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1	The climate reconstruction in Shandong Peninsula, North
2	China, during the last millennia based on stalagmite
3	laminae together with its a comparison to δ^{18} O
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Abstract

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Stalagmite ky1, with a length of 75 mm and the upper part (from top to 42.769 mm depth) consisting of 678 laminae, was collected from Kaiyuan Cave in the coastal areas of Shandong Peninsula, northern China, located at in a warm temperate zone in the East Asia monsoon area. Based on high precision dating with the U-230Th technique and continuous lamina counting of laminae, it can be confirmed that the 1st and 678th laminae are have been confirmed to be 1894±20_AD and 1217±20_AD from top to bottom, respectively. By the measurement of laminae thickness and $\delta^{18}O$ ratios, we have gotobtained the time series data of laminae thickness of laminae and δ^{18} O ratios from 1217±20 AD to 1894±20 AD, analyzed the climatic-environmental meanings of variations in the laminae thickness of laminae, variations which have a good correspondenceing relation with the cumulative departure curve of the drought-waterlog index in the historical period. The results shows that, in the ~678 years from 1217±20_AD to 1894±20_AD, both the laminae thickness of the laminae and the degree of fluctuation in the laminae thickness of the laminae fluctuation degree of stalagmite ky1 have obvious staged stages of variation characteristic, and are completely synchronized with the contemporaneous intensity of the summer monsoons intensity and precipitation as time changed. Among, tThere is a negative correlation between the laminae thickness of the laminae and the summer monsoon intensity and precipitation. On the other hand, tThere is a positive correlation between the degree of fluctuation in the laminae thickness of the laminae fluctuation degree and both the intensity of the summer monsoons intensity and the precipitation. Therefore, for the Kaiyuan Cave in the coastal area both of both the warm temperate zone and the East Asia monsoon area, the variations of in the laminae thickness of the laminae are not only relatived to the change of in the climatic factors themselves, but also relatived to the degree of climatic stability degree. For to achieve this, in the coastal area belonging to the warm temperate zone and the East Asia monsoon area, the climate change between the LIA (Little Ice Age) and the MWP (Medieval Warm Period), besides in addition to less precipitation and low temperatures, that is to say,(a type of dry and cold climate), also shows an obviously decreasing trend of in the degree of climatic stability degree.

4 5 **k**

Keywords

Little Ice Age, <u>laminae</u> thickness <u>of laminae</u>, <u>degree of</u> climatic stability <u>degree</u>, Kaiyuan Cave in Shandong Peninsula of CHINA, the coastal area <u>of in the</u> warm temperate zone, East Asia monsoon area

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1 Introduction

Calcareous speleothems, which have advantages in for precisely dating and 11 high-resolution sampling, are becoming one of the best geological record carriers 12 of for major climate changes (Burns et al., 2003; Cheng et al., 2009; Dykoski et al., 13 2005; Genty et al., 2003; Fairchild et al., 2006; Wang et al., 2001; Wang et al., 2008; 14 15 Qin et al., 1999; Yuan et al., 2004) and high resolution reconstruction of the paleoclimate environment (Committee 16 and on Surface Temperature Reconstructions for the Last 2,000 Years and National Research Council, 2006; 17 Fleitmann et al., 2003; Hou et al., 2003; McDermott et al., 2001; Paulsen et al., 2003; 18 Tan et al., 2003; Tan, 2007; Wang et al., 2005; Zhang et al., 2008). Besides In 19 addition to the most widely-dely used carbon_(C) and oxygen_(O) stable isotopes 20 and trace elements, laminae and the growth rate of stalagmites could also be used 21 22 as proxies for the paleoclimate environment. However, different authors haved very different climate and environment interpretations about relative to laminae thickness 23 of laminae based on different stalagmites from different climatic regions. For 24 instance, the stalagmite laminane were confirmed as annual laminae at in the 25 earliest studies (Baker et al., 1993), the structure of the laminane reflected the 26 intensity of the ancient rainfall (Baker et al., 1999), and there was a positive 27 correlation between the growth rate of stalagmites and precipitation_(Brook et al., 28 29 1999). However, there was a negative correlation between the growth rate of stalagmites and precipitation (Proctor et al., 2000; Proctor et al., 2002), there was a 30 responsive relationship between the growth rate of the stalagmites and the winter 31

temperature (Frisia et al., 2003), and the growth rate of the stalagmites was influenced by the vegetation density on the top of the cave (Baldini et al., 2005). There was a well-understood relationship between the speleothem growth rate and climate (Baldini, 2010; Mariethoz et al., 2012). The situations are is more complex in humid and semi-humid regions because other factors, such as drip rate, atmospheric Pco2 in the cave and the seasonality of the climate, may also affect speleothem growth rates (Cai et al., 2011; Duan et al., 2012). The research investigation of stalagmite laminae in the middle reach of the Yangtze River indicates that the thickness of stalagmite lamianelaminae may be regarded as the a substitute index of for the summer monsoon intensity of in East Asia (Liu et al., 2005). There was a good response relationship with between the variations in the thickness of the laminane thickness variations and the variations of in rainfall (Tan et al., 1997; Ban et al., 2005). There was a response relation ship between the growth rate of the stalagmites and the temperature in summer; therefore, so the laminae thickness of the laminae may be regarded as a substitute index of for East Asia monsoon intensity (Tan et al., 2004). The δ^{18} O record of ZJD-21 indicates that δ¹⁸O in the stalagmite was mainly influenced mainly by the rainfall amount of rainfall and/or the summer/winter rainfall ratio, with lighter lower values corresponding to wetter conditions and/or more summer monsoonal rains_(Kuo et al., 2011). The Wanxiang Cave WX42B record indicates that the stalagmite δ¹⁸O has recorded local/regional moisture change_(Li et al., 2011). The growth rate and the observed temperature had a significant positive correlation (Tan et al., 2013). The upper part of ky1(from the top to a depth of 42.769 mm-depth, 0-42.769 mm) consists of 678 continuous clearly transmitting annual laminae clearly, because because the transmitting laminae of the stalagmite ky1 are very similar to the annual

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mm) consists of 678 continuous <u>clearly</u> transmitting annual laminae <u>clearly</u>, <u>because</u> <u>because</u> the transmitting laminae of <u>the</u> stalagmite ky1 are very similar to the annual laminae of <u>Shihua Cave</u> in Beijing, and have <u>all the all of the typical characteristics</u> of the latter laminae, which consist of so-called <u>Nn</u>orthern type laminae (<u>Zhou et al.</u>, 2010). There are <u>clearly</u> very thin opaque laminae between stalagmite laminae <u>clearly</u>, but the calcite laminae were thick and transmitting between <u>the</u> stalagmite laminae (<u>Tan et al.</u>, 1999; <u>Tan et al.</u>, 2002). Because <u>of</u> stalagmite ky1, with a very

have stopped growing not long ago₂; ilts deposition time may be the past several centuries or one millennium, which has recorded the climatic-environmental information of the Shandong Peninsula since the late MWP (Medieval Warm Period), including the late MWP, the whole LIA (Little Ice Age); and the early CWP (Current Warm Period) (Lamp, 1965; Lamp, 1972; Matthews, 2005; Ogilvie and Jónsson, 2001). In this research, on the basis of high precision dating with the U-230Th technique, we have observed and measured the laminae thickness of the laminae and dated all the all of the laminae in the upper part of stalagmite ky1, obtained and researched the time series data one laminae thickness of laminae; and compared it these data with the time series data for both the efoxygen (O) stable isotope value and the drought-waterlog index, and we discussed the climatic and environmental evolution of the coastal part of the warm temperate zone as well as the East Asia monsoon area since the LIA, especially in the transition periods of MWP/LIA and LIA/CWP.

2 Geological setting and sample description

Stalagmite ky1 was collected in 2008_AD from Kaiyuan Cave (36°24'32.20"N, 118°02'3.06"E) in western Shandong Peninsula, the coastal area of northern China (Fig.1Fig. 1, 2). The cave is located in the northwest hilly area of Lushan Mountain in Zibo eCity, Shandong_Perovince, with an elevation of 175_m above sea level (a.s.l.) (Fig.2). As the largest peninsula in China, it-the Shandong Peninsula wais located between the Bohai eSea and the Yellow eSea, and in its western region, the Cambrian Middle Zhangxia formation (mainly the oolitic shale, shale in clip to thin-layer limestone, oolitic limestone, algal clot limestone) and the Ordovician ef Badou formation and Gezhuang formation (mainly for the gray-dark gray thick layer of mud, wafer-thin limestone, dolomitic limestone and marl), awere widely distributed with the thickness of 24-238_m, including its-the lower section integrated with the Gezhuang eGroup and the upper section disconformity in contacted with the Carboniferous Benxi formation) (Shandong Provincial Bureau of Geology &

Minerals, 1991), which awere the main components of the Lushan Mountain, Yishan Mountain, and Mengshan Mountain with the highest elevation (1108_m,1031 m and 1150_m respectivelym, respectively). According to field investigation, the landforms development of the carbonate rocks in montanic caves are well developed, there are many caves outcroppings on the surface, secondary carbonate sedimentary bodies are developing well with typical morphological characteristics.

Kaiyuan eCave developed in the dolomite of the Ordovicia Zhifangzhuang formation with the a total strata thickness of the strata of about approximately 110 m. The total length of the cave is 1280_m, the overall distribution is a northwest-southeast strike along with twists and turns, and the space width inside the cave is generally 2 to 8_m and can be up to 30m. At the top of the cave, the surface of the bedrock is covered by soil with a general thickness of 50-80_cm, and the biggest thicknessthickest soil was more than 1.0_m.; tThe soil types are calcareous rocky soil and drab soil (The sSoil and fertilizer Fertilizer workstation of Shandong Province, 1994). The area of Kaiyuan eCave is currently influenced by both summer and winter monsoons with annual precipitation of ~620 mm and an annual mean temperature of ~13°C, and Ssummer monsoons prevails during July and August, contributing to half of the annual precipitation (Fig. 3).

3 Analytical methods and data processing

3.1 Establishment of a Ttime scale establishment

The stalagmite ky1 is conical in shape and consists of very pure calcite-_(Fig. 4). The polished surface of the stalagmite and laminae observation of the laminae by microscope show that stalagmite ky1 hads no hiatus during the growing process. the upper part (0-42.769 mm) is made up of comprises 678 laminane overlain by continuous deposits. aAll lamianelaminae were typical transmitting annual laminane. The stalagmite ky1 lt-has 232Th concentrations ranging from 704.6±5.1 ppt to 1245.2±5.0 ppt (Table 1), which was conducted determined at the High-

precision Mass Spectrometry and Environment Change Laboratory (HISPEC) of the National Taiwan University using high precision dating with the U-²³⁰Th technique (Shen et al., 2002).

3.2 Measurement of the thickness of the Llaminae thickness measurement

First of all, tThe stalagmite ky1 was first cut along the growth axis, and a slice was picked selected from the profile of the stalagmite and then polished. Secondly, under the LEIKA DMRX microscope (magnification of 200x, eyepiece of 10x, objective of 20x), we used transmission light to observe characteristics of the laminae characteristic along the growth axis layer-b_by-_layer. Thirdly, we measured the thickness of 678 laminae along three different paths layer-_by-_layer, calculated the thickness of every one of the laminane thickness—on average according to the three data points for of the each of the laminae. Fourthly, we had dated every one of the laminae layer-_by-_layer and get_determined the time series data of for the thickness of the stalagmite laminae of the stalagmite thickness. At lastFinally, we contrasted the time series data and the δ^{18} O ratio data series, analyzed the paleoclimate environment characteristic of the different stages; and discussed the climatic-environmental meanings of the variations of in the thickness

of the laminae thickness.

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3.3 δ^{18} O isotopes test

Firstly, perpendicular to the growth axis and along the horizons position of 9.5 mm and 18.5 mm from the top, we collected respectively four 4 samples equally spaced in at 20 mm from the growth center equally spaced that were used for the Hendy test. Secondly, along the growing direction of growth, we collected a 4 mm depth×5 mm width×75 mm length stone strip along growing axis, and scraped 330 samples using medical scalpel from top to bottom with a sampling density of 7-8 samples/mm_(separation distance of 0.1296_mm on the average). In-From the 330 samples above, we chose 175 samples to measure their δ^{18} O ratios, basically following the principle of an interval test in order to avoid the mixed pollution between adjacent samples. Next, we confirmed the sedimentation time according to their horizons positions and formed the time series data of for δ^{18} O ratios. The δ^{18} O ratios were measured using an automated individual- carbonate reaction (Kiel) device coupled with a Thermo-Fisher MAT 253 mass spectrometer at the State Key Laboratory of Palaeobiology and Stratigraphy of the Nanjing Institute of Geology and Palaeontology, Chinese Academy of Sciences. Each powdered sample (~0.08 to 0.1 mg of carbonate) was reacted with 103% H₃PO₄ at 9090°€°C to liberate sufficient CO₂ for isotopic analysis. The standard used is NBS19, and one standard was analyzed with every ten samples. One sample out of ten was duplicated to check the replication. All isotope ratios are reported in per mil (%) deviations relative to the Vienna Peedee Belemnite (VPDB) standard in the conventional manner. The standard deviation (1σ) for replicate measurements on NBS-19 is <±0.10‰.

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4 Results and discussion

4.1 The <u>laminae</u> thickness of <u>the</u> stalagmite <u>laminae</u> and <u>the</u> dating results <u>of</u> dating

In the upper part (0-42.769 mm) of stalagmite ky1, the dating result of for ages

1 corrected in Table 1 show that the three samples in the horizons positions of the 6 mm, 15_mm and 25_mm are <u>dated at</u> 1761.9±20.3_AD, 1696.6±13.6_AD and 2 1556.4±13 .6 AD, respectively (Table 1). Altogethermong, there are 221 laminae 3 between the horizons positions of 6 mm and 25 mm, and their age intervals are 206 4 years according to the U-²³⁰Th dating results. † The difference of in age between the 5 laminae determined by counting and by U-230Th dating results is only 15 years. But 6 7 However, there are 109 laminane between the horizons positions of 6 mm and 15 mm, and their age intervals are 65 years according to the result of the U-230Th 8 9 dating., and tThere are 112 laminae between the horizons positions of 15 mm and 25 mm, and their age intervals are 141 years according to the results of U-230Th 10 dating... On the other hand, ilf we use the horizon position of 6 mm as a datum to 11 for calculatione, the ages of the 1st and 678th laminae are 1894±20.3 and 1217±20.3 12 AD, respectively. If we use the horizon position of 25 mm as a datum to for 13 calculatione, the ages of the 1st and 678th laminae are 1909±13.6 AD and 1232±13.6 14 AD, respectively, the age intervals are only 14 years differents. Finally, in 15 considering of the error of the measurement of the thickness of the laminae 16 17 thickness measurement accumulating downward layer by layer, we chose the 133th-133rd of the laminae corresponding to the horizon position of 6 mm as the a 18 datum to calculate the age of the other laminane in the upper part of stalagmite ky1. 19 The results show that the deposition times of the 1st and 678th laminae are 20 21 1894±20.3 and 1217±20.3_AD_(the dating error is ±20.3 years, similarly hereinafter for the AD ages in this paper), respectively, the age of the other laminane were 22 calculated by analogy., hereby Thus, we got obtained the time series data of for the 23 thickness of the laminane thickness of stalagmite ky1_(Fig. 5). 24

4.2 Characteristics of the Laminae shape of the laminae characteristic

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Stalagmite ky1 <u>obviously</u> developed continuous transmitting laminae obviously (Fig._4). Under the microscope, first_ly, the laminae thickness <u>of the laminae</u> were was rather changeable_-, <u>tT</u>he maximum thickness <u>reached was more than 800 μm, and the minimum thickness reached was less than 15 μm (Fig._6a). Because the</u>

variations in the thickness of the laminae thickness variations may corresponding 1 to the climatic environmental changes when the laminane were growing, it shows 2 the potential value of these transmitting laminane in for reconstructing the 3 paleoclimate environment is illustrated (Genty et al., 1996; Baker et al., 1999; Tan 4 et al., 2004; Ban et al., 2005; Liu et al., 2005; Zhang et al., 2008; Muangsong et al., 5 2014; Liu et al., 2015). Secondly, most of the boundariesy of the lamiane laminae 6 7 are straight, but some lamianelaminae are obviously curved (Fig._6b). Hereby wWhen we analyzed the climatic-environmental meanings of the thickness of the 8 9 stalagmite laminae thickness, we acquired the laminae thickness values of a-the same laminae in different paths and calculated their average values along multiple 10 paths in order to getdetermine the substituted index information of for climatic-11 environmental change that had statistical significances. Thirdly, colors in some of 12 the boundaries of the transmitting lamiane aminae were deeper obviously deeper 13 14 (Fig. 6c). † These laminae had a special structure similar to supera annual laminae. This special structure may indicate that climatic-environmental changes not only 15 have seasonal changes, but also have multi-interannual changes. Fourthly, the light 16 transmission of some transmitting lamianelaminae is obviously different from the 17 light transmission of adjacent lamianelaminae:, the color is deeper, and there are 18 dark spots (Fig._6a, d). Whether these dark laminae have some mineralogy and 19 geochemistry characteristics different from other transmitting laminae, and 20 21 what their climatic-environmental meanings significance are may be, these dark laminae may need further and special research in the future. 22

4.3 Variations in the thickness of the Llaminae thickness variations

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The variation-range of variation in the thickness of the 678 laminae thickness of stalagmite ky1 (upper part) were 13.03~872.8_µm._The formed age determined of the maximum thickness (872.8_µm) of the laminae was 1551_AD, the formed age determined for of the minimum thickness (13.03_µm) of the laminae was 1245 AD, and the average value of for all laminae were was 63.08_µm (Fig._7a). In the 678 years from 1217_AD to 1894_AD, the laminae thickness of the laminae of from

1 stalagmite ky1 have obvious stagesd of variation.s, Stalagmite ky1 it had undergone the transition from low values to high values and again to low values, and both the 2 laminae thickness of the laminae and the fluctuationg degree of variations of in the 3 laminae thickness of the laminae had obvious stagesed of variations (Fig. 7a). From 4 1217 AD to 1471 AD, it was the low value period of laminae thickness of the laminae 5 with the an average value of 46.08 µm. Among Then, the period from 1217 AD to 6 7 1372 AD was a relatively low fluctuation period relatively, the period from 1372 AD to 1471 AD was a period of relatively high fluctuation. period relatively, the two 8 9 periods above presented the trend of rising firstly and then falling then. From 1471 AD to 1744_AD, it was a period of high value-high fluctuation period of the laminae 10 thickness of the laminae, with the average value of 88.8307 µm. Among, tThis period 11 could be divided into three secondary high value-high fluctuation periods, 1471 AD-12 1548_AD, 1548_AD-1637_AD and 1637_AD-1744_AD____, e<u>E</u>very period <u>has shows</u> the 13 14 trend of increasing firstly and then decreasing. then, the average values of for the thickness of the laminae thickness were 82.2027 µm, 82.5491 µm and 98.8252 µm, 15 successively. From 1744_AD to 1894_AD, it_there was a period of relatively low 16 17 values period of the laminae thickness of the laminae, with a group of peak values appeared appearing in about approximately 1776 AD and with the an average value 18 of 45.1164_um. Among, tThe period from 1217_AD to 1372_AD was a period of 19 <u>relatively</u> low fluctuation. <u>period relatively</u>, tThe period from 1744_AD to 1831_AD 20 21 was a <u>period</u> of relatively high fluctuation. <u>period relatively</u>, <u>t</u>The two periods above presented the trend of rising firstly and then falling. then, tThe period from 1831 AD 22 to 1880 AD was a period of relatively high fluctuation, without period relatively and 23 not have a trend of obviously rising or falling. obviously, it This period of rising was 24 25 a short, rising period from 1880_AD to 1894_AD.

4.4 <u>Variations in the δ^{18} O ratio variations</u>

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The variation range of δ^{18} O ratios of <u>in</u> the 172 samples above was -6.247%--8.599%, <u>with</u> the maximum value_(-6.247%) appeared appearing in 1603_AD, and the minimum value (-8.599%) appeared appearing in 1460 AD., the value of

1 all of the samples was -7.674% on average (Fig. 7c). In the 678 years from 1217 AD to 1894 AD, δ^{18} O ratios had obvious stagesd of variation. s, it-The ratios had 2 undergone the a transition from low values to high values and again to low values, 3 and both the δ^{18} O ratios and the <u>degree</u> of fluctuation degree of δ^{18} O ratios had 4 obvious stagesd of variation s (Fig. 7c). From 1217 AD to 1480 AD, it there was a 5 period of low values period of δ^{18} O ratios with an average value of -8.104%. 6 Among, tThe period from 1217_AD to 1384_AD was a period of relatively low 7 fluctuation. period relatively, this period hads a trend of decreasing slowly. in total, 8 9 t The period from 1384 AD to 1480 AD was a period of relatively high fluctuation period relatively, and this period presented showed the trend of rising firstly and then 10 falling then. From 1480_AD to 1746_AD, it was the a period of high value-high 11 fluctuation period with the an average value of -7.301%. Among, tThis period could 12 be divided into three secondary high value-high fluctuation periods: 1480_AD-1542 13 14 AD, 1542 AD-1633 AD and 1633 AD-1746 AD., eEvery secondary period hads the trend of increasing firstly and then decreasing then or decreasing firstly and 15 then increasing. then, tTheir inflection points appeared in the ages of 1498 AD, 1603 16 AD and 1663 AD, respectively. Their average values of the δ^{18} O ratios were 17 -7.393%,--6.953% and -7.513%, successively. From 1764_AD to 1894_AD, it 18 was a low value period with an average value of -8.199\\(\text{-}\)Among, tThe period 19 from 1746_AD to 1831_AD was a period of relatively high fluctuation period 20 21 relatively, tThis period presented showed a trend of rising firstly and then falling. then, tThe period from 1831 AD to 1880 AD was a period of relatively low fluctuation 22 period relatively and did not have a trend of obviously rising or falling. obviously, it 23 There was a short rising period from 1880 AD to 1894 AD. 24

4.5 Drought/waterlog index variations

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In order tTo show the relationship between the variations of in the laminae thickness of the laminae, and the δ¹8O ratios and the changes of in climate, we calculated cumulative departure values of for the drought/water log index in the area of Kaiyuan Cave from 1470 AD to 1894 AD. The data source iwas the Yearly

1 eCharts of dDryness/wWetness in China for the Last 500-year Period., the charts are compiled by the Chinese aAcademy of mMeteorological Sciences of the China 2 Meteorological Administration according to lots extensive of Chinese historical 3 literature and published by the China Cartographic Publishing House (Chinese 4 aAcademy of mMeteorological sSciences of the China 5 Administration, 1981). In the charts, the degree of drought/waterlog— is represented 6 7 by the drought/waterlog index which that has five values including 1, 2, 3, 4 and 5, with 1 representing the waterlog and 5 representing drought, and its distribution is 8 9 represented through the index isolines. On the basis of Yearly- eCharts of <u>dD</u>ryness/<u>wW</u>etness in China for the <u>|Last 500-yYear pPeriod</u>, we acquired the 10 drought/waterlog indexes indices in for the area near Kaiyuan eCave according to 11 its geographical coordinates, and we checked the drought/waterlog indexes indices 12 again refer<u>ring</u> to <u>the</u> local chronicles. We dreaw a cumulative departure curve from 13 14 1470 to 1894_AD with a rising trend representing the changes of associated with becoming dryer and a declining trend representing the change of associated with 15 becoming waterlogged (Fig. 7b). Based on the cumulative departure 16 17 curve, it there was a period of less precipitation period in this area from 1480 to 1744 AD.; tThis period is startsed with the transition of MWP/LIA and endsed with the 18 transition of LIA/CWP.; its The primary fluctuations of this period were corresponding 19 to the curve of the thickness of the laminae thickness. (Fig. 7b). So tThe high value-20 21 high fluctuation period of the laminae thickness of stalagmite ky1 laminae above occurred under the background of drought and less precipitation. On the other 22 hand However, there is a correlation between the δ^{18} O ratios of stalagmite ky1 and 23 the change of in the summer monsoon intensity and precipitation (Cheng et al., 24 25 2009). This indicatinges that stronger So, there is a correlation between the summer monsoon intensity/precipitation and the growth of stalagmites, weaker summer 26 monsoon intensity and less precipitation may be of benefit to the growth of 27 28 stalagmites growing in LIA.

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laminae thickness variations

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Owing to Because of the difference of in homologous laminae thickness stages 2 of the laminae and δ^{18} O varios ratios stage ranged ranging from 2 years to 14 years, 3 in consideration of the error of the dating technique was ± 20 years (the time series 4 data we got infrom section 4.1) and the resolution of the δ^{18} O sample was 3.9 years, 5 we could say the two synchronize with time-_variationying, mean that i.e., the low 6 7 value period and the high value period of the δ^{18} O ratios were corresponding to the low value period and the high value period of the thickness of the stalagmite laminae. 8 9 thickness, tThe low fluctuation period and the high fluctuation period of for the δ^{18} O ratios were corresponding to the low fluctuation period and high fluctuation period 10 of thickness of stalagmite laminae thickness (Fig. 7a,ca, c). On the other hand, tThe 11 analysis result of for the δ^{18} O varios variations showed that, δ^{18} O ratios of for the 12 four4 samples were -7.506\%, -7.753\%, -7.981\% and -7.691\% which for the 13 14 samples that were collected at a 9.5 mm distance from the top of the stalagmite and the 5, 10, 15 and 20 mm distance from the growing axis of growth, - respectively. 15 The δ^{18} O ratios of for the four4 samples that were collected at an 18.5 mm distance 16 from the top of the stalagmite were -6.571%, -6.671%, -6.540% and -6.542%. 17 which were collected 18.5mm distance from the top of stalagmite and At 5, 10, 15 18 and 20_mm distances from the growing axis of growth, respectively, and the δ^{18} O 19 ratios were similar in-for the same laminae (Table 2). Hence, the Hendy Test carried 20 21 out about for Ky1 ky1 indicates that calcite in Kky1 should be deposited under isotopic equilibrium conditions.; tThe possibility of its-the dynamic fractionation of 22 the calcite in the sedimentary process is small; therefore, so the stalagmite δ^{18} O 23 mainly mainly reflectsed mainly the original external climate signal (Hendy, 1971). 24 25 Therefore, the stalagmite δ^{18} O can be used to collect and reconstruct the information onf climate change (Tan et al., 2009; Kuo et al., 2011; Li et al., 2011; Tan et al., 2013; 26 Liu et al., 2015). 27 The obvious synchronization relationship between the variations in the laminae 28 thickness of the laminae variations and the δ^{18} O ratios variations of in stalagmite 29

ky1 shows a closely relationship between the variations in the -deposition rate

variations of the stalagmite and climate change (Fig. 7). Because Kaiyuan Cave is located at in a warm temperate zone influenced by the East Asia monsoon, its rainy season coincides with high temperatures. The precipitation, carried by the summer monsoon from the low latitude of the Pacific Ocean of low latitude, concentrates in summer. However, when the winter monsoon from the interior Asian continent with at a high latitude is prevailsing, there is rare precipitation. In this research, we interpreted the climatic meanings of the stalagmite ky1 δ^{18} O ratios, based on the relationship between the cumulative departure of the drought/waterlog index and the curves of the δ^{18} O ratios. The characteristics of contemporary warm temperate weather, also referring to the assumption of the Asia monsoon intensity by Cheng et al. (2009) and the precipitation as is assumed by Zhang et al. (2008) about the climatic meanings of stalagmite δ^{18} O records, with lower δ^{18} O ratios representing a stronger summer monsoon and higher δ^{18} O ratios representing <u>a</u> weaker summer monsoon, that is to say, the $\delta^{18}O$ ratios are anti-correlating correlative with precipitation (Fig._7). Hereby, it There was a strong summer monsoon-more precipitation period from 1217_AD to1480_AD, a weak summer monsoon-less precipitation period from 1480 AD to 1746 AD, and a strong summer monsoon-more precipitation period again from 1746 AD to 1894 AD again. On the other hand, tThe degree of fluctuation degree of the summer monsoon intensity and precipitation is not the same or similar in different periods. As a whole, the <u>degree</u> of fluctuation degree-was lower when the summer monsoon was stronger and the precipitation was moregreater., tThe degree of fluctuation degree was higher when the summer monsoon was weaker and the precipitation was less. Among, tThe period from 1217 AD to 1480_AD can be divided into one low fluctuation period and one high fluctuation period. The period from 1480 AD to 1746 AD can be divided into three high fluctuation periods. The period from 1746 AD to 1894 AD was included a high fluctuation period, a low fluctuation period and a weaker-less fluctuation period, successively.

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According to the <u>thickness</u> of the laminae thickness and the δ^{18} O record of stalagmite ky1, the <u>laminae</u> thickness of the <u>laminae</u> and both summer monsoon

intensity and precipitation haves a negative correlation. The higher value period of 1 laminae the thickness of the laminae is correspondsing to weaker summer 2 monsoon-less precipitation, and the lower value is correspondsing to stronger 3 summer monsoon-more precipitation. On the other hand, tThe thickness of the 4 laminae thickness and the degree of fluctuation degree of the summer monsoon 5 intensity-precipitation have a positive correlation. The period of the higher values 6 7 period offor the laminae thickness of the laminae is correspondsing to a high degree of fluctuation degree of the summer monsoon intensity-precipitation, and a lower 8 9 value is correspondsing to a low degree of fluctuation degree of in the summer monsoon-precipitation. Therefore, Kaiyuan Cave, in the coastal area both of a 10 warm temperate zone and the East Asia monsoon area, demonstrates that the 11 variations in the thickness of the laminae thickness are not only relative to the 12 summer monsoon intensity-precipitation, but also relative to their degree of 13 14 fluctuation degree. This is because karstic water cyclesd faster and residence time wais shorter in the fracture of rock., the dissolution was insufficient and weak, so; 15 therefore, the deposition rate and the thickness of the laminae thickness of from the 16 stalagmite wereas low in the period with more precipitation period. But However, in 17 the period of less precipitation period, the karstic water cycled slower, and the 18 residence time was longer in the fracture of the rock; the dissolution was sufficient 19 and strong; therefore, so the deposition rate and the thickness of the laminae 20 21 thickness of the stalagmite were high. However, karstic water would be reduced or dry up if the period of less precipitation lasted for a long time., this The period of less 22 precipitation is also bad for water dissolution and growth of the stalagmite laminae 23 growing. So uUnder the background of weaker summer monsoons and less 24 25 precipitation, the <u>degree</u> of fluctuation <u>degree</u> of the summer monsoon intensityprecipitation goes becomes higher, this is beneficial to increasing the average value 26 of the thickness of the laminae thickness of the stalagmite, but the degree of the 27 28 fluctuation degree also goes becomes higher. Because of the degree of fluctuation degree of the summer monsoon intensity-precipitation reflecting the degree of 29 30 climatic stabilized stabilization degree, according to both the thickness of the laminae thickness and the δ18O record of stalagmite ky1 from the Kaiyuan Cave, the climate change between MWP and LIA in the coastal area both of both a warm temperate zone and the East Asia monsoon area, besides in addition to less precipitation and a lower temperature, also shows that the degree of climatic stability obviously degree decreased obviously.

5 Conclusions

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The upper part of stalagmite ky1 (0-42.769_mm) clearly consists of 678 continuously transmitting annual laminae. clearly, its The deposition time of deposition is ranges from 1217 ± 20 AD to 1894 ± 20 AD; therefore, so the laminae contains the climatic-environmental change information of for the late MWP, the whole LIA and the early CWP. The analysis shows that both the variations of in the thickness of the laminae thickness themselves and the fluctuating degree of variations of in the thickness of the laminae thickness fluctuation degree of stalagmite ky1 have obviously staged characteristics from 1217_AD to 1894_AD_; # Both the variations in the thickness of the laminae themselves and the fluctuating degree of variation in the thickness of the laminae of stalagmite ky1 had undergone the transition from low values to high values and again to low values, which was synchronized with the contemporaneous variations of in the δ^{18} O ratios and the degree of fluctuation degree of the δ^{18} O ratios. According to the comparison among the thicknesses of the laminae thickness, the drought/waterlog index and the synchronous δ¹⁸O ratios of stalagmite ky1, the thickness of the laminae thickness and the summer monsoon intensity-precipitation have a negative correlation. The higher value periods of the thickness of the laminae thickness are corresponding to weaker summer monsoon-less precipitation, and low value periods are corresponding to stronger summer monsoon-more precipitation. On the other hand, tThe thickness of the laminae thickness and the degree of fluctuation degree of the summer monsoon intensity-precipitation have a positive correlation. The higher value periods of thickness of the laminae thickness are corresponding to a high degree of fluctuation degree of summer monsoon intensity/precipitation, and the lower value periods are corresponding to a low degree of fluctuation degree of in the summer monsoon-precipitation. Therefore,__Kaiyuan Cave, in the coastal area both of a warm temperate zone and the East Asia monsoon area, with the relationship between the variations of the laminae thickness and climate change, besides in addition to the effects of climate factor variations like such as temperature and precipitation on the thickness of the laminae thickness, also reflects closely the <u>degree</u> of <u>fluctuation</u> degree of the summer monsoon intensity and the degree of climatic stability degree in addition. As a On the whole, it-there was a period of stronger summer monsoons from 1217 AD to 1470 AD. Among, the climatic stability was high from 1217_AD to 1370_AD firstly, and was reduced from 1370_AD to 1470_AD. From 1470_AD to 1740_AD, it there was a period of weaker summer monsoon-lower degree of stability that degree period, could be divided into three secondary periods with a trend of stronger firstly and then weaker then or weaker firstly and then stronger then divided by 1550_AD and 1640_AD. Since 1640_AD, the summer monsoon has again entered a strong period again. Among, tThe degree of stability degree was high from 1740 AD to 1830 AD, and the degree of stability wadegree is reduced from 1830 AD to 1880 AD., tThe summer monsoon became weaker for a short time since 1880 AD.

The conclusions of this research can enrich the acquaintance knowledge about the climatic-environmental meaning of the thickness of the laminae thickness of a stalagmite, and contribute to the comprehensiond of the specific manifestation of the MWP and LIA in the coastal area both of a warm temperate zone and the East Asia monsoon area of northern China, especially the transition time of MWP/LIA and the lasting time period of that the LIA lasted and the climatic characteristics of the LIA, and that may also deepen the research of into the climate change in the Asian summer monsoon area based on the secondary carbonate record in the karst cave.

Acknowledgements

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Table 1. U-series isotopic results and ages for stalagmite ky1 from Kaiyuan Cave, Shandong peninsula, Northern China.

Sample ID	1	2	3
Dist. from top (mm)	6.0	15.0	25.0
²³⁸ U ppb ^a	347.47± 0.63	434.45± 0.92	334.58± 0.61
²³² Th ppt	1245.2± 5.0	959.9± 4.9	704.6± 5.1
δ^{234} Umeasured	1457.9± 5.5	1341.2± 5.1	1320.3± 4.6
[²³⁰ Th/ ²³⁸ U] activity ^c	0.00652± 0.00014	0.00732± 0.00011	0.01021± 0.00013
[²³⁰ Th/ ²³² Th] ppm ^d	30.0 ± 0.68	54.63± 0.89	79.9± 1.2
Age uncorrected BP ^f	289.6± 6.5	341.4± 5.4	480.6± 6.3
Age corrected ^{c,e} BP ^f	251.1± 20.3	316.4± 13.6	456.6± 13.6
Age corrected ^{c,e} AD	1761.9± 20.3	1696.6± 13.6	1556.4± 13.6
$\delta^{234} U_{\text{initial}} corrected^b$	1458.9± 5.5	1342.4± 5.1	1322.1± 4.6

Chemistry was performed on July. 8, 2013 with the analysis method of Shen et al. (2003), and instrumental analysis on MC-ICP-MS (Shen et al., 2012) \pm . Analytical errors are 2σ of the mean.

 $a[^{238}U] = [^{235}U] \times 137.818 \ (\pm 0.65\%) \ (Hiess et al., 2012)^{\frac{1}{2}} \delta^{234}U = ([^{234}U/^{238}U]_{activity} - 1) \times 1000.$

 $^b\delta^{234}$ U_{initial} corrected was calculated based on 230 Th age (T), i.e., δ^{234} U_{initial} = δ^{234} U_{measured}X e $^{\lambda 234^{\circ}T}$, and T is the corrected age.

 $^{\circ}$ [230Th/238U]_{activity} = 1 - $^{\circ}$ e^{- λ 230T} + (δ ²³⁴U_{measured}/1000)[λ 230/(λ 230 - λ 234)](1 - $^{\circ}$ e^{-(λ 230 - λ 234) $^{\circ}$), where $^{\circ}$ is the age. Decay constants are 9.1705 x 10⁻⁶ yr⁻¹ for ²³⁰Th, 2.8221 x 10⁻⁶ yr⁻¹ for ²³⁴U (Cheng et al., 2013, EPSL), and 1.55125 x 10⁻¹⁰ yr⁻¹ for ²³⁸U (Jaffey et al., 1971).}

^oThe degree of detrital ²³⁰Th contamination is indicated by the [²³⁰Th/²³²Th] atomic ratio instead of the activity ratio.

 $^{\rm e}$ Age corrections for samples were calculated using an estimated atomic 230 Th/ 232 Th ratio of 4 ± 2 ppm. Those are the values for a material at secular equilibrium, with the crustal 232 Th/ 238 U value of 3.8. The errors are arbitrarily assumed to be 50%.

^fBP_(Before Present)², "present" in this table refers to 2013 AD.

- Table- 2. The results of the Hendy tests conducted along two growth lamianelaminae of ky1 at depths of 9.5 mm
- 2 and 18.5 mm individually, which indicate that calcite in ky1 was deposited under isotopic equilibrium conditions
- according to the Hendy Test rules (Hendy,1971).

Sample <u>nN</u> umber	Distance from the <u>*T</u> op/	Distance from the eCenter of gCrowth/	δ ¹⁸ Ο/ ‰
Sample HIV	mm	mm	0 0//00
KY1-9/10-5	9.5	5.0	-7.506
KY1-9/10-10		10.0	-7.753
KY1-9/10-15		15.0	-7.981
KY1-9/10-20		20.0	-7.691
KY1-18/19-5		5.0	-6.571
KY1-18/19-10		10.0	-6.671
KY1-18/19-15	18.5	15.0	-6.540
KY1-18/19-20		20.0	-6.542

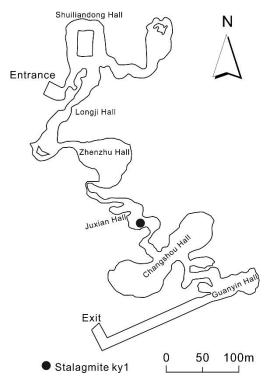


Fig.1Fig. 1. The map of Kaiyuan Cave. The black point is the location where we collected the sample in the Cave. The cave has an entrance and an exit, and consists of six6 small malls.



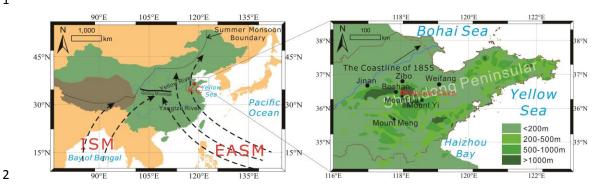
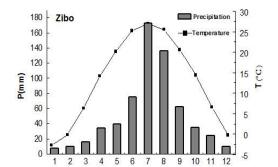


Fig._2. Location of Kaiyuan Cave and Shandong Peninsula in monsoonal China. KC: Kaiyuan Cave (36°24′32.20″N, 118°02′3.06″E). ISM: India Summer Monsoon; EASM: East Asia Summer Monsoon. The dashed black thin line indicates the northwestern boundary of the Asian summer monsoon. The dashed black lines with arrows indicate the routes of the summer monsoon. The dashed black lines with arrows ion the left indicate the routes of the summer monsoon. The brown area is the Qinghai-Tibet Plateau. The green area is China, and the yellow area is the other area.





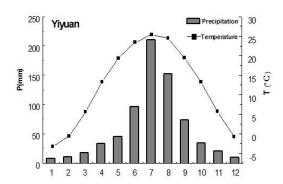


Fig. 3. Monthly mean temperature (T) and precipitation (P) atof Zibo at Zibo Station (1952-1980) and Yiyuan (1958-2005) at the Yiyuan Station, two meteorological stations close to the study site (Fig. 1).



Fig. 4. Polished longitudinal cross-section of stalagmites ky1

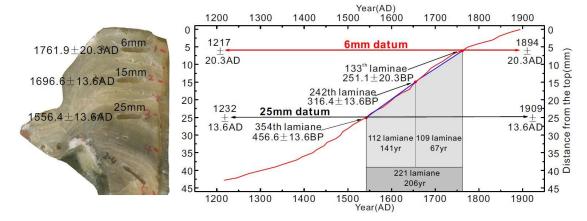
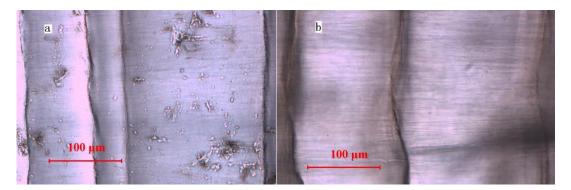


Fig._5. The age model for stalagmite ky1 established by laminae counting of laminae and high precision dating results with the U-230Th technique. It-This figure is the photo of stalagmite ky1, and the age label was based on high precision dating results with the U-230Th technique ion the left. The blue line is the high precision dating results with the U-230Th technique and their connecting lines. The red line is the age scale established by this article. The age of other laminaes were determined by annually laminae counting upward and downward based on the 133th-133rd of the laminae corresponding to the horizon-position of 6 mm, the age of which is 1762±20.3 AD that is decided by high precision dating results with the U-230Th technique.



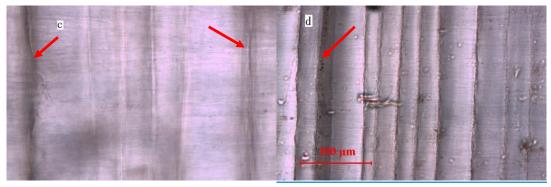


Fig._6. The characteristics of the transmitting laminae in the upper part of stalagmite ky1, show that the thickness of the laminae thickness hasve obvious variations. The boundary was curved, and the color near the boundary was deeper and because of the dark transmitting laminae. The thickness of the Leaminae thickness haveshows obvious variations (a), the curve of the boundary of transmitting laminae (b), the color variations of the boundary of transmitting laminae, the arrows indicating the darker boundaries, the boundaries in the middle were obviously whiter obviously—(c), dark transmitting laminae (d) (the arrows indicated in the figure).

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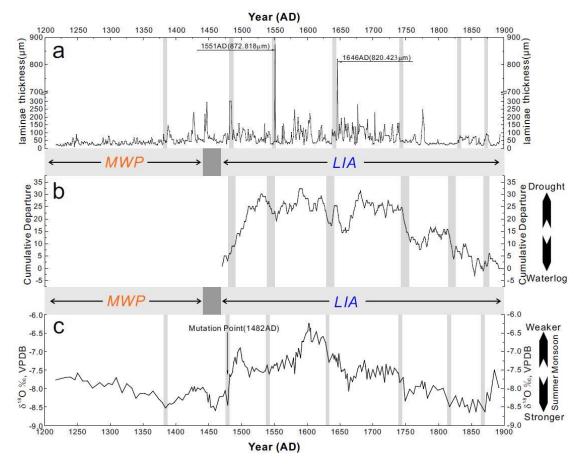


Fig._7. The formed year of formation and the thickness data series of the 678 laminae in the upper part (0-42.769 mm) of stalagmite ky1 (a), the cumulative departure curve (b) and the δ^{18} O ratio data series of for 172 samples (c). The thickness of the laminae formed in 1551 AD and 1646 AD were up to 872.818 µm and 820.423 µm, respectively, they are, much higher than other laminae. The cumulative departure curve (b) is drown drawn by drought/waterlog indexes indices on the basis of the Yearly charts of Dryness/Wetness in China for the Last 500yYear pPeriod (Chinese aAcademy of mMeteorological sSciences of the China Meteorological Administration, 1981), the curve has a rising trend representing less precipitation and the climate becoming drier, and the curve has a declining trend representing precipitation climate becominges more and the waterloggingwaterlogged.