Editor Comments:

The Title is a bit awkward. I think it should read "Estimates of late middle Eocene pCO2 based on stomatal density of modern and fossil Nageia leaves". This better incorporates the contents of the manuscript with correct grammar. Also, I think ~40 Ma is better described as late middle Eocene, and I do not think Gaertner needs inclusion as the genus, while not well known, is well established.

The title has been changed as suggested.

Lines 44-47: This is over-referenced and arguably incorrect. I would just put "In particular, there are few reconstructions for the late middle Eocene." (And then cite those papers that show this aspect).

Changed as suggested.

Line 128: I think this should be "… into two "sections", …" (I assuming here that sections is a biological classification term).

Changed as suggested.

Line 134: Needs space.

Changed as suggested.

Lines 161-162: I think this should be "… surfaces of leaves might be used …"

Changed as suggested.

Lines 173-174: This should be "… hours, the reaction was stopped when specimen fragments … and transparent. The …" (There should be no semicolons in this case).

Changed as suggested.

Line 175: Should be "After, the ..."

Changed as suggested.

Line 178: Needs period as already has an "and" (i.e., "… 30 min). The …")

Changed as suggested.

A Lines 192-197: This needs rewriting as important information is missing. It should be something like “Further information on the sections is provided by Lie et al. (2015). Importantly, the formations span a depositional age of approximately XX to YY Ma, or late middle Eocene (REF). This has been determined HOW? (1-2 sentences). Now start new paragraph “Macrofossil …” (Are other approaches the same as for modern leaves? This is not crystal clear right now).

Lines 192-194 are changed as suggested. The approaches for fossil are different from modern ones, so we add “directly” in front of “treated” to make the sentence clearer.

Line 216: Remove space.

Changed as suggested.

Line 229: Spell-out “Figure”

Changed as suggested.

Line 245: Indent paragraph.

Changed as suggested.

Line 293: Should be “For modern Nageia, we find that … increase, but that …”

Changed as suggested.

Lines 295-296: Should be “… case has been observed for some flora.”

Changed as suggested.

Line 300: Can remove “herein …”. Just end the sentence after “cells.”

Changed as suggested.

Line 301: Start new paragraph and indent.

Changed as suggested.

Lines 312-313: Tense. Should be “… was young or mature, or grew in a sunny or shady environment”.

Changed as suggested.

Lines 321-322: Should be “…, which is fairly close to GEOCARB III predictions …”

Changed as suggested.

Lines 326-327: Should be "… were generally lower than during much of the Cretaceous, but probably also decreased significantly from the early to late Eocene (REF). Here, I would fix the references. For example, two sets of references not needed, and Zachos et al. 2001 does not really discuss pCO2.

Changed as suggested.

Lines 330-334: I would change to "However, there is a wide range of estimates for the Eocene (REF)." (Then remove the rest, as it is not really pertinent).

Changed as suggested.

B Lines 336-380: This needs to be condensed and rewritten. All that should be here are past estimates for the middle to late Eocene, and how the Nageia estimates compare. However, there is a critical piece of information that needs to be tied in with the comment above. The world was dynamic in the Paleogene, including in the late middle Eocene, when the MECO occurred. Thus, the exact age matters, and it is possible that a values may differ because of slight offsets in time.

Changed as suggested.

In addition, we made a few changes as follows:

Line 237-238, add two spaces after "=".

Line 349, we use "to" substitute for "and".

~~The~~ Estimates of late middle Eocene pCO_2 ~~estimates of the late Eocene in South China~~ based on stomatal density of modern and fossil *Nageia* ~~Gaertner~~ leaves

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Abstract:

Atmospheric pCO_2 concentrations have been estimated for intervals of the Eocene using various models and proxy information. Here we reconstruct late middle Eocene (42.0-38.5~~~ 40.3~~ Ma) pCO_2 based on the fossil leaves of *Nageia maomingensis* Jin et Liu collected from the Maoming Basin, Guangdong Province, China. We first determine relationships between atmospheric pCO_2 concentrations, stomatal density (SD) and stomatal index (SI) using "modern" leaves of *N. motleyi* (Parl.) De Laub, the nearest living species to the Eocene fossils. This work indicates that the SD inversely responds to pCO_2, while SI has almost no relationship with pCO_2. Eocene pCO_2 concentrations can be reconstructed based on a regression approach and the stomatal ratio method by using the SD. The first approach gives a pCO_2 of 351.9 ± 6.6 ppmv, whereas the one based on stomatal ratio gives a pCO_2 of 537.5 ± 56.5 ppmv. Here, we

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explored the potential of *N. maomingensis* in pCO_2 reconstruction and obtained different results according to different methods, providing a new insight for the reconstruction of paleoclimate and paleoenvironment in conifers.

Keywords: pCO_2, late middle Eocene, *Nageia*, Maoming Basin, South China.

1 Introduction

The Eocene (55.8-33.9 Ma) generally was much warmer than present-day, although temperatures varied significantly across this time interval (Zachos et al., 2008). Climate of the early Eocene was extremely warm, particularly during the early Eocene Climatic Optimum (EECO; 51 to 53 Ma), and the Paleocene-Eocene Thermal Maximum (PETM; ~55.9 Ma). However, global climatic conditions cooled significantly by the middle to late Eocene (40 to 36 Ma). Indeed, small, ephemeral ice-sheets and Arctic sea ice likely existed during the latest Eocene (Moran et al., 2006; Zachos et al., 2008).

Many authors have suggested that changes in temperature during the Phanerozoic were linked to atmospheric pCO_2 (Petit et al., 1999; Retallack, 2001; Royer, 2006). Central to these discussions are records across the Eocene, as this epoch spans the last major change from a “greenhouse” world to an “icehouse” world. The Eocene pCO_2 record remains incomplete and debated (Kürschner et al., 2001; Royer et al., 2001; Beerling et al., 2002; Greenwood et al., 2003; Royer, 2003). Most pCO_2

reconstructions have focused on the Cretaceous-Tertiary and Paleocene-Eocene boundaries (65 to 50 Ma~~; Koch et al., 1992; Stott, 1992; Sinha and Stott, 1994; Royer et al., 2001; Beerling and Royer, 2002; Nordt et al., 2002; Royer, 2003; Fletcher et al., 2008; Roth-Nebelsick et al., 2012; 2014; Grein et al., 2013; Huang et al., 2013; Maxbauer et al.,2014~~) and the middle Eocene. ~~(Maxbauer et al., 2014), while few reconstructions were conducted at the late Eocene~~ In particular, there are few reconstructions for the late middle Eocene (Pagani et al., 2005; Maxbauer et al., 2014). In addition, the pCO_2 reconstruction results have varied based on different proxies. Various methods having been used in pCO_2 reconstruction mainly include the computer modeling methods: GEOCARB-I, GEOCARB-II, GEOCARB-III, GEOCARB-SULF and the proxies: ice cores, paleosol carbonate, phytoplankton, nahcolite, Boron, and stomata parameters.

The abundance of stomatal cells can be measured on modern leaves and well-preserved fossil leaves. Various plants show a negative correlation between atmospheric CO_2 concentration and stomatal density (SD), stomatal index (SI), or both. As such, these parameters have been determined in fossil leaves to reconstruct past pCO_2; examples include *Ginkgo* (Retallack, 2001, 2009a; Beerling et al., 2002; Royer, 2003; Kürschner et al., 2008; Smith et al., 2010), *Metasequoia* (Royer, 2003; Doria et al., 2011), *Taxodium* (Stults et al., 2011), *Betula* (Kürschner et al., 2001; Sun et al., 2012), *Neolitsea* (Greenwood et al., 2003), and *Quercus* (Kürschner et al., 1996, 2001), *Laurus* and *Ocotea* (Kürschner et al., 2008). Recently, positive correlations between stomatal index or stomatal frequency and pCO_2 have been reported based on fossil *Typha* and *Quercus* (Bai et al., 2015; Hu et al., 2015). However, the tropical and subtropical moist broadleaf forest conifer tree *Nageia* has not been used previously in paleobotanical estimates of pCO_2 concentration.

Herein, we firstly document correlations between stomatal properties and atmospheric CO_2 concentrations using leaves of the extant species *Nageia motleyi* (Parl.) De Laub. that were collected over the last two centuries. This provides a training dataset for application to fossil representatives of *Nageia*. We secondly measure stomatal parameters on fossil *Nageia* leaves from ~~the~~ late middle Eocene of South China to estimate past CO_2 levels. The work provides further insights for discussing Eocene climate change.

2 Background

2.1 Stomatal proxy in pCO_2 research

Stomatal information gathered from careful examination of leaves has been widely used for reconstructions of past pCO_2 concentrations (Beerling and Kelly, 1997; Doria et al., 2011). The three main parameters are stomatal density (SD), which is expressed as the total number of stomata divided by area, epidermal density (ED), which is expressed as the total number of epidermal cells per area, and the stomatal index (SI), which is defined as the percentage of stomata among the total number of cells within an area [SI = SD × 100 / (SD+ED)]. Woodward (1987) considered that both SD and SI had inverse relationships with atmospheric CO_2 during the development of the leaves. Subsequently, McElwain (1998) created the stomatal ratio (SR) method to reconstruct pCO_2. SR is a ratio of the stomatal density or index of a fossil [$SD_{(f)}$ or $SI_{(f)}$] to that of corresponding nearest living equivalent [$SD_{(e)}$ or $SI_{(e)}$], expressed as follows:

$$SR = SI_{(e)} / SI_{(f)}) \quad (1)$$

The stomatal ratio method is a semi-quantitative method of reconstructing pCO_2 concentrations under certain standardizations. An example is the “Carboniferous standardization” (Chaloner and McElwain, 1997), where one stomatal ratio unit equals two RCO_2 units:

$$SR = 2\ RCO_2 \quad (2)$$

and the value of RCO_2 is the pCO_2 level divided by the pre-industrial atmospheric level (PIL) of 300 ppm (McElwain, 1998) or that of the year when the nearest living equivalent (NLE) was collected (Berner, 1994; McElwain, 1998):

$$RCO_2 = C_{(f)} / 300 \text{ or } RCO_2 = C_{(f)} / C_{(e)} \quad (3)$$

The estimated pCO_2 level can then be expressed as follows:

$$C_{(f)} = 0.5 \times C_{(e)} \times SD_{(e)} / SD_{(f)} \text{ or } C_{(f)} = 0.5 \times C_{(e)} \times SI_{(e)} / SI_{(f)} \quad (4)$$

where $C_{(f)}$ is the pCO_2 represented by the fossil leaf, and $C_{(e)}$ is the atmospheric CO_2 of the year when the leaf of the NLE species was collected (McElwain and Chaloner, 1995, 1996; McElwain 1998). The equation adapts to the pCO_2 concentration prior to Cenozoic.

Another standardization, the “Recent standardization” (McElwain, 1998), is expressed as one stomatal ratio unit being equal to one RCO_2 unit:

$$SR = 1\ RCO_2 \quad (5)$$

According to the equations stated above, the pCO_2 concentration can be expressed as:

$$C_{(f)} = C_e \times SD_{(e)} / SD_{(f)} \text{ or } C_{(f)} = C_e \times SI_{(e)} / SI_{(f)} \quad (6)$$

This standardization is usually used for reconstruction based on Cenozoic fossils (Chaloner and McElwain, 1997; McElwain, 1998; Beerling and Royer, 2002).

Kouwenberg et al. (2003) proposed some special stomatal quantification methods for conifer leaves with stomata arranged in rows. The stomatal number per Length (SNL) is expressed as the number of abaxial stomata plus the number of adaxial stomata divided by leaf length in millimeters. Stomatal rows (SRO) is expressed as the number of stomatal rows in both stomatal bands. Stomatal density per length (SDL) is expressed as the equation SDL = SD × SRO. True stomatal density per length (TSDL) is expressed as the equation TSDL = SD × band width (in millimeters). The band width on *Nageia motleyi* leaves was measured as leaf blade width.

2.2 Review of extant and fossil *Nageia*

The genus *Nageia*, including seven living species, is a special group of Podocarpaceae, a large family of conifers mainly distributed in the southern hemisphere. *Nageia* has broadly ovate-elliptic to oblong-lanceolate, multiveined (without a midvein), spirally arranged or in decussate, and opposite or subopposite leaves (Cheng et al., 1978; Fu et al., 1999). Generally, *Nageia* is divided into two sections, *Nageia* Sect. *Nageia* and *Nageia* Sect. *Dammaroideae* (Mill 1999, 2001). Both sections are mainly distributed in southeast Asia and Australasia from north latitude 30 ° to nearly the equator (Fu, 1992; Fig. 1). Four species of the *N*. section *Nageia* -- *Nageia nagi* (Thunberg) O. Kuntze, *N. fleuryi* (Hickel) De Laub., *N. formosensis* (Dummer) C. N. Page, and *N. nankoensis* (Hayata) R. R. Mill -- have hypostomatic leaves where stomata only occur on the abaxial side. One species of this section -- *N. maxima* (De Laub.) De Laub. -- is characterized by amphistomatic leaves, but where only a few stomata are found on the adaxial side (Hill and Pole, 1992; Sun, 2008). Both *N. wallichiana* (Presl) O. Kuntze and *N. motleyi* of the *N*. section *Dammaroideae* are amphistomatic with abundant stomata distributed on both sides of the leaf. This is especially true for *N. motleyi*, which has approximately equal stomata numbers on both surfaces (Hill and Pole, 1992; Sun, 2008).

The fossil record of *Nageia* can be traced back to the Cretaceous. Krassilov (1965) described *Podocarpus* (*Nageia*) *sujfunensis* Krassilov from the Lower Cretaceous of Far East Russia. Kimura et al. (1988) reported *Podocarpus* (*Nageia*) *ryosekiensis* Kimura, Ohanaet Mimoto, an ultimate leafy branch bearing a seed, from the Early Barremian in southwestern Japan. In China, a Cretaceous petrified wood, *Podocarpus* (*Nageia*) nagi Pilger, was discovered from the Dabie Mountains in central Henan, China (Yang et al., 1990). Jin et al. (2010) reported a upper Eocene *Nageia* leaf named *N. hainanensis* Jin, Qiu, Zhu et Kodrul from the Changchang Basin of Hainan Island, South China. Recently, Liu et al. (2015) found another leaf species *N. maomingensis* Jin et Liu from upper middle Eocene of Maoming Basin, South China.

Although some of the *Nageia* fossil materials described in the above studies (Krassilov, 1965; Jin et al., 2010; Liu et al., 2015) have well-preserved cuticles, these studies are mainly concentrated on morphology, systematics and phytogeography.

Here we try to reconstruct the pCO_2 concentration based on stomatal data of *Nageia maomingensis* Jin et Liu. Among the modern *Nageia* species mentioned above, *N. motleyi* was considered as the NLE species of *N. maomingensis* (Liu et al., 2015). However, because of the species-specific inverse relationship between atmospheric CO_2 partial pressure and SD (Woodward and Bazzaz, 1988), it is necessary to explore whether the SD and SI of *N. motleyi* show negative correlations with the CO_2 concentration before applying the stomatal method. Both *N. maomingensis* and *N. motleyi* are amphistomatic, suggesting that both upper and lower surfaces of ~~the leaf are needed~~leaves might be used to estimate the pCO_2 concentrations.

3 Material and methods

3.1 Extant leaf preparation

We examined 12 specimens of extant *Nageia motleyi* from different herbaria (Table 1). We removed one or two leaves from each specimen, and took three fragments (0.25 mm^2) from every leaf (Fig. 2a) and numbered them for analysis.

The numbered fragments were boiled for 5-10 min in water. Subsequently, after being macerated in a mixed solution of 10% acetic acid and 10% H_2O_2 (1:1) and heated in the thermostatic water bath at 85 C for 8.5 hours~~;~~, the reaction was stopped when the specimen~~s~~ fragments turned white and semitransparent.~~;~~ The cuticles were then rinsed with distilled water until the pH of the water became neutral. After, ~~that~~ the cuticles were treated in Schulze's solution (one part of potassium chlorate saturated solution and three part of concentrated nitric acid) for 30 min, rinsed in water, and then treated with 8% KOH (up to 30 min). ~~and the~~ The abaxial and adaxial cuticles were separated with a hair mounted on needle. Finally, the cuticles were stained with 1% Safranin T alcoholic solution for 5 min, sealed with Neutral Balsam and observed under LM.

3.2 Fossil leaf preparation

Maoming Basin (21°42'33.2"N, 110°53'19.4"E) is located in southwestern Guangdong, South China including Cretaceous and Tertiary strata. Tertiary strata are fluvial and lacustrine sedimentary units, divided into the Gaopengling, Laohuling, Shangcun, Huangniuling and Youganwo formations in descending order, aged from late Eocene to early Oligocene (Wang et al., 1994).

Four fossil leaves of *Nageia maomingensis* were recovered from the Youganwo (MMJ1-001) and Huangniuling (MMJ2-003, MMJ2-004 and MMJ3-003) formations of Maoming Basin, South China ~~(Fig. 1B, 1C in Liu et al., 2015)~~. Further information on the sections is provided by Liu et al. (2015). Importantly, the formations span a depositional age of approximately 42.0 to 38.5 Ma which was considered as late

Eocene by Wang et al. (1994), but it can be recognized as late middle Eocene according to Walker and Geissman (2009). The age from Youganwo to Huangniuling formations is late Eocene (~ 40.3 Ma). Precise information regarding locations is provided by Liu et al., (2015).

Macrofossil cuticular fragments were taken from the middle part of each fossil leaf (Fig. 2c) and directly treated with Schulze's solution for approximately 1h and 5–10% KOH for 30 min (Ye, 1981). The cuticles were observed and photographed under a Carl Zeiss Axio Scope A1 light microscope (LM). All fossil specimens and cuticle slides are housed in the Museum of Biology of Sun Yat-sen University, Guangzhou, China.

3.3 Stomatal counting strategy and calculation methods

The basic stomatal parameters, SD, ED and SI, were counted based on analyzing pictures taken with a light microscope (LM). A total of 2816 pictures (200× magnification of Zeiss LM) of cuticles from 21 leaves of *N. motleyi* were counted. Each counting field was 0.366 mm^2. We used a standard sampling protocol (Poole and Kürschner, 1999), counting all full stomata in the image plus stomata straddling the left and top margins, as presented in Figure 2(b), and (d).

The SNL, SRO, SDL, and TSDL were also determined based on LM images. A total of 2293 pictures (200× magnification of Zeiss LM) of the cuticles from 21 leaves of *N. motleyi* were counted. Each counting field was 0.366 mm^2. None of the aforementioned counting areas overlapped and they were larger than the minimum area (0.03 mm^2) for statistics (Poole and Kürschner, 1999). In this study, the stomatal data of both surfaces are applied in pCO_2 reconstruction because both the fossil and NLE species are amphistomatic.

4 Results

4.1 Correlations between the CO_2 concentrations and stomatal parameters of *Nageia motleyi*

The SD and SI data of the adaxial sides of *N. motleyi* leaves are presented in Table 2. The SDs and SIs average 62.28 mm^{-2} and 3.30 %, respectively. However, the SDs and SIs data of the abaxial sides, summarized in Table 3, give higher average values (70.03 mm^{-2} in SDs and 3.90 % in SIs) than those from the adaxial sides. The combined SD and SI of the adaxial and abaxial surfaces average 66.14 $mm^{-2}$and 3.60 %, respectively (table 4).

Figure 3 shows the relationships between the stomatal parameters (SD and SI) of modern *N. motleyi* and the atmospheric CO_2 concentration (SD-CO_2 relationship and SI-CO_2 relationship). R^2 values in the SD-CO_2 relationship from the adaxial and abaxial surfaces of *N. motley* are up to 0.4667 and 0.3824 (Fig. 3a, b), suggesting that the stomatal densities of *N. motleyi* are inverse to the CO_2 concentrations. However,

Fig. 3c and d indicate no relationship between the SIs and CO_2 concentrations for the extremely low level of the R^2 values (0.2558 and 0.0248). Figs. 3e and 3f based on the combined data also show that SD inversely responds to the atmospheric CO_2 concentration (R^2 = 0.4421), while SI has almost no relationship with the atmospheric CO_2 concentration (R^2 = 0.1177).

The mean values of SNL, SDL and TSDL are 9.81, 326.39 and 1226.93 no.·mm^{-1}, respectively (Table 5). Fig. 4 shows the relationships between SNL (SDL, TSDL) and CO_2 concentrations. The low R^2 values in the Fig. 4a and 4c indicate that SNL (R^2 = 0.0643) and TSDL (R^2 = 0.0788) have no relationship with the CO_2 concentration in this study. Fig. 4b shows that there is a weak reverse relevance between SDL and the CO_2 concentration (R^2 = 0.3154).

Compared with the SDL method, the SD-based method shows a larger R^2 value, indicating a stronger relevance between the SD and CO_2 concentrations. In this study, the pCO_2 is reconstructed based on the regression equations of SD-CO_2 relationship. Additionally, the stomatal ratio method can be also used in estimating pCO_2 concentration of the late middle Eocene based on stomatal densities (SDs) of the fossil species *N. maomingensis* and extant species *N. motleyi*. The SD results of specimen No. 18328 are selected to reconstruct the pCO_2 concentration, because they are closest to the fitted equations in Fig. 3. This specimen was collected by Neth. Ind. For. Service from Riau on Ond. Karimon, Archipel. Ind., Malaysia, in 1934 at an altitude of 5 m and CO_2 concentration of 306.46 ppmv (Brown, 2010).

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4.2 The pCO_2 estimates results

4.2.1 The regression approach

The summary of stomatal parameters of the fossil *Nageia* and reconstruction results are provided in Tables 6–8. The mean SD and SI values of the adaxial surface are 44.5 mm^{-2} and 1.8 %, respectively (Table 6). The mean SD and SI values of the abaxial surface are 49.8 mm^{-2} and 2.07 %, respectively (Table 7).

Based on the regression approach, the pCO_2 was reconstructed as 351.9 ±6.6 ppmv and 365.6 ±7.6 ppmv according to the SD of adaxial and abaxial sides. The combined SD value is an average of 46.6 mm^{-2} (Table 8), giving the reconstructed pCO_2 of 358.1 ±5.0 ppmv.

4.2.2 The stomatal ratio method

Mean SR value of the adaxial side (SR=1.69 ±0.18) is a little larger than that of the abaxial side (SR=1.60 ±0.11) in fossil *Nageia* leaves (Tables 6 and 7). The pCO_2 reconstruction results are537.5 ±56.5 ppmv (Table 6) and 496.1 ±35.7 ppmv (Table 7) based on the adaxial and abaxial cuticles, respectively. Based on the combined SD of both leaf sides, the pCO_2 result is 519.9 ±35.0 ppmv.

The partial pressure of CO_2 decreases with elevation (Gale, 1972). Jones (1992) proposed that the relationship between elevation and partial pressure in the lower atmosphere can be expressed as $P = -10.6E + 100$, where E is elevation in kilometers and P is the percentage of partial pressure relative to sea level. Various studies corroborate that SI and SD of many plants have positive correlations with altitude (Körner and Cochrane, 1985; Woodward, 1986; Woodward and Bazzaz, 1988; Beerling et al., 1992; Rundgren and Beerling, 1999) while they are negatively related to the partial pressure of CO_2 (Woodward and Bazzaz, 1988). Therefore, it is essential to take elevation calibration into account during pCO_2 concentration estimates. However, Royer (2003) pointed out that it is unnecessary to provide this conversion when trees lived at <250 m in elevation. In this paper, the nearest living equivalent species, *Nageia motleyi*, grows at 5 m in elevation with P =99.9, suggesting that CO_2 concentration estimates were only underestimated by 0.1%. Consequently, no correction is needed for the reconstruction result in this study. After being projected into a long-term carbon cycle model (GEOCARB III; Berner and Kothavalá, 2001), the results of this study compares well with CO_2 concentrations for corresponding age within their error ranges (Fig. 5).

5 Discussion

5.1 Stomatal parameters response to CO_2

For modern *Nageia*, ~~Here,~~ we find that SD decreases as atmospheric CO_2 concentrations increase, ~~however,~~but that SI does not. Generally, SI is more sensitive in response to the atmospheric CO_2 concentration than SD (Beerling, 1999; Royer, 2001). However, the reverse case has been observed for some flora.~~is not unfound.~~

For example, Kouwenberg et al. (2003) reported that SD is better than SI in reflecting the negative relationships with CO_2 in conifer needles, accounting for the special paralleled mode of the ordinary epidermal and stomatal formation. Although *Nageia* is broad-leaved rather than needle-leaved, it also has well paralleled epidermal cells. ~~herein showing the different relationships between CO_2 and SD or SI.~~

Compared with SD, the SDL has weaker correlation with CO_2 at a smaller R^2. The SNL and TSDL have no response to the change of CO_2. The insensitivity of SNL, SDL and TSDL might account for the characters of broad-leaved leaf shape and paralleled epidermal cells. The SNL should be applied to conifer needles with single file of stomata (Kouwenberg et al., 2003). The SDL and TSDL were considered as the most appropriate method when the stomatal rows grouped in bands in a hypo- or amphistomatal conifer needle species (Kouwenberg et al., 2003). Considering all the stomatal parameters above, SD appears to be the most sensitive to CO_2.

The SD-CO_2 correlation shows one value from leaf No. 40798 offset from the others. The SI-CO_2 correlation shows different offset values in different leaf sides. The offset values might be affected by leaf maturity and light intensity. However, it is hard to distinguish whether a fossil leaf ~~is~~ was young or mature, or grew in a sunny or shady environment~~or live in the sunny or shady light regimes~~.

The R^2 value (0.5) of SD-CO_2 based on the adaxial side is higher than from the abaxial side and the combination of both sides, indicating that the correlation of SD-CO_2 is stronger than the others parameters herein. Therefore, the SD on the adaxial side is the best in reconstructing pCO_2. The reconstruction result based on the regression approach is 351.9 ±6.6 ppmv lower than the one based on the stomatal ratio method (Table 6), and it is relatively lower than the results based on the other proxies (Fig. 6; Freeman and Hayes, 1992; Pagani et al., 2005; Maxbauer et al., 2014). However, the result based on stomatal ratio method is 537.5 ±56.5 ppmv, which is fairly close~~closest~~ to GEOCARB III predictions (Fig. 5) and historical reconstruction trends (Fig. 6).

5.2 Paleoclimate reconstructed history

The pCO_2 levels throughout the Cenozoic were ~~relatively~~ generally lower than during much of~~through~~ the Cretaceous, but probably also decreased significantly from the early to late Eocene. However, there is a wide range of estimates for the Eocene (Koch et al., 1992; Sinha and Stott, 1994; Ekart et al., 1999; Greenwood et al., 2003; Royer, 2003; Pagani et al., 2005; Wing et al., 2005; Lowenstein and Demicco, 2006; Fletcher et al., 2008; Zachos et al., 2008; Beerling et al., 2009; Bijl et al., 2010; Smith et al., 2010; Doria et al., 2011; Kato et al., 2011; Maxbauer et al., 2014). ~~(Ekart et al., 1999), but had an overall decreasing trend with some significant increases on short-time scales (e.g. in the earliest Eocene and middle Miocene, Zachos et al., 2001, 2008; Wing et al., 2005; Lowenstein and Demicco, 2006; Fletcher et al., 2008; Bijl et al., 2010; Kato et al., 2011). There is a wide range in pCO_2 estimates for the Paleogene, reflecting problems in the various proxies. Both the fractionation of carbon isotopes by phytoplankton (Freeman and Hayes, 1992) and analysis of~~

paleosol (fossil soil) carbonates (Ekart et al., 1999) demonstrate that carbon dioxide levels were less than 1000 ppmv before the Cretaceous-Tertiary boundary and have been decreasing since the Paleocene.

Based on the measurements of palaeosol carbon isotopes, Cerling (1991) reported that pCO_2 levels for the Eocene and Miocene through to the present was lower than 700 ppmv. Fletcher et al. (2008) also showed that atmospheric CO_2 levels were approximately 680 ppmv by 60 million years ago. However, Stott (1992) reconstructed pCO_2 as 450–550 ppmv for the early Eocene based on phytoplankton. Additionally, reconstructions using the stomatal ratio method based on *Ginkgo*, *Metasequoia*, and Lauraceae leaves also revealed a low pCO_2 level between 300 and 500 ppmv during the early Eocene (Kürschner et al., 2001; Royer et al., 2001; Greenwood et al., 2003; Royer, 2003) except a single high estimate of about 800 ppmv near the Paleocene/Eocene boundary (Royer et al., 2001).

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Subsequently, Smith et al. (2010) reconstructed the value of the early Eocene pCO_2 ranging from 580 ± 40 to 780 ± 50 ppmv using the stomatal ratio method (recent standardization) based on both SI and SD. A climatic optimum occurred in the middle Eocene (MECO): the reconstructed CO_2 concentrations are mainly between 700 and to 1000 ppmv during the late middle Eocene climate transition (42–38 Ma) using stomatal indices of fossil *Metasequoia* needles, but concentrations declined to 450 ppmv toward the top of the investigated section (Doria et al., 2011). Jacques et al. (2014) used CLAMP to calibrate climate change in Antarctica during the early-middle Eocene, suggesting a seasonal alternation of high- and low-pressure systems over

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Antarctica during the early-middle Eocene. Spicer et al. (2014) also reconstructed a relatively lower cool temperature than $\delta^{18}O$ records (Keating-Bitonti et al., 2011) in the middle Eocene of Hainan Island, South China using CLAMP, indicating a not uniformly warm climate in the low latitude during the Eocene. An overall decreasing trend of the pCO_2 level was presented after the middle Eocene (Fig. 6; Retallack, 2009b).

The ice-sheets started to appear in the Antarctic during the Late Eocene (Zachos et al., 2001), then the temperature suffered an apparent further decrease from the late Eocene ~~to the early Oligocene~~onwards (Fig. 6). ~~(Roth-Nebelsick et al., 2004), which resulted in the Antarctic being almost fully covered by ice-sheets. Subsequently, the climate variation was comparatively stable with a little wobbling in temperature during the Oligocene period (Fig. 6), while a small and ephemeral Late Oligocene Warming was present in the latest part of the Oligocene, resulting in reducing the Antarctic ice sheets to a minimum and forming a brief period of glaciation at that time (Zachos et al., 2001). During the Middle Miocene, a quick rise in temperature was shown, which was followed by a small glaciation (Fig. 6; Zachos et al., 2001; Roth-Nebelsick et al., 2004; Beerling and Royer, 2011). Subsequently, the CO_2 concentration decreased gradually and reached 280 ppmv until the period of the industrial revolution (Fig. 6). Since then, however, the CO_2 concentration rebounded to present day level.~~

In conclusion, although various results were made by different pCO_2 reconstruction proxies at the same time, their entire decreasing tendency of pCO_2 level are remarkably consistent with each other since the Eocene (Fig. 6). Fig. 6 shows that during the Eocene the temperature was higher than at present. Comparing to the estimates of late middle Eocene pCO_2 by Doria et al. (2011), the present result~~The reconstructed pCO_2~~ of 351.9 ±6.6 ppmv based on the regression approach ~~is~~ shows a remarkably lower pCO_2 level~~, during the early late Eocene. The result~~ while the one based on the stomatal ratio method of 537.5 ±56.5 ppmv is within the variation range of 500–1000 ppmv, which is closely consistent with the pCO_2 changes over the geological ages (Fig. 6). The world was dynamic in the Paleogene, including in the late middle Eocene, when the MECO occurred. Thus, the exact age matters, and it is possible that the values may differ because of slight offsets in time.

6 Conclusion

In this study, we reconstructed ~~the~~ late middle Eocene pCO_2 based on the fossil leaves of *Nageia maomingensis* Jin et Liu from ~~the~~ late middle Eocene of Maoming Basin, Guangdong Province, China. *Nageia* is a special element in conifers by its broad multi-veined leaf that lacks mid-vein. The stomatal data analysis suggests that only stomatal densities (SD) from both sides of *Nageia motleyi* leaves have significant negative correlations with the atmospheric CO_2 concentration. The SD from the

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adaxial side gives the best correlation to the CO_2. Based on SDs, the pCO_2 concentration is reconstructed using both the regression approach and the stomatal ratio method. The pCO_2 result based on the regression approach is 351.9 ± 6.6 ppmv, showing a relatively lower CO_2 level. The reconstructed result based on the stomatal ratio method is 537.5 ± 56.5 ppmv consistent with the variation trends based on the other proxies. Here, we explored the potential of *N. maomingensis* in pCO_2 reconstruction and obtained different results according to different methods, providing a new insight for the reconstruction of paleoclimate and paleoenvironment in conifers.

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References

Bai, Y. J., Chen, L. Q., Ranhotra, S. P., Wang, Q., Wang, Y. F., Li, C. S.: Reconstructing atmospheric CO_2 during the Plio–Pleistocene transition by fossil *Typha*. Global Change Biology, 21, 874–881, doi:10.1111/gcb.12670, 2015.

Beerling, D. J.: Stomatal density and index: theory and application, in: Jones, T. P., and Rowe, N. P., (Eds.), Fossil Plants and Spores: Modern Techniques, Geological Society, London, 251–256, 1999.

Beerling, D. J., and Kelly, C. K.: Stomatal density responses of temperate woodland plants over the past seven decades of CO_2 increase: A comparison of salisbury (1927) with contemporary data, American Journal of Botany, 84, 1572–1583, 1997.

Beerling, D. J., and Royer, D. L.: Reading a CO_2 signal from fossil stomata, New Phytologist, 153, 387–397, doi:10.1046/j.0028-646X.2001.00335.x, 2002.

Beerling, D. J., and Royer, D. L.: Convergent Cenozoic CO_2 history, Natural Geoscience, 4, 418–420, doi:10.1038/ngeo1186, 2011.

Beerling, D. J., Chaloner, W. G., Huntley, B., Pearson, J. A., Tooley, M. J., and Woodward, F. I.: Variations in the stomatal density of *Salix herbacea* L. under the changing atmospheric CO_2 concentrations of late- and post-glacial time, Philosophical Transactions of the Royal Society of London, ser. B. 336, 215–224, doi:10.1098/rstb.1992.0057, 1992.

Beerling, D. J., Fox, A., and Anderson, C. W.: Quantitative uncertainty analyses of ancient atmospheric CO_2 estimates from fossil leaves, American Journal of

Science, 309, 775–787, doi:10.2475/09.2009.01, 2009.

Beerling, D. J., Lomax, B. H., Royer, D. L., Upchurch Jr., G. R., and Kump, L. R.: An atmospheric pCO_2 reconstruction across the Cretaceous-Tertiary boundary from leaf megafossils, Proceedings of the National Academy of Sciences of the United States of America, 99, 7836–7840, doi:10.1073/pnas.122573099, 2002.

Berner, R. A.: GEOCARB II: A revised model of atmospheric CO_2 over Phanerozoic time, American Journal of Science, 294, 56–91, doi:10.2475/ajs.294.1.56, 1994.

Berner, R. A., and Kothavalá, Z.: GEOCARB III: A revised model of Atmospheric CO_2 over Phanerozoic time, American Journal of Science, 301, 182–204, doi:10.2475/ajs.301.2.182, 2001.

Bijl, P. K., Houben, A. J. P., Schouten, S., Bohaty, S. M., Sluijs, A., Reichart, G., Sinninghe Damsté, J. S., and Brinkhuis, H.: Transient Middle Eocene atmospheric CO_2 and temperature variations, Science, 330, 819–821, doi:10.1126/science.1193654, 2010.

Brown, L. R.: Atmospheric carbon dioxide concentration, 1000-2009 (Supporting data), in: Brown, L. R., (Ed.), World on the Edge: How to Prevent Environmental and Economic Collapse. Chapter 4 Data: Rising Temperatures, Melting Ice, and Food Security, Earth policy institute, Norton, W.W. & Company, New York, London (http://www.earth-policy.org/books/wote/wote_data), 2010.

~~Cerling, T. E.: Carbon dioxide in the atmosphere: evidence from Cenozoic and Mesozoic palaeosols, American Journal of Science, 291, 377–400, doi:10.2475/ajs.291.4.377, 1991.~~

Cerling, T. E.: Use of carbon isotopes in paleosols as an indicator of the P(CO_2) of the paleoatmosphere, Global Biogeochemical Cycles, 6, 307–314, doi:10.1029/92GB01102, 1992.

Chaloner, W. G., and McElwain, J. C.: The fossil plant record and global climate change, Review of Palaeobotany and Palynology, 95, 73–82, doi:10.1016/S0034-6667(96)00028-0, 1997.

Cheng, W. C., Fu, L. K., and Chao, C. S.: *Podocarpus* (Podocarpaceae), in: Cheng, Wanchün, and Fu, Likuo, (Eds.), Flora of China, Science Press, Beijing, 7, 398–422, 1978 (in Chinese).

Doria, G., Royer, D. L., Wolfe, A. P., Fox, A., Westgate, J. A., and Beerling, D. J.: Declining atmospheric CO_2 during the Late Middle Eocene climate transition, American Journal of Science, 311, 63–75, doi:10.2475/01.2011.03, 2011.

Ekart, D. D., Cerling, T. E., Montanez, I. P., and Tabor, N. J.: A 400 million year carbon isotope record of pedogenic carbonate: implications for paleoatmospheric carbon dioxide, American Journal of Science, 299, 805–827, doi:10.2475/ajs.299.10.805, 1999.

Fletcher, B. J., Brentnall, S. J., Anderson, C. W., Berner, R. A., and Beerling, D. J.: Atmospheric carbon dioxide linked with Mesozoic and Early Cenozoic climate change, Nature Geoscience, 1, 43–48, doi:10.1038/ngeo.2007.29, 2008.

Freeman, K. H., and Hayes, J. M.: Fractionation of carbon isotopes by phytoplankton and estimates of ancient CO_2 levels, Global Biogeochemical Cycles, 6, 185–198, doi:10.1029/92GB00190, 1992.

Fu, D. Z.: Nageiaceae – a new gymnosperm family, Acta Phytotaxonomica Sinica, 30, 515–528, 1992 (in Chinese with English summary).

Fu L. K., Li Y., and Mill, R. R.: Podocarpaceae, in: Wu Z. Y., and Raven, P. H., (Eds.), Flora of China, Science Press, Beijing, 4, 78–84, 1999.

Gale, J.: Availability of carbon dioxide for photosynthesis at high altitudes: theoretical considerations, Ecology, 53, 494–497, doi:10.2307/1934239, 1972.

Greenwood, D. G., Scarr, M. J., and Christophel, D. C.: Leaf stomatal frequency in the Australian tropical rain forest tree *Neolitseadealbata* (Lauraceae) as a proxy measure of atmospheric pCO_2, Palaeogeography, Palaeoclimatology, Palaeoecology, 196, 375–393, doi:10.1016/S0031-0182(03)00465-6, 2003.

Grein, M., Oehm, C., Konrad, W., Utescher, T., Kunzmann, L., and Roth-Nebelsick, A.: Atmospheric CO_2 from the late Oligocene to early Miocene based on photosynthesis data and fossil leaf characteristics, Palaeogeography, Palaeoclimatology, Palaeoecology, 374, 41–51, doi:10.1016/j.palaeo.2012.12.025, 2013.

Henderiks, J., and Pagani, M.: Coccolithophore cell size and the Paleogene decline in atmospheric CO_2, Earth and Planetary Science Letters 269, 575–583, doi:10.1016/j.epsl.2008.03.016, 2008.

Hill, R. S., and Pole, M. S.: Leaf and shoot morphology of extant *Afrocarpus*, *Nageia* and *Retrophyllum* (Podocarpaceae) species, and species with similar leaf arrangement, from Tertiary sediments in Australasia, Australian Systematic Botany, 5, 337–358, doi:10.1071/SB9920337, 1992.

Hu, J. J., Xing, Y. W., Turkington, R., Jacques, F. M. B. Su, T., Huang, Y. J. and Zhou Z. K.: A new positive relationship between pCO_2 and stomatal frequency in *Quercus guyavifolia* (Fagaceae): a potential proxy for palaeo-CO_2 levels, Annals of Botany, 1–12, doi:10.1093/aob/mcv007, 2015.

Huang, C., Retallack, G. J., Wang, C., and Huang, Q.: Paleoatmospheric pCO_2 fluctuations across the Cretaceous-Tertiary boundary recorded from paleosol carbonates in NE China, Palaeogeography, Palaeoclimatology, Palaeoecology, 385, 95–105, doi.org/10.1016/j.palaeo.2013.01.005, 2013.

Jacques, F. M. B., Shi, G. L., Li, H. M., and Wang, W. M.: An Early-Middle Eocene Antarctic summer monsoon: Evidence of ‘fossil climates’, Gondwana Research, 25, 1422–1428, doi:10.1016/j.gr.2012.08.007, 2014.

Jin, J. H., Qiu, J., Zhu, Y. A., and Kodrul, T. M.: First fossil record of the genus *Nageia* (Podocarpaceae) in South China and its phytogeographic implications, Plant Systematics and Evolution, 285, 159–163, doi:10.1007/s00606-010-0267-4, 2010.

Jones, H. G.: Plants and microclimate, Cambridge UK Cambridge University Press, 1–428, 1992.

Kato, Y., Fujinaga, K., and Suzuki, K.: Marine Os isotopic fluctuations in the Early Eocene greenhouse interval as recorded by metalliferous umbers from a Tertiary ophiolite in Japan, Gondwana Research, 20, 594–607, doi:10.1016/j.gr.2010.12.007, 2011.

Keating-Bitonti, C. R., Ivany, L. C., Affek, H. P., Douglas, P., and Samson, S. D.:

Warm, not super-hot, temperatures in the early Eocene subtropics, Geology 39, 771–774, doi: 10.1130/G32054.1, 2011.

Kimura, T., Ohana, T., and Mimoto, K.: Discovery of a podocarpaceous plant from the Lower Cretaceous of Kochi Prefecture, in the outer zone of southwest Japan, Proceedings of the Japan Academy, ser. B, 64, 213–216, doi:10.2183/pjab.64.213, 1988.

Koch, P. L., Zachos, J. C., and Gingerich, P. D.: Correlation between isotope records in marine and continental carbon reservoirs near the Palaeocene/Eocene boundary, Nature, 358, 319–322, doi:10.1038/358319a0, 1992.

Kouwenberg, L. L. R., McElwain J. C., Kürschner, W. M., Wagner, F., Beerling, S. J., Mayle, F. E., and Visscher, H.: stomatal frequency adjustment of four conifer species to historical changes in atmospheric CO_2, American Journal of Botany, 90, 610–619, 2003.

Körner, Ch., and Cochrane, P. M.: Stomatal responses and water relations of *Eucalyptus pauciflora* in summer along an elevational gradient, Oecologia, 66, 443–455, doi:10.1007/BF00378313, 1985.

Krassilov, V. A.: New coniferales from Lower Cretaceous of Primorye, Botanical Journal, 50, 1450–1455 (in Russia), 1965.

Kürschner, W. M., van der Burgh, J., Visscher, H., and Dilcher, D. L.: Oak leaves as biosensors of Late Neogene and Early Pleistocene paleoatmospheric CO_2 concentrations, Marine Micropaleontology, 27, 299–312, doi:10.1016/0377-8398(95)00067-4, 1996.

Kürschner, W. M., Wagner, F., Dilcher, D. L., and Visscher, H.: Using fossil leaves for the reconstruction of Cenozoic paleoatmospheric CO_2 concentrations, in: Gerhard, L. C., Harrison, W. E., Hanson, B. M., (Eds.), Geological Perspectives of Global Climate Change, APPG Studies in Geology, 47, Tulsa, 169–189, 2001.

Kürschner, W. M., Kvaček, Z., and Dilcher, D. L.: The impact of Miocene atmospheric carbon dioxide fluctuations on climate and the evolution of terrestrial ecosystems, Proceedings of the National Academy of Sciences of the United States of America, 105, 449–453, doi:10.1073/pnas.0708588105, 2008.

Liu, X. Y., Gao, Q., and Jin, J. H.: Late Eocene leaves of *Nageia* Gaertner (section *Dammaroideae* Mill) from Maoming Basin, South China and their implications on phytogeography, Journal of Systematics and Evolution, 53, 297–307, doi:10.1111/jse.12133, 2015.

Lowenstein, T. K., and Demicco, R. V.: Elevated Eocene atmospheric CO_2 and its subsequent decline, Science, 313, 1928, doi:10.1126/science.1129555, 2006.

McElwain, J. C.: Do fossil plants signal palaeoatmospheric carbon dioxide concentration in the geological past, Philosophical Transactions of the Royal Society, Lond B, 353, 83–96, doi:10.1098/rstb.1998.0193, 1998.

McElwain, J. C., and Chaloner, W. G.: Stomatal density and index of fossil plants track atmospheric carbon dioxide in the Palaeozoic, Annals of Botany, 76, 389–395, doi:10.1006/anbo.1995.1112, 1995.

McElwain, J. C., and Chaloner, W. G.: The fossil cuticle as a skeletal record of environmental changes, Palaios, 11, 376–388, doi: 10.2307/3515247, 1996.

Mill, R. R.: A new combination in *Nageia* (Podocarpaceae): Novon, 9, 77–78, 1999.

Mill, R. R.: A new sectional combination in *Nageia* Gaertn (Podocarpaceae), Edinburgh Journal of Botany, 58, 499–501, doi:10.1017/S0960428601000804, 2001.

Maxbauer, D. P., Royer, D. L., and LePage, B. A.: High Artic forests during the middle Eocene supported by moderate levels of atmospheric CO_2, Geology, 42, 1027–1030, doi:10.1130/G36014.1, 2014.

Moran, K., Backman, J., Brinkhuis, H., Clemens, S. C., Cronin, T., Dickens, G. R., Eynaud, F., Gattacceca, J., Jakobsson, M., Jordan, R. W., Kaminski, M., King, J., Koc, N., Krylov, A., Martinez, N., Matthiessen, J., McInroy, D., Moore, T.C., Onodera, J., O'Regan, M., Pälike, H., Rea, B., Rio, D., Sakamoto, T., Smith, D. C., Stein, R., St John, K., Suto, I., Suzuki, N., Takahashi, K., Watanabe, M., Yamamoto, M., Farrell, J., Frank, M., Kubik, P., Jokat, W., and Kristoffersen, Y.: The Cenozoic palaeoenvironment of the Arctic Ocean, Nature, 441, 601–605, doi:10.1038/nature04800, 2006.

Nordt, L., Atchley, S., and Dworkin, S. I.: Paleosol barometer indicates extreme fluctuations in atmospheric CO_2 across the Cretaceous-Tertiary boundary, Geology, 30, 703–706, doi:10.1130/0091-7613(2002)030<0703:PBIEFI>2.0.CO;2, 2002.

Pagani, M., Arthur, M. A., and Freeman, K. H.: Miocene evolution of atmospheric carbon dioxide, Paleoceanography, 14, doi:10.1029/1999PA900006, 273–292, 1999.

Pagani, M., Zachos, J. C., Freeman, K. H., Tipple, B., and Bohaty, S.: Marked decline in atmospheric carbon dioxide concentrations during the Paleocene, Science, 309, 600–603, doi:10.1126/science.1110063, 2005.

Pearson, P. N., Foster, G. L., and Wade, B. S.: Atmospheric carbon dioxide through the Eocene-Oligocene climate transition, Nature, 461, 1110–1113, doi:10.1038/nature08447, 2009.

Petit, J. R., Jouzel, J., Raynaud, D., Barkov, N. I., Barnola, J.-M., Basile, I., Bender, M., Chappellaz, J., Davis, M., Delaygue, G., Delmotte, M., Kotlyakov, V. M., Legrand, M., Lipenkov, V. Y., Lorius, C., PÉpin, L., Ritz, C., Saltzman, E., and Stievenard, M.: Climate and atmospheric history of the past 420,000 years from the Vostokicecore, Antarctica, Nature, 399, 429–436, doi:10.1038/20859, 1999.

Pieter, T., and Keeling, R.: Recent monthly average Mauna Loa CO_2, NOAA/ESRL, www.esrl.noaa.gov/gmd/ccgg/trends/ (accessed March 2015), 2015.

Poole, I., and Kürschner, W. M.: Stomatal density and index: the practice, in: Jones, T.P., and Rowe, N.P., (Eds.), Fossil Plants and Spores: Modern Techniques, Geological Society, London, 257–260, 1999.

Retallack, G. J.: A 300-million-year record of atmospheric carbon dioxide from fossil plant cuticles, Nature, 411, 287–290, doi:10.1038/35077041, 2001.

Retallack, G. J.: Greenhouse crises of the past 300 million years, Geological Society of America Bulletin, 121, 1441–1455, doi:10.1130/B26341.1, 2009a.

Retallack, G. J.: Refining a pedogenic-carbonate CO_2 paleobarometer to quantify a Middle Miocene greenhouse spike, Palaeogeography, Palaeoclimatology, Palaeoecology, 281, 57–65, doi:10.1016/j.palaeo.2009.07.011, 2009b.

~~Roth-Nebelsick, A., Utescher, T., Mosbrugger, V., Diester-Haass, L., and Walther, H.: Changes in atmospheric CO_2 concentrations and climate from the Late Eocene to Early Miocene: palaeobotanical reconstruction based on fossil floras from Saxony, Germany, Palaeogeography, Palaeoclimatology, Palaeoecology, 205, 43–67, doi:10.1016/j.palaeo.2003.11.014, 2004.~~

Roth-Nebelsick, A., Grein, M., Utescher, T., and Konrad, W.: Stomatal pore length change in leaves of *Eotrigonobalanus furcinervis* (Fagaceae) from the Late Eocene to the Latest Oligocene and its impact on gas exchange and CO_2 reconstruction, Review of Palaeobotany and Palynology, 174, 106–112, doi:10.1016/j.revpalbo.2012.01.001, 2012.

Roth-Nebelsick, A., Oehm, C., Grein, M., Utescher, T., Kunzmann, L., Friedrich, J.-P., and Konrad, W.: Stomatal density and index data of *Platanus neptuni* leaf fossils and their evaluation as a CO_2 proxy for the Oligocene, Review of Palaeobotany and Palynology, 206, 1–9, doi:10.1016/j.revpalbo.2014.03.001, 2014.

Royer, D. L.: Stomatal density and stomatal index as indicators of paleoatmospheric CO_2 concentration, Review of Palaeobotony and Palynology, 114, 1–28, doi:10.1016/S0034-6667(00)00074-9, 2001.

Royer, D. L.: Estimating Latest Cretaceous and Tertiary atmospheric CO_2 from stomatal indices, in: Wing, S. L., Gingerich, P. D., Schmitz, B., and Thomas, E., (Eds.), Causes and Consequences of Globally Warm Climates in the Early Paleocene, Geological Society of America Special Paper, 79–93, 2003.

Royer, D. L.: CO_2-forced climate thresholds during the Phanerozoic, Geochimica et Cosmochimica Acta, 70, 5665–5675, doi:10.1016/j.gca.2005.11.031, 2006.

Royer, D. L., Wing, S. L., Beerling, D. J., Jolley, D. W., Koch, P. L., Hickey, L. J., and Berner, R. A.: Paleobotanical evidence for near present-day levels of atmospheric CO_2 during part of the Tertiary, Science, 292, 2310–2313, doi:10.1126/science.292.5525.2310, 2001.

Rundgren, M., and Beerling, D. J.: A Holocene CO_2 record from the stomatal index of subfossil *Salix herbacea* L. leaves from northern Sweden, The Holocene, 9, 509–513, doi:10.1191/095968399677717287, 1999.

Seki, O, Foster, G. L., Schmidt, D. N., Mackensen, A., Kawamura, K., and Pancost, R. D.: Alkenone and boron-based Pliocene pCO_2 records, Earth and Planetary Science Letters, 292, 201–211, doi:10.1016/j.epsl.2010.01.037, 2010.

Sinha, A., and Stott, L. D.: New atmospheric pCO_2 estimates from paleosols during the late Paleocene/early Eocene global warming interval, Global and Planetary Change, 9, 297–307, doi:10.1016/0921-8181(94)00010-7, 1994,

Smith, R. Y., Greenwood, D. R., and Basinger, J.F.: Estimating paleoatmospheric pCO_2 during the Early Eocene Climatic Optimum from stomatal frequency of *Ginkgo*, Okanagan Highlands, British Columbia, Canada, Palaeogeography, Palaeoclimatology, Palaeoecology, 293, 120–131, doi:10.1016/j.palaeo.2010.05.006, 2010.

Spicer, A. R., Herman, A. B., Liao, W. B., Spicer, T. E. V., Kodrul, T. M., Yang, J., and Jin, J. H.: Cool tropics in the middle Eocene: Evidence from the Changchang

Flora, Hainan Island, China, Palaeogeography, Palaeoclimatology, Palaeoecology, 412, 1–16, doi:10.1016/j.palaeo.2014.07.011, 2014.

Stott, L. D.: Higher temperatures and lower oceanic pCO_2: A climate enigma at the end of the Paleocene Epoch, Paleoceanography, 7, 395–404, doi:10.1029/92PA01183, 1992.

Stults, D. Z., Wagner-Cremer, F., and Axsmith, B. J.: Atmospheric paleo-CO_2 estimates based on *Taxodium distichum* (Cupressaceae) fossils from the Miocene and Pliocene of Eastern North America, Palaeogeography Palaeoclimatology Palaeoecology, 309, 327–332, doi:10.1016/j.palaeo.2011.06.017, 2011.

Sun, B. N., Ding, S. T., Wu, J. Y., Dong, C., Xie, S. P., and Lin, Z. C.: Carbon isotope and stomatal data of Late Pliocene Betulaceae leaves from SW China: Implications for palaeoatmospheric CO_2-levels, Turkish Journal of Earth Sciences, 21, 237–250, doi:10.3906/yer-1003-42, 2012.

Sun, T. X.: Cuticle micromorphology of *Nageia*, Journal of Wuhan Botanical Research, 26, 554–560, doi:10.3969/j.issn.2095-0837.2008.06.002, 2008 (in Chinese with English abstract).

Tripati, A. K., Roberts, C. D., and Eagle, R. A.: Coupling of CO_2 and ice sheet stability over major climate transitions of the last 20 million years, Science, 326, 1394–1397, doi:10.1126/science.1178296, 2009.

Van der Burgh, J., Visscher, H., Dilcher, D. L., and Kürschner, W. M.: Paleoatmospheric signatures in Neogene fossil leaves, Science, 260, 1788–1790, doi:10.1126/science.260.5115.1788, 1993.

Wang, J. D., Li, H. M., and Zhu Z. Y.: Magnetostratigraphy of Tertiary rocks from Maoming Basin, Guangdong province, China, Chinese Journal of Geochemistry, 13, 165–175, doi:10.1007/BF02838516, 1994.

Walker, J. D., and Geissman, J. W.: Geologic Time Scale, Geological Society of America, doi:10.1130/2009.CTS004R2C, 2009.

Wing, S. L., Harrington, G. J., Smith, F. A., Bloch, J. I., Boyer, D. M., and Freeman, K. H.: Transient floral change and rapid global warming at the Paleocene-Eocene boundary, Science, 310, 993–996, doi:10.1126/science.1116913, 2005.

Woodward, F. I.: Ecophysiological studies on the shrub *Vaccinium myrtillus* L. taken from a wide altitudinal range, Oecologia, 70, 580–586, doi:10.1007/BF00379908, 1986.

Woodward, F. I.: Stomatal numbers are sensitive to increases in CO_2 concentration from pre-industrial levels, Nature, 327, 617–618, doi:10.1038/327617a0, 1987.

Woodward, F. I., and Bazzaz, F. A.: The responses of stomatal density to CO_2 partial pressure, Journal of Experimental Botany, 39, 1771–1781, doi:10.1093/jxb/39.12.1771, 1988.

Yang, J. J., Qi, G. F., and Xu, R. H.: Studies on fossil woods excavated from the Dabie mountains, Scientia Silvae Sinicae, 26, 379–386, 1990 (in Chinese with English abstract).

Ye, M. N.: On the preparation methods of fossil cuticle. Palaeontological Society of China (Ed.), Selected papers of the 12th Annual conference of the Palaeontological Society of China, Science Press, Beijing, 170–179, 1981 (in

Chinese).

Zachos, J., Pagani, M., Sloan, L., Thomas, E., and Billups, K.: Trends, rhythms, aberrations in global climate 65 Ma to present, Science, 292, 686–693, doi:10.1126/science.1059412, 2001.

Zachos, J., Dickens, G. R. and Zeebe, R. E.: An early Cenozoic perspective on greenhouse warming and carbon-cycle dynamics, Nature, 451, 279–283, doi:10.1038/nature06588, 2008.

Figure 1. Map showing the distribution of extant and fossil *Nageia* and their mean annual temperature (Modified after the map from http://www.sage.wisc.edu/atlas/maps.php?datasetid=35&includerelatedlinks=1&dataset=35).

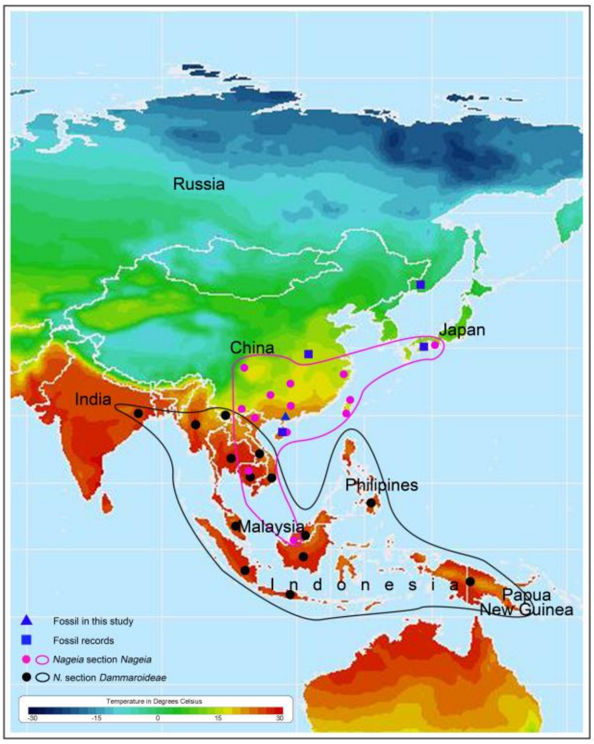

Figure 2. Sampling areas and counting rules are shown. (a) *Nageia motleyi* (Parl.) De Laub.leaf. Black squares in the middle of the leaf show the sampling areas for preparing the cuticles. (b) The abaxial side of the cuticle from *N. motleyi* leaf. Black circles show the counted stomatal complexes. (c) *N. maomingensis* Jin et Liu. Red squares in the middle of the leaf indicate the sampling areas. (d) The abaxial side of the fossil cuticle. Red circles show the counted stomatal complexes. Scale bars: (a) and (c) = 1 cm; (b) and (d) = 50 μm.

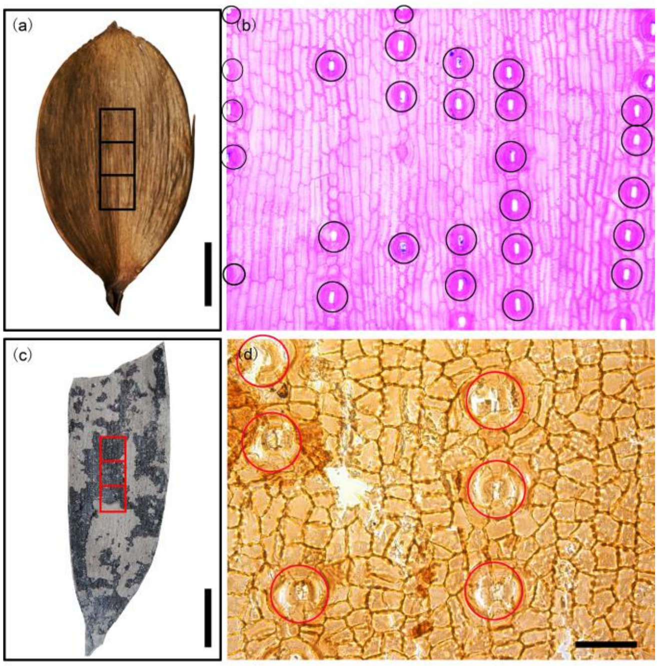

Figure 3. Correlation between SD and SI versus CO_2 concentration for modern *Nageia motleyi*. (a) Trends of SD with CO_2 concentration for the adaxial surface. (b) Trends of SD with CO_2 concentration for the abaxial surface. (c) Trends of SI with CO_2 concentration for the adaxial surface. (d) Trends of SI with CO_2 concentration for the abaxial surface. (e) Trends of SD with CO_2 concentration for the combined data of both leaf surfaces. (f) Trends of SI with CO_2 concentration for the combined data of both leaf surfaces.

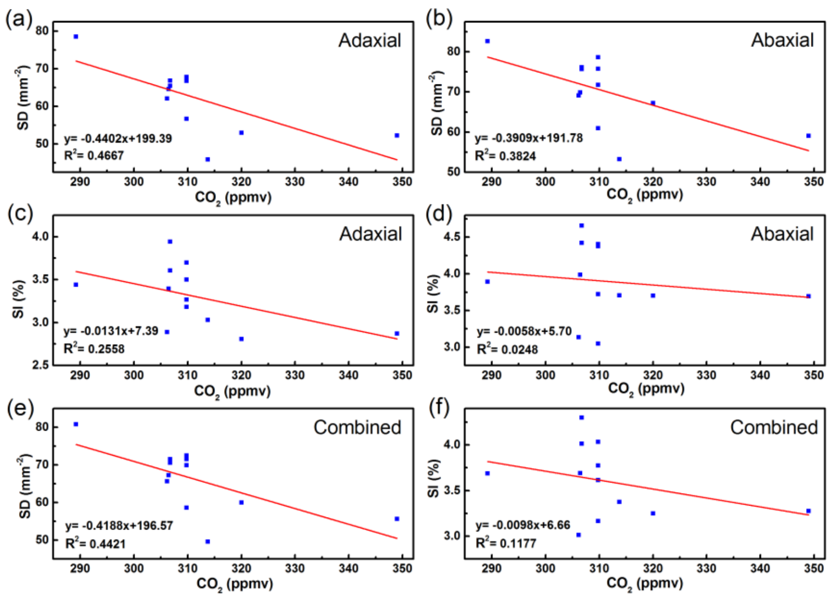

Figure 4. Correlation between SNL, SDL and TSDL versus CO_2 concentration for modern *Nageia motleyi*. (a) Trends of SNL with CO_2 concentration for the adaxial surface. (b) Trends of SDL with CO_2 concentration for the adaxial surface. (c) Trends of TSDL with CO_2 concentration for the adaxial surface.

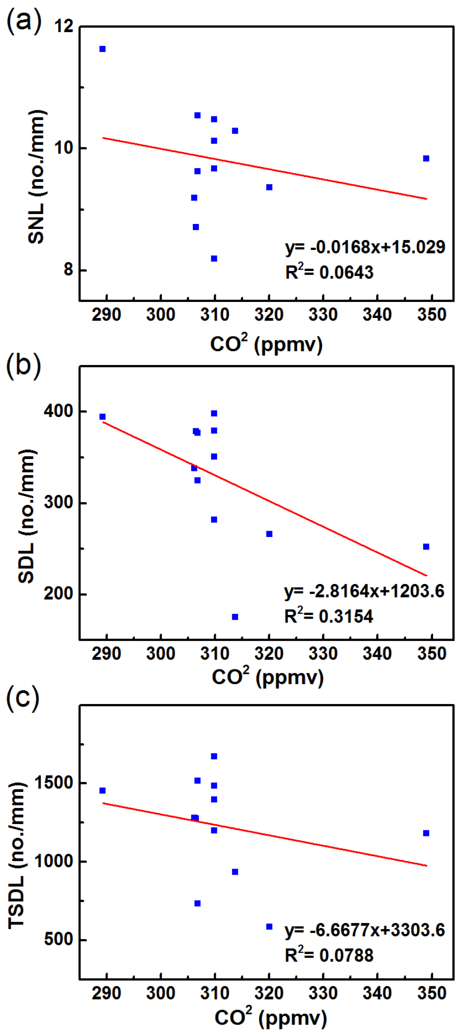

Figure 5. The pCO_2 reconstruction results and extant CO_2 concentrations are projected onto the long-term carbon cycle model (GEOCARB III; Berner and Kothavalá, 2001). The pCO_2 results based on the regression approach and stomatal ratio method are represented by red and blue squares, respectively.

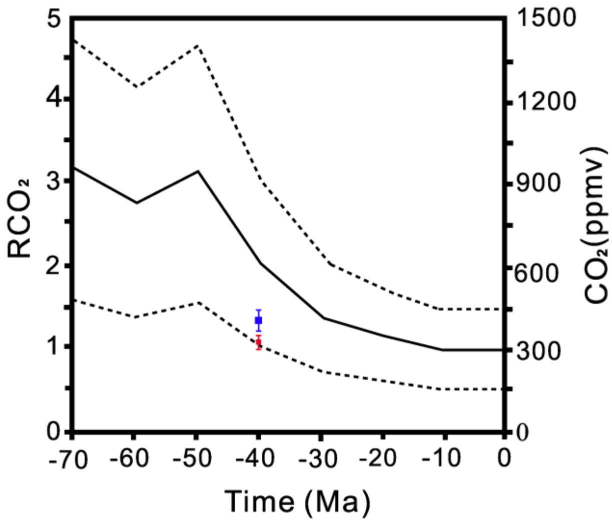

Figure 6. Atmospheric CO_2 estimates from proxies over the past 60 million years. The horizontal dashed line indicates monthly atmospheric CO_2 concentration for March 2015 at Mauna Loa, Hawaii (401.5 ppmv) (Pieter and Keeling, 2015). The vertical lines show the error bars. The data are from the supporting data of Beerling and Royer (2011) and references in Table 9. The lower blue star shows the reconstructed result based on the regression approach. The higher one presents the result of stomatal ratio method.

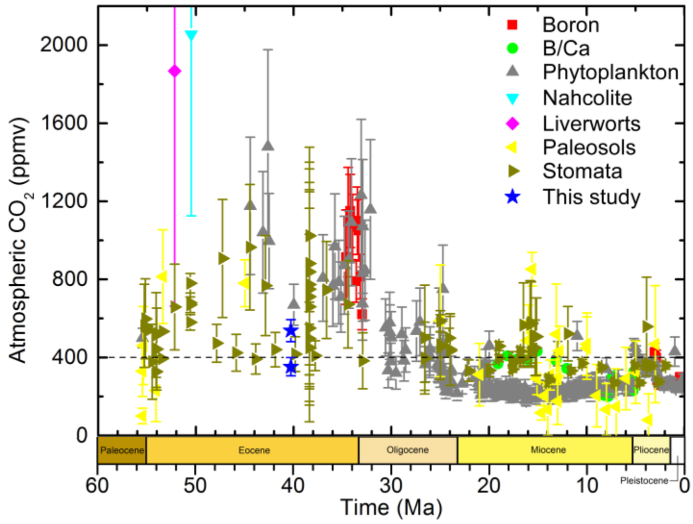

Table 1. Modern *Nageia motleyi* (Parl.) De Laub samples and atmospheric CO_2 values of their collection dates from ice core data (Brown, 2010).

Herbarium	Collection number	Collecting locality	Collectors	Number of leaf samples	Collection date	CO_2 (ppmv)
LE	No. 2649	Malaysia	Beccari, O.	1	1868	289.23
A/GH	No. bb. 17229	150 m, Riau on Ond. Karimon, Archipel. Ind.	Neth. Ind. For. Service	2	1932	306.19
A/GH	No. bb. 18328	5 m, Z. O. afd. v. Borneo Tidoengsche Landen, Archipel. Ind.	Neth. Ind. For. Service	2	1934	306.46
A/GH	No. bb. 21151	500 m, Z. O. afd. Borneo, Poeroek Tjahoe Tahoedjan, Archipel. Ind.	Neth. Ind. For. Service	2	1936	306.76
KEP	No. 30887	Kata Tinggi, Johor, Malaysia	Corner, E.J.H.	1	1936	306.76
KEP	No. 57329	Batang Padang, Perak, Malaysia	Unkonwn	2	1947	309.82
KEP	No. 57330	Batang Padang, Perak, Malaysia	Unkonwn	2	1947	309.82
KEP	No. 55897	Batang Padang, Perak, Malaysia	Unkonwn	2	1947	309.82
KEP	No. 61064	Batang Padang, Perak, Malaysia	Syed Woh	2	1947	309.82
E	No. bb. 40798	51 m, Kuala Trengganu-Besut Road, Bukit Bintang Block, Gunong Tebu Forest reserve, Malaysia	Sinclair, J. and Kiah bin, Salleh	2	1955	313.73
KEP	No. 80548	Gombak, Selangor, Malaysia	Rahim	1	1965	320.04
KEP	No. 33343	Jelebu, Negeri Sembilan, Malaysia	Yap, S.K.	2	1987	348.98

Note: A/GH—Harvard University Herbarium, Harvard University, 22 Divinity Avenue, Cambridge, Massachusetts 02138, USA (www.huh.harvard.edu).

E—The Herbarium of Royal Botanic Garden, Edinburgh EH3 5LR, Scotland, UK (www.rbge.org.uk).

LE—The Herbarium of the V.L. Komarov Botanical Institute of the Russian Academy of Sciences, Prof. Popov Street 2, Saint Petersburg 197376, Russia (www.binran.ru).

KEP—Kepong Herbarium, Forest Research Institute Malaysia, 52109 Kepong, Selangor, Malaysia (http://www.frim.gov.my/).

Table 2. Summary of stomatal parameters of the adaxial surface from modern *Nageia motleyi* (Parl.) De Laub.

Collection number	Collection date	CO_2 (ppmv)	SD (mm^{-2})					SI (%)				
			x	σ	s.e.	t*s.e.	n	*x*	σ	s.e.	t*s.e.	n
No.2649	1868	289.23	78.60	15.44	1.41	2.76	120	3.44	0.66	0.06	0.12	120
No.bb.17229	1932	306.19	62.14	17.20	1.78	3.50	93	2.89	0.68	0.07	0.14	93
No.bb.18328	1934	306.46	64.57	15.05	1.58	3.11	90	3.39	1.01	0.11	0.21	90
No.bb.21151	1936	306.76	65.45	11.14	1.17	2.30	90	3.94	0.74	0.08	0.15	90
No.SFN30887	1936	306.76	66.90	16.10	1.27	2.49	161	3.61	0.92	0.07	0.14	161
No.61064	1947	309.82	56.71	16.81	1.95	3.83	74	3.27	1.26	0.15	0.29	74
No.57330	1947	309.82	67.37	15.97	2.04	4.01	61	3.70	0.82	0.10	0.20	61
No.57329	1947	309.82	67.85	15.61	1.70	3.34	84	3.50	0.90	0.10	0.20	84
No.55897	1947	309.82	66.74	14.10	1.78	3.48	63	3.18	0.66	0.08	0.16	63
No.40798	1955	313.73	45.89	13.81	1.12	2.20	151	3.03	0.87	0.07	0.14	151
No.KEP80548	1965	320.04	52.94	11.25	0.85	1.67	175	2.81	0.61	0.05	0.09	175
No.FRI33343	1987	348.98	52.25	12.05	0.77	1.51	242	2.87	0.69	0.04	0.09	242
Mean	–	–	62.28	14.54	1.45	2.85	117	3.30	0.52	0.08	0.16	117

Note: *x*—mean; σ—standard deviation; s.e. —standard error of mean; n— numbers of photos counts (40×); t·s.e.— 95% confidence interval.

Table 3. Summary of stomatal parameters of the abaxial surface from modern *Nageia motleyi* (Parl.) De Laub.

Collection number	Collection date	CO_2 (ppmv)	SD (mm^{-2})					SI (%)				
			x	σ	s.e.	t*s.e.	n	*x*	σ	s.e.	t*s.e.	n
No.2649	1868	289.23	82.71	12.23	1.02	2.00	144	3.89	0.58	0.05	0.09	144
No.bb.17229	1932	306.19	69.16	14.23	1.48	2.90	93	3.13	0.58	0.06	0.12	93
No.bb.18328	1934	306.46	69.92	14.38	1.52	2.97	90	3.99	1.08	0.11	0.22	90
No.bb.21151	1936	306.76	75.68	15.74	1.66	3.25	90	4.66	0.88	0.09	0.18	90
No.SFN30887	1936	306.76	76.18	12.51	0.99	1.93	161	4.42	0.89	0.07	0.14	161
No.61064	1947	309.82	60.93	11.02	1.39	2.72	63	3.05	0.62	0.08	0.15	63
No.57330	1947	309.82	75.82	14.14	1.82	3.58	60	4.38	0.84	0.11	0.21	60
No.57329	1947	309.82	71.74	16.84	1.75	3.42	93	3.72	0.62	0.06	0.13	93
No.55897	1947	309.82	78.63	13.41	1.75	3.42	59	4.41	1.00	0.13	0.26	59
No.40798	1955	313.73	53.22	13.88	1.12	2.19	155	3.71	0.93	0.07	0.15	155
No.KEP80548	1965	320.04	67.22	13.97	1.07	2.09	171	3.70	0.80	0.06	0.12	171
No.FRI33343	1987	348.98	59.09	12.10	0.79	1.55	233	3.69	0.86	0.06	0.11	233
Mean	–	–	70.03	13.70	1.36	2.67	118	3.90	0.81	0.08	0.16	118

Note: *x*—mean; σ—standard deviation; s.e. —standard error of mean; n— numbers of photos counts (40×); t･s.e.— 95% confidence interval.

Table 4. Summary of stomatal parameters of the combined data of the adaxial and abaxial surfaces from modern *Nageia motleyi* (Parl.) De Laub.

Collection number	Collection date	CO_2 (ppmv)	SD (mm^{-2})					SI (%)				
			x	σ	s.e.	t*s.e.	n	*x*	σ	s.e.	t*s.e.	n
No.2649	1868	289.23	80.84	13.74	0.85	1.66	264	3.69	0.66	0.04	0.08	264
No.bb.17229	1932	306.19	65.65	16.13	1.18	2.32	186	3.01	0.64	0.05	0.09	186
No.bb.18328	1934	306.46	67.24	14.92	1.11	2.18	180	3.69	1.08	0.08	0.16	180
No.bb.21151	1936	306.76	70.57	14.53	1.08	2.12	180	4.30	0.89	0.07	0.13	180
No.SFN30887	1936	306.76	71.54	15.12	0.84	1.65	322	4.01	0.99	0.05	0.11	322
No.61064	1947	309.82	58.65	14.54	1.24	2.43	137	3.17	1.02	0.09	0.17	137
No.57330	1947	309.82	71.56	15.61	1.42	2.78	121	4.03	0.89	0.08	0.16	121
No.57329	1947	309.82	69.90	16.33	1.23	2.41	177	3.62	0.77	0.06	0.11	177
No.55897	1947	309.82	72.49	14.95	1.35	2.65	122	3.77	1.04	0.09	0.18	122
No.40798	1955	313.73	49.60	14.31	0.82	1.60	306	3.37	0.96	0.05	0.11	306
No.KEP80548	1965	320.04	60.00	14.53	0.78	1.53	346	3.25	0.84	0.05	0.09	346
No.FRI33343	1987	348.98	55.61	12.53	0.58	1.13	475	3.28	0.88	0.04	0.08	475
Mean	–	–	66.14	14.77	1.04	2.08	235	3.60	0.89	0.06	0.12	235

Note: *x*—mean; σ—standard deviation; s.e. —standard error of mean; n— numbers of photos counts (40×); t･s.e.— 95% confidence interval.

783 Table 5. Summary of stomatal parameters from modern *Nageia motleyi* (Parl.) De Laub (Kouwenberg et al., 2003).

Collection number	Collection date	CO_2 (ppmv)	SNL	SDL	TSDL	n
No.2649	1868	289.23	11.64	394.38	1455.10	264
No.bb.17229	1932	306.19	9.19	337.98	1280.12	186
No.bb.18328	1934	306.46	8.71	378.92	1277.63	180
No.bb.21151	1936	306.76	9.62	376.93	1517.21	180
No.SFN30887	1936	306.76	10.55	325.08	735.38	240
No.61064	1947	309.82	8.19	282.04	1200.66	133
No.57330	1947	309.82	9.67	397.83	1397.33	119
No.57329	1947	309.82	10.13	350.98	1672.50	176
No.55897	1947	309.82	10.48	379.06	1486.13	122
No.40798	1955	313.73	10.29	175.14	933.85	305
No.KEP80548	1965	320.04	9.36	266.16	585.72	263
No.FRI33343	1987	348.98	9.84	252.20	1181.51	125
Mean	–	–	9.81	326.39	1226.93	191

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Table 6. Summary of stomatal parameters of the adaxial surface of fossil *Nageia* and pCO_2 [$C_{(f)}$] estimates results.

Species	Age	SD (mm^{-2})				SI (%)				SR		pCO_2(ppmv)		$C_{(f)}$ (ppmv)	
		x	σ	s.e.	n	*x*	σ	s.e.	n	*x*	t*s.e	*x*	t*s.e	*x*	t*s.e
MMJ1-001	Late Eocene	52.5	17.1	3.1	30	2.08	0.7	0.1	30	1.35	0.19	333.6	13.9	412.1	62.0
MMJ2-003	Late Eocene	42.3	12.9	2.4	30	1.80	0.6	0.1	30	1.75	0.39	356.8	10.5	536.1	126.2
MMJ2-004	Late Eocene	39.9	13.6	2.5	30	1.66	0.6	0.1	30	1.81	0.32	362.4	11.0	554.3	101.9
MMJ3-003a	Late Eocene	43.2	17.7	3.2	30	1.67	0.7	0.1	30	1.84	0.43	354.8	14.4	564.6	135.7
Mean	Late Eocene	44.5	16.3	1.5	120	1.80	0.7	0.1	120	1.69	0.18	351.9	6.6	516.8	56.5

Note: *x*—mean; σ—standard deviation; s.e. —standard error of mean; n— numbers of photos counts (400×); t･s.e.— 95% confidence interval. pCO_2— the result based the regression approach; $C_{(f)}$— the result based on the stomatal method.

Table 7. Summary of stomatal parameters of the abaxial surface of fossil *Nageia* and pCO_2 [$C_{(f)}$] estimates results.

Species	Age	SD (mm^{-2})				SI (%)				SR		pCO_2(ppmv)		$C_{(f)}$ (ppmv)	
		x	σ	s.e.	n	*x*	σ	s.e.	n	*x*	t*s.e	*x*	t*s.e	*x*	t*s.e
MMJ1-001	Late Eocene	47.7	17.7	3.2	30	2.11	0.8	0.2	30	1.66	0.23	368.6	16.2	515.6	72.3
MMJ2-003	Late Eocene	50.9	18.3	3.3	30	2.12	0.8	0.1	30	1.57	0.23	360.9	16.6	486.0	70.7
MMJ2-004	Late Eocene	48.2	15.8	2.9	30	2.14	0.7	0.1	30	1.63	0.25	367.4	14.5	504.6	77.3
MMJ3-003a	Late Eocene	48.9	12.6	2.7	22	1.85	0.5	0.1	22	1.52	0.19	365.4	13.5	472.3	59.0
Mean	Late Eocene	48.9	16.2	1.5	112	2.07	0.7	0.1	112	1.60	0.11	365.6	7.6	496.1	35.7

Note: *x*—mean; σ—standard deviation; s.e. —standard error of mean; n— numbers of photos counts (400×); t･s.e.— 95% confidence interval. pCO_2— the result based the regression approach; $C_{(f)}$— the result based on the stomatal method.

Table 8. Summary of stomatal parameters of the combined data of the adaxial and abaxial surfaces of fossil *Nageia* and pCO_2 [$C_{(f)}$] estimates results.

Species	Age	SD (mm^{-2})				SI (%)				SR		pCO_2(ppmv)		$C_{(f)}$ (ppmv)	
		x	σ	s.e.	n	*x*	σ	s.e.	n	*x*	t*s.e	*x*	t*s.e	*x*	t*s.e
MMJ1-001	Late Eocene	50.1	17.5	2.3	60	2.09	0.8	0.1	60	1.50	0.15	349.7	10.6	471.2	47.8
MMJ2-003	Late Eocene	46.5	16.3	2.1	60	1.96	0.7	0.1	60	1.67	0.24	358.3	9.8	524.1	75.7
MMJ2-004	Late Eocene	44.0	15.8	2.0	60	1.90	0.7	0.1	60	1.73	0.17	364.3	9.5	542.9	52.6
MMJ3-003a	Late Eocene	45.6	16.1	2.2	52	1.75	0.6	0.1	52	1.73	0.28	360.5	10.4	544.6	88.3
Mean	Late Eocene	46.6	16.4	1.1	232	1.93	0.7	0.1	232	1.66	0.11	358.1	5.0	519.9	35.0

Note: *x*—mean; σ—standard deviation; s.e. —standard error of mean; n— numbers of photos counts (400×); t·s.e.— 95% confidence interval. pCO_2— the result based the regression approach; $C_{(f)}$— the result based on the stomatal method.

Table 9. pCO_2 estimates proxies and corresponding references.

Proxies	References
Boron	Pearson et al., 2009; Seki et al., 2010
B/Ca	Tripati et al., 2009
Phytoplankton	Freeman and Hayes, 1992; Stott, 1992; Pagani et al., 1999, 2005; Henderiks and Pagani, 2008; Seki et al., 2010
Nahcolite	Lowenstein and Demicco, 2006
Liverworts	Fletcher et al., 2008
Paleosols	Cerling, 1992; Koch et al., 1992; Ekart et al., 1999; Royer et al., 2001; Nordt et al., 2002; Retallack, 2009b; Huang et al. 2013
Stomata	Van der Burgh et al., 1993; Kürschner et al., 1996, 2001, 2008; McElwain, 1998; Royer et al., 2001, 2003; Greenwood et al., 2003; Beerling et al., 2009; Retallack, 2009a; Smith et al., 2010; Doria et al., 2011; Roth-Nebelsick et al., 2012; 2014; Grein et al., 2013; Maxbauer et al., 2014