemieux-Dudon, B., Duval, P., Loutre, M.-F., and Lhomme, N.:

hypothesis on the insolation/climate phase and to reach an accuracy within 1 kyr in the coming years.

Volcanic horizons bring important age markers, provided that they are precisely dated. This is the case for the last millenium (Traufetter et al., 2004), but beyond that limit, only a few of them have accurate absolute ages (Narcisi et al., 2006). In Antarctic ice cores, comparison to absolutely dated paleoclimatic records is particularly relevant for the dating of the D-O events, which have been accurately dated in several archives, and whose rapid transitions can be localized with a high precision in the Antarctic CH₄ record. The transfer of those age markers to the Antarctic ice matrix requires the evaluation of the ice/gas bubbles difference with a firn densification model (e.g., Goujon et al., 2003, and references therein).

In this article, we present EDC3, the new 800 kyr age scale of the EPICA Dome C ice core, which is based on a combination of various age markers and of a glaciological model. It is constructed in three steps. First, an age scale is derived by applying an ice flow model at Dome C. Independent age markers are used to control some poorly known parameters of this model (such as the conditions at the base of the glacier). Second, the age scale is synchronised onto the new Greenlandic GICC05 age scale during three time periods: the last 6 kyr, the last deglaciation, and the Laschamp event (around 41 kyr BP). Third, the age scale is corrected in the bottom ~500 m (corresponding to the time period 400–800 kyr BP), where the model is unable to capture the complex ice flow pattern.

In Sect. 2, we first present the different age markers that can be derived on the EDC ice core. We then describe in Sect. 3 the construction of EDC3. In Sect. 4, we compare it with other age scales from the late Quaternary. Finally, we discuss the accuracies of this new time scale in Sect. 5.

In this paper, the notation “yr BP” refers to “years before AD1950”.

The local insolation signature of air content in Antarctic ice: A new step toward an absolute dating of ice records, submitted, 2007.

CPD

3, 575–606, 2007

The EDC3 chronology

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

EGU

2 Age markers

In this section, we describe the dated horizons (so-called age markers) which can be derived on the EDC ice core.

2.1 Dated volcanic eruptions during the last millenium

5 Using the sulphate data (Castellano et al, 2005), several volcanic eruptions of known age have been identified in the EDC96 ice core during the Holocene. Among these, only a few of the most recent are independently absolutely dated (Traufetter et al., 2004): Krakatau, 8.35 m⁵, AD1884±1; Tambora, 12.34 m, AD1816±1; Huaynaputina, 23.20 m, AD1601±1; Kuwae, 29.77 m, AD1460±5; Unknown (El Chichon?), 38.12 m,
10 AD1259±5; Unidentified, 39.22 m, AD1228±5; Unknown, 41.52 m, AD1171±6.

2.2 Synchronisation onto GICC05 and INTCAL with ¹⁰Be for the last 6 kyr

¹⁰Be and ¹⁴C are cosmogenic radionuclides, and their production rates are modulated by solar activity and by the strength of the Earth's magnetic field. Therefore ¹⁰Be records in Greenland and Antarctica, as well as atmospheric ¹⁴C reconstructions (INTCAL, Reimer et al., 2004) show common variations.
15

Three methods were used independently to construct age scales for EDC over the last 6 kyr. The first two are obtained by wiggle matching the EDML ¹⁰Be record to either the GRIP ¹⁰Be record dated by layer counting (GICC05, Vinther et al., 2006), or with the INTCAL atmospheric ¹⁴C reconstruction (Reimer et al., 2004). These age scales
20 have been transferred to EDC96 by volcanic synchronisation (Severi et al., 2007⁶). The third time scale is obtained by wiggle matching to the Vostok ¹⁰Be record with INTCAL

⁵The identification of this volcanic eruption has actually been revised since the study by Castellano et al. (2005).

⁶Severi, M., Castellano, E., Morganti, A., et al.: Synchronisation of the EDML1 and EDC3 timescales for the last 52 kyr by volcanic signature matching, in preparation, 2007.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

atmospheric ^{14}C reconstructions (Raisbeck et al., 1998). The resulting Vostok age scale (more precisely the VK-BH1 core age scale) was then transferred to EDC96 via the VK-BH7 core by volcanic matching (Udisti et al., 2004).

We derive two age markers from these chronologies, at periods of high $^{10}\text{Be}/^{14}\text{C}$ variations (during which the synchronisation is robust). The three chronologies give similar ages within 30 years for these two periods and we calculated average ages of : 2716 yr BP and 5279 yr BP for the age markers at 107.83 m and 181.13 m, respectively.

2.3 Match to GICC05 with CH_4 during the last deglaciation

During the last deglaciation, synchronisation to the NGRIP GICC05 chronology (Rasmussen et al., 2006) is possible with the transitions (Björck et al., 1998) that are common to the Greenland and Antarctic high resolution methane records, and the Greenland climate record (Severinghaus and Brook, 1999): GS-2a/GI-1e (Oldest Dryas/Bølling), GI-1a/GS-1 (Allerød/Younger Dryas) and GS-1/Holocene (Younger Dryas/Holocene). In that way an age for the CH_4 transitions can be obtained. This age for the gas record then has to be transferred to an age for the ice. However, the uncertainty in the estimation of this age difference (Δage) is large at EDC because of the low accumulation rate and the low temperature (typical model estimates of Δage at EDC are 2200 yr for the Holocene and 5500 yr for the LGM). This forces us to make a detour via the EDML ice core where accumulation rate and temperature are higher (typical model estimates of Δage at EDML are 700 yr for the Holocene and 1800 yr for the LGM). For the rapid warmings at the GS-2a/GI-1e and GS-1/Holocene transitions, the EDML CH_4 data were matched to the NGRIP stable isotope record (NGRIP project members, 2004). The corresponding GICC05 ages were transferred first from the EDML gas depth-scale to the EDML ice depth-scale by subtracting the calculated Δdepth (depth difference between gas bubbles and ice with the same age). Δdepth was obtained by multiplying the modelled close off depth (in ice equivalent, Loulergue

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

et al., 2007⁷) with the EDML mechanical thinning function (Huybrecht et al., 2007⁸). This age was further transferred to EDC via the volcanic match between both cores (Severi et al., 2007⁶).

The two derived age markers are 11.65 ± 0.32 and 14.64 ± 0.35 kyr BP for respectively 355.34 m and 421.15 m EDC96-depth. The uncertainty is estimated as the root mean square sum of the GICC05 age error (the number of uncertain layers given by Rasmussen et al., 2006, is taken as 2σ) and of a 300 yr 2σ uncertainty resulting from the uncertainty in the Δ depth estimate at EDML ($2\sigma=10$ m).

2.4 Match to GICC05 during the Laschamp event

The Laschamp geomagnetic excursion gives rise to a structured peak in the ^{10}Be records from Greenland (Yiou et al., 1997) and Antarctica (Raisbeck et al., 2002), which allows EDC96 to be synchronised to GRIP (Raisbeck et al., 2007⁹), and in turn, to NGRIP dated by layer counting (GICC05, Krogh Andersen et al., 2007; Svensson et al., 2007). Two of the ^{10}Be sub-peaks have been localized in the EDC96 ice core at 735.35 m and 744.81 m, and at 2231.9 m and 2246.2 m at GRIP. The corresponding GICC05 age for the middle of these two peaks is 4 200 yr BP (max counting error of 1627 yr), corresponding to a depth of 740.08 m at EDC (Raisbeck et al., 2002) and we adopt this age.

This age of the Laschamp event is compatible (within the uncertainties) with K-Ar and ^{40}Ar - ^{39}Ar ages from contemporaneous lava flow (40.4 ± 2.0 kyr BP, Guillou et al.,

⁷Loulergue, L., Parrenin, F., Blunier, T., Barnola, J.M., Spahni, R., Schilt, A., Raisbeck, G., Chappellaz, J.: Gas age-ice age differences along the EPICA DC and EDML ice cores, 0–50 kyr, submitted, 2007.

⁸Huybrechts, P., Rybak, O., Pattyn, F., and Steinhage, D.: Ice thinning and non-climatic biases of the upper 2500 m of the EDML $\delta^{18}\text{O}$ record from a nested model of the Antarctic ice sheet, in preparation, 2007.

⁹Raisbeck, G., Yiou, F., Jouzel, J., and Stocker, T: Direct North-South Synchronization of Abrupt Climate Change Records in Ice Cores Using Beryllium 10, in preparation, 2007.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

2004). During this time period, which corresponds to D-O event 10 (Yiou et al., 1997; Raisbeck et al., 2002), GICC05 is also in good agreement with the Hulu Cave U-Th chronology (41.4 kyr BP, Wang et al., 2001), or with the Cariaco basin record (Hughen et al., 2004) when its ^{14}C ages are calibrated following the Fairbanks et al. (2005) curve (we obtain again an age of 41.2 kyr BP for the middle of the ^{10}Be peak corresponding to the middle of D-O 10). Genty et al. (2003) also found a compatible U-Th age of 40.0 kyr BP for the middle of D-O 10, though the identification of D-O 10 in this record is more ambiguous.

2.5 The Mont Berlin ash layer

Thanks to geochemical data (major elements and trace elements), Narcisi et al. (2006) identified a volcanic ash layer originating from a Mt Berlin (Antarctica) eruption. This event has been dated at 92.5 ± 2 kyr BP by an Ar/Ar method applied on ash material collected close to the volcano.

2.6 Timing of termination II

The rapid CH_4 event marking the end of termination II can be dated by comparison to speleothem records, dated by the U-Th method, assuming that these fast transitions are synchronous. We obtain 129.3 kyr BP from Dongge cave in China (Yuan et al., 2004), and 130.9 kyr BP from Pekiin cave in Northern Israel (Bar-Matthews et al., 2003). We took the average of these two ages (130.1 kyr BP) and assumed a confidence interval of 2 kyr. We used the Δ depth estimate from the EDC2 age scale to export the CH_4 depth of 1723 m to an ice depth of 1699 m on EDC99. The uncertainty introduced by this ice/gas depth difference evaluation is only a few hundreds of years, and we neglect it here in front of the uncertainty in the absolute age.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

2.7 Air content age markers 0–440 kyr BP

The total air content of the polar ice may be interpreted as a marker of the local summer insolation (Raynaud et al., 2007⁴). Indeed, the solar radiative power absorbed at the surface influences the snow structure in the first upper meters and, in turn, the porosity of snow in the bubble close-off layer. Even if the detailed physical mechanism is still under debate, the presence of a strong 41 kyr obliquity frequency in the air content signal makes it appropriate for the application of an orbital tuning method. We used 19 age markers from the air content age scale available back to 440 kyr BP. Each age marker corresponds to a minima of obliquity, and the assumed uncertainty is 4 kyr.

2.8 $^{18}\text{O}_{\text{atm}}$ age markers for stages 300–800 kyr BP

A relationship between the isotopic composition of atmospheric oxygen ($\delta^{18}\text{O}$ of O_2 , noted $\delta^{18}\text{O}_{\text{atm}}$) and daily northern hemisphere summer insolation has been observed at Vostok for the youngest four climate cycles. This property has been exploited to construct various orbital age scales for Vostok (Petit et al., 1999; Shackleton, 2000). Dreyfus et al. (2007)² used a similar approach to derive an age scale for the bottom part of the EDC ice core (300–800 kyr BP) by assuming that $^{18}\text{O}_{\text{atm}}$ lags the precession variations by 5 kyr with an estimated uncertainty of 6 kyr. The selected age markers are placed at each mid-transition of $\delta^{18}\text{O}_{\text{atm}}$ (see Dreyfus et al., 2007² for more details).

2.9 The Brunhes-Matuyama reversal

The most recent of the geomagnetic inversions, referred to as the Brunhes-Matuyama reversal, has been localized between 3161 and 3170 m in the EDC ^{10}Be record (Raisbeck et al., 2006). This reversal has been dated radiometrically to have occurred 776 ± 12 kyr BP (Coe et al., 2004), taking into account decay constant and calibration uncertainties. This transition has also been orbitally dated to be 778 kyr ago (Tauxe et al., 1996). Several authors have also reported evidence for a “precursor” event, 15 kyr

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

before the B-M boundary (Brown et al., 2004), supported by the EDC ^{10}Be record.

3 Construction of the time scale

3.1 The EDC3 age scale

The EDC3 age scale was constructed in three stages.

5 First, a preliminary dating was obtained by purely ice flow modelling (Parrenin et al., 2007¹). The ice flow model has two components. 1) The initial annual layer thickness (i.e. the accumulation rate) is evaluated from the deuterium content of the ice, assuming an exponential relationship between accumulation rate and deuterium content, the later
10 being corrected for variations in isotopic composition of the mean ocean. 2) The vertical compression of the layers, or total thinning ratio, is evaluated with a mechanical model. The age at a depth z is then given by:

$$\text{age}(z) = \int_0^z \frac{1}{T(z') a(z')} dz'. \quad (1)$$

where $a(z)$ is the initial annual layer thickness and $T(z)$ is the compression factor. This ice flow model contains several poorly known parameters: the average Holocene accu-
15 mulation rate, the slope between deuterium and logarithm of accumulation, the basal melting, and two parameters for the vertical profile of velocity (basal sliding and internal deformation). The values of these parameters have been determined by independent age markers, using a Monte Carlo Markov Chain (MCMC) inverse method. 21 age
20 markers have been selected and are listed in Table 3 in Parrenin et al. (2007)¹. Not all those listed in Sect. 2 have been selected, in order to prevent over-tuning the model in certain parts which would be a detriment to other parts. There are 8 age markers during the last climatic cycle, and in particular 3 during the Holocene. It is important to understand how this “modelled” age scale is dependent on these age markers. The

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

**The EDC3
chronology**

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

average Holocene accumulation rate impacts the Holocene ages and is mainly constrained by the Holocene age markers (dated volcanoes and ^{10}Be age markers). Then the deuterium – accumulation slope impacts the glacial ages and is mainly constrained by the age of the Laschamp event. The basal melting impacts the total duration of the record and is mainly constrained by the bottom age markers. Finally, the two parameters related to the vertical profile of velocity only induce general trends in the age scale and are constrained by all the other age markers. Hence, the resulting age scale does not match perfectly the age markers obtained by comparison to insolation variations (obtained from the air content record).

The second stage is an a posteriori strict match of the age scale to dated volcanoes and to the NorthGRIP GICC05 time scale in the top part. In this part, the total thinning function is close to 1 and thought to be well constrained by the ice flow model. For this reason we expect the main sources of uncertainties to come from the accumulation model. Consequently we modified the modelled accumulation rate so that the resulting age scale fits perfectly with: 1) the dated volcanoes of the last millenium; 2) the two ^{10}Be age markers in the last 6 kyr (Sect. 2.2); 3) one methane age marker during the last deglaciation (Sect. 2.3); 4) the Laschamp age marker at 41.2 kyr BP (Sect. 2.4). These age markers are listed in Table 1.

The third stage is a correction of the modelled thinning function in the bottom 500 m of the core (beyond MIS11, ~ 400 kyr BP), where the ice flow model is unable to fit the $\delta^{18}\text{O}_{\text{atm}}$ age markers (Dreyfus et al., 2007²). This problem had first been detected by Lisiecki and Raymo (2005), who suggested a problem in the accumulation estimate. However, Dreyfus et al. (2007)² showed, by a comparison of deuterium and CO_2 variations, that this anomaly is principally due to the presence of ice flow irregularities. They a posteriori corrected the total thinning function so that the resulting age scale fits those $\delta^{18}\text{O}_{\text{atm}}$ age markers within their confidence interval. See Dreyfus et al. (2007², Table 1) for a complete list of the age markers used and for more details on the method.

By following this procedure, we have used the best available chronological information for each section of the core, while still allowing the model to provide a smooth

interpolation of all unmarked periods.

As stated in the introduction, two different cores have been drilled at EDC: EDC96 extending to 788 m depth (approximately back to 45 kyr BP), and EDC99 drilled down to the bedrock. For the first ~45 kyr, the majority of paleoclimatic reconstructions have been obtained from EDC96. Therefore, EDC3 has been defined on EDC96 depths on the shallow part and on EDC99 depths in the bottom part. The age scale has then been transferred to EDC99 in the shallow part thanks to a volcanic synchronisation of the two cores (Wolff et al., 2005).

Estimates of the gas-ice age difference and related discussions can be found in Loulergue et al. (2007)⁷.

3.2 EDC3 exported to EDML, Dome Fuji and Vostok

The EDC3 age scale was then exported to EDML, Dome Fuji and Vostok thanks to synchronisation of these ice cores with EDC. The EDC-EDML synchronisation and the resulting EDML1 chronology are fully described in Severi et al. (2007)⁶ and Ruth et al. (2007)¹⁰. The EDC-DF and EDC-VK synchronisations are done by matching isotopic records, and by using common volcanic horizons (Narcisi et al., 2005¹¹). See supplementary material (<http://www.clim-past-discuss.net/3/575/2007/cpd-3-575-2007-supplement.zip>) for a list of synchronisation markers used.

¹⁰Ruth, U., Barnola, J.-M., Beer, J., et al.: EDML1: A chronology for the EDML ice core, Antarctica, over the last 150 000 years, in preparation, 2007.

¹¹We did not use the EDC-VK volcanic synchronisation obtained by Udisti et al. (2004), because it concerns the 5G VK core, and not the 3G core on which the deuterium measurements have been performed.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

4 EDC3 compared to other age scales

4.1 Comparison with EDC2

The former EDC1 time scale for EDC96 (Schwander et al., 2001), and the extended EDC2 for the last 740 kyr (EPICA community members, 2004) were also built on a combination of age markers and modelling information. As for EDC3, a one dimensional ice flow model was controlled by a set of age markers. There are however several important differences. For EDC1, the time scale extended only back to 45 kyr BP, and different glaciological parameters were used for different time periods covered by the time scale. EDC2 extended only back to 740 kyr BP and there was no a posteriori correction of the age scale, neither in the top part, nor in the bottom part where the ice flow is complex. Moreover, the ice flow model did not take into account basal sliding and variations in ice sheet thickness, and the age markers were mainly obtained by comparison to the oceanic Bassinot stack (Bassinot et al., 1994).

Figure 1 compares EDC1 and EDC3 on the last 45 kyr. EDC2 is also shown for convenience, but EDC1 was still the official age scale for the top part of the core. EDC3 is younger by a few decades for the last 2 kyr. Then it is older by less than 100 yr between 2 and 8 kyr BP. The difference increases to ~ 200 yr for the early Holocene period, around 10 kyr BP. Then the difference becomes positive (EDC1 is older) with a maximum of ~ 600 yr at the LGM. The difference then decreases roughly linearly down to -700 yr at 45 kyr BP.

Figure 2 compares EDC2 and EDC3. The difference ranges approximately between $+3$ kyr and -1.5 kyr for the last 400 kyr. Then EDC3 is older during the last glacial period, with a difference of ~ 3 kyr for MIS 5.5. This is due in particular to the use of the Mt Berlin and Termination II age markers. The difference then slowly decreases back to MIS 10.

For the period 400–800 kyr BP, the difference is much larger, and reaches $+20$ and -7 kyr. This is due to the a posteriori correction in EDC3 of ice flow irregularities in the bottom part of the core. The largest differences are for MIS 13–14 (where EDC3 is

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

older by 15–20 kyr), MIS 15.3 (where EDC3 is younger by ~5 kyr) and MIS 16 (where EDC3 is older by 10–15 kyr). Duration of MIS 15.1 has been considerably shortened in EDC3, while duration of MIS 12 is now larger.

4.2 Comparison with LR04

5 The LR04 marine stack is composed of benthic $\delta^{18}\text{O}$ records from 57 globally distributed sites aligned by an automated graphic correlation algorithm (Lisiecki and Raymo, 2005). The LR04 age model is derived from tuning the $\delta^{18}\text{O}$ stack to a simple ice model based on 21 June insolation at 65°N , with additional constraints from the sedimentation to prevent overtuning.

10 On Fig. 3, we compared the EDC deuterium record on EDC3 with the LR04 stack on its own time scale, shifted by 2.5 kyr towards older ages. This 2.5 kyr phase has been deduced from the last deglaciation. The overall agreement between both time scales is good, with differences never exceeding 5 kyr. In contrast, the previous EDC2 time scale showed disagreements up to 20 kyr with LR04 in the part older than 400 kyr BP
15 (Lisiecki and Raymo, 2005; Dreyfus et al., 2007²).

The age difference is particularly small during the last 400 kyr (back to MIS11), oscillating between -2.5 kyr and 1.5 kyr. This age difference may reflect either errors in the synchronisation, or may be due to phases in the climatic system, i.e. related to the fact that both curves do not represent the same climatic proxy. The fact that this
20 difference is stable is reassuring because it shows a certain consistency between both time scales which were derived completely independently. The glaciological modelling method thus seems appropriate for Dome C back to MIS11 without any additional distortion. The age difference decreases to approximately -5 kyr between 450 and 600 kyr BP, then reaches its maximum at termination VII (from MIS16 to MIS15) with 3.5 kyr, decreases to -5 kyr at MIS18, and finally increases to around 2.5 kyr at termination 9
25 (from MIS20 to MIS19). This bottom interval (beyond MIS11) where the age difference is less stable, is where the ice flow model becomes inaccurate (Dreyfus et al., 2007²).

The EDC3 chronology

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

4.3 Comparison with DF and VK glaciological chronologies

On Fig. 4 and Fig. 5, we compare the EDC isotopic record on the EDC3 time scale with the Dome Fuji and Vostok isotopic record, on their respective glaciological age scales DFGT-2006 (Parrenin et al., 2007¹) and VK-FGT1 (Parrenin et al., 2004).

5 The age differences are always less than 1 kyr for the last ~90 kyr. This good agreement is especially remarkable because very few age markers were used for the last glacial part. We interpret this good agreement as the fact that the glaciological models used are robust for the upper part of the ice sheets where the mechanical ice flow is still predictable. It is also a proof that the assumed relationship between isotopic content
10 of the ice and surface accumulation rate is valid within a few percent.

The situation is more complex for the second climatic cycle, where EDC3 is significantly older than both DFGT-2006 and VK-FGT1, the difference reaching around 5 kyr. The agreement is again better for the third and fourth climatic cycles, with differences never exceeding 2 kyr.

15 Figure 6 compares the duration of climatic events in EDC3 and DFGT-2006, or in EDC3 and VK-FGT1. These three time scales are consistent, generally within 20%. It should be noted that differences depicted on this figure may either reflect a real difference in the age scales, or an error in the synchronisation process. The agreement is very good back to ~90 kyr BP, but then the situation for MIS5.4 to 6 is more complex.
20 MIS5.4 is significantly shorter in EDC3 than in DFGT-2006 and VK-FGT1. Then, the duration of MIS5.5 (taken between the onset of the deuterium rise and the end of its fall) is intermediate in EDC3 between its duration in DFGT-2006 (~1 kyr shorter than in EDC3) and its duration in VK-FGT1 (~1 kyr longer than in EDC3). Finally, the duration of MIS6 is significantly shorter in EDC3 than in both other age scales. We do not know
25 at this stage if these discrepancies are due to errors in the accumulation models or in the mechanical thinning models.

CPD

3, 575–606, 2007

The EDC3 chronology

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

EGU

5 Confidence interval of the age scale

The confidence interval determination is a difficult task when no robust statistical information is available. Here, we evaluate it subjectively by using the comparison with the other age scales and with the age markers.

5 Back to AD1600, the error in EDC3 mainly comes from the interpolation of the dated volcanoes which we estimate (2σ) to be 3 yr. Between AD1100-1600, the age error of the volcanoes increases to 5 yr, and adding an interpolation error we estimate the total error at 8 yr. The accuracy is then constrained by the precision of the ^{10}Be age markers, which we estimate at 100 yr. We thus estimate that the 2σ error on EDC3
10 increases to 100 yr at 2000 yr BP and stays stable back to 6000 yr BP. The accuracy of EDC3 then increases to 400 yr at 14 kyr BP, which is roughly the error on the CH_4 age markers. By comparison to the Dome Fuji and Vostok chronologies and to the GICC05 age of the Laschamp, we estimate the confidence interval to increase to 1 kyr at 18 kyr (Last Glacial Maximum), 1.5 kyr at 40 kyr, and finally 3 kyr at 100 kyr BP. Our estimated
15 confidence interval is constrained by the quality of the orbital tuning age markers from air content or $^{18}\text{O}_{\text{atm}}$ records; we estimate it to increase to 6 kyr at 130 kyr and to stay stable down to the bottom of the record.

In terms of event durations, we estimate the accuracy to be 20% for the top part of the record (back to MIS11), by comparison to Vostok and Dome Fuji glaciological age
20 scales. For the bottom part (below MIS11), a more conservative estimate of 40% is more appropriate because of the flow anomalies (Dreyfus et al., 2007²).

6 Conclusion and perspectives

We derived an EDC3 chronology for the EPICA Dome C ice core, which was then exported to EDML, Dome Fuji and Vostok ice cores by synchronisation of these ice
25 cores. This chronology has been obtained by using a combination of age markers and ice flow modelling. The good agreement between EDC, Vostok and Dome Fuji

CPD

3, 575–606, 2007

The EDC3 chronology

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

EGU

ice flow models points to the good accuracy of EDC3 in terms of event durations, which we estimate to be better than 20% for the last 400 kyr. This is a significant improvement with respect to marine age scales where the resolution is rougher and where the sedimentation is less regular.

5 Apart from ice flow modelling improvements, further developments need to be done concerning the inverse method used for the conjunction of models and age markers. The method used for EDC3 is based on a so-called deterministic approach, where the uncertainties in the ice flow models are supposed to originate from poorly known physical parameters. In reality, there are other non-identified sources of uncertainty in these
10 models which need to be taken into account in a statistical way. A second potential improvement is to apply this inverse method to several drilling sites simultaneously, to obtain a common and optimal age scale for several ice cores, as has been done in the marine world (Lisiecki and Raymo, 2005).

We also hope that the precision of the age markers will increase in the coming years. The number of U-Th dated speleothems for the last climatic cycles should increase in
15 the future (Henderson, 2006). New local insolation proxies such as O_2/N_2 and air content are also a promising source of accurate age markers, but the physical mechanisms involved need to be better understood and the precision of these age scales needs to be independently confirmed.

20 *Acknowledgements.* It is a contribution to the European Project for Ice Coring in Antarctica (EPICA), a joint European Science Foundation/European Commission scientific programme, funded by the EU (EPICA-MIS) and by national contributions from Belgium, Denmark, France, Germany, Italy, the Netherlands, Norway, Sweden, Switzerland and the United Kingdom. The main logistic support was provided by IPEV and PNRA (at Dome C) and AWI (at Dronning
25 Maud Land). This work was partly funded by the French ANR projects MIDIGA and PICC. This is EPICA publication no. XX.

The EDC3 chronology

F. Parrenin et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[I◀](#)[▶I](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

References

- Andersen, K. K., Svensson, A., Johnsen, S. J., Rasmussen, S. O., Bigler, M., Röthlisberger, R., Ruth, U., Siggaard-Andersen, M.-L., Steffensen, J. P., Dahl-Jensen, D., Vinther, B. M. and Clausen, H. B: The Greenland Ice Core Chronology 2005, 15-42 kyr. Part 1: Constructing the time scale, *Quat. Sci. Rev.*, 25, 3246–3257, 2006.
- Bar-Matthews, M., Ayalon, A., Gilmour, M., Matthews, A., and Hawkesworth, C. J: Sea-land oxygen isotopic relationships from planktonic foraminifera and speleothems in the Eastern Mediterranean region and their implication for paleorainfall during interglacial intervals, *Chem. Geol.*, 67, 3181–3199, 2003.
- Bassinot, F. C., Labeyrie, L. D., Vincent, E., Quidelleur, X., Shackleton, N. J., and Lancelot, Y: The astronomical theory of climate and the age of the Brunhes-Matuyama magnetic reversal, *Earth Planet. Sci. Lett.*, 126(1–3), 91–108, 1994.
- Bender, M. L: Orbital tuning chronology for the Vostok climate record supported by trapped gas composition, *Earth Planet. Sci. Lett.*, 204, 275–289, 2002.
- Björck, S., Walker, M. J., Cwynar, L. C., Johnsen, S., Knudsen, K.-L., Lowe, J. J., Wohlfarth, B., and intimate members: An event stratigraphy for the Last Termination in the North Atlantic region based on the Greenland ice-core record: A proposal by the INTIMATE group, *J. Quat. Sci.*, 13(4), 283–292, 1998.
- Blunier, T., Chappellaz, J., Schwander, J., Dällenbach, A., Stauffer, B., Stocker, T., Raynaud, D., Jouzel, J., Clausen, H., Hammer, C., and Johnsen, S: Asynchrony of Antarctic and Greenland climate change during the last glacial period, *Nature*, 394(6695), 739–743, 1998.
- Brown, L. L., Singer, B. S., Pickens, J. C. and Jicha, B. R: Paleomagnetic directions and $^{40}\text{Ar}/^{39}\text{Ar}$ ages from the Tataro-San Pedro volcanic complex, Chilean Andes: Lava record of a Matuyama-Brunhes precursor?, *J. Geophys. Res. (Solid Earth)*, 109(B12), B12101, doi:10.1029/2004JB003007, 2004.
- Castellano, E., Becagli, S., Hansson, M., Hutterli, M., Petit, J. R., Rampino, M. R., Severi, M., Steffensen, J. P., Traversi, R., and Udisti, R: Holocene volcanic history as recorded in the sulfate stratigraphy of the European Project for Ice Coring in Antarctica Dome C (EDC96) ice core, *J. Geophys. Res.*, 116, D06114, doi:10.1029/2004JD005259, 2005.
- Coe, R. S., Singer, B. S., Pringle, M. S., and Zhao, X: Matuyama-Brunhes reversal and Kamikatsura event on Maui: paleomagnetic directions, $^{40}\text{Ar}/^{39}\text{Ar}$ ages and implications, *Earth Planet. Sci. Lett.*, 222, 667–684, 2004.

CPD

3, 575–606, 2007

The EDC3 chronology

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

EGU

**The EDC3
chronology**

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

- Ekaykin, A. A., Lipenkov, V. Y., Barkov, N. I., Petit, J.-R., and Masson-Delmotte, V: Spatial and temporal variability in isotope composition of recent snow in the vicinity of Vostok station, Antarctica: implications for ice-core record interpretation, *Ann. Glaciol.*, 35, 181–186, 2002.
- EPICA community members: One-to-one coupling of glacial climate variability in Greenland and Antarctica, *Nature*, 444, 195–198, 2006.
- EPICA community members: 8 glacial cycles from an Antarctic ice core, *Nature*, 429, 623–628, 2004.
- Fairbanks, R. G., Mortlock, R. A., Chiu, T.-C., Cao, L., Kaplan, A., Guilderson, T. P., Fairbanks, T. W., Bloom, A. L., Grootes, P. M., and Nadeau, M.-J: Radiocarbon calibration curve spanning 0 to 50 000 years BP based on paired $^{230}\text{Th}/^{234}\text{U}/^{238}\text{U}$ and ^{14}C dates on pristine corals, *Quat. Sci. Rev.*, 24, 1781–1796, 2005.
- Genty, D., Blamart, D., Ouhadi, R., Gilmour, M., Baker, A., Jouzel, J., and Van-Exter, S: Precise dating of Dansgaard-Oeschger climate oscillations in western Europe from stalagmite data, *Nature*, 421, 833–837, 2003.
- Goujon, C., Barnola, J.-M., and Ritz, C: Modeling the densification of polar firn including heat diffusion: application to close-off characteristics and gas isotopic fractionation for Antarctica and Greenland sites, *J. Geophys. Res.*, 108(D24), ACL10/1-10, 2003.
- Guillou, H., Singer, B. S., Laj, C., Kissel, C., Scaillet, S., and Jicha, B. R: On the age of the Laschamp geomagnetic excursion, *Earth Planet. Sci. Lett.*, 227, 331–343, 2004.
- Henderson, G. M: Caving In to New Chronologies, *Science*, 313, 620–622, 2006.
- Henderson, G. M. and Slowey, N. C: Evidence from U-Th dating against northern hemisphere forcing of the penultimate deglaciation, *Nature*, 404(6773), 61–66, 2000.
- Hughen, K., Lehman, S., Southon, J., Overpeck, J., Marchal, O., Herring, C. and Turnbull, J: ^{14}C activity and global carbon cycle changes over the past 50 000 years, *Science*, 303, 202–205, 2004.
- Jouzel, J., Vaikmae, R., Petit, J., Martin, M., Duclos, Y., Stievenard, M., Lorius, C., Toots, M., Mélières, M., Burckle, L., Barkov, N., and Kotlyakov, V: The two-step shape and timing of the last deglaciation in antarctica, *Clim. Dyn.*, 11, 151–161, 1995.
- Jouzel, J., Masson, V., Cattani, O., Falourd, S., Stievenard, M., Stenni, B., Longinelli, A., Johnsen, S., Steffensen, J., Petit, J., Schwander, J., Souchez, R., and Barkov, N. I: A new 27 kyr high resolution East Antarctic climate record, *Geophys. Res. Lett.*, 28(16), 3199–3202, 2001.
- Laskar, J: The chaotic motion of the solar system: a numerical estimate of the size of the

**The EDC3
chronology**

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

chaotic zones, *Icarus*, 88, 266–291, 1990.

Lisiecki, L. E. and Raymo, M. E: A Plio-Pleistocene Stack of 57 Globally Distributed Benthic $\delta^{18}\text{O}$ Records, *Paleoceanography*, 20(1), PA1003, doi:10.1029/2004PA001071, 2005.

Martinson, D. G., Pisias, N. G., Hays, J. D., Imbrie, J., Moore, T. C., and Shackleton, N. J: Age dating and the orbital theory of the ice ages: development of a high-resolution 0–300 000 years chronostratigraphy, *Quat. Res.*, 27, 1–30, 1987.

Monnin, E., Indermuhle, A., Dallenbach, A., Fluckiger, J., Stauffer, B., Stocker, T. F., Raynaud, D., and Barnola, J.-M: Atmospheric CO_2 concentrations over the last glacial termination, *Science*, 291(5501), 112–114, 2001.

Narcisi, B., Petit, J. R., Delmonte, B., Basile-Doelsch, I., and Maggi, V: Characteristics and sources of tephra layers in the EPICA-Dome C ice record (East Antarctica): Implications for past atmospheric circulation and ice core stratigraphic correlations [rapid communication], *Earth Planet. Sci. Lett.*, 239, 253–265, 2005.

Narcisi, B., Petit, J.-R., and Tiepolo, M: A volcanic marker (92 kyr) for dating deep East Antarctic cores, *Quat. Sci. Rev.*, 25, 2682–2687, 2006.

NorthGRIP project members: High-resolution record of Northern Hemisphere climate extending into the last interglacial period, *Nature*, 431, 147–151, 2004.

Parrenin, F., Remy, F., Ritz, C., Siegert, M., and Jouzel, J: New modelling of the Vostok ice flow line and implication for the glaciological chronology of the Vostok ice core, *J. Geophys. Res.*, 109, D20102, doi:10.1029/2004JD004561, 2004.

Petit, J. R., Jouzel, J., Raynaud, D., Barkov, N. I., Barnola, J.-M., Basile, I., Bender, M., Chappellaz, J., Devis, M., Delaygue, G., Delmotte, M., Kotlyakov, V. M., Legrand, M., Lipenkov, V. Y., Lorius, C., Pepin, L., Ritz, C., Saltzman, E., and Stievenard, M: Climate and atmospheric history of the past 420 000 years from the Vostok ice core, *Antarctica, Nature*, 399(6735), 429–436, 1999.

Raisbeck, G. M., Yiou, F., Bard, E., Dollfus, D., Jouzel, J., and Petit, J. R: Absolute dating of the last 7000 years of the Vostok ice core using ^{10}Be , *Mineral. Mag.*, 62A, 1228–1228, 1998.

Raisbeck, G. M., Yiou, F., Cattani, O., and Jouzel, J: ^{10}Be evidence for the Matuyama-Brunhes geomagnetic reversal in the EPICA Dome C ice core, *Nature*, 444(7115), 82–84, 2006.

Raisbeck, G., Yiou, F. and Jouzel, J: Cosmogenic ^{10}Be as a high resolution correlation tool for climate records, *Geochimica Et Cosmochimica Acta*, 66(15A), A623, 2002.

Rasmussen, S. O., Andersen, K. K., Svensson, A. M., Steffensen, J. P., Vinther, B. M., Clausen, H. B., Siggaard-Andersen, M.-L., Johnsen, S. J., Larsen, L. B., Dahl-Jensen, D., Bigler, M.,

**The EDC3
chronology**

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Rothlisberger, R., Fischer, H., Goto-Azuma, K., Hansson, M. E., and Ruth, U: A new Greenland ice core chronology for the last glacial termination, *J. Geophys. Res.*, 111, D06102, doi:10.1029/2005JD006079, 2006.

Reimer, P., Baillie, M., Bard, E., Bayliss, A., Beck, J., Bertrand, C., Blackwell, P., Buck, C., Burr, G., Cutler, K., Damon, P., Edwards, R., Fairbanks, R., Friedrich, M., Guilderson, T., Hogg, A., Hughen, K., Kromer, B., McCormac, G., Manning, S., Ramsey, C., Reimer, R., Remmele, S., Southon, J., Stuiver, M., Talamo, S., Taylor, F., van der Plicht, J., and Weyhenmeyer, C: IntCal04 terrestrial radiocarbon age calibration, 0–26 cal kyr BP, *RADIOCARBON*, 46(3), 1029–1058, 2004.

Schulz, M.: On the 1470-year pacing of Dansgaard-Oeschger warm events, *Paleoceanography*, 17(2), 1014, doi:10.1029/2000PA000571, 2002.

Schwander, J., Jouzel, J., Hammer, C. U., Petit, J.-R., Udisti, R. and Wolff, E: A tentative chronology for the EPICA Dome Concordia ice core, *Geophys. Res. Lett.*, 28(22), 4243–4246, 2001.

Severinghaus, J. P. and Brook, E. J: Abrupt climate change at the end of the last glacial period inferred from trapped air in polar ice, *Science*, 286, 930–934, 1999.

Shackleton, N. J: The 100 000-year ice-age cycle identified and found to lag temperature, carbon dioxide, and orbital eccentricity, *tScience*, 289(5486), 1897–1902, 2000.

Siegenthaler, U., Stocker, T., Monnin, E., Luthi, D., Schwander, J., Stauffer, B., Raynaud, D., Barnola, J.-M., Fischer, H., Masson-Delmotte, V., and Jouzel, J: Stable Carbon Cycle–Climate Relationship During the Late Pleistocene, *Science*, 310, 1313–1317, 2005.

Spahni, R., Chappellaz, J., Stocker, T., Loulergue, L., Hausammann, G., Kawamura, K., Flückiger, J., Schwander, J., Raynaud, D., Masson-Delmotte, V., and Jouzel, J: Atmospheric Methane and Nitrous Oxide of the Late Pleistocene from Antarctic Ice Cores, *Science*, 310, 1317–1321, 2005.

Svensson, A., Andersen, K. K., Bigler, M., Clausen, H. B., Dahl-Jensen, D., Davies, S. M., Johnsen, S. J., Muscheler, R., Rasmussen, S. O., Röthlisberger, R., Steffensen, J. P., and Vinther, B. M: The Greenland Ice Core Chronology 2005, 15–42 kyr. Part 2: Comparison to other records, *Quat. Sci. Rev.*, 25, 3258–3267, 2006.

Tauxe, L., Herbert, T., Shackleton, N. J. and Kok, Y. S: Astronomical calibration of the Matuyama-Brunhes boundary: Consequences for magnetic remanence acquisition in marine carbonates and the Asian loess sequences, *Earth Planet. Sci. Lett.*, 140, 133–146, 1996.

- Traufetter, F., Oerter, H., Fischer, H., Weller, R., and Miller, H: Spatio-temporal variability in volcanic sulphate deposition over the past 2 kyr in snow pits and firn cores from Amundsenisen, Antarctica, *J. Glaciol.*, 50(168), 137–146, 2004.
- 5 Udisti, R., Becagli, S., Castellano, E., Delmonte, B., Jouzel, J., Petit, J.-R., Schwander, J., Stenni, B., and Wolff, E. W.: Stratigraphic correlations between the EPICA-Dome C and Vostok ice cores showing the relative variations of snow accumulation over the past 45 kyr, *J. Geophys. Res.*, 109(D8), D08101, doi:10.1029/2003JD004180, 2004.
- 10 Vinther, B. M., Clausen, H. B., Johnsen, S. J., Rasmussen, S. O., Andersen, K. K., Buchardt, S. L., Dahl-Jensen, D., Seierstad, I. K., Siggaard-Andersen, M.-L., Steffensen, J. P., Svensson, A., Olsen, J., and Heinemeier, J.: A synchronized dating of three Greenland ice cores throughout the Holocene, *J. Geophys. Res.*, 111, 13102, doi:10.1029/2005JD006921, 2006.
- Wolff, E., Cook, E., Barnes, P. and Mulvaney, R.: Signal variability in replicate ice cores, *J. Glaciol.*, 51(174), 462-468, 2005.
- 15 Wolff, E. W., Fischer, H., Fundel, F., Ruth, U., Twarloh, B., Littot, G. C., Mulvaney, R., Röthlisberger, R., de Angelis, M., Boutron, C. F., Hansson, M., Jonsell, U., Hutterli, M. A., Lambert, F., Kaufmann, P., Stauffer, B., Stocker, T. F., Steffensen, J. P., Bigler, M., Siggaard-Andersen, M. L., Udisti, R., Becagli, S., Castellano, E., Severi, M., Wagenbach, D., Barbante, C., Gabrielli, P., and Gaspari, V.: Southern Ocean sea-ice extent, productivity and iron flux over the past eight glacial cycles, *Nature*, 440, 491–496, 2006.
- 20 Yiou, F., Raisbeck, G. M., Baumgartner, S., Beer, J., Hammer, C., Johnsen, J., Jouzel, J., Kubik, P. W., Lestringuez, J., Stievenard, M., Suter, M., and Yiou, P.: Beryllium 10 in the Greenland Ice Core Project ice core at Summit, Greenland, *J. Geophys. Res.*, 102, 26 783–26 794, 1997.
- 25 Yuan, D., Cheng, H., Edwards, R. L., Dykoski, C. A., Kelly, M. J., Zhang, M., Qing, J., Lin, Y., Wang, Y., Wu, J., Dorale, J. A., An, Z., and Cai, Y.: Timing, Duration, and Transitions of the Last Interglacial Asian Monsoon, *Science*, 304(5670), 575–578, 2004.

**The EDC3
chronology**F. Parrenin et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Table 1. Age markers used for the construction of the EDC3 age scale. They fall into 3 categories: 1) Age markers used to control the poorly known parameters of the modelling; 2) Age markers used for a posteriori correction in the top part of the core (EDC3 is required to pass exactly through those age markers); 3) Age markers used to correct for ice flow anomalies in the bottom part.

| age marker | depth (m) | age (kyr BP) | error bar (kyr BP) | model control | top correction | bottom correction |
|--------------------------------|-----------|--------------|--------------------|---------------|----------------|-------------------|
| Krakatua | 8.35 | 0.066 | 0.001 | | X | |
| Tambora | 12.34 | 0.134 | 0.001 | | X | |
| Huaynaputina | 23.20 | 0.349 | 0.001 | | X | |
| Kuwae | 29.27 | 0.492 | 0.005 | | X | |
| El Chichon? | 38.12 | 0.691 | 0.005 | X | X | |
| Unidentified | 39.22 | 0.722 | 0.006 | | X | |
| Unknown | 41.52 | 0.779 | 0.006 | | X | |
| $^{10}\text{Be}/^{14}\text{C}$ | 107.83 | 2.716 | 0.05 | X | | |
| $^{10}\text{Be}/^{14}\text{C}$ | 181.12 | 5.28 | 0.05 | | X | |
| YD/Holocene | 361.5 | 11.65 | 0.18 | X | | |
| PB/BO | 427.2 | 15.0 | 0.24 | X | X | |
| ^{10}Be peak | 740.08 | 41.2 | 1 | X | X | |
| Mt Berlin erupt. | 1265.10 | 92.5 | 2 | X | | |
| term. II | 1698.91 | 130.1 | 2 | X | | |
| air content | 1082.34 | 70.6 | 4 | X | | |
| air content | 1484.59 | 109.4 | 4 | X | | |
| air content | 1838.09 | 147.6 | 4 | X | | |
| air content | 2019.73 | 185.3 | 4 | X | | |
| air content | 2230.71 | 227.3 | 4 | X | | |
| air content | 2387.95 | 270.4 | 4 | X | | |
| air content | 2503.74 | 313.4 | 4 | X | | |
| air content | 2620.23 | 352.4 | 4 | X | | |
| air content | 2692.69 | 390.5 | 4 | X | | |
| air content | 2789.58 | 431.4 | 4 | X | | |
| $^{18}\text{O}_{\text{atm}}$ | 2714.32 | 398.4 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2749.04 | 408.6 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2772.27 | 422.0 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2799.36 | 441.0 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2812.69 | 454.3 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2819.2 | 464.6 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2829.36 | 474.8 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2841.75 | 485.3 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2856.27 | 495.9 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2872.56 | 506.6 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2890.33 | 517.6 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2913.3 | 532.0 | 6 | | | X |

The EDC3 chronology

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

The EDC3
chronology

F. Parrenin et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

I◀

▶I

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Table 1. Continued.

| age marker | depth (m) | age (kyr BP) | error bar (kyr BP) | model control | top correction | bottom correction |
|------------------------------|--------------|-----------------|--------------------|---------------|----------------|-------------------|
| $^{18}\text{O}_{\text{atm}}$ | 2921.99 | 545.3 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2938.24 | 556.4 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2968.08 | 567.6 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 2998.96 | 578.6 | 6 | X | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3008.93 | 589.5 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3017.25 | 600.1 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3027.54 | 610.9 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3035.41 | 622.1 | 6 | X | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3043.01 | 634.4 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3048.51 | 649.1 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3056.77 | 660.8 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3065.93 | 671.7 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3077.74 | 682.3 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3093.51 | 693.2 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3112.43 | 704.0 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3119.57 | 714.4 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3124.27 | 724.4 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3136.18 | 733.9 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3143.2 | 741.9 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3152.25 | 749.2 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3158.91 | 758.1 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3166.87 | 767.7 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3174.81 | 777.6 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3180.6 | 787.7 | 6 | | | X |
| $^{18}\text{O}_{\text{atm}}$ | 3189.83 | 797.5 | 6 | | | X |
| B-M reversal | 3165 | 785 | 20 | X | | |

The EDC3 chronology

F. Parrenin et al.

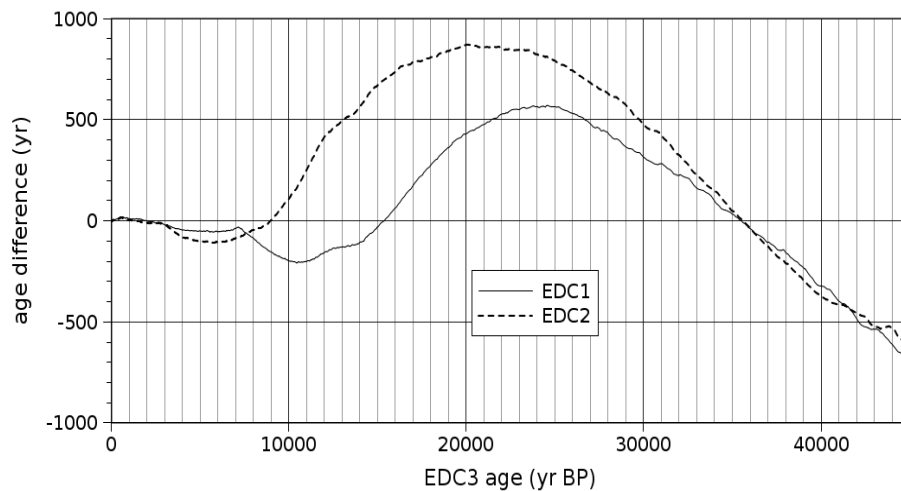


Fig. 1. Age difference between the EDC1 (resp. EDC2) and EDC3 time scales for the last 45 kyr.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**The EDC3
chronology**

F. Parrenin et al.

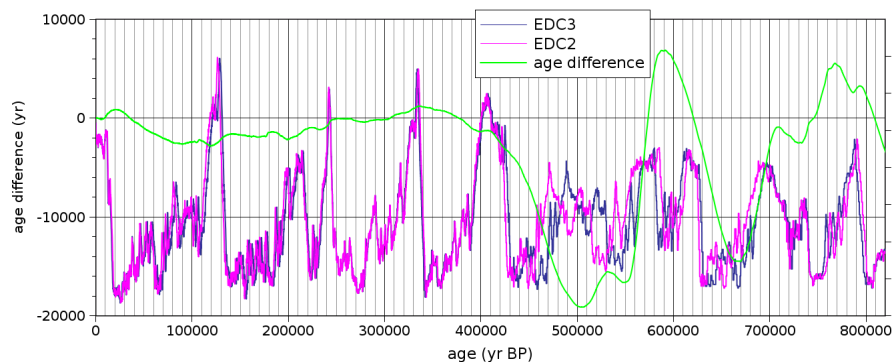


Fig. 2. Comparison of the EDC deuterium record on the EDC2 and EDC3 time scales. The green curve represents the difference in age between EDC2 and EDC3. Y-axes for isotopic records are normalised.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

The EDC3 chronology

F. Parrenin et al.

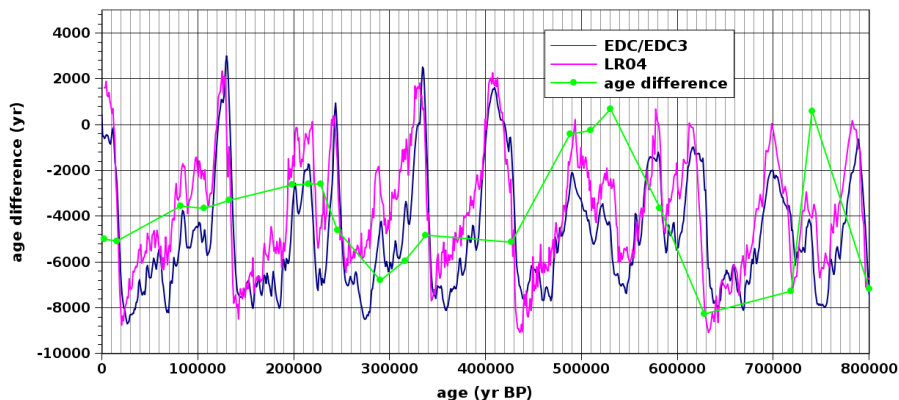


Fig. 3. Comparison of the EDC deuterium record on the EDC3 time scale with the LR04 marine stack on its own time scale, shifted by 2.5 kyr towards older ages. The green curve represents the difference in age between LR04 (+2.5 kyr) and EDC3 assuming both records are synchronous. Y-axes for isotopic records are normalised.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

The EDC3
chronology

F. Parrenin et al.

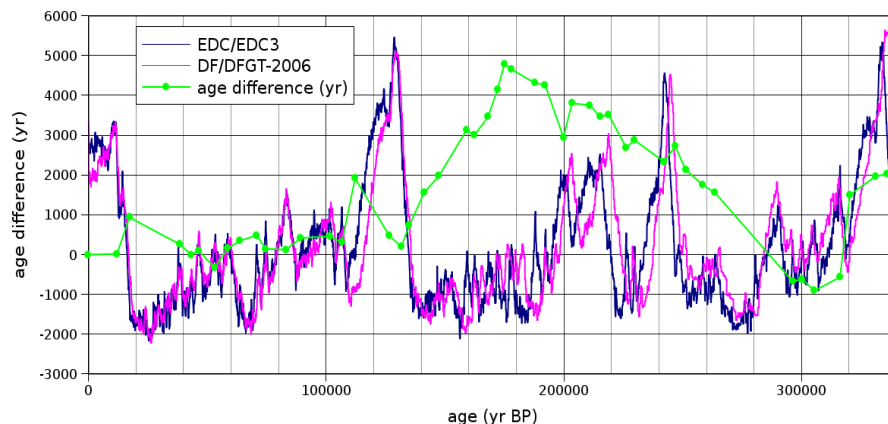


Fig. 4. Comparison of the EDC deuterium record on the EDC3 time scale with the Dome Fuji $\delta^{18}\text{O}$ record on the DFGT-2006 time scale (Parrenin et al., 2007¹). The green curve represents the difference in age between DFGT-2006 and EDC3 at the depth of the synchronisation markers. Y-axes for isotopic records are normalised.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

The EDC3 chronology

F. Parrenin et al.

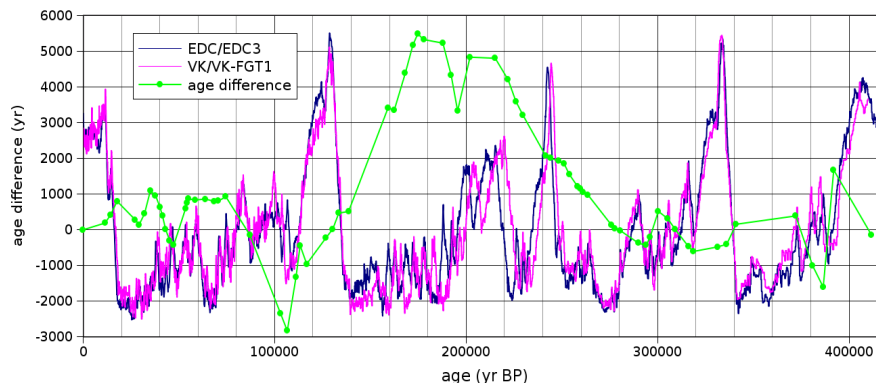


Fig. 5. Comparison of the EDC deuterium record on the EDC3 time scale with the Vostok deuterium record on the VK-FGT1 time scale (Parrenin et al., 2004). The green curve represents the difference in age between VK-FGT1 and EDC3 at the depth of the synchronisation markers. Y-axes for isotopic records are normalised.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**The EDC3
chronology**

F. Parrenin et al.

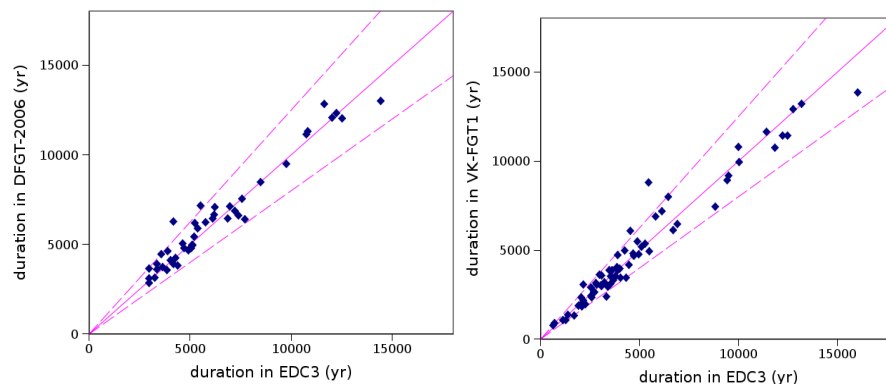


Fig. 6. Durations between two consecutive synchronisation markers in EDC3 compared to durations in DFGT-2006 (left panel) or in VK-FGT1 (right panel). Plain pink line is the 1:1 curve. Dashed pink lines represent the 1:0.8 and 1:1/0.8 lines.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)