

Professor Alan M. Haywood Director of Research and Innovation School of Earth and Environment University of Leeds Leeds LS2 9JT West Yorkshire, UK

February 15, 2016.

Response to the editor

Dear Wing-Le,

Please find the uploaded and revised version of the PlioMIP2 experimental design paper for CP.

As well as undertaking a number of very minor changes (as requested by the reviewers and which are highlighted in the track changes version of the document), we have completed the following more substantial changes:

1) We have added a map of modern lakes distribution to Figure 5

2) We have added a map of modern soil distribution to Figure 6

3) We have modified the list of experiments in line with our response to reviewer 2 in Table 3 and these changes have followed through to the manuscript.

4) We have enhanced the identification of the core experiments in Table3 and added the experimental nomenclature to Figure 2.

5) We have added a new figure to the supplementary information to enhance the description of the forcing factorisation experiments.

We assess that these changes are adequate to address the substantive issues of scientific relevance raised by the reviewers.

Sincere regards,

Harson

Alan Haywood (for all co-authors).



The Pliocene Model Intercomparison Project (PlioMIP) Phase 2: Scientific Objectives and **Experimental Design**

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5	Alan M. Haywood ^{1*} , Harry, J. Dowsett ² , Aisling M. Dolan ¹ , David Rowley ³ , Ayako
6	Abe-Ouchi ⁴ , Bette Otto-Bliesner ⁵ , Mark A. Chandler ⁶ , Stephen J. Hunter ^{1,} Daniel
7	J. Lunt ⁷ , Matthew Pound ⁸ , Ulrich Salzmann ⁸ .
8	
9	1. School of Earth and Environment, University of Leeds, Woodhouse Lane, Leeds, LS29JT,
10	UK: *earamh@leeds.ac.uk
11	2. Eastern Geology & Paleoclimate Science Center, U.S. Geological Survey, MS 926A, 12201
12	Sunrise Valley Drive, Reston, VA 20192, USA
13	3. Department of Geophysical Sciences, University of Chicago, 5734 S. Ellis Avenue, Chicago,
14	IL 60637, USA
15	4. Center for Climate System Research (CCSR), University of Tokyo, Japan
16	5. CCR, CGD/NCAR, PO Box 3000, Boulder, CO 80307-3000, USA
17	6. NASA Goddard Institute for Space Studies, 2880 Broadway, New York, NY 10025 USA
18	7. School of Geographical Sciences, University of Bristol, University Road, Bristol, BS8 1SS,
19	UK
20	8. Department of Geography, Faculty of Engineering and Environment, Northumbria
21	University, Ellison Building, Newcastle upon Tyne, NE1 8ST, UK.
22	
23	Abstract
24	The Pliocene Model Intercomparison Project (PlioMIP) is a co-ordinated international climate
25	modelling initiative to study and understand climate and environments of the Late Pliocene,

and their potential relevance in the context of future climate change. PlioMIP operates under

the umbrella of the Palaeoclimate Modelling Intercomparison Project (PMIP), which examines
 multiple intervals in Earth history, the consistency of model predictions in simulating these
 intervals and their ability to reproduce climate signals preserved in geological climate
 archives.

This paper provides a thorough model intercomparison project description, and documents
 the experimental design in a detailed way. Specifically, this paper describes the experimental
 design and boundary conditions that will be utilised for the experiments in Phase 2 of PlioMIP.

- 34 The Pliocene Model Intercomparison Project (PlioMIP) is a co-ordinated international
- 35 climate modelling initiative to study and understand climate and environments of the Late
- 36 Pliocene, and their potential relevance in the context of future climate change. PlioMIP
- 37 examines the consistency of model predictions in simulating Pliocene climate, and their
- 38 ability to reproduce climate signals preserved by geological climate archives. Here we
- 39 provide a description of the aim and objectives of the next phase of the model
- 40 intercomparison project (PlioMIP Phase 2), and we present the experimental design and
- 41 boundary conditions that will be utilised for climate model experiments in Phase 2.
- 42 Following on from PlioMIP Phase 1, Phase 2 will continue to be a mechanism for sampling
- 43 structural uncertainty within climate models. However, Phase 1 demonstrated the
- 44 requirement to better understand boundary condition uncertainties as well as uncertainty
- 45 in the methodologies used for data-model comparison. Therefore, our strategy for Phase 2
- 46 is to utilise state-of-the-art boundary conditions that have emerged over the last 5 years.
- 47 These include a new palaeogeographic reconstruction, detailing ocean bathymetry and
- 48 <u>land/ice surface topography. The ice surface topography is built upon the lessons learned</u>
- 49 <u>from offline ice sheet modelling studies. Land surface cover has been enhanced by recent</u>
- 50 additions of Pliocene soils and lakes. Atmospheric reconstructions of palaeo-CO₂ are
- 51 emerging on orbital timescales and these are also incorporated into PlioMIP Phase 2. New
- 52 records of surface and sea surface temperature change are being produced that will be
- 53 more temporally consistent with the boundary conditions and forcings used within models.
- 54 Finally we have designed a suite of prioritized experiments that tackle issues surrounding
- 55 the basic understanding of the Pliocene and its relevance in the context of future climate
- 56 <u>change in a discrete way.</u>

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58 1. Introduction to PlioMIP

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60 1.1 PlioMIP Phase 1 Design and Objectives

The PlioMIP project was initiated in 2008 and is closely aligned with the U.S. Geological Survey Projectgram known as PRISM (Pliocene Research Interpretation and Synoptic Mapping). The PRISM₇ project which has spent more than 25 years focusing on the reconstruction and understanding of the mid-Pliocene climate (~3.3 to 3 million years ago), as well as the production of boundary condition data sets suitable for use with numerical climate models.

66 Phase 1 of the PlioMIP project commenced in 2008 and was concluded in 2014. In Phase 1 67 two mid-Pliocene experiments were performed. Experiment 1 used atmosphere-only General 68 Circulation Models (GCMs) with prescribed surface boundary conditions (sea-surface 69 temperatures, sea-ice, and vegetation) derived from the PRISM3D data set (Dowsett et al., 70 2010).__Land/sea distribution and topography were also prescribed from PRISM3D. 71 Experiment 2 used coupled ocean-atmosphere GCMs where sea-surface temperatures and 72 sea-ice were predicted dynamically by the models; vegetation, land/sea distribution, and 73 topography remained fixed to PRISM3D estimates.

74 The scientific objectives in Phase 1 were to:

- Examine large-scale features of mid-Pliocene climate that are consistent across
 models.
- Determine the dominant components of mid-Pliocene warming derived from the
 imposed boundary conditions.
- Examine first order changes in ocean circulation between the mid-Pliocene and
 present-day.
- Examine the behaviour of the <u>m</u>Monsoons (e.g. their intensity).
- Compare model results with proxy data to determine the performance of models
 simulating a warm climate state.
- Use the mid-Pliocene as a tool to evaluate the long term sensitivity of the climate
 system to near modern concentrations of atmospheric CO₂.

87 1.2 PlioMIP Phase 1 Accomplishments

88 In the context of co-ordinated international model intercomparison projects, PlioMIP 89 achieved a number of firsts. For example, it was the first palaeoclimate modelling 90 intercomparison project to require altered vegetation distributions to be modified in climate 91 models, facilitating vegetation-climate feedbacks to be incorporated into the model 92 intercomparison. It was also the first intercomparison project that required individual groups 93 to fully document the implementation of palaeo-boundary conditions within their models, 94 along with the basic climatological responses. This was designed to facilitate the 95 intercomparison itself by enabling artefacts of individual methodologies of boundary 96 condition implementation to be separated from robust model responses to imposed Pliocene 97 boundary conditions. Through PlioMIP, a spin off project known as PLISMIP (Pliocene Ice 98 Sheet Model Intercomparison Project; Dolan et al. 2011) was initiated and has focused on 1) 99 assessing ice sheet model dependency of Greenland Ice Sheet reconstructions during the 100 Pliocene using shallow ice approximation ice sheet models (Dolan et al., 2011; Koenig et al., 101 2014), 2) examining the effect of different GCM climatological forcing on predicted ice sheet 102 configurations (Dolan et al. 2014) and 3) using shallow shelf ice sheet models for Antarctica 103 to test both ice sheet model and climate model dependency on predicted ice sheet 104 reconstructions (de Boer et al. 2015).

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107 Outputs from PlioMIP Phase 1 include:

- Identified consistency in surface temperature change across models in the tropics.<u>and</u> <u>r</u>
 <u>in contrast</u><u>a</u> lLack of consistency identified in <u>the model simulated temperature</u> responses
 at high latitudes (Haywood et al., 2013a).
- <u>In contrast mM</u>odel predictions are inconsistent in terms of total precipitation rate in the
 tropics (Haywood et al., 2013a).
- Global annual mean surface temperatures increased by 1.84°C to 3.6°C and show a greater
 range for Experiment 2 using coupled ocean-atmosphere models than Experiment 1 using
 fixed sea-surface temperatures (Haywood et al., 2013a).

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116	• There was no clear indication-signal signal in the model ensemblemodel predictions to
117	support <u>either</u> enhanced <u>or weaker</u> Atlantic Meridional Overturning Circulation and Ocean
118	Heat Transport to the high latitudes (Z. Zhang et al., 2013).

- Model predictions of enhanced Atlantic Meridional Overturning Circulation and Ocean
 Heat Transport to high latitudes are inconsistent, in sign as well as strength (Z. Zhang et al., 2013).
- Clear sky albedo and greenhouse gas emissivity dominate polar amplification of surface
 temperature warming during the Pliocene. This demonstrated the importance of specified
 ice sheet and high latitude vegetation boundary conditions and simulated sea ice and snow
 albedo feedbacks. Furthermore, the dominance of greenhouse gas emissivity in driving
- 126 surface temperature changes in the tropics was identified (Hill et al., 2014).
- The simulated weakened mid-Pliocene East Asian winter winds in north monsoon China
 and intensified East Asian summer winds in monsoon China agreed well with geological
 reconstructions (<u>R.</u> Zhang et al., 2013).
- Data-model comparison using both sea surface and surface temperature proxies indicate that climate models potentially underestimate the magnitude of polar amplification.
 However, current limitations in age control and correlation make interpreting <u>data-modelmodeldata</u> discrepancies challenging (Dowsett et al., 2012, Dowsett et al., 2013a, Salzmann et al., 2013).
- Model results indicate that longer term climate sensitivity (Earth System Sensitivity) is
 greater than Charney Sensitivity (best estimate ESS/CS ratio of 1.5: Haywood et al., 2013a).
- 137 138

139 1.3 PlioMIP - Emerging Challenges/Opportunities

One of the key findings in PlioMIP Phase 1 was the potential underestimation of model predicted surface temperature warming in the high latitudes. Understanding <u>datamodel <u>modeldata</u> discord is non-trivial and can rarely be attributed to a single factor. The complexity
 of understanding <u>datamodel-modeldata</u> discord is highlighted by the PMIP Triangle (Figure
 1), which illustrates three possible contributions to <u>datamodel-modeldata</u> discrepancy, and
</u>

has at its vertices model physics (structural and parameter uncertainty), model boundaryconditions and proxy data uncertainty.

147 Following on from PlioMIP Phase 1, Phase 2 will continue to be a mechanism for sampling 148 structural uncertainty within climate models as a suite of different models will take part in 149 PlioMIP. However, Phase 1 demonstrated the requirement to better understand boundary 150 condition uncertainties as well as weaknesses in the methodologies used for data-model 151 comparison which largely stemmed from the time averaged nature of proxy data used in 152 previous data--model comparisons (Dowsett et al., 2013a; Salzmann et al., 2013). Therefore, 153 our strategy for Phase 2 is to utilise state-of-the-art boundary conditions that have emerged 154 over the last 5 years. These include a new palaeogeography reconstruction detailing ocean 155 bathymetry and land/ice surface topography, and new data sets describing the distribution 156 of Pliocene soils and lakes. The ice surface topography is built upon the lessons learned during 157 the PLISMIP project (Dolan et al., 2014). Land surface cover will be enhanced by recent 158 additions of Pliocene soils and lakes (Pound et al., 2014). -Atmospheric reconstructions of 159 palaeo-CO₂ are emerging on orbital timescales (e.g. Bartoli et al., 2011; Badger et al., 2013) 160 and these will also be incorporated into PlioMIP Phase 2.

161 It was recognised during Phase 1, that a key influence on datamodel-modeldata discord stems 162 from uncertainties associated with the derivation of the proxy_-data sets used to assess the 163 climate models. Although certainty surrounding any proxy data set is limited by analytical, 164 spatial and temporal uncertainty, Phase 1 highlighted temporal uncertainty as an important 165 constraint on more robust methodologies for data-model comparison (DMC: Dowsett et al., 166 2013a; Haywood et al., 2013b; Salzmann et al., 2013). The concept of climate stability during 167 the Pliocene is overly simplistic both in geological climate archives and climate modelling 168 approaches.

Due to the increasing recognition of climate variability in the Pliocene, time averaged approaches to palaeoenvironmental reconstruction have reached their ultimate potential to evaluate climate models. Therefore, enhancing the temporal resolution of data collection in order to more adequately understand climate variation in the Pliocene is required, and along <u>with</u> developing a more strategic approach to the choice of relevant Pliocene event(s) to reconstruct and model<u>is needed</u>. One of PlioMIP's guiding principles is to utilise palaeoenvironments to better inform us of likely scenarios for future global change. To this 176 end, the event chosen for PlioMIP Phase 2 focuses on the identification of a 'time slice' 177 centred on an interglacial peak (MIS KM5c; 3.205 Ma) that has near-modern orbital forcing, 178 and yet retains many of the characteristics of Pliocene warmth on which we have focussed in 179 the past (Dowsett et al., 2013b; Haywood et al., 2013b; Salzmann et al., 2013; Prescott et al., 180 2014). Discussions surrounding potential modification of the LR04 benthic isotope stack 181 (Lisiecki and Raymo, 2005) are currently ongoing, which may lead to a modification of the 182 assigned Marine Isotope Stage assigned - KM5c to the astrochronological age of 3.205 in the 183 future.

PRISM and the wider Pliocene data community are rising to the challenge to obtain higher
resolution proxy_-data that will inform the models about the chosen time slice (e.g. Dowsett
et al., 2013b; see also Haywood et al., 2013b). The key differences between the PRISM data
that underpinned PlioMIP Phase 1 and the new direction for data collection include:

- Expanding to a community-wide effort, new data generation will focus on key
 locations and specific regions that have been identified by PlioMIP Phase 1 as
 important for understanding Pliocene climate variability and model performance.
- In order to increase our understanding of temporal changes in Pliocene climate, time
 series data <u>are being will be</u>produced as standard, which will in essence increase
 previous temporal resolution by two orders of magnitude and lead to enhanced
 methods of data-/model comparison (Dowsett et al. 2013b).-
- We will encourage the use of multi-proxy methods of data generation. This will enable
 us to derive more robust and holistic palaeoenvironmental reconstructions.
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198 1.4 Pliocene for Future and Pliocene for Pliocene

The utilization of the mid-Pliocene as a means to understand future global change ("Pliocene for Future") remains a priority in Phase 2. It is our intention to forge even stronger links between PlioMIP, PMIP, CMIP and the next IPCC assessment. However, we recognise that many researchers are primarily interested in the Pliocene because it represents a considerable challenge to our understanding of the operation of the Earth System ("Pliocene for Pliocene"). Furthermore, a number of scientific requirements and priorities do not fit exclusively within a Pliocene for Future mandate. For example, state of the art palaeographic reconstructions are indicating more substantial regional variations in palaeogeography than
were <u>appreciatedknown</u> in the past (Hill, 2014). Due to the differing requirements identified,
in PlioMIP Phase 2 we have designed a portfolio of model experiments that effectively address
both the "Pliocene for Future" and "Pliocene for Pliocene" agendas. This is illustrated in the
following CMIP-style diagram (e.g. Taylor et al., 2012) where priorities for both agendas are
highlighted, with both agendas sharing a common core experiment_z which will be promoted
as the PlioMIP Phase 2 experiment within CMIP.

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214 2. Strategy and Methodology

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216 2.1 Naming Convention and Summary of the Experimental Design for PlioMIP Phase 2

The experiments in PlioMIP Phase 2 have been grouped into hal<u>vesf's</u> "**Pliocene4Pliocene**" and "**Pliocene4Future**" and <u>wouldshould</u> ideally be completed by all participating groups. However, <u>onlythe the</u> core experiments <u>must be completed</u> by all groups. Each half of the project is divided into two 'tiers' (Fig. 2). After the core experiments, tier 1 experiments are identified as a higher priority for completion than tier 2.

222 We describe several model simulations, which essentially consist of various combinations of 223 boundary conditions associated with prescribed CO₂, orography, soils, lakes, and ice sheets. 224 To simplify the experimental descriptions, we use the following nomenclature: Ex^c, where c 225 is the concentration of CO_2 in ppmv, and x are any boundary conditions which are Pliocene as 226 opposed to pre-industrial, where x can be any or none of o,i, where o is orography and i is ice 227 sheets. For example, a pre-industrial simulation with 280 ppmv CO₂ we denote E²⁸⁰. A Pliocene simulation with 400 ppmv is Eoi⁴⁰⁰, and a simulation with Pliocene ice sheets, but 228 preindustrial orography, and at 560 ppmv, is Ei⁵⁶⁰. Note that in all our simulations, orography 229 230 and lakes and soils are modified in unison, and so 'o' denotes changes to orography, 231 bathymetry, land-sea mask, lakes and soils combined.

- 232 Within the Pliocene4Future agenda, given the uncertainty in total greenhouse gas forcing for
- 233 the KM5c time slice, we have proposed $\frac{1}{2}$ -simulations using $\frac{350 \text{ and }}{450}$ ppmv CO₂ (Eoi³⁵⁰,
- 234 Eoi⁴⁵⁰). <u>Both these experiments will facilitate model evaluation using proxy data. Eoi⁴⁵⁰ This</u>

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235 also enables the experimental design to accommodate other Earth System processes that 236 may have an effect on radiative forcing, besides greenhouse gas concentrations. For example, 237 Unger and Yue (2014) have demonstrated that chemistry-climate feedbacks, in terms of their 238 radiative forcing, may play as important, or even more important, role as CO₂ during the 239 Pliocene. With a 450 ppmv experiment we also aim to address how uncertainty in radiative 240 forcing can account for high latitude data-/model mismatches that were revealed in PlioMIP 241 Phase 1 (Haywood et al. 2013a; Dowsett et al., 2012 and 2013a; Salzmann et al., 2013). We 242 have also specified a pre-industrial experiment with 560 ppmvPliocene CO2 as a tier 1 243 experiment (E560400)_. This is to facilitate an investigation into Climate (Charney) and Earth 244 System Sensitivity.

Within tier 2 we have proposed two experiments that are designed to assess the dependence
of climate sensitivity on the background climate and boundary condition states. Here we wish
to-to compare the response of the system to CO₂ forcing, between the Pliocene and the
modern, by specifying a <u>Pliocene experiment with 280-560</u> ppmv CO₂ (Eoi²⁸⁰⁾, as well as a preindustrial experiment using 400 ppmv CO₂ (E⁴⁰⁰) concentration. in both a Pliocene (Eoi⁵⁶⁰) as
well as pre-industrial experiment (E⁵⁶⁰).

251 For our Pliocene4Pliocene agenda we have within tier 1 focused on the atmospheric CO₂ 252 uncertainty by specifying a higher and lower CO₂ experiment at 450 and 350 ppmv (Eoi⁴⁵⁰ and Eoi^{350}), which provides a 100 ppmv uncertainty bracket around our KM5c core experiment 253 254 (using 400 ppmv CO₂). Within tier 2 we have specified a series of experiments designed to 255 identify the individual contribution of boundary condition changes to the overall modelled Pliocene climate response (E⁴⁰⁰, E²⁸⁰, Eo⁴⁰⁰, Eoi⁴⁰⁰). To assess non-linearity in the factorization 256 of the forcings, we have specified an enhanced factorization methodology (E^{400} , E^{280} , Eo^{400} , 257 258 Eo²⁸⁰, Ei⁴⁰⁰, Ei²⁸⁰, Eoi⁴⁰⁰, Eoi²⁸⁰: see section 3.2).

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260 2.2 Standard and <u>e</u>Enhanced boundary conditions

All required boundary conditions can be accessed from the United States Geological Survey PlioMIP2 website (see: <u>http://geology.er.usgs.gov/egpsc/prism/7_pliomip2.html</u>). For the Pliocene experiment two versions of the palaeogeography (including land/sea mask (LSM), topography, bathymetry and ice distribution) are provided. The **standard** boundary condition

265 data package does not require a modelling group to have the ability to alter the LSM or 266 bathymetry (apart for selected critical regions of the Bering Strait, Canadian Archipelago and 267 Hudson Bayi.e. specified ocean gateways). The enhanced boundary condition requires the ability to change the model's LSM and ocean bathymetry more generally. The standard 268 269 boundary conditions data set -data package using an approximately modern LSM-is provided 270 in order to maximise the potential number of participating modelling groups. If groups are 271 unable to make any changes to their models LSM then they may use their own LSM from their 272 pre-industrial simulation in PlioMIP Phase 2, since it is difficult in some climate models to 273 successfully alter the LSM. Groups that are not able to change their LSMs at all (i.e. altering 274 selected ocean gateways) are required to use their own modern LSM. A PRISM4/PlioMIP 275 Phase 2 modern land/sea mask is provided to help guide the implementation of Pliocene 276 topography into different climate models. Groups are asked to make every effort to 277 implement as many of the boundary conditions in the enhanced data packages as possible; 278 however, we recognise that this will not be possible for all groups.

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280 2.3 Core Experimental Design and Boundary Conditions

281 2.3.1 Integration, atmospheric gases/aerosols, solar constant/orbital configuration

The experimental design for the core Pliocene KM5c time slice experiment is summarised in Table 1 (standard and enhanced boundary conditions). **Integration length is to be set to at least 500 years** in accordance with CMIP guidelines-(Coupled Model Intercomparison Project Phase) for <u>equilibrated</u> coupled model experiments (see: Taylor et al., 2012). The **concentration of CO₂ in the atmosphere is to be set to 400 ppmv.** In the absence of proxy data, all other trace gases and aerosols are specified to be identical to the individual group's pre-industrial control experiment.

<u>It is difficult to reconstruct the concentration of atmospheric CO₂ during the PlioceneWhen trying to reconstruct Pliocene CO₂ uncertainty is inevitable. While Pliocene CO₂ reconstruction is <u>difficult, it is</u> an important ongoing area of research with new records and syntheses due to emerge over the next few years. Current evidence for Pliocene CO₂ comes from a number of sources: (1) the stomatal density of fossil leaves (Kürschner et al., 1996), (2) carbon isotope analyses (e.g. Raymo et al., 1996), (3) alkenone-based estimates (Pagani et al., 2010; Seki et</u>

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295 al., 2010; Badger et al., 2014) and (4) boron isotope analyses (e.g. Seki et al., 2010). For the 296 warm intervals of the Pliocene values of CO₂ from each of these proxies vary, but within error 297 they may overlap (Bartoli et al., 2011). The stomatal density records support a CO_2 298 concentration of 350 to 380 ppmv. The average of the Raymo et al. (1996) carbon isotope 299 analyses is similar to the stomatal-based estimates, but peaks above that value (beyond 425 300 ppmv) occur. The Pagani et al. (2010) study reconstructed CO₂ from a number of different 301 marine records, and in three of the six marine records a CO₂ value of 400 is reasonable and 302 within the range of 365 to 415 ppmv. In the Seki et al. (2010) study the alkenone-based CO₂ 303 record is consistent with a value around 400 ppmv. Badger et al. (2014), have demonstrated 304 that while absolute alkenone-based CO₂ reconstructions are influenced by a number of 305 factors including productivity, cell size, SST, other local palaeoceanographic conditions as well 306 as secondary effects of alkenone δ^{13} C, assessments of the degree of variability in CO₂ (rather 307 than absolute concentration) are likely to be more robust, and indicate less than 55 ppmv of 308 variation between 3.3 and 2.8 million years ago. Atmospheric CO₂ is an obvious choice for 309 sensitivity tests as part of PlioMIP Phase 2 and is addressed within the experimental design 310 for PlioMIP Phase 2. Information on the concentration of other greenhouse gasses such as 311 Methane and Nitrogen Dioxide is absent for the Pliocene and must therefore be prescribed 312 at a pre-industrial level. The CO₂ concentrations specified within PlioMIP Phase 2 are therefore designed to account for the total greenhouse gas forcing derived from all sources. 313

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315 The solar constant is to be specified as the same as in each participating group's pre-316 industrial control run. In the pastprevious versions, the PRISM boundary conditions (Dowsett 317 et al. 2010) represented an average of the warm intervals during the time slab (~3.3 to 3 318 million yr), rather than conditions occurring during a discrete time slice. This made it 319 impossible to prescribe an orbital configuration which that would be representative of the 320 entire 300,000 year interval. However, due to the new focus within PRISM4 and PlioMIP Phase 321 2 to increase the temporal resolution of proxy records, and to concentrate on a smaller 322 interval of time approaching a time slice reconstruction for MIS KM5c, it is now possible to 323 provide climate models with more certain values for astronomical and orbital forcing. The 324 KM5c time slice was selected partly on the basis of a strong similarity in orbital forcing to 325 present-day. Therefore, in the interests of simplicity of the experimental design,

astronomical/orbital forcing in Pliocene experiments (eccentricity, obliquity, and
 precession) is to remain unchanged from each models pre-industrial control simulation.

328

329 2.3.2 Palaeogeography (land/sea mask, topography, bathymetry, ocean gateways, land ice) 330 The PRISM4 palaeogeography provides a consistent reconstruction of topography, 331 bathymetry, ice sheets and the land-sea mask that can be implemented in PlioMIP Phase 2 332 models. The PRISM4 Pliocene palaeogeography data set is provided in NetCDF format at a 1° 333 $\times\,1^\circ$ resolution. The PRISM4 palaeogeography includes components, such as the contribution 334 of dynamic topography caused by changes in the mantle flow (e.g. Rowley et al., 2013) and 335 the glacial isostatic response of loading specific Pliocene ice sheets (e.g. Raymo et al., 2010), 336 that were not previously considered in the PRISM3D reconstruction of Sohl et al. (2009). In 337 the sstandard boundary condition data set all ocean gateways remain the same as the 338 modern except for the Bering Strait, which that should be closed, and the Canadian Arctic 339 Archipelago which should also be closed (isolating Baffin Bay and the Labrador Sea from the 340 Arctic Ocean). In the enhanced boundary condition data set the Bering Strait and Canadian 341 Arctic Archipelago are also closed, but there are other required changes in the Torres Strait, 342 Java Sea, South China Sea, Kara Strait as well as a West Antarctic Seaway.

343 The approach taken to derive PRISM4 ice sheets in the palaeogeography reconstruction is 344 different to PRISM3D (Dowsett et al., 2010). The results of PLISMIP have shown that ice sheet 345 model dependency over Greenland is low. However, the initial climatological forcing has a 346 large impact on the predicted Greenland ice sheet configuration (Dolan et al., 2014; Koenig 347 et al., 2014). Using a compilation of the results presented in Koenig et al. (2014), we have 348 implemented an ice sheet configuration over Greenland in PRISM4 where we have the 349 highest-confidence in the possibility of ice sheet location during the warmest parts of the Late 350 Pliocene (see Fig. 6b in Koenig et al. 2014). The reconstruction of Keonig et al. (2014) was 351 modified by removing ice from Southern Greenland. The presence of ice in that region is 352 inconsistent with palynological studies that suggest that Southern Greenland was vegetated 353 during warm intervals of the Pliocene (e.g. de Vernal and Mudie, 1989). - The PRISM4 354 Greenland Ice Sheet configuration is smaller than in PRISM3D and ice is limited to high 355 elevations in the Eastern Greenland Mountains (Fig. 4).

357 Over Antarctica, work in PLISMIP is still ongoing (de Boer et al. 2015); therefore we have 358 decided to use an ice sheet that best agrees with the available proxy_-data. Based on evidence 359 from the ANDRILL core data and ice sheet modelling (Naish et al., 2009; Pollard and DeConto, 360 2009) that suggests that, in specific warm periods of the Late Pliocene, there was no ice 361 present in West Antarctica, this region remains ice free in the PRISM4 palaeogeography 362 reconstruction (Fig. 4). Over East Antarctica, Cook et al. (2013) show that the Wilkes 363 subglacial basin may have been highly dynamic during the warmest parts of the Late Pliocene 364 and they infer significant potential for ice sheet retreat in this region. Additionally, Young et 365 al. (2011) highlight the Aurora subglacial basin as an area which may have been subject to 366 marine ice sheet instabilities in the past (potentially in the Pliocene). Therefore, over East 367 Antarctica PlioMIP Phase 2 uses the PRISM3D ice sheet reconstruction (Hill et al., 2007; Hill, 368 2009; Dowsett et al., 2010), as this remains consistent with more recently available data. In 369 this reconstruction (Fig. 4) large portions of the East Antarctic ice sheet show little change or 370 a small increase in surface altitude with respect to modern, and significant ice sheet retreat 371 is limited to the low-lying Wilkes and Aurora subglacial basins.

372

For the Pliocene experiments, two versions of the palaeogeography will be provided toclimate modelling groups:

375 Standard: For the models where altering the LSM and bathymetry is problematic, we 376 provide a palaeogeography with a modern land-sea configuration and bathymetry \$77 (apart from in the Hudson Bay, Bering Strait and Canadian Archipelago). In this 378 instance the Late Pliocene topographic elevations were extended to the modern 379 coastline, and the bathymetry remained at modern values. Groups that are unable to 380 change their land-sea mask or bathymetry at all are asked to use their local modern 381 boundary conditions; however guidance on the implementation of Pliocene 382 topography in this case should be taken from the standard palaeogeography data set. 383 Enhanced: This presents the full palaeogeographic reconstruction including all • 384 changes to topography, bathymetry, ice sheets and the LSM.

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385 To ensure that the climate anomalies (Pliocene minus present day) from all PlioMIP Phase 2 386 climate models are directly comparable, i.e. that they reflect differences in the models 387 themselves rather than the differences of modern boundary conditions, it has been decided 388 to implement Pliocene topography (and bathymetry) as an anomaly to whatever standard 389 modern topographic data set is used by each modelling group in their own model. To create 390 the Pliocene topography (and bathymetry) the difference between the PRISM4 Pliocene and 391 PRISM4 Modern topography (bathymetry) should be calculated and added to the modern 392 topographic (bathymetric) data sets each participating modelling group employs within their 393 own standard-pre-industrial control simulations.

394 Such that:

395

Plio^{TOPO} = (PRISM4^{PlioTOPO} – PRISM4^{ModernTOPO}) + Modern^{TOPO} Local

396 and

397 Plio^{BATH}= (PRISM4^{PlioBATH} – PRISM4^{ModernBATH}) + Modern^{BATH} Local

With this formulation it is possible that on occasion grid cells may become land where the intention is for an ocean cell to be specified and vice-versa. In this case the specified Pliocene LSM takes precedence, in other words ensure that the integrity of Pliocene LSM boundary condition data is always preserved. Data_sets to be provided at a 1° × 1° resolution for the core experiments can be found in Table 1.

403

404 2.3.3 Vegetation, Lakes, Soils and Rivers

405 A global data set of vegetation for the KM5c time slice is not available. A number of climate 406 models now have the ability to simulate predict the type and distribution of vegetation using 407 dynamic vegetation models. In PlioMIP Phase 2 vegetation models should be initialised with 408 pre-industrial vegetation cover and spun up until an equilibrium condition is reached. If 409 Pliocene vegetation cannot be predicted dynamically, modelling groups can prescribe vegetation using the Salzmann et al. (2008) PRISM3 vegetation reconstruction used within 410 411 PlioMIP Phase 1 (Haywood et al. 2010 and Haywood et al. 2011), and provided as a mega 412 biome reconstruction in the PlioMIP Phase 2 boundary condition files. An equivalent potential 413 natural vegetation data set is also provided to guide how groups implement prescribed Pliocene vegetation. Further details on correctly approaching the implementation of
prescribed Pliocene vegetation for PlioMIP Phase 2 can be found in Haywood et al. (2010:
Section 3.5).

417 Due to lack of information covering the distribution of lakes and soils during PlioMIP Phase 1, 418 lakes were absent from the land cover boundary conditions. Since PlioMIP Phase 1, the global 419 distribution of Late Pliocene soils and lakes have been reconstructed through a synthesis of 420 geological data (Pound et al. 2014). Initial experiments using the Hadley Centre Coupled 421 Climate Model Version 3 (HadCM3) indicate regionally confined changes of local climate and 422 vegetation in response to the new lakes and soils boundary condition (Pound et al. 2014). 423 When combined (lakes plus soils), the feedbacks on climate from Late Pliocene lakes and soils 424 improve the proxy data-model fit in western North America as well as the southern part of 425 northern Africa (Pound et al. 2014).

426 In PlioMIP Phase 2 all modelling groups should implement the Pound et al. (2014) data sets 427 for global lake (Fig. 5) and soils distribution (Fig. 6). If lake distribution is a dynamically 428 predicted variable within a model (i.e. lake distributions can change as a result of predicted 429 changes in climate), prescribing the Pound et al. (2014) lake data set is not necessary. The 430 lake data set provides information on both lake size as well as the fractional coverage of lakes 431 within model grid boxes. Figure 5 also shows how the lake distribution and sizes differ from 432 modern, most notably the absence of post-glacial lakes in North America and the presence of 433 large lakes in Central Africa (Pound et al., 2014).

The colour (for albedo) and texture translations for the nine soil orders used in the modelling of Late Pliocene soils and lakes are provided to guide the implementation of soil type and distribution in models. This translation is based upon the definition of soils with the HadCM3 (Table 2).

Groups should **implement Pliocene lakes using the anomaly method** (the anomaly between the provided Pliocene and modern lake data sets added to each groups local modern lake distribution data set), and ensure that minimum lake fractions do not fall below 0 and the maximum do not exceed 1 (100%). Groups may implement the Pliocene soils using whatever method they deem most appropriate for their model. This may be by applying the provided Pliocene soil properties directly in their Pliocene simulation (i.e. as an absolute), or by calculating an anomaly from the provided modern soils data, and adding this to the local 445 modern control soil properties. Alternatively, groups may choose to develop a regression of 446 the provided modern soil properties with their local modern control soil properties, and then 447 apply the resulting regression formulae to the provided Pliocene soil properties. 448 449 With regard to river routing the required solution is to follow modern river routes except 450 where this would be inappropriate due to the appearance of new land grid cells in the 451 Pliocene land/sea mask, in which case rivers should be routed to the nearest ocean grid box 452 or most appropriate river outflow point. 453 454 3. Sensitivity experiments and forcing factorization 455 **3.1 Sensitivity Experiments** 456 3.1.1 Pliocene for Future Tier 1 and 2 457 Within the Pliocene for Future agenda a pre-industrial experiment with 560 ppmv Pliocene 458 CO_2 has been selected as a tier 1 experiment (E^{560400}). This is to facilitate an investigation into 459 Climate (Charney) and Earth System Sensitivity. Also given the uncertainty in total greenhouse 460 gas forcing for the KM5c time slice, we have proposed a simulation using 450 and 350 ppmv 461 CO₂ (Eoi⁴⁵⁰, Eoi³⁵⁰). Within tier 2 we have proposed two experiments that are designed to 462 assess the similarity how similar of Pliocene and future climate feedbacks to higher CO2 are 463 blevelsetween the Pliocene and the future by specifying a Pliocene experiment using 280560 464 ppmv CO₂ concentration in both a Pliocene (Eoi²⁸⁰⁵⁶⁰) as well as pre-industrial experiment using 400 ppmv (E⁴⁰⁰⁵⁶⁰). 465 466

467 3.1.2 Pliocene for Pliocene Tier 1

For the Pliocene for Pliocene agenda we have within tier 1 focused on the atmospheric CO₂
uncertainty by specifying a high and low CO₂ experiment at 450 and 350 ppmv (Eoi⁴⁵⁰ and
Eoi³⁵⁰ respectively), which provides a 100 ppmv uncertainty bracket around our KM5c core
experiment (using 400 ppmv CO₂).

472

473 3.2 Pliocene for Pliocene Tier 2 Forcing Factorization Experiments

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The primary aim of the Pliocene for Pliocene Tier 2 forcing factorisation experiments is to assess the relative importance of various boundary condition changes which contribute to Pliocene warmth. Following a similar methodology adopted in Lunt et al. (2012) we intend to partition the total Pliocene warming (or temperature change; ΔT) into three components, each due to the change in one of the following boundary conditions: CO₂, topography and ice sheets. Our factorisation, which is that proposed by Lunt et al. (2012), can be written:

 $480 \qquad \Delta T = dT_{CO2} + dT_{topo} + dT_{ice}$

481 $dT_{CO2} = \frac{1}{4} \left[(E^{400} - E^{280}) + (Eo^{400} - Eo^{280}) + (Ei^{400} - Ei^{280}) + (Eoi^{400} - Eoi^{280}) \right]$

482 $dT_{orog} = \frac{1}{4} [(Eo^{280} - E^{280}) + (Eo^{400} - E^{400}) + (Eoi^{280} - Ei^{280}) + (Eoi^{400} - Ei^{400})]$

483 $dT_{ice} = \frac{1}{4} [(Ei^{280} - E^{280}) + (Ei^{400} - E^{400}) + (Eoi^{280} - Eo^{280}) + (Eoi^{400} - Eo^{400})]$

484

This gives a total of 8 simulations required (2^N, where N is the number of processes factorised, a in this case), although only 5 of them (Eo⁴⁰⁰, Eo²⁸⁰, Ei⁴⁰⁰, Ei²⁸⁰, Eoi²⁸⁰) are in addition to simulations already in Tier 1 or the Core. This method, although more computationally demanding than the linear approach (e.g. Broccoli and Manabe, 1987; von Deimling et al., 2006), has the advantage that it takes into account non-linear interactions, is symmetric, and is unique (Table 3).

491 If groups do not have the computational resource to carry out the full factorisation, they may492 carry out a linear factorisation, as follows:

493

 $494 \quad dT_{CO2} = E^{400} - E^{280}$

495 $dT_{orog} = Eo^{400} - E^{400}$

- 496 $dT_{ice} = Eoi^{400} Eo^{400}$
- 497

This is a total of 4 simulations, but only 1 of them (Eo^{400}) in addition to simulations already in

499 Tier 1 or the Core. Further guidance on boundary condition implementation for the forcing

500 <u>factorization experiment can be found in Figure 1 of the Supplementary Information.</u>

501

502 4. Proxy data for the evaluation of model outputs

503 Short, high-resolution time series extending from MIS M2 through KM3 will be necessary to 504 meet the evaluation requirements of PlioMIP Phase 2. Marine sequences will depend upon 505 chronology from the Lisiecki and Raymo 20054 (LR04) time scale and should have multiple 506 palaeoenvironmental proxies (Dowsett et al. 2013a). Previous work from the palaeoclimate 507 data community suggests a number of sites potentially suitable for evaluation of PlioMIP 508 Phase 2 model outputs (e.g. Dowsett et al., 2012; 2013a, 2013b; Fedorov et al., 2013; 509 Salzmann et al., 2013, Brigham-Grette et al., 2013). Well-dated, high resolution records from 510 the continental interior are scarce, and terrestrial reconstructions will be mostly based on 511 marine and marginal marine sequences. The primary areas of discord between simulated and 512 estimated Pliocene palaeoclimate conditions identified in PlioMIP Phase 1 include the mid-513 to-high latitude North Atlantic, tropics and upwelling regions (Dowsett et al., 2012). The 514 PRISM4 marine and terrestrial contribution to the PlioMIP Phase 2 community evaluation 515 data set has been initially concentrated in the North Atlantic region (Fig. 7).

516

517 5. Variables, output format, data processing and storage

518

510 519 If the PlioMIP Phase 2 core experiment is adopted as a CMIP6 simulation, model data for this 520 experiment must use the Climate Model Output Rewriter (CMOR) format and stored on an 521 ESGF node (The Earth System Grid Federation). The CMOR library has been specially 522 developed to help meet the requirements of the Model Intercomparison. Further details of 523 CMIP6 experiments and require<u>d</u> outputs<u>/-and-required</u>-CMOR file formats will be made

available on the CMIP6 website (http://www.wcrp-climate.org/index.php/wgcm-

525 526

524

cmip/wgcm-cmip6).

527 If the PlioMIP Phase 2 core experiment is specified as a PMIP core experiments the same guidelines for output format and storage of data detailed for CMIP6 applies. For PlioMIP 529 Phase 2 experiments listed within Tiers 1 and 2 more flexibility in terms of data storage and 530 file formats is available. PlioMIP Phase 2 has modified the established variables list outlined 531 by the 3rd Phase of the PMIP project. The list of required variables can be found listed on the 532 PlioMIP Phase 2 website (http://geology.er.usgs.gov/egpsc/prism/7_pliomip2.html). All

533	model outputs will be submitted initially to a data repository at the University of LeedsLeeds								
534	(including the PlioMIP Phase 2 core experiment, which may have data replicated in CMOR								
535	format on an ESGF node). Requests for access to should be sent to A Haywood. In general								
536	(CMIP6 guidelines aside) PlioMIP project requires participants to prepare their data files so								
537	that they meet the following constraints (regardless of the way their models produce and								
538	store their results).								
539									
540	• The data files have to be in the (now widely used) <u>NeetCDF</u> binary file format and								
541	conform to the CF (Climate and Forecast) metadata convention (outlined on the								
542	website http://cf-pcmdi.llnl.gov/).								
543	• There must be only one output variable per file.								
544	• For the data that are a function of longitude and latitude, only regular grids (grids								
545	representable as a Cartesian product of longitude and latitude axes) are allowed.								
546	• The file names have to follow the PMIP2 file name convention and be unique (see the								
547	PMIP2 website).								
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550	List of Tables								
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552	Table 1: Details of NetCDF data packages provided to facilitate PlioMIP Phase 2								
553	experiments.								
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555	Table 2: The colour (for albedo) and texture translations for the soil orders used in the								
556	modelling of Late Pliocene soils, based upon HadCM3 classification.								
557									
558	Table 3: Details of all experiments proposed in PlioMIP Phase 2. *By ice sheet regions we								
559	mean the land masses of Greenland and Antarctica (not the areas of ice specified within the								
560	ice-masks). ** Where ice retreat (i.e. the change from pre-industrial ice to Pliocene ice) leaves								
561	information gaps in soils, please extrapolate modern soil values from nearest grid square. ¹ For	For For							
562	experiment Eoi ⁴⁰⁰ , Eoi ³⁵⁰ and Eoi ⁴⁵⁰ this may be using the standard or enhanced Pliocene LSM.	For							
563	² For simplicity of approach we assume that all forcing factorisation experiments will only use	For							
		FOL							

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564	the standard rather than enhanced data sets. ³ Prescribed static vegetation is also an option.	(Formatted: Superscript
565	Red = Core experiment (compulsory), Blue = Tier 1 and 2 sensitivity experiments (optional).		
566	4P4F = Pliocene for Future; P4P = Pliocene for Pliocene. See also Appendix 1 in Supplementary		
567	Information. Details of all experiments proposed in PlioMIP Phase 2. *By ice sheet regions we		
568	mean the land masses of Greenland and Antarctica (not the areas of ice specified within the		
569	ice-masks). ** Where ice retreat (i.e. the change from pre-industrial ice to Pliocene ice) leaves		
570	information gaps in soils, please extrapolate modern soil values from nearest grid square. ⁴ For	(Formatted: Superscript
571	experiment Eoi400 this may be using the standard or enhanced Pliocene LSM. ² For simplicity	(Formatted: Superscript
572	of approach we assume that all forcing factorisation experiments will only use the standard		
573	rather than enhanced datasets. ³ Prescribed static vegetation is also an option. Red = Core		Formatted: Superscript
574	experiment, Blue = Tier 1 and 2 sensitivity experiments. $^{4}_{A}$ P4F = Pliocene for Future; P4P =	(Formatted: Superscript
575	Pliocene for Pliocene. See also Appendix 1 in Supplementary Information.		
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581	List of Figures		
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583	Figure 1. The PMIP Triangle which illustrates three possible contributions to datamodel-		
584	modeldata discrepancy, and has at its vertices model physics (structural and parameter		
585	uncertainty), model boundary conditions and proxy data uncertainty (Haywood et al.,		
586	2013a).		
587			
588	Figure 2: Experimental design strategy adopted for PlioMIP Phase 2. Core experiments will		
589	be completed by all model groups. Tier 1 and Tier 2 in either "Pliocene4Future" or		
590	"Pliocene4Pliocene" describe a series of sensitivity tests (Tier 1 being a higher priority for		
591	completion than Tier 2). Please note that Pliocene4Future Tier 1 experiment Pre-Industrial		
592	CO ₂ 400 also appears as a Tier 2 Pliocene4Pliocene experiment (Pre-Ind+PlioCO ₂). See Table		
593	3 for the naming convention and further details of all PlioMIP Phase 2 experiments, as well		
594	as Appendix 1 in Supplementary Information.		

595		
596 507	Figure 3: PRISM4 palaeogeography (enhanced) including topography/bathymetry (m) over	
597	the ice sheets (left). PRISM4 topographic and bathymetric anomaly (m) from modern	
598	(ETOPO1: right). Red boxes highlight the Canadian archipelago and Bering Strait as closed in	
599 600	both the standard and enhanced boundary condition data sets.	
601	Figure 4: PRISM4 land-sea mask (enhanced version) showing Greenland and Antarctic Ice	
602	Sheets distribution. Canadian archipelago and Bering Strait closed (red boxes) in both the	
603	standard and enhanced boundary condition data sets.	
604		
605	Figure 5: Modern and Pliocene (PRISM4) fractional lake coverage data set (Pound et al.	
606	2014). Modern data is based upon the FAO/UNESCO modern soil map (Version 3.6).	
607	Figure 5: PRISM4 fractional lake coverage data set (Pound et al. 2014).	
608		
609	Figure 6: Pound et al. (2014) data set of global modern and Pliocene soil types (shown on	
610	the enhanced PlioMIP2 land-sea mask). Modern data is based upon the FAO/UNESCO	
611	modern soil map (Version 3.6).	
612	Figure 6: Pound et al. (2014) data set of global Pliocene soil types shown on the enhanced	
613	PlioMIP2 land sea mask.	
614		
615	Figure 7: Initial PRISM4 sites being investigated to generate time slice proxy data for model	
616	evaluation in PlioMIP Phase 2.	
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650 Tables

Data_set Name		Description		
Plio_std.zip	Plio_std_topo_v1.0.nc	PRISM4	Pliocene	palaeogeography
	Plio_std_LSM_v1.0.nc	reconstruction including new topograp		w topography and

	Plio_std_soil_v1.0.nc Plio_std_lake_v1.0.nc Plio_std_mbiome_v1.0.nc (only for models that cannot predict vegetation) Plio_std_icemask_v1.0.nc	ice sheets; however a modern land-sea mask has been applied. No information on bathymetry is provided. Fractional coverage of lakes as well as the global distribution of soil characteristics is also provided. Salzmann et al. (2008) Pliocene biome reconstruction is also available and has been adapted to fit the new ice mask.
Plio_enh.zip	Plio_enh_topo_v1.0.nc Plio_enh_LSM_v1.0.nc Plio_enh_soil_v1.0.nc Plio_enh_lake_v1.0.nc Plio_enh_mbiome_v1.0.nc (only for models that cannot predict vegetation) Plio_enh_icemask_v1.0.nc	Full PRISM4 Pliocene palaeogeography reconstruction including new topography, bathymetry, ice sheets and land-sea mask. Fractional coverage of lakes as well as the global distribution of soil characteristics also provided (soil distributions altered to match enhanced land-sea mask). Salzmann et al. (2008) Pliocene biome reconstruction is also available and has been modified to fit the new palaeogeographic and ice reconstruction.
Modern_std.zip	Modern_std_topo_v1.0.nc Modern_std_LSM_v1.0.nc Modern_std_soil_v1.0.nc Modern_std_mbiome_v1.0.nc	Modern files for reference purposes only. Full modern palaeogeography reconstruction including present-day topography, bathymetry, ice sheets and land-sea mask derived from ETOPO1. Global distribution of soil and vegetation characteristics using the same descriptors as the Pliocene reconstruction provided to aid the implementation of Pliocene soil and vegetation characteristics. Soil file also contains the lake distribution and ice-mask information.

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 Table 1: Details of NetCDF data packages provided to facilitate PlioMIP Phase 2

654 experiments.

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Soil Group

Soil Colour

Texture

Albedo

658		Gelisol (31)	Intermediate	Medium	0.17	
659		Histosol (32)	Dark	Fine	0.11	
660		Spodosol (33)	Intermediate	Medium/Coarse	0.17	
661		Oxisol (34)	Intermediate	Fine/Medium	0.17	
662		Vertisol (35)	Dark	Fine	0.11	
663		Aridisol (36)	Light	Coarse	0.35	
664		Ultisol (37)	Intermediate	Fine/Medium	0.17	
665		Mollisol (38)	Dark	Medium	0.35	
666		Alfisol (39)	Intermediate	Medium	0.17	
667						
668	Table 2: The	colour (for albe	do) and texture transla	ations for the soil orde	ers used in the	Formatted: Left
669	modelling of	Late Pliocene so	oils, based upon HadCl	V3 classification.		
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ID	Description	LSM ^{1,2}	торо.	SOILS	LAKES	ICE	VEGETATION ³	CO2	STATUS <u>:</u> Tier 1 or 2 (T) & P4F/P4P ⁴	Formatted Table
E ²⁸⁰	Pre-industrial experiment as per control simulation in PlioMIP2 experiment.	Modern	Modern	Modern	Modern	Modern	Dynamic	280	CORE	Formatted: Font: Bold
E ⁴⁰⁰	Pre-industrial experiment as per control simulation in core PlioMIP2 experiment - CO_2 400 ppmv.	Modern	Modern	Modern	Modern	Modern	Dynamic	400	T1: P4F - Tier 2: P4P	
E ⁵⁶⁰	Pre-industrial experiment as per control simulation in core PlioMIP2 experiment - CO ₂ 560 ppmv.	Modern	Modern	Modern	Modern	Modern	Dynamic	560	T2: P4F	
0 ²⁸⁰	Pre-industrial experiment as per control simulation in core PlioMIP2 experiment, however topography (including soils and lakes) is set to Pliocene values outside of ice sheet regions. <u>The land masses of Greenland and Antarctica should have pre-industrial boundary conditions (see Fig. Stal</u> ^a .	Modern	Pliocene	Pliocene	Pliocene	Modern	Dynamic	280	T2: P4P	
Ei ²⁸⁰	Pre-industrial experiment as per control simulation in <u>core PlioMIP2</u> experiment, however the ice configurations on Greenland and Antarctica are set to be Pliocene.**	Modern	Modern	Modern	Modern	Pliocene	Dynamic	280	T2: P4P	
0 ⁴⁰⁰	Pliocene experiment as per control simulation in core PlioMIP2 experiment, however ice sheets on Greenland and Antarctica set to modern.	Modern	Pliocene	Pliocene	Pliocene	Modern	Dynamic	400	T2: P4P	
Ei ⁴⁰⁰	Pliocene experiment as per control simulation in Core PlioMIP2 experiment. Topography outside of the ice sheet regions set to modern. Soils and lakes are also modern in this experiment.	Modern	Modern	Modern	Modern	Pliocene	Dynamic	400	T2: P4P	
oi ²⁸⁰	Pliocene experiment as per control simulation in Core PlioMIP2 experiment $- CO_2 280$ ppmv	Modern	Pliocene	Pliocene	Pliocene	Pliocene	Dynamic	280	T2: P4P	
0i ⁴⁰⁰	Pliocene experiment as per control simulation in Core PlioMIP2 experiment.	Pliocene <mark>/ or</mark> Modern	Pliocene	Pliocene	Pliocene	Pliocene	Dynamic	400	CORE	Formatted: Font: Bold Formatted: Font: Bold
0i ⁴⁵⁰	Pliocene experiment as per control simulation in Core PlioMIP2 experiment - CO_2 @ 450 ppmv)	Pliocene <mark>/-or</mark> Modern	Pliocene	Pliocene	Pliocene	Pliocene	Dynamic	450	T1: P4F - T1: P4P	ronnatter ront. Doit
0i ³⁵⁰	Pliocene experiment as per control simulation in Core PlioMIP2 experiment , but with CO_2 set to 350 ppmv)	Pliocene <mark>/-or</mark> Modern	Pliocene	Pliocene	Pliocene	Pliocene	Dynamic	350	<u>T1: P4F -</u> T1: P4P	
oi⁵⁶⁰	Pliocene experiment as per control simulation in Core PlioMIP2 experiment , but with CO ₂ set to 560 ppmv)	Pliocene or Modern	Pliocene	Pliocene	Pliocene	Pliocene	Dynamic	560	T2: P4F	

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687 Table 3: Details of all experiments proposed in PlioMIP Phase 2. *By ice sheet regions we mean the land masses of Greenland and Antarctica (not the areas of ice specified

688 within the ice-masks). ** Where ice retreat (i.e. the change from pre-industrial ice to Pliocene ice) leaves information gaps in soils, please extrapolate modern soil values

from nearest grid square. ¹For experiments Eoi⁴⁰⁰, Eoi³⁵⁰ and Eoi⁴⁵⁰ this may be using the standard or enhanced Pliocene LSM. ²For simplicity of approach we assume that all

690 forcing factorisation experiments will only use the standard rather than enhanced data_sets. ³Prescribed static vegetation is also an option. Red background and bold text =

691 Core experiments (compulsory), Blue = Tier 1 and 2 sensitivity experiments (optional). ⁴P4F = Pliocene for Future; P4P = Pliocene for Pliocene. See also Appendix 1 in

692 Supplementary Information.

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693 Figures



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Figure 1. The PMIP Triangle which illustrates three possible contributions to <u>datamodel-modeldata</u> discrepancy, and has at its vertices model physics (structural and parameter uncertainty), model boundary conditions and proxy data uncertainty (Haywood et al., 2013a).
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- 706 be completed by all model groups. Tier 1 and Tier 2 in either "Pliocene4Future" or
- 707 "Pliocene4Pliocene" describe a series of sensitivity tests (Tier 1 being a higher priority for
- 708 completion than Tier 2). Please note that Pliocene4Future Tier 1 experiment Pre-Industrial

- 709 CO₂ 400 also appears as a Tier 2 Pliocene4Pliocene experiment (Pre-Ind+PlioCO₂). See Table
- 710 3 for the naming convention and further details of all PlioMIP Phase 2 experiments, as well
- 711 as Appendix 1 in the Supplementary Information.
- 712
- 713



714

- 715 **Figure 3**: PRISM4 palaeogeography (enhanced) including topography/bathymetry (m) over
- 716 the ice sheets (left). PRISM4 topographic and bathymetric anomaly (m) from modern
- 717 (ETOPO1: right). Red boxes highlight the Canadian archipelago and Bering Strait as closed in
- 718 both the standard and enhanced boundary condition data sets.

719



721 Figure 4: PRISM4 land-sea mask (enhanced version) showing Greenland and Antarctic Ice

722 Sheets distribution. Canadian archipelago and Bering Strait closed (red boxes) in both the

723 standard and enhanced boundary condition data sets.



725







PRISM4 Lake Fraction (enhand



Figure 6: Pound et al. (2014) data set of global <u>modern and</u> Pliocene soil types (shown on
the enhanced PlioMIP2 land-sea mask). <u>Modern data is based upon the FAO/UNESCO</u>
modern soil map (Version 3.6).



735

Figure 7: Initial PRISM4 sites being investigated to generate time slice proxy data for model

- 737 evaluation in PlioMIP Phase 2.
- 738
- 739

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756	D. J. Lunt acknowledges NERC grant NE/H006273/1.	
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