1. Need approximate age

Changed as suggested.

2. Last part of sentence not clear.

Changed as suggested.

3. Last sentence needs rewriting.

Rewritten as suggested.

4. Do not see the point of the original opening sentence and references for paragraph 2.

Changed as suggested.

5. This paragraph needs rewriting to place into proper context. Right now, it does not convey a good understanding of the debate and problems.

We added a context. Hopefully, it is proper.

6. Seems awkward

Changed as suggested.

7. Why not include these in the above list??

Because these two papers reported the positive correlation between stomatal index or stomatal frequency and pCO_2 .

8. Probably two good summary papers would suffice here.

Changed as suggested.

9. Not sure what this last sentence means.

We deleted some words of this sentence.

10. The last few bits can be condensed as sort of a repeat from a above.

The repeated sentence was removed.

11. Correct?

Yes, it is cited from Kouwenberg et al. (2003).

12. Things get problematic here as some of this seems to be methods. The parts on *Nageia* seemingly belong later.

The sentences should belong to the later part were deleted.

13. Split this convoluted sentence and rewrite the second bit as difficult to follow. Does this apply to *Podacarpus* or *Nageia*?

Changed as suggested.

14. Both sections? Not clear.

Yes, we are mean to "both sections".

15. Avoid Latin abbreviations in the middle of sentences.

Changed as suggested.

16. This needs slight expansion and conformity. I would list specific locations, types of material and approximate ages.

Changed as suggested.

17. As above, not quite clear what is being suggested.

We changed the whole sentence making it easily understandable.

18. Between ** This important point – species specific relationships -- needs to come earlier.

Changed as suggested.

19. Between *** I think can removed as long as clearly documneted in Tabel 1

Changed as suggested.

** Needs a brief paragraph here noting the Maoming Basin, rock units, depositional environment, and criticially, the age.**

Changed as suggested.

20. This is not crystal clear to me.

We fixed this sentence and corresponding figure 2b to make it clearly to you are readers.

21. See above. The specifics to *Nageia* should be here, but the more general procedure should be presented above.

We think it's more proper to put the specifics to *Nageia* in front of the sentence of "Both *N. maomingensis* and *N. motleyi* are amphistomatic".

Hello Xiaoyan:

In fixing the "Discussion", I think important and interesting to fully compare various stomatal indices in the modern analog to the known pCO_2 .

For example, is use of one leaf side better than the other, or is the combination of both sides more accurate?

Discussed as suggested.

Basically, which of multiple possible methods gives the best correlation to pCO₂?

SD of the adaxial side gives the best correlation to pCO₂.

You also want to briefly discuss why some of the leaves give values offset from others.

Mentioned in the discussion part.

1	The pCO ₂ estimates of the late Eocene in South China based on	
2	stomatal density of Nageia Gaertner leaves	
3		
4	XIAO-YAN LIU ^{1,2} , QI GAO ¹ , MENG HAN ¹ and JIAN-HUA JIN ^{1,2} *	
5		
6	⁴ State Key Laboratory of Biocontrol and Guangdong Provincial Key Laboratory of Plant	
7	Resources, School of Life Sciences, Sun Yat-sen University, Guangzhou 510275, China	
8	² State Key Laboratory of Palaeobiology and Stratigraphy, Nanjing Institute of Geology and	
9	Palaeontology, Chinese Academy of Sciences, Nanjing 210008, China	
10		
11	Abstract:	
12	Late Eocene <u>Atmospheric</u> pCO ₂ concentrations is haves been estimated for intervals of	
13	the Eocene based onusing various models and proxy informationies. Here we	
14	reconstructed the late Eocene (~ 40.3 XX Ma) pCO ₂ based on the fossil species leaves	批注 [GD1]: Need approximate age
15	of Nageia maomingensis Jin et Liu from collected from the late Eocene of Maoming	
16	Basin, Guangdong Province, China. This is the first paleoatmospheric estimates for	
17	the late Eocene of South China using stomatal data. Studies of We first determine	
18	relationships between atmospheric pCO ₂ concentrations, stomatal density (SD) and	
19	stomatal index (SI) with-using "modern" leaves of N. motleyi (Parl.) De Laub-, the	
20	nearest living equivalent species of to the Eocene fossils, This work indicates that	
21	the SD inversely responds to $\underline{pCO_2}$ atmospheric CO ₂ concentration, while SI has	

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*

22	almost no relationships with pCO2 atmospheric CO2 concentration. Therefore, the	
23	Eocene pCO ₂ concentrations iscan be reconstructed based on both thae regression	
24	approach- and the stomatal ratio method by using the SD is reconstructed based on-	批注 clear
25	the SD of the fossil leaves in comparison with <i>N. motleyi</i> . <u>The pCO₂ result based on</u>	
26	the regression The first approach gives a pCO ₂ of (351.9 \pm 6.6 ppmv, whereas) is more	
27	reasonable and much lower than the one based on stomatal ratio method gives a pCO ₂ .	
28	of 537.5 \pm 56.5 (ppmv). Here, we explored the potential of <i>N</i> . maomingensis in pCO ₂	
29	reconstruction and obtained different results according to different methods, providing	
30	a new insight for the reconstruction of paleoclimate and paleoenvironment in conifers.	批注 rewr
31	Results suggest that the mean CO ₂ concentration was 391.0 \pm 41.1 ppmv or 386.5 \pm	
32	27.8 ppmv during the late Eocene, which is significantly higher than the CO_2 -	
33	concentrations documented from 1968 to 1955 but similar to the values for current-	
34	atmosphere indicating that the Carbon Dioxide levels during that the late Eocene at-	
35	that time may have been similar to today	
36		
37	Keywords: pCO ₂ , late Eocene, <i>Nageia</i> , Maoming Basin, South China.	
38		
39	1 Introduction	
40		
41	The Eocene (55.8-33.9 Ma) generally was characterized by much warmer	
42	temperatures-than present-day, although temperatures also-varied significantly across	
43	this time interval (Zachos et al., 2008). Climate of the early Eocene to middle Eocene-	

批注 [GD2]: Last part of sentence not clear.

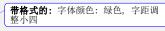
批注 [GD3]: Last sentence needs rewriting.

44	was extremely warm, particularly during the early Eccene with the evidence of the	
45	early Eocene Climatic Optimum (EECO; 5 <u>31</u> 2 to <u>50-53</u> Ma; Zachos et al., 2001), also-	
46	and known as the Paleocene-Eocene Thermal Maximum (PETM-event; ~55.98 Ma).;	
47	Wing et al., 2005; Kato et al., 2011). A series of sudden and extreme global warming-	
48	events (hyperthermals; Deconto et al., 2012) have been described ~ 55.5 to 52 Ma;-	
49	and a mid-Eocene warming has been recognized, the middle Eocene Climate-	
50	Optimum (MECO; ~40 Ma; Bijl et al., 2010). However, the global climatic conditions	
51	cooled significantly bye occurred colder conditions during the early-middle (50 to 48-	
52	Ma) and the late Eocene (40 to 36 Ma; Zachos et al., 2001). The appearance of Indeed.	
53	small, -ephemeral ice-sheets and Arctic sea ice likely existed during the latest Eocene	
54	suggests the coldest climate of the Eocene (Moran et al., 20046; Zachos et al., 20081).	/
55	Atmospheric CO2 concentrations have been well correlated with global surface-	
56	temperature change (Mann et al., 1998; Crowley, 2000; Barnett et al., 2001; Harries et	
57	al., 2001; Levitus et al., 2001; Mitchell et al., 2001). Mostany authors have link-	
58	suggested the that changes in temperature to atmospheric CO2 concentration during	
59	the entire Phanerozoic were linked to atmospheric (pCO ₂) (Petit et al., 1999; Retallack,	
60	2001; Royer, 2006). However, Central to these discussions are records across the	
61	Eocene, as this epoch spans the last major change from a "greenhouse" world to an	
62	<u>"icehouse" world. The Eocene</u> pCO ₂ record remains incomplete and debated	
63	<u>**(Beerling et al., 2002;</u> K ürschner et al., 2001; Royer et al., 2001; Beerling et al.,	
64	2002; Greenwood et al., 2003; Royer, 2003; Wing et al., 2005; Kato et al., 2011).	
65	Most pCO ₂ <u>reconstructions</u> estimates works have focused on the <u>Cretaceous-Tertiary</u>	

批注 [GD4]: Do not see the point of the original opening sentence and references for paragraph 2.

66	and Paleocene-Eocene boundary boundaries (65 to 50 Ma; Koch et al., 1992; Stott,	
67	1992; Sinha and Stott, 1994; Royer et al., 2001; Beerling and Royer, 2002; Nordt et	
68	al., 2002; Royer, 2003; Fletcher et al., 2008; Roth-Nebelsick et al., 2012; 2014; Grein	
69	et al., 2013; Huang et al., 2013; Maxbauer et al., 2014Beerling et al., 2002; Wing et al.,	
70	2005; Kato et al., 2011) and the middle Eocene (MaxbauerDeconto et al., 20122014),	
71	while few reconstructions were conducted at the late Eocene. In addition, the pCO ₂	
72	reconstruction results have varied based on different proxies. Various methods having	
73	been used in pCO ₂ reconstruction mainly include the computer modeling methods:	
74	GEOCARB-I, GEOCARB-II, GEOCARB-III, GEOCARB-SULF and the proxies: ice	
75	cores, paleosol carbonate, phytoplankton, nahcolite, Boron, and stomata	
76	parameters. <u>**</u>	/
77	<u>The abundance of Generally, ss</u> tomatal data (stomatal density and index)cells can	
78	be easily and accurately obtainmeasured from on modern leaves and well-preserved	
79	fossil <u>landeaves-modern leaves</u> . Various plants showing the <u>a</u> negative correlation	
80	between <u>atmospheric CO₂ concentration and stomatal density</u> (SD), or stomatal index	
81	(SI). or both. As such, these parameters and atmospheric CO2 concentration have been	
82	determined in fossil leaves used to reconstruct past pCO2; ,-for-examples include,	
83	including Ginkgo (Retallack, 2001, 2009a; Beerling et al., 2002; Royer, 2003;	
84	Retallack, 2001, 2009a; <u>K ürschner et al., 2008;</u> Smith et al., 2010), Metasequoia	
85	(Royer, 2003; Doria et al., 2011), Taxodium (Stults et al., 2011), Betula (Kürschner et	
86	al., 2001; Sun et al., 2012), Neolitsea (Greenwood et al., 2003), and Quercus	
87	(K ürschner et al., 1996, 2001) <u>, Laurus and Ocotea (K ürschner et al., 2008) and</u>	

批注 [GD5]: This paragraph needs rewriting to place into proper context. Right now, it does not convey a good understanding of the debate and problems.



88	multiple trees (K ürschner et al., 2008). Recently, positive correlations between
89	stomatal index or stomatal frequency and pCO_2 have been reported based on fossil
90	<i>Typha</i> and <i>Quercus</i> (Bai et al., 2015; Hu et al., 2015). However, the tropical and
91	subtropical moist broadleaf forest conifer tree Nageia has <u>not</u> been overlooked used
92	<u>previously</u> in paleobotanical estimates of pCO_2 concentration.
93	Herein, we firstly document the correlations between the SD and SIstomatal
94	properties and atmospheric CO ₂ concentrations for of using Nageia using leaves of
95	the extant species Nageia motleyi (Parl.) De Laub. that were leaves and collected
96	atmospheric CO ₂ -concentrationover the last two centuries. This to-provides a training
97	dataset for application to fossil representatives of <i>Nageia</i> . Furthermore, wWe
98	secondly estimate a new pCO ₂ level for the late Eocene using the stomatal density
99	method based-measure stomatal parameters on fossil Nageia leaves from the late
100	Eocene of South China to estimate past CO ₂ levels. The work, provides furthering
101	significant implications on insights fordiscussing Eocene the climate change.
102	rhythm throughout the Eocene.
103	
104	2 Background
105	
106	2.1 Stomatal proxy in pCO ₂ research
107	
108	Stomatal proxy-information gathered from careful examination of leaves is has
109	been widely used in for reconstructions of past pCO ₂ concentrations (Beerling and

 批注 [GD6]: Seems awkward

 带格式的:字体颜色:绿色,字距调整小四

批注 [GD7]: Why not include these in the above list??

110	Kelly, 1997; Doria et al., 2011 REF). The three main parameters are stomatal density	
111	(SD), which is expressed as the total number of stomata divided by area, epidermal	
112	density (ED), which is expressed as the total number of epidermal cells per area, and	
113	the stomatal index (SI), which is calculated by the calculating defined as the	
114	percentage of stomata among the total number of cells within an area [SI = SD×100 /	
115	(SD+ED)]. Woodward (1987) considered that both SD and SI had inverse	
116	relationships with atmospheric CO ₂ during the development of the leaves.	
117	Subsequently, McElwain (1998) created the stomatal ratio (SR) method to reconstruct	
118	pCO ₂ . SR is a ratio of the stomatal density or index of a fossil $[SD_{(f)} \text{ or } SI_{(f)}]$ to that of	
119	corresponding nearest living equivalent $[SD_{(e)} \text{ or } SI_{(e)}]$, expressed as follows:	
120	$SR = SI_{(e)} / SI_{(f)} $ (1)	
121	The stomatal ratio method is a semi-quantitative method of reconstructing pCO_2	
122	concentrations under certain standardizations. One <u>An example</u> is the "Carboniferous	
123	standardization" (Chaloner and McElwain, 1997), indicating that where one stomatal	
124	ratio unit was equals to two RCO2 units and :	
125	$SR = 2 RCO_2 $ (2)	
126	and Where the value of RCO_2 is the pCO ₂ level divided by the pre-industrial	
127	atmospheric level (PIL) of 300 ppm (McElwain, 1998) or that of the year when the	
128	nearest living equivalent (NLE) was collected (Berner, 1994; McElwain, 1998):	
129	$\text{RCO}_2 = C_{(f)} / 300 \text{ or } \text{RCO}_2 = C_{(f)} / C_{(e)}$ (3)	
130	Then tThe estimated pCO ₂ level is can then be expressed as follows:	
131	$C_{(f)} = 0.5 \times C_{(e)} \times SD_{(e)} / SD_{(f)} \text{ or } C_{(f)} = 0.5 \times C_{(e)} \times SI_{(e)} / SI_{(f)} $ (4)	

批注 [GD8]: Probably two good summary papers would suffice here.

<u>w</u> Where $C_{(f)}$ is the pCO ₂ are presented by the fossil leaf, and $C_{(e)}$ is the atmospheric		
CO_2 of the year when the <u>leaf of the NLE</u> species was collected (McElwain and		
Chaloner, 1995, 1996; McElwain 1998). The equation adapts to the pCO_2		
concentration prior to Cenozoic era.		批注 [GD9]: Not sure what this last sentence means.
Another standardization, the "Recent standardization" (McElwain, 1998), is		
expressed as one stomatal ratio unit being equal to one RCO ₂ unit:		
$SR = 1 RCO_2 $ (5)		
According to the equations stated above, the pCO ₂ concentration can be expressed		
as:		
$C_{(f)} = C_e \times SD_{(e)} / SD_{(f)} \text{ or } C_{(f)} = C_e \times SI_{(e)} / SI_{(f)}$ (6)		
Where $C_{(f)}$, $C_{(e)}$, $SD_{(e)}$, $SD_{(f)}$, $SI_{(e)}$ and $SI_{(f)}$ are stated above. This standardization is		
usually used for reconstruction based on Cenozoic fossils (Chaloner and McElwain,		
1997; McElwain, 1998; Beerling and Royer, 2002).		批注 [GD10]: The last few bits can be condensed as sort of a repeat from a
Kouwenberg et al. (2003) proposed some special stomatal quantification methods		above.
for the conifers leaves with stomata arranged in rows. The main terms are as follows:		
stomatal nmumber per Length (SNL) is expressed as (the number of abaxial stomata		
plus the number of adaxial stomata) divided by leaf length in millimeters. Stomatal	\langle	批注 [GD11]: Correct?
rows (SRO) is expressed as the number of stomatal rows in both stomatal bands.		批注 [GD12]: Things get problematic here as some of this seems to be methods.
Stomatal density per length (SDL) is expressed as the equation $SDL = SD \times SRO$.		The parts on Nageia seemingly belong later.
<u>Turue stomatal density per length (TSDL) is expressed as the equation TSDL = SD \times</u>		
<u>Turue stomatal density per length (TSDL) is expressed as the equation $TSDL = SD \times B$</u> band width (in millimeters). The band width on <i>Nageia motleyi</i> leaves was measured		
	Chaloner, 1995, 1996; McElwain 1998). The equation adapts to the pCO ₂ concentration prior to Cenozoic era. Another standardization, the "Recent standardization" (McElwain, 1998), is expressed as one stomatal ratio unit being equal to one RCO ₂ unit: $SR = 1 \text{ RCO}_2$ (5) According to the equations stated above, the pCO ₂ concentration can be expressed as: $C_{(f)} = C_e \times SD_{(e)} / SD_{(f)}$ or $C_{(f)} = C_e \times SI_{(e)} / SI_{(f)}$ (6) Where $C_{(f)} = C_{(e)} - SD_{(e)} - SD_{(e)} - SI_{(e)} - and SI_{(f)}$ are stated above. This standardization is usually used for reconstruction based on Cenozoic fossils (Chaloner and McElwain, 1997; McElwain, 1998; Beerling and Royer, 2002). Kouwenberg et al. (2003) proposed some special stomatal quantification methods. for the conifers leaves with stomata arranged in rows. The main terms are as follows: stomatal number per Length (SNL) is expressed as (the number of abaxial stomata plus the number of adaxial stomata) divided by leaf length in millimeters. Stomatal rows (SRO) is expressed as the number of stomatal rows in both stomatal bands.	Chaloner, 1995, 1996; McElwain 1998). The equation adapts to the pCO ₂ concentration prior to Cenozoic era, Another standardization, the "Recent standardization" (McElwain, 1998), is expressed as one stomatal ratio unit being equal to one RCO ₂ unit: $SR = 1 RCO_2$ (5) According to the equations stated above, the pCO ₂ concentration can be expressed as: $C_{(j)} = C_e \times SD_{(e)} / SD_{(j)}$ or $C_{(j)} = C_e \times SI_{(e)} / SI_{(j)}$ (6) Where $C_{(j)} - SD_{(e)} - SD_{(j)} - SI_{(e)} - and SI_{(g)}$ are stated above. This standardization is usually used for reconstruction based on Cenozoic fossils (Chaloner and McElwain, 1997; McElwain, 1998; Beerling and Royer, 2002). Kouwenberg et al. (2003) proposed some special stomatal quantification methods. for the conifers leaves with stomata arranged in rows. The main terms are as follows: stomatal number per Length (SNL) is expressed as the number of abaxial stomata plus the number of adaxial stomata) divided by leaf length in millimeters. Stomatal rows (SRO) is expressed as the number of stomatal rows in both stomatal bands.

155 2.2 Review of extant and fossil Nageia

156

157	The genus Nageia-, including only seven living species, is an special group of
158	Podocarpaceae-, which is a large family of conifers mainly distributed in the southern
159	hemisphere. <u>Nageia has with broadly ovate-elliptic to oblong-lanceolate</u> , multiveined
160	(without a midvein), spirally arranged or in decussate, and opposite or subopposite
161	leaves (Cheng et al., 1978; Fu et al., 1999). Generally, <i>Nageia</i> is divided into <i>Nageia</i>
162	Sect. Nageia and Nageia Sect. Dammaroideae (Mill 1999, 2001).; Both sections are
163	XX -mainly distributed in <u>south</u> east ern Asia <u>and Australasia</u> from north latitude 30 ° to
164	nearly the equator and coastal mountain areas and island areas of the western Pacific-
165	Ocean, including South China, South Japan, Malaya and Indonesia, New Guinea, and
166	other Pacific islands (Fu, 1992; Fig. 1a). Four species of the N. section Nageia, ie.,
167	Nageia nagi (Thunberg) O. Kuntze, N. fleuryi (Hickel) De Laub., N. formosensis
168	(Dummer) C. N. Page, and N. nankoensis (Hayata) R. R. Mill, of the N. section-
169	Nageia, have hypostomatic leaves where (the stomata only distributed occur on the
170	abaxial side.) leaves, with a single exceptionOne species of this section N. maxima
171	(De Laub.) De Laub. which is characterized by amphistomatic leaves, but where
172	only a few stomata are found on the adaxial side (Hill and Pole, 1992; Sun, 2008).
173	Both N. wallichiana (Presl) O. Kuntze and N. motleyi., of the N. section
174	Dammaroideae, are amphistomatic with abundant stomata distributed on both sides of
175	the leaf. This is, especially true for N. motleyimotleyi, which has ving approximately

批注 [GD13]: Split this convoluted sentence and rewrite the second bit as difficult to follow. Does this apply to Podacarp or Nageia?

批注 [GD14]: Both sections? Not clear.

批注 [GD15]: Avoid Latin abbreviations in the middle of sentences

176	equal stomata <u>numbers</u> on both surfaces (Hill and Pole, 1992; Sun, 2008).
177	The fossil records of <i>Nageia</i> can be traced back to the Cretaceous. <u>Krassilov (1965)</u>
178	described Podocarpus (Nageia) sujfunensis Krassilov from the Lower Cretaceous of
179	Far East Russia. Kimura et al. (1988) reported Podocarpus (Nageia) ryosekiensis
180	Kimura, Ohanaet Mimoto, an ultimate leafy branch bearing a seed, from the Early
181	Barremian in southwestern Japan. In China, a Cretaceous petrified wood, Podocarpus
182	(Nageia) nagi Pilger, was discovered from the Dabie Mountains in central Henan,
183	China (Yang et al., 1990). Jin et al. (2010) reported a upper Eocene Nageia leaf
184	named N. hainanensis Jin, Qiu, Zhu et Kodrul from the Changchang Basin of Hainan
185	Island, South China. Recently, Liu et al. (2015) found another leaf species N.
186	maomingensis Jin et Liu from upper Eocene of Maoming Basin, South China.with the
187	evidence from Far East Russia, Japan and Henan, China (Krassilov, 1965; Matsuo,-
188	1977; Kimura et al., 1988; Yang, 1990) and be extended into the Eocene of Hainan-
189	Island and Guangdong (Maoming), South China (Jin et al., 2010; Liu et al., 2015; Fig.
190	1a). Although some of the <i>Nageia</i> fossil materials described <u>in the above studies</u>
191	(Krassilov, 1965; Jin et al., 2010; Liu et al., 2015) have well-preserved cuticles (e.g.,
192	Jin et al., 2010; Liu et al., 2015), the above _ these studies are mainly concentrated on
193	the morphology, systematics and phytogeography.
194	Here we try to reconstruct the pCO ₂ concentration based on stomatal data of
195	Nageia maomingensis Jin et Liu <u>. was reported described on the basised of</u> n four-
196	leaves with well preserved cuticles recovered fromlate upper Eocene sedimentary
197	rocks of South China (Liu et al., 2015). Among the modern Nageia species mentioned

批注 [GD16]: This needs slight expansion and conformity. I would list specific locations, types of material and approximate ages.

批注 [GD17]: As above, not quite clear what is being suggested.

198	above, <i>N. motleyi</i> was considered as the nearest extant livingNLE (NEL) species of <i>N</i> .	
199	maomingensis (Liu et al., 2015). However, because of the species-specific inverse	
200	relationship between atmospheric CO ₂ partial pressure and SD (Woodward and	
201	Bazzaz, 1988), it is necessary to explore whether the SD and SI of N. motleyi show	
202	negative correlations with the CO_2 concentration before applying the stomatal method.	
203	Both N. maomingensis and N. motleyi are amphistomatic, indicating suggesting that	
204	both upper and lower surfaces of the leaf are needed to estimate the pCO ₂	
205	concentration-sduring the late Eocene. <u>**However, because of the species specific</u>	
206	inverse relationship between atmospheric CO2-partial pressure and SD (Woodward-	
207	and Bazzaz, 1988), it is necessary to explore whether the SD and SI of the N. motleyi-	
208	show the negative correlations with the CO ₂ concentration before applying the-	
209	stomatal method. <u>**</u>	
210		
211	3 Material and methods	
212		
213	3.1 Extant leaf preparation	
214		
215	We examined five- <u>12</u> specimens of extant Nageia motleyi from different herbaria	
216	(Table 1): <u>***(1) the specimen numbered 2649 (the herbarium of the V. L. Komarov</u>	
217	Botanical Institute of the Russian Academy of Sciences) was collected by Beccari O	
218	from Malaysia in 1868; (2) the specimen numbered bb. 17229 (Harvard University-	
210		
219	Herbarium) was collected by Neth. Ind. For. from Riau on Ond. Karimon, Archipel.	•

批注 [GD18]: Between ** This important point – species specific relationships -needs to come earlier

220	Ind. in the latitude of 150 m in 1932; (3) the specimen numbered bb. 18328 (Harvard-
221	University Herbarium) was collected from Z. O. afd. v. Borneo Tidoengsche Landen,
222	Archipel. Ind., Malaysia in the latitude of 5 m in 1934; (4) the specimen numbered bb.
223	21151 (Harvard University Herbaria) was collected from Z. O. afd. Borneo, Poeroek-
224	Tjahoe Tahoedjan, Archipel. Ind. in the latitude of 500 m in 1936; (5) the specimen-
225	numbered bb. 40798 (Herbarium of Royal Botanic Garden, Edinburgh) was collected
226	by Sinclair, J. and Kiah bin, Salleh from Gunong Tebu Forest reserve, Malaysia in the-
227	latitude of 51 m in 1955 (Table 1). *** We removed one or two leaves from each
228	specimen, and took three fragments (0.25 mm ²) from every leaf (Fig. 2a) and
229	numbered them for analysis.
230	The numbered fragments were boiled for 5-10 min in water. Subsequently, after
231	being macerated in a mixed solution of 10% acetic acid and 10% $H_2O_2(1:1)$ and
232	heated in the thermostatic water bath at 85 C for 8.5 hours; the reaction was stopped
233	when the specimens fragments turned white and semitransparent; The cuticles were
234	then rinsed with distilled water until the pH of the water became neutral. After that the
235	cuticles were treated in Schulze's solution (one part of potassium chlorate saturated
236	solution and three part of concentrated nitric acid) for 30 min, rinsed in water, and
237	then treated with 8% KOH (up to 30 min) and the abaxial and adaxial cuticles were
238	separated with a hair mounted on needle. Finally, the cuticles were stained by with 1%
239	Safranin T alcoholic solution for 5 min, sealed with Neutral Balsam and observed
240	under the -LM.

批注 [GD19]: Between *** I think can removed as long as clearly documneted in Tabel 1

241

3.2 Fossil leaf preparation

243	
244	** Needs a brief paragraph here noting the Maoming Basin, rock units, depositional
245	environment, and criticially, the age.**
246	Maoming Basin (21 °42'33.2"N, 110 °53'19.4"E) is located in southwestern
247	Guangdong, South China including Cretaceous and Tertiary strata. Tertiary strata are
248	fluvial and lacustrine sedimentary units, divided into the Gaopengling, Laohuling,
249	Shangcun, Huangniuling and Youganwo formations in descending order, aged from
250	late Eocene to early Oligocene (Wang et al., 1994).
251	Four fossil leaves of Nageia maomingensis were recovered from the late Eocene of
252	the Youganwo Formation (MMJ1-001) and the Huangniuling Formation (MMJ2-003,
253	MMJ2-004 and MMJ3-003) formations of Maoming Basin, South China (Fig. 1B, 1C
254	in Liu et al., 2015). The age from Youganwo to Huangniuling formations is late
255	Eocene (~ 40.3 Ma). the layersPrecise information wasregarding locations areis
256	<u>provided shown in Fig. 1b-c inby Liu et al., (2015</u>). Macrofossil cuticular fragments
257	were taken from the middle part of each fossil leaf (Fig. 2c) and treated by with
258	Schulze's solution for approximately 1h and 5–10% KOH for 30 min (Ye, 1981). The
259	cuticles were observed and photographed under a Carl Zeiss Axio Scope A1 light
260	microscope (LM). All fossil specimens and cuticle slides are housed in the Museum of
261	Biology of Sun Yat-sen University, Guangzhou, China.
262	

3.3 Stomatal counting strategy and calculation methods

The basic stomatal parameters, SD, ED and SI, weare counted based on analyzing
the pictures taken with the <u>a</u> light microscope (LM) using the standard sampling
protocol, only counting those stomata touching or straddling the left hand side and top
including the corner between them, provided by Poole and Kürschner (1999; Fig. 2b,
$\frac{2d}{2d}$. A total of $\frac{1116-2816}{2816}$ pictures (200× magnification of Zeiss LM) of the cuticles
from <u>9-21</u> leaves of <i>N. motleyi</i> were counted. Each counting field was 0.366 mm ² . We
used a standard sampling protocol (Poole and K ürschner, 1999), and only counted
counting all full stomata in the image plus stomata straddling the left and top margins,
as presented in Figure 2(b), and (d). those stomata touching or straddling the left-hand-
side and top including the corner between them, provided by Fig. 2b, 2d) In <i>Nageia</i> -
maomingensis, 112 views (400×magnification) of the abaxial side and 150 views-
$(400 \times magnification)$ of the adaxial side of cuticles were counted with an area of
0.092 mm ² . None of the counting areas above overlapped and they were larger than-
the minimum area (0.03 mm ²) for statistics (Poole and K ürschner, 1999). In this study,
the stomatic data of both surfaces are applied in pCO ₂ reconstruction because both our
fossil species and the NLE species are amphistomatic.
The SNL, SRO, SDL, and TSDL weare also countdetermined based on the pictures
taken with the light microscope (LM images.) using the strategies stated in the
background**A total of 2293 pictures (200×magnification of Zeiss LM) of the
cuticles from 21 leaves of <i>N. motleyi</i> were counted. Each counting field was 0.366
mm ² . None of the aforementioned counting areas overlapped and they were larger

批注 [GD20]: This is not crystal clear to me

批注 [GD21]: See above. The specifics to Nageia should be here, but the more general procedure should be presented above.

286	than the minimum area (0.03 mm ²) for statistics (Poole and Kürschner, 1999). In this
287	study, the stomatal data of both surfaces are applied in pCO ₂ reconstruction because
288	both the fossil and NLE species are amphistomatic.
289	
290	
291	4 Results
292	
293	4.1 Correlations between the CO ₂ concentrations and stomatal parameters of
294	Nageia motleyi
295	
296	The SD and SI data of the adaxial sides of N. motleyi leaves are shown-presented in
297	Table 2. The SDs and SisSIs average _ range from 62.28 mm ⁻² and 3.30 % 45.89 to-
298	78.6 (mm ⁻²) and from 2.89 to 3.94 (%), respectively. However, the SDs and SIs data
299	of the abaxial sides, summarized in Table 3, give significantly higher average values
300	(<u>70.03 mm⁻²53.22 82.71</u> in SDs and <u>3.13 4.663.90 % in</u> SIs) than those from the
301	adaxial sides. The combined SD and SI of the adaxial and abaxial surfaces average
302	66.14 mm ⁻² and 3.60 %, respectively (table 4).
303	Figure Fig. 3 shows the relationships between the stomatal parameters (SD and SI)
304	of modern <i>N. motleyi</i> and the atmospheric CO ₂ concentration (SD-CO ₂ relationships
305	and SI-CO ₂ relationships). R^2 values in the SDCO ₂ relationships from both the
306	adaxial and abaxial surfaces of <i>N. motley</i> are up to 0.841-0.4667 and 0.7250.3824 (Fig.
307	3a, b), suggesting that the stomatal densities of <i>N. motleyi</i> are in significant inverse

308	proportion to the CO ₂ concentrations. However, the figure Fig. $3(c)$ and (d) indicate no
309	relationships between the SIs and CO_2 concentrations for the extremely low level of
310	the R ² values (0.0030.2558 and 0.06080.0248). Figs. 3e and 3f based on the combined
311	data also show that SD inversely responds to the atmospheric CO_2 concentration (R^2)
312	=0.4421), while SI has almost no relationship with the atmospheric CO_2 concentration
313	$(\mathbf{R}^2 = 0.1177).$
314	The mean values of SNL, SDL and TSDL are 9.81, 326.39 and 1226.93 no. \cdot mm ⁻¹ ,
315	respectively (Table 5). Fig. 4 shows the relationships between SNL (SDL, TSDL) and
316	<u>CO₂ concentrations. The low R² values in the Fig. 4a and 4c indicate that SNL (R² = </u>
317	0.0643) and TSDL ($R^2 = 0.0788$) have no relationship with the CO ₂ concentration in
318	this study. Fig. 4b shows that there is a weak reverse relevance between SDL and the
319	$\underline{CO_2}$ concentration (R ² = 0.3154).
319 320	<u>CO₂ concentration (R² = 0.3154).</u> Compared with the SDL method, the SD-based method shows a larger R ² value,
320	Compared with the SDL method, the SD-based method shows a larger R ² value,
320 321	<u>Compared with the SDL method, the SD-based method shows a larger R^2 value,</u> indicating a stronger relevance between the SD and CO ₂ concentrations. In this study,
320 321 322 323	Compared with the SDL method, the SD-based method shows a larger R^2 value, indicating a stronger relevance between the SD and CO ₂ concentrations. In this study, the pCO ₂ is reconstructed based on the regression equations of SD-CO ₂ relationship.
320 321 322	Compared with the SDL method, the SD-based method shows a larger R ² value, indicating a stronger relevance between the SD and CO ₂ concentrations. In this study, the pCO ₂ is reconstructed based on the regression equations of SD-CO ₂ relationship. Additionally,
320 321 322 323 324	Compared with the SDL method, the SD-based method shows a larger R ² value, indicating a stronger relevance between the SD and CO ₂ concentrations. In this study, the pCO ₂ is reconstructed based on the regression equations of SD-CO ₂ relationship. Additionally, According to the results stated above, the stomatal ratio method can be also used in
320 321 322 323 324 325	Compared with the SDL method, the SD-based method shows a larger R ² value, indicating a stronger relevance between the SD and CO ₂ concentrations. In this study. the pCO ₂ is reconstructed based on the regression equations of SD-CO ₂ relationship. Additionally, According to the results stated above, the stomatal ratio method can be also used in estimating pCO ₂ concentration of the late Eocene based on the stomatal densities
 320 321 322 323 324 325 326 	Compared with the SDL method, the SD-based method shows a larger R ² value, indicating a stronger relevance between the SD and CO ₂ concentrations. In this study, the pCO ₂ is reconstructed based on the regression equations of SD-CO ₂ relationship. Additionally, According to the results stated above, the stomatal ratio method can be also used in estimating pCO ₂ concentration of the late Eocene based on the stomatal densities (SDs) of the fossil species <i>N. maomingensis</i> and the extant species <i>N. motleyi</i> .

330	(Beerling, 1999) and more accurate in responding to the variation of pCO ₂ -
331	concentration (Royer, 2001). However, the study of Kouwenberg et al. (2003)
332	indicated that the SD better reflects the negative relationships with atmospheric CO ₂ -
333	concentration.
334	The SD results of specimen No. 18328 are selected to reconstruct the pCO ₂
335	concentration, because they are closest to the fitted equations in Fig. 3. This specimen
336	was collected by Neth. Ind. For. Service from Riau on Ond. Karimon, Archipel. Ind.,
337	Malaysia, in 1934 at an altitude of 5 m and CO ₂ concentration of 306.46 ppmv
338	(Brown, 2010). The SD results of specimen No. 40798 are closest to the fitted-
339	equations in Fig. 3a and 3b and therefore are selected to reconstruct the pCO_2 -
340	concentration. The specimen was collected by J. Sinclair. & Salleh Kiah Bin from-
341	Gunong Tebu Forest Reserve, Malaysia, in 1955 at an altitude of 51 m and a CO ₂ -
342	concentration of 313.73 ppmv during that time (Brown, 2010). Therefore, the SD-
343	from the adaxial and abaxial surfaces of <i>N. maomingensis</i> and its NLE species <i>N</i> .
344	<i>motleyi</i> are used to recover pCO ₂ concentrations based on the stomatal ratio method.
345	
346	4.2 Stomatal parameters and The pCO ₂ estimates results
347	4.2.1 The regression approach
348	After being projected into a long-term carbon cycle model (GEOCARB III; Berner-
349	and Kothaval á, 2001), the results of this study compares well with the CO2-
350	concentrations for corresponding age within their error ranges (Fig. 4). The summary
351	of stomatal parameters of the extant and fossil Nageia and reconstruction results are
	1

352	provided in Tables 2 <u>6–8. and 3, respectively. SD and SI values were calculated for all-</u>
353	samples of the extant and fossil Nageia. The mean SD and SI values of the adaxial
354	surface are <u>44.5 mm⁻²44.5 ± 2.9</u> and <u>1.8 % 1.80 ± 0.12</u> , respectively (Table <u>56</u>). The
355	mean SD and SI values of the abaxial surfacevalues of the abaxial and abaxial surface
356	are $49.8 \text{ mm}^{-2}48.9 \pm 3.0$ and $2.07 \% 53.22 \pm 2.2$, respectively (Table 2, 37).
357	Based on the regression approach, the pCO ₂ was reconstructed as 351.9 ± 6.6 ppmv
358	and 365.6 \pm 7.6 ppmv according to the SD of adaxial and abaxial sides. The combined
359	SD value is an average of 46.6 mm ⁻² (Table 8), giving the reconstructed pCO ₂ of
360	<u>358.1 ± 5.0 ppmv.</u>
361	
362	4.2.2 The stomatal ratio method
363	The mMean SR values of both sides the adaxial side (SR=1.69 ± 0.18) are is a little
364	larger quite similar with 1.24 ± 0.13 than that in adaxial of the abaxial side (SR=1.60 \pm
365	<u>0.11) in fossil Nageia leaves and 1.23 ± 0.09 in abaxial side</u> (Tables 4- <u>6</u> and <u>57</u>). The
366	pCO ₂ average reconstruction results are of pCO ₂ concentration in the late Eocene of
367	Maoming Basin is 391.0 $\pm 41.1537.5 \pm 56.5$ ppmv (Table 46) and $\frac{386.5 \pm 27.8496.1}{27.8496.1}$
368	± 35.7 ppmv (Table 57) with a 95% confidence interval based on the adaxial and
369	abaxial cuticles, respectively. Based on the combined SD of both leaf sides, the pCO_2
370	result is 519.9 \pm 35.0 ppmv. Clearly the two estimates are rather similar with a
371	difference of 5 ppmv in mean value, which is clearly less than their own standard-
372	error, indicating that the reconstructions based on both sides are consistent in this-
373	fossil species. Table 4 shows gradually increasing pCO ₂ level from the lower layer to-

the upper ones, while the pCO₂ estimated results based on the abaxial side are random with the highest result in lowest layer (Table 5).

376 The partial pressure of CO_2 decreases with elevation (Gale, 1972). Jones (1992) 377 proposed that the relationship between elevation and partial pressure in the lower 378 atmosphere can be expressed as P = -10.6E + 100, where E is elevation in kilometers and P is the percentage of partial pressure relative to sea level. Various studies 379 380 corroborate that SI and SD of many plants have positive correlations with altitude (Körner and Cochrane, 1985; Woodward, 1986; Woodward and Bazzaz, 1988; 381 382 Beerling et al., 1992; Rundgren and Beerling, 1999) while they are negatively related 383 to the partial pressure of CO_2 (Woodward and Bazzaz, 1988). Therefore, it is essential 384 to take elevation calibration into account during the pCO₂ concentration estimates. 385 However, Royer (2003) pointed out that it is unnecessary make thisto provide this conversion when the trees lived at <250 m in elevation. In this paper, the nearest 386 387 living equivalent species, Nageia motleyi, grows at $\frac{51}{5}$ m in elevation with P =99.599.9, suggesting that CO_2 concentration estimates were only underestimated by 388 0.50.1%. Consequently, no correction is needed for the reconstruction result in this 389 390 study. After being projected into a long-term carbon cycle model (GEOCARB III; Berner and Kothaval á 2001), the results of this study compares well with CO₂ 391 concentrations for corresponding age within their error ranges (Fig. 5). 392 393 **5** Discussion 394

396	5.1 Stomatal parameters response to CO ₂
397	Here, we find that SD decreases as atmospheric CO ₂ concentrations increase,
398	however, SI does not. Generally, SI is more sensitive in response to the atmospheric
399	CO ₂ concentration than SD (Beerling, 1999; Royer, 2001). However, the reverse case
400	is not unfound. For example, Kouwenberg et al. (2003) reported that SD is better than
401	SI in reflecting the negative relationships with CO ₂ in conifer needles, accounting for
402	the special paralleled mode of the ordinary epidermal and stomatal formation.
403	Although Nageia is broad-leaved rather than needle-leaved, it also has well paralleled
404	epidermal cells herein showing the different relationships between CO ₂ and SD or SI.
405	Compared with SD, the SDL has weaker correlation with CO_2 at a smaller R^2 . The
406	SNL and TSDL have no response to the change of CO ₂ . The insensitivity of SNL,
407	SDL and TSDL might account for the characters of broad-leaved leaf shape and
408	paralleled epidermal cells. The SNL should be applied to conifer needles with single
409	file of stomata (Kouwenberg et al., 2003). The SDL and TSDL were considered as the
410	most appropriate method when the stomatal rows grouped in bands in a hypo- or
411	amphistomatal conifer needle species (Kouwenberg et al., 2003). Considering all the
412	stomatal parameters above, SD appears to be the most sensitive to CO ₂ .
413	The SD-CO ₂ correlation shows one value from leaf No. 40798 offset from the
414	others. The SI-CO ₂ correlation shows different offset values in different leaf sides.
415	The offset values might be affected by leaf maturity and light intensity. However, it is
416	hard to distinguish whether a fossil leaf is young or mature, or live in the sunny or
417	shady light regimes.

6 5.1 Stomatal parameters response to CO

418	The R^2 value (0.5) of SD-CO ₂ based on the adaxial side is higher than from the
419	abaxial side and the combination of both sides, indicating that the correlation of
420	SD-CO ₂ is stronger than the others parameters herein. Therefore, the SD on the
421	adaxial side is the best in reconstructing pCO ₂ . The reconstruction result based on the
422	regression approach is 351.9 ± 6.6 ppmv lower than the one based on the stomatal
423	ratio method (Table 6), and it is relatively lower than the results based on the other
424	proxies (Fig. 6; Freeman and Hayes, 1992; Pagani et al., 2005; Maxbauer et al., 2014).
425	However, the result based on stomatal ratio method is 537.5 ± 56.5 ppmv which is
426	closest to GEOCARB III (Fig. 5) and historical reconstruction trends (Fig. 6).
427	
428	5.1-2 Paleoclimate reconstructed history
429	
429 430	The pCO ₂ <u>levels</u> throughout the Cenozoic was were relatively lower than the levels
	The pCO ₂ <u>levels</u> throughout the Cenozoic was were relatively lower than the levels through the Cretaceous (Ekart et al., 1999), but it is had an overall decreasing trend
430	
430 431	through the Cretaceous (Ekart et al., 1999), but it is had an overall decreasing trend
430 431 432	through the Cretaceous (Ekart et al., 1999), but it is had an overall decreasing trend with some significantly changesincreases on short-time scales (e.g. in the earliest
430 431 432 433	through the Cretaceous (Ekart et al., 1999), but it is had an overall decreasing trend with some significantly changesincreases on short-time scales (e.g. in the earliest Eocene and middle Miocene, Zachos et al., 2001, 2008; Wing et al., 2005; Lowenstein
430 431 432 433 434	through the Cretaceous (Ekart et al., 1999), but it is had an overall decreasing trend with some significantly changesincreases_on short-time scales (e,g. in the earliest Eocene and middle Miocene, Zachos et al., 2001, 2008; Wing et al., 2005; Lowenstein and Demicco, 2006; Fletcher et al., 2008; Zachos et al., 2008; Bijl et al., 2010; Kato et
430 431 432 433 434 435	through the Cretaceous (Ekart et al., 1999), but it is had an overall decreasing trend with some significantly changesincreases_on short-time scales (e.g. in the earliest Eocene and middle Miocene, Zachos et al., 2001, 2008; Wing et al., 2005; Lowenstein and Demicco, 2006; Fletcher et al., 2008; Zachos et al., 2008; Bijl et al., 2010; Kato et al., 2011). There is a wide range in pCO ₂ estimates for the Paleogene, reflecting both-
430 431 432 433 434 435 436	through the Cretaceous (Ekart et al., 1999), but <u>it is had an overall decreasing trend</u> with some significant ly changesincreases_on short-time scales (e.g. in the earliest Eocene and middle Miocene, Zachos et al., 2001, 2008; Wing et al., 2005; Lowenstein and Demicco, 2006; Fletcher et al., 2008; Zachos et al., 2008; Bijl et al., 2010; Kato et al., 2011). There is a wide range in pCO ₂ estimates for the Paleogene, reflecting both- problems in the various proxies. Both the fractionation of carbon isotopes by

the Paleocene.

441	Based on the measurements of palaeosol carbon isotopes, Cerling (1991) reported
442	that pCO_2 levels for the Eocene and Miocene through to the present was lower than
443	700 ppmv. Fletcher et al. (2008) also showed that an atmospheric CO_2 levels of were
444	approximately 680 ppmv by 60 million years ago. However, Stott (1992)
445	reconstructed pCO ₂ as 450 —550 ppmv for the early Eocene based on phytoplankton.
446	Additionally, the reconstructions using the stomatal ratio method based on the leaves-
447	of Ginkgo, Metasequoia, and Lauraceae <u>leaves</u> also revealed a low pCO_2 level
448	between 300 and 500 ppmv during the early Eocene (Kürschner et al., 2001; Royer et
449	al., 2001; Greenwood et al., 2003; Royer, 2003) except a single high estimate of about
450	800 ppmv near the Paleocene/Eocene boundary (Royer et al., 2001).
451	Subsequently, Smith et al. (2010) reconstructed the value of pCO_2 ranging from 580
452	± 40 to 780 ± 50 ppmv using the stomatal ratio method (recent standardization) based
453	on both SI and SD. A climatic optimum occurred in the middle Eocene (MECO): the
454	reconstructed CO ₂ concentrations are mainly between 700 and 1000 ppmv during the
455	late middle Eocene climate transition (42-38 Ma) using stomatal indices of fossil
456	Metasequoia needles, but concentrations declined to 450 ppmv toward the top of the
457	investigated section (Doria et al., 2011). Jacques et al. (20122014) used CLAMP to
458	calibrate climate change in Antarctica during the early-middle Eocene, suggesting a
459	seasonal alternation of high- and low-pressure systems over Antarctica during the
460	early-middle Eocene. Spicer et al. (2014) also reconstructed a relatively lower cool
461	temperature than δ^{18} O records (Keating-Bitonti et al., 2011) in the middle Eocene of

462	Hainan Island, South China using CLAMP, indicating a not uniformly warm climate
463	in the low latitude during the Eocene. The two results of our study agree well with the
464	estimates of approximately the same period based on phytoplankton (Freeman and
465	Hayes, 1992; Ekart et al., 1999; Pagani et al., 2005) and other stomatal studies-
466	(McElwain, 1998) (Fig. 5). The pCO ₂ level showed an <u>An</u> overall decreasing trend <u>of</u>
467	the pCO ₂ level was presented after the MECO periodmiddle Eocene, indicating the
468	consistance with the pCO ₂ -levels herein (Fig. 56; Retallack, 2009-b).
469	The ice-sheets started to appear in the Antarctic during the Late Eocene (Zachos et
470	al., 2001), then the temperature suffered an apparent further decrease from the late
471	Eccene to the early Oligocene (Roth-Nebelsick et al., 2004), which resulted in the
472	Antarctic being almost fully covered by ice-sheets. Subsequently, the climate
473	variation was comparatively stable with a little wobbling in temperature during the
474	Oligocene period (Fig. 56), while a small and ephemeral Late Oligocene Warming
475	was present in the latest part of the Oligocene, resulting in reducing the ice sheets in-
476	Antarctic ice sheets to a minimum and forming a brief period of glaciation at that time
477	(Zachos et al., 2001). During the Middle Miocene, a quick rise in temperature was
478	shown, which was followed by a small glaciation (Fig. 56; Zachos et al., 2001;
479	Roth-Nebelsick et al., 2004; Beerling and Royer, 2011). Subsequently, the CO ₂
480	concentration decreased gradually and reached 280 ppmv until the period of the
481	industrial revolution (Fig. $\frac{56}{2}$). Since then, however, the CO ₂ concentration rebounded_
482	to present day level.

483 In conclusion, although various results were made by different pCO₂ reconstruction

484	proxies at the same time, their entire decreasing tendency of pCO_2 level are
485	remarkably consistent with each other since the Eocene (Fig. 56). Furthermore, ig. 6
486	shows that during the Eocene the temperature was higher than at present. The
487	reconstructed pCO ₂ concentration of 351.9 \pm 6.6 ppmv based on the regression
488	approach SD of fossil Nageia are 391.0 \pm 41.1 ppmv and 386.5 \pm 27.8 is ppmv,
489	showing shows a remarkably low pCO ₂ level during the early late Eocene. The result
490	based on the stomatal ratio method of 537.5 \pm 56.5 ppmv is closely consistent with
491	the pCO ₂ changes over the geological ages (Fig. 6).
492	
493	5.2 Implications from Nageia motleyi ecology
494	
495	Nageia motleyi is restrictedly distributed in the southern half of Malay Peninsula,
496	adjacent Sumatra, and southern Borneo (Fig. 1a) with the mean annual temperature of
497	ca. 25–30 ℃ which is higher than South China (ca. 20–25 ℃; Fig. 1a). This species is-
498	generally scattered in the canopy of primary and secondary rainforests on massive-
499	substrates and situations from well-drained, even arid, slopes to waterlogged peat-
500	swamps at elevations of 15-500 (~1000) m (Eckenwalder, 2009) and in Borneo-
501	surviving where there is deep peat in a mixed ramin-peat swamp, ridges, and hill sides
502	in bindang-dipterocarp forest, and 1,000 m on podsolic sandy loam (Coomes and
503	Bellingham, 2011). All the living ecological characteristics of N. motleyi provide a
504	significant implication that the temperature during the Late Eocene might have been-
505	similar to that in the area where N. motleyi grows today.

506	Palynological assemblages from the late Eocene of Maoming Basin of Guangdong-
507	(Aleksandrova et al., 2012) suggest that the Youganwo Formation was humid, and the-
508	Huangniuling Formation had an increase of average annual temperatures and
509	humidity during this period. Additionally, according to the winged fruits Shorea-
510	maomingensis Feng, Kodrul et Jin (Dipterocarpaceae) recovered from the late Eocene-
511	of the Huangniuling Formation of the Maoming Basin and the living conditions of
512	modern Shorea, Feng et al. (2013) point out the occurrence of seasonally dry climate-
513	at that time and a temperature higher than today.
514	In this article, we reconstructed the pCO ₂ of the late Eocene as 391.0 ± 41.1 ppmv
515	and 386.5 \pm 27.8 ppmv, which are distinctly higher than the CO ₂ level of
516	289.23-313.73 ppmv from extant leaves collected from 1968 to 1955 (Table 1), but
517	similar to the extant CO ₂ concentration of 387.35–401.52 ppmv from 2009 to 2015-
518	(Brown, 2010; Pieter and Keeling, 2015). Compares with the reconstruction results in-
519	figure 5, our estimates show comparatively low pCO ₂ concentration during the late-
520	Eccene. Combined with the low pCO2 and the living conditions of N. sect.
521	Dammaroideae (adapted to warm areas of East Asia) (Fig. 1), we conclude that the
522	other factors may have played a role in the global climate changing process. Owing to
523	the totally decreasing trend of the global climate change from the late Eocene-
524	reconstructed based on the proxies of stomata, paleosols, phytoplankton and B/Ca-
525	(Fig. 5), the plants of N. Sect. Demmaroideae migrated toward south and ultimately-
526	disappeared from South China (Fig. 1).
527	

528 6 Conclusion

525	
530	In this study, we reconstructed the late Eocene pCO_2 based on the fossil leaves of
531	Nageia maomingensis Jin et Liu from the late Eocene of Maoming Basin, Guangdong
532	Province, China. Nageia is a special element in conifers by its broad multi-veined leaf
533	that lacks mid-vein. The stomatal data analysis suggests that only stomatal densities
534	(SD) from both sides of Nageia motleyi leaves have significant negative correlations
535	with the atmospheric CO ₂ concentration. The SD from the adaxial side gives the best
536	correlation to the CO ₂ . Based on SDs, the pCO ₂ concentration is reconstructed using
537	both the regression approach and the stomatal ratio method. The pCO ₂ result based on
538	the regression approach is 351.9 ± 6.6 ppmv, showing a relatively lower CO ₂ level.
539	The reconstructed result based on the stomatal ratio method is 537.5 ± 56.5 ppmv
540	consistent with the variation trends based on the other proxies. Here, we explored the
541	potential of <i>N. maomingensis</i> in pCO ₂ reconstruction and obtained different results
542	according to different methods, providing a new insight for the reconstruction of
543	paleoclimate and paleoenvironment in conifers. The stomatal data analysis suggests
544	only the stomatal densities from both sides of Nageia motleyi leaves have significant
545	negative correlations with the atmospheric CO ₂ concentration, suggesting that we
546	can estimate the pCO2 of the Eocene in South China based on the stomatal densities-
547	of the Eocene fossil leaves of N. maomingensis and their nearest living equivalent
548	species N. motleyi. Based on the stomatal ratio method, pCO ₂ concentration of the late
549	Eccene of Macming Basin, Guangdong Province, is reconstructed as 391.0 ± 41.1

550	ppmv (based on the adaxial side of leaf cuticles) and 386.5 \pm 27.8 ppmv (based on the
551	abaxial side of leaf cuticles), showing low pCO ₂ levels during the globally warm
552	epoch of the Eocene, which is significantly higher than the historical CO ₂ .
553	concentrations from 1868 to 1955 (around the industrial atmospheric level, 300 ppmv)
554	and similar to the concentration of today.
555	
556	Acknowledgements. This study was supported by the National Natural Science
557	Foundation of China (Grant No. 41210001, 41572011), the National Basic Research-
558	Program of China (973 Program) (Grant No. 2012CB822003), State Key Laboratory-
559	of Palacobiology and Stratigraphy (Nanjing Institute of Geology and Palacontology,
560	CAS) (Grant No. 123110), the Fundamental Research Funds for the Central
561	Universities (Grant No.121gjc04), and the Guangdong Provincial Natural Science-
562	Foundation of China (Grant No.10151027501000020), the Key Project of Sun Yat-sen
563	University for inviting foreign teachers , the Scientific Research Fund, Hongda Zhang,
564	Sun Yat-sen University, and the State Scholarship Fund of China Scholarship Council-
565	(CSC) (File No. 201306380046). We greatly thank the Sun Yet-sen (SYS) University-
566	Herbarium and the Herbarium of the V.L. Komarov Botanical Institute of the Russian
567	Academy of Sciences (LE) for their permission to examine and collect extant Nageia
568	specimens. We also express sincere gratitude to Prof. Sun Tongxing (Yancheng
569	Teachers University), Dr. David Boufford (Harvard University) and Dr. Richard
570	Chung Cheng Kong (Forest Research Institute Malaysia) for providing extant N.
571	motleyi leaves from the herbarium of the Royal Botanic Garden at Edinburgh (E), the

572	Harvard University Herbaria (A/GH) and the herbarium of Forest Research Institute
573	Malaysia (KEP). We sincerely appreciate the guidance of Chengqian Wang (Harbin
574	Institute of Technology) on preparing Figs. $3-\underline{56}$. We also offer sincere gratitude to
575	Professor Prof. Steven R. Manchester and Mr. Terry Lott (Florida Museum of Natural
576	History, University of Florida) for suggestions and modificationand Ms. Margaret-
577	Joyner (US) for editing.

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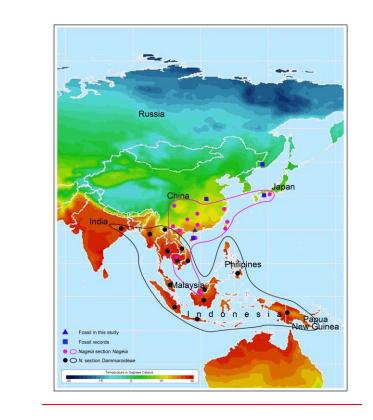
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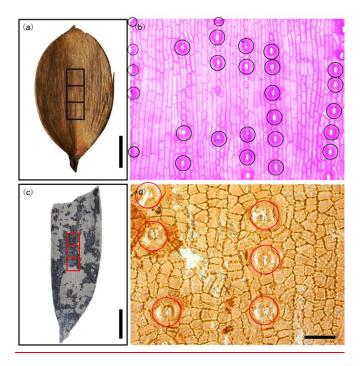
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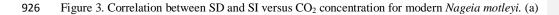
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- 913 Figure 1. Map showing the distribution of extant and fossil *Nageia* and their mean annual
- 914 temperature (Modified after the map from
- 915 <u>http://www.sage.wisc.edu/atlas/maps.php?datasetid=35&includerelatedlinks=1&dataset=35</u>).



- 918 Figure 2. Sampling areas and counting rules are shown. (a) *Nageia motleyi* (Parl.) De Laub.leaf.
- 919 Black squares in the middle of the leaf show the sampling areas for preparing the cuticles. (b) The
- 920 abaxial side of the cuticle from *N. motleyi* leaf. Black circles show the counted stomatal
- 921 complexes. (c) N. maomingensis Jin et Liu. Red squares in the middle of the leaf indicate the
- 922 sampling areas. (d) The abaxial side of the fossil cuticle. Red circles show the counted stomatal
- 923 complexes. Scale bars: (a) and (c) = 1 cm; (b) and (d) = 50 μ m.





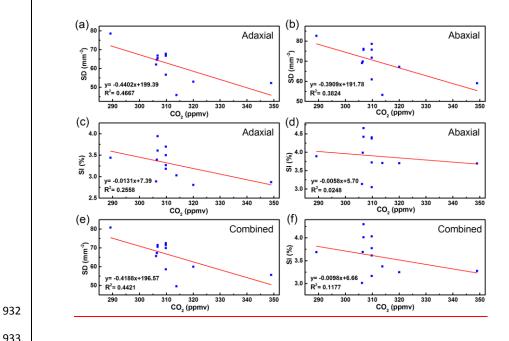
Trends of SD with CO₂ concentration for the adaxial surface. (b) Trends of SD with CO₂ 927

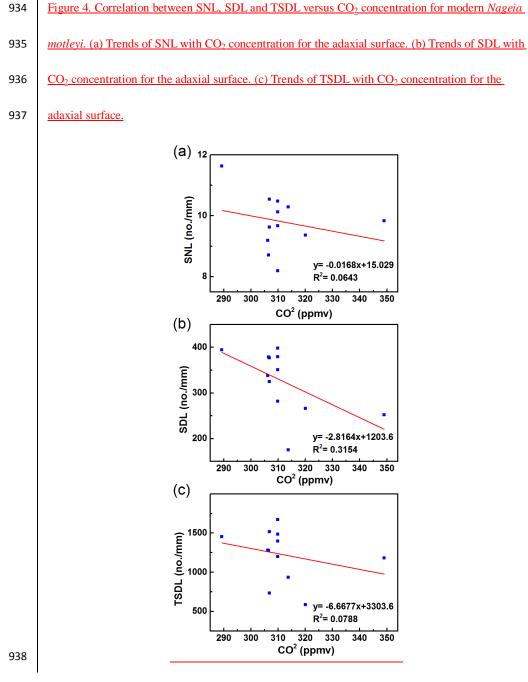
928 concentration for the abaxial surface. (c) Trends of SI with CO₂ concentration for the adaxial

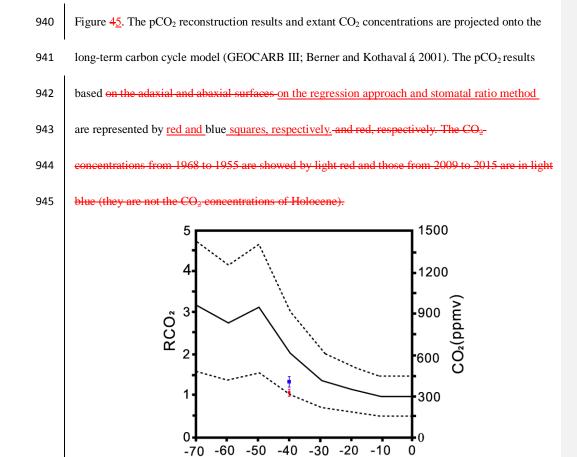
929 surface. (d) Trends of SI with CO₂ concentration for the abaxial surface. (e) Trends of SD with

930 CO₂ concentration for the combined data of both leaf surfaces. (f) Trends of SI with CO₂

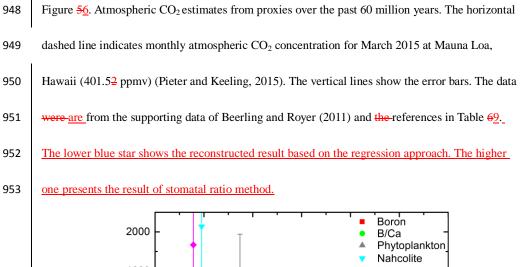


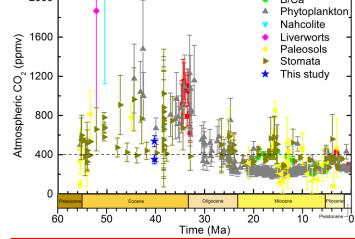






Time (Ma)





955 Table 1. Modern <i>Nageia motleyi</i> (Parl.) De Laub samples and atmospheric CO ₂ valu	lues of their collection dates from ice core data (Brown, 2010).
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Herbarium	Collection-	Collecting levelity	Collectors	Number of	Collection	CO 2-
	number	Collecting locality	Concetors	leaf samples	date	(ppmv)
LE	No. 2649	Malaysia	Beccari, O.	4	1868	289.23
A/GH	No. bb. 17229	150 m, Riau on Ond. Karimon, Archipel. Ind.	Neth. Ind. For. Service	2	1932	306.19
A/GH	No. bb. 18328	5 m, Z. O. afd. v. Borneo Tidoengsche Landen, Archipel. Ind.	Neth. Ind. For. Service	2	1934	306.46
A/GH	No. bb. 21151	500 m, Z. O. afd. Borneo, Poeroek Tjahoe Tahoedjan,	Neth. Ind. For. Service	2	1936	306.76
		Archipel. Ind.				
Đ	No. bb. 40798	51 m, Kuala Trengganu-Besut Road, Bukit Bintang Block,	Sinclair, J. and Kiah	2	1955	313.73
		Gunong Tebu Forest reserve, Malaysia	bin, Salleh			

Note: A/GH - Harvard University Herbarium, Harvard University, 22 Divinity Avenue, Cambridge, Massachusetts 02138, USA (www.huh.harvard.edu).

E The Herbarium of Royal Botanic Garden, Edinburgh EH3 5LR, Scotland, UK (www.rbge.org.uk).

LE - The Herbarium of the V.L. Komarov Botanical Institute of the Russian Academy of Sciences, Prof. Popov Street 2, Saint Petersburg 197376, Russia (www.binran.ru).

<u>Herbarium</u>	Collection number	Collecting locality	Collectors	<u>Number of</u> leaf samples	Collection date	<u>CO2</u> (ppmv)
<u>LE</u>	<u>No. 2649</u>	<u>Malaysia</u>	Beccari, O.	<u>1</u>	<u>1868</u>	289.23
<u>A/GH</u>	<u>No. bb. 17229</u>	150 m, Riau on Ond. Karimon, Archipel. Ind.	Neth. Ind. For. Service	<u>2</u>	<u>1932</u>	<u>306.19</u>
<u>A/GH</u>	<u>No. bb. 18328</u>	5 m, Z. O. afd. v. Borneo Tidoengsche Landen, Archipel. Ind.	Neth. Ind. For. Service	<u>2</u>	<u>1934</u>	<u>306.46</u>
<u>A/GH</u>	<u>No. bb. 21151</u>	500 m, Z. O. afd. Borneo, Poeroek Tjahoe Tahoedjan,	Neth. Ind. For. Service	<u>2</u>	<u>1936</u>	<u>306.76</u>
		Archipel. Ind.				
<u>KEP</u>	<u>No. 30887</u>	<u>Kata Tinggi, Johor, Malaysia</u>	Corner, E.J.H.	<u>1</u>	<u>1936</u>	<u>306.76</u>
<u>KEP</u>	<u>No. 57329</u>	Batang Padang, Perak, Malaysia	<u>Unkonwn</u>	<u>2</u>	<u>1947</u>	<u>309.82</u>
<u>KEP</u>	<u>No. 57330</u>	Batang Padang, Perak, Malaysia	<u>Unkonwn</u>	<u>2</u>	<u>1947</u>	<u>309.82</u>
<u>KEP</u>	<u>No. 55897</u>	Batang Padang, Perak, Malaysia	<u>Unkonwn</u>	<u>2</u>	<u>1947</u>	<u>309.82</u>
<u>KEP</u>	<u>No. 61064</u>	Batang Padang, Perak, Malaysia	Syed Woh	<u>2</u>	<u>1947</u>	<u>309.82</u>

<u>E</u>	<u>No. bb. 40798</u>	51 m, Kuala Trengganu-Besut Road, Bukit Bintang Block,	Sinclair, J. and Kiah	<u>2</u>	<u>1955</u>	<u>313.73</u>		
		Gunong Tebu Forest reserve, Malaysia	bin, Salleh					
<u>KEP</u>	<u>No. 80548</u>	Gombak, Selangor, Malaysia	<u>Rahim</u>	<u>1</u>	<u>1965</u>	<u>320.04</u>		
<u>KEP</u>	<u>No. 33343</u>	Jelebu, Negeri Sembilan, Malaysia	<u>Yap, S.K.</u>	<u>2</u>	<u>1987</u>	<u>348.98</u>		
Note: A/GH—Harvard University Herbarium, Harvard University, 22 Divinity Avenue, Cambridge, Massachusetts 02138, USA (www.huh.harvard.edu).								
<u>E—</u> T	E-The Herbarium of Royal Botanic Garden, Edinburgh EH3 5LR, Scotland, UK (www.rbge.org.uk).							
<u>LE</u> —	LE—The Herbarium of the V.L. Komarov Botanical Institute of the Russian Academy of Sciences, Prof. Popov Street 2, Saint Petersburg 197376, Russia (www.binran.ru).							

KEP-Kepong Herbarium, Forest Research Institute Malaysia, 52109 Kepong, Selangor, Malaysia (http://www.frim.gov.my/).

Collection	Collection	CO ₂ (ppmv)		<u> </u>	$SD (mm^{-2})$					<u>SI (%)</u>		
number	<u>date</u>	<u>eo₂ (ppmv)</u>	<u>x</u>	<u></u>	<u>s.e.</u>	<u>t*s.e.</u>	<u>n</u>	<u>x</u>	<u></u>	<u>s.e.</u>	<u>t*s.e.</u>	<u>n</u>
<u>No.2649</u>	<u>1868</u>	<u>289.23</u>	<u>78.60</u>	<u>15.44</u>	<u>1.41</u>	<u>2.76</u>	<u>120</u>	<u>3.44</u>	<u>0.66</u>	<u>0.06</u>	<u>0.12</u>	<u>120</u>
No.bb.17229	<u>1932</u>	<u>306.19</u>	<u>62.14</u>	<u>17.20</u>	<u>1.78</u>	<u>3.50</u>	<u>93</u>	<u>2.89</u>	<u>0.68</u>	<u>0.07</u>	<u>0.14</u>	<u>93</u>
<u>No.bb.18328</u>	<u>1934</u>	<u>306.46</u>	<u>64.57</u>	<u>15.05</u>	<u>1.58</u>	<u>3.11</u>	<u>90</u>	<u>3.39</u>	<u>1.01</u>	<u>0.11</u>	<u>0.21</u>	<u>90</u>
<u>No.bb.21151</u>	<u>1936</u>	<u>306.76</u>	<u>65.45</u>	<u>11.14</u>	<u>1.17</u>	<u>2.30</u>	<u>90</u>	<u>3.94</u>	<u>0.74</u>	<u>0.08</u>	<u>0.15</u>	<u>90</u>
<u>No.SFN30887</u>	<u>1936</u>	<u>306.76</u>	<u>66.90</u>	<u>16.10</u>	<u>1.27</u>	<u>2.49</u>	<u>161</u>	<u>3.61</u>	<u>0.92</u>	<u>0.07</u>	<u>0.14</u>	<u>161</u>
<u>No.61064</u>	<u>1947</u>	<u>309.82</u>	<u>56.71</u>	<u>16.81</u>	<u>1.95</u>	<u>3.83</u>	<u>74</u>	<u>3.27</u>	<u>1.26</u>	<u>0.15</u>	<u>0.29</u>	<u>74</u>
<u>No.57330</u>	<u>1947</u>	<u>309.82</u>	<u>67.37</u>	<u>15.97</u>	<u>2.04</u>	<u>4.01</u>	<u>61</u>	<u>3.70</u>	<u>0.82</u>	<u>0.10</u>	<u>0.20</u>	<u>61</u>
<u>No.57329</u>	<u>1947</u>	<u>309.82</u>	<u>67.85</u>	<u>15.61</u>	<u>1.70</u>	<u>3.34</u>	<u>84</u>	<u>3.50</u>	<u>0.90</u>	<u>0.10</u>	<u>0.20</u>	<u>84</u>
<u>No.55897</u>	<u>1947</u>	<u>309.82</u>	<u>66.74</u>	<u>14.10</u>	<u>1.78</u>	<u>3.48</u>	<u>63</u>	<u>3.18</u>	<u>0.66</u>	0.08	<u>0.16</u>	<u>63</u>
<u>No.40798</u>	<u>1955</u>	<u>313.73</u>	<u>45.89</u>	<u>13.81</u>	<u>1.12</u>	<u>2.20</u>	<u>151</u>	<u>3.03</u>	<u>0.87</u>	<u>0.07</u>	<u>0.14</u>	<u>151</u>
No.KEP80548	<u>1965</u>	<u>320.04</u>	<u>52.94</u>	<u>11.25</u>	<u>0.85</u>	<u>1.67</u>	<u>175</u>	<u>2.81</u>	<u>0.61</u>	<u>0.05</u>	<u>0.09</u>	<u>175</u>
No.FRI33343	<u>1987</u>	<u>348.98</u>	<u>52.25</u>	<u>12.05</u>	<u>0.77</u>	<u>1.51</u>	<u>242</u>	<u>2.87</u>	<u>0.69</u>	<u>0.04</u>	<u>0.09</u>	<u>242</u>
Mean	=	=	<u>62.28</u>	<u>14.54</u>	<u>1.45</u>	<u>2.85</u>	<u>117</u>	<u>3.30</u>	<u>0.52</u>	0.08	<u>0.16</u>	<u>117</u>
<u>Note: x</u> —mean; σ–	standard devia	tion; s.e. —stand	ard error of	mean; n—	numbers of	photos cou	nts (40×); t•	s.e.— 95% co	onfidence ir	nterval.		

957 Table 2. Summary of stomatal parameters of the adaxial surface form from modern *Nageia motleyi* (Parl.) De Laub.

Collection	Collection	<u>CO₂ (ppmv)</u> –		<u>SI</u>	<u>D (mm⁻²)</u>				<u>SI (%)</u>					
number	<u>date</u>	<u>CO₂ (ppiiiv)</u> –	<u>x</u>	<u></u>	<u>s.e.</u>	<u>t*s.e.</u>	<u>n</u>	<u>x</u>	<u></u>	<u>s.e.</u>	<u>t*s.e.</u>]		
<u>No.2649</u>	<u>1868</u>	<u>289.23</u>	<u>82.71</u>	<u>12.23</u>	<u>1.02</u>	<u>2.00</u>	<u>144</u>	<u>3.89</u>	<u>0.58</u>	<u>0.05</u>	<u>0.09</u>	1		
No.bb.17229	<u>1932</u>	<u>306.19</u>	<u>69.16</u>	<u>14.23</u>	<u>1.48</u>	<u>2.90</u>	<u>93</u>	<u>3.13</u>	<u>0.58</u>	<u>0.06</u>	<u>0.12</u>			
No.bb.18328	<u>1934</u>	<u>306.46</u>	<u>69.92</u>	<u>14.38</u>	<u>1.52</u>	<u>2.97</u>	<u>90</u>	<u>3.99</u>	<u>1.08</u>	<u>0.11</u>	<u>0.22</u>	1		
<u>No.bb.21151</u>	<u>1936</u>	<u>306.76</u>	<u>75.68</u>	<u>15.74</u>	<u>1.66</u>	<u>3.25</u>	<u>90</u>	<u>4.66</u>	<u>0.88</u>	<u>0.09</u>	<u>0.18</u>			
<u>No.SFN30887</u>	<u>1936</u>	<u>306.76</u>	<u>76.18</u>	<u>12.51</u>	<u>0.99</u>	<u>1.93</u>	<u>161</u>	<u>4.42</u>	<u>0.89</u>	<u>0.07</u>	<u>0.14</u>	1		
<u>No.61064</u>	<u>1947</u>	<u>309.82</u>	<u>60.93</u>	<u>11.02</u>	<u>1.39</u>	<u>2.72</u>	<u>63</u>	<u>3.05</u>	<u>0.62</u>	<u>0.08</u>	<u>0.15</u>			
<u>No.57330</u>	<u>1947</u>	<u>309.82</u>	<u>75.82</u>	<u>14.14</u>	<u>1.82</u>	<u>3.58</u>	<u>60</u>	<u>4.38</u>	<u>0.84</u>	<u>0.11</u>	<u>0.21</u>			
<u>No.57329</u>	<u>1947</u>	<u>309.82</u>	<u>71.74</u>	<u>16.84</u>	<u>1.75</u>	<u>3.42</u>	<u>93</u>	<u>3.72</u>	<u>0.62</u>	<u>0.06</u>	<u>0.13</u>			
<u>No.55897</u>	<u>1947</u>	<u>309.82</u>	<u>78.63</u>	<u>13.41</u>	<u>1.75</u>	<u>3.42</u>	<u>59</u>	<u>4.41</u>	<u>1.00</u>	<u>0.13</u>	<u>0.26</u>			
<u>No.40798</u>	<u>1955</u>	<u>313.73</u>	<u>53.22</u>	<u>13.88</u>	<u>1.12</u>	<u>2.19</u>	<u>155</u>	<u>3.71</u>	<u>0.93</u>	0.07	<u>0.15</u>	-		
No.KEP80548	<u>1965</u>	<u>320.04</u>	<u>67.22</u>	<u>13.97</u>	<u>1.07</u>	<u>2.09</u>	<u>171</u>	<u>3.70</u>	<u>0.80</u>	<u>0.06</u>	<u>0.12</u>	-		
No.FRI33343	<u>1987</u>	<u>348.98</u>	<u>59.09</u>	<u>12.10</u>	<u>0.79</u>	<u>1.55</u>	<u>233</u>	<u>3.69</u>	<u>0.86</u>	<u>0.06</u>	<u>0.11</u>			
Mean	=	=	<u>70.03</u>	<u>13.70</u>	<u>1.36</u>	<u>2.67</u>	<u>118</u>	<u>3.90</u>	<u>0.81</u>	<u>0.08</u>	<u>0.16</u>	-		

959Table 3. Summary of stomatal parameters of the abaxial surface form from modern Nageia motleyi (Parl.) De Laub.

960	Table 4

Table 4. Summar	ry of stomatal	parameters of	the combi	ned data of	the adaxia	<u>l and abax</u>	ial surface	es from mode	ern <i>Nageic</i>	<u>ı motleyi (</u>	Parl.) De	Laub.		
Collection	Collection	<u>CO₂ (ppmv)</u> –		<u>S1</u>	$D (\text{mm}^{-2})$				<u>SI(%)</u>					
number	date	<u>CO₂ (ppinv)</u> –	<u>x</u>	<u></u>	<u>s.e.</u>	<u>t*s.e.</u>	<u>n</u>	<u>x</u>	<u></u>	<u>s.e.</u>	<u>t*s.e.</u>	<u>n</u>		
<u>No.2649</u>	<u>1868</u>	<u>289.23</u>	<u>80.84</u>	<u>13.74</u>	<u>0.85</u>	<u>1.66</u>	<u>264</u>	<u>3.69</u>	<u>0.66</u>	<u>0.04</u>	<u>0.08</u>	<u>264</u>		
No.bb.17229	<u>1932</u>	<u>306.19</u>	<u>65.65</u>	<u>16.13</u>	<u>1.18</u>	<u>2.32</u>	<u>186</u>	<u>3.01</u>	<u>0.64</u>	<u>0.05</u>	<u>0.09</u>	<u>186</u>		
No.bb.18328	<u>1934</u>	<u>306.46</u>	<u>67.24</u>	<u>14.92</u>	<u>1.11</u>	<u>2.18</u>	<u>180</u>	<u>3.69</u>	<u>1.08</u>	<u>0.08</u>	<u>0.16</u>	<u>180</u>		
<u>No.bb.21151</u>	<u>1936</u>	<u>306.76</u>	<u>70.57</u>	<u>14.53</u>	<u>1.08</u>	<u>2.12</u>	<u>180</u>	<u>4.30</u>	<u>0.89</u>	<u>0.07</u>	<u>0.13</u>	<u>180</u>		
<u>No.SFN30887</u>	<u>1936</u>	<u>306.76</u>	<u>71.54</u>	<u>15.12</u>	<u>0.84</u>	<u>1.65</u>	<u>322</u>	<u>4.01</u>	<u>0.99</u>	<u>0.05</u>	<u>0.11</u>	<u>322</u>		
<u>No.61064</u>	<u>1947</u>	<u>309.82</u>	<u>58.65</u>	<u>14.54</u>	<u>1.24</u>	<u>2.43</u>	<u>137</u>	<u>3.17</u>	<u>1.02</u>	<u>0.09</u>	<u>0.17</u>	<u>137</u>		
<u>No.57330</u>	<u>1947</u>	<u>309.82</u>	71.56	<u>15.61</u>	<u>1.42</u>	<u>2.78</u>	<u>121</u>	<u>4.03</u>	<u>0.89</u>	<u>0.08</u>	<u>0.16</u>	<u>121</u>		
<u>No.57329</u>	<u>1947</u>	<u>309.82</u>	<u>69.90</u>	<u>16.33</u>	<u>1.23</u>	<u>2.41</u>	<u>177</u>	<u>3.62</u>	<u>0.77</u>	<u>0.06</u>	<u>0.11</u>	<u>177</u>		
<u>No.55897</u>	<u>1947</u>	<u>309.82</u>	<u>72.49</u>	<u>14.95</u>	<u>1.35</u>	<u>2.65</u>	<u>122</u>	<u>3.77</u>	<u>1.04</u>	<u>0.09</u>	<u>0.18</u>	<u>122</u>		
<u>No.40798</u>	<u>1955</u>	<u>313.73</u>	<u>49.60</u>	<u>14.31</u>	<u>0.82</u>	<u>1.60</u>	<u>306</u>	<u>3.37</u>	<u>0.96</u>	<u>0.05</u>	<u>0.11</u>	<u>306</u>		
No.KEP80548	<u>1965</u>	<u>320.04</u>	<u>60.00</u>	<u>14.53</u>	<u>0.78</u>	<u>1.53</u>	<u>346</u>	<u>3.25</u>	<u>0.84</u>	<u>0.05</u>	<u>0.09</u>	<u>346</u>		
<u>No.FRI33343</u>	<u>1987</u>	<u>348.98</u>	<u>55.61</u>	<u>12.53</u>	<u>0.58</u>	<u>1.13</u>	<u>475</u>	<u>3.28</u>	<u>0.88</u>	<u>0.04</u>	<u>0.08</u>	<u>475</u>		
Mean	=	=	<u>66.14</u>	<u>14.77</u>	<u>1.04</u>	<u>2.08</u>	<u>235</u>	<u>3.60</u>	<u>0.89</u>	<u>0.06</u>	<u>0.12</u>	<u>235</u>		

Table 4. Summary of stomatal parameters of the adaxial surface of fossil Nageia and pCO₂ [C_{\oplus}] estimates results.

Note: x—mean; σ—standard deviation; s.e. —standard error of mean; n— numbers of photos counts (40×); t · s.e. — 95% confidence interval.

964	Table 5. Summary of stomatal parameters of the abaxial surface of fossil Nageia and pCO ₂ [$C_{(f)}$] estimates results.

		-		
965	Table 5. Summary of stomatal	parameters from modern Nageia m	otleyi (Parl.) De La	ub (Kouwenberg et al., 2003).

			•		· · · ·		
	Collection number	Collection date	<u>CO2</u> (ppmv)	<u>SNL</u>	<u>SDL</u>	<u>TSDL</u>	<u>n</u>
	<u>No.2649</u>	<u>1868</u>	<u>289.23</u>	<u>11.64</u>	<u>394.38</u>	<u>1455.10</u>	<u>264</u>
	No.bb.17229	<u>1932</u>	<u>306.19</u>	<u>9.19</u>	<u>337.98</u>	<u>1280.12</u>	<u>186</u>
	<u>No.bb.18328</u>	<u>1934</u>	<u>306.46</u>	<u>8.71</u>	<u>378.92</u>	<u>1277.63</u>	<u>180</u>
	No.bb.21151	<u>1936</u>	<u>306.76</u>	<u>9.62</u>	<u>376.93</u>	<u>1517.21</u>	<u>180</u>
	<u>No.SFN30887</u>	<u>1936</u>	<u>306.76</u>	<u>10.55</u>	<u>325.08</u>	<u>735.38</u>	<u>240</u>
	<u>No.61064</u>	<u>1947</u>	<u>309.82</u>	<u>8.19</u>	282.04	<u>1200.66</u>	<u>133</u>
	<u>No.57330</u>	<u>1947</u>	<u>309.82</u>	<u>9.67</u>	<u>397.83</u>	<u>1397.33</u>	<u>119</u>
ĺ	<u>No.57329</u>	<u>1947</u>	<u>309.82</u>	<u>10.13</u>	<u>350.98</u>	<u>1672.50</u>	<u>176</u>
	<u>No.55897</u>	<u>1947</u>	<u>309.82</u>	<u>10.48</u>	<u>379.06</u>	<u>1486.13</u>	<u>122</u>
	<u>No.40798</u>	<u>1955</u>	<u>313.73</u>	<u>10.29</u>	<u>175.14</u>	<u>933.85</u>	<u>305</u>
	No.KEP80548	<u>1965</u>	<u>320.04</u>	<u>9.36</u>	<u>266.16</u>	<u>585.72</u>	<u>263</u>
	No.FRI33343	<u>1987</u>	<u>348.98</u>	<u>9.84</u>	<u>252.20</u>	<u>1181.51</u>	<u>125</u>
	Mean	=	=	<u>9.81</u>	<u>326.39</u>	<u>1226.93</u>	<u>191</u>

Proxies							+	Referen	ces						
Boron	Pearson et a	1., 2009; S	seki et al.	, 2010											
B/Ca	Tripati et al.	, 2009													
Phytoplankton	Freeman and	- veman and Hayes, 1992; Stott, 1992; Pagani et al., 1999, 2005; Henderiks and Pagani, 2008; Seki et al., 2010													
Nahcolite	Lowenstein-	owenstein and Demicco, 2006													
Liverworts	Fletcher et a	etcher et al., 2008													
Paleosols	Cerling, 199	orling, 1992; Koch et al., 1992; Ekart et al., 1999; Retallack, 2009b; Royer et al., 2001													
Stomata	Van der Bur	der Burgh et al., 1993; Kürschner et al., 1996; McElwain, 1998; Kürschner et al., 2001; Greenwood et al., 2003; Royer, 2003; Kürschner													
	at al 2008.	D oorling	at al 200		ar at al.	2001; Retal	llock 20	000.50	nith at al	2010· D	oria et al	2011			
	et al., 2008,	Deerning	et al., 200	$\mathcal{P}, \mathbf{Roy}$	<i>i ct ai.</i> , <i>2</i>	2001, Reta	nack, 20	07a, 51	inui ci ai	., 2010, D	or in or un	, 2011			
Table 6. Summa				-											
Table 6. Summa				he aday				<i>eia</i> an			timates r		<u>ppmv)</u>	<u>C_(f) (p</u>	opmv
Table 6. Summa			ters of t	he aday			ssil Nag	<i>eia</i> an		$[C_{(f)}]$ est	timates r	esults.	<u>ppmv)</u> <u>t*s.e</u>	<u><u>C_(f) (p</u> <u>x</u></u>	
<u>Species</u>	ary of stomata	l parame	ters of t SD (mi	he adax m^{-2}	<u>xial surf</u>	ace of fos	ssil Nag SI (%	<u>eia an</u>	d pCO ₂	[C _(f)] es	timates r <u>R</u>	<u>pCO₂(j</u>			<u>opmv</u> <u>t</u>
Species MMJ1-001 I	ary of stomata	l parame	<u>ters of t</u> <u>SD (m</u> <u>σ</u>	he adax m ⁻²) <u>s.e.</u>	<u>xial surf</u>	$\frac{2}{x}$	ssil Nag SI (% <u>o</u>	<u>eia an</u> 5) <u>s.e.</u>	<u>d pCO₂ n</u>	$\frac{[C_{(f)}] \text{ est}}{\underline{x}}$	timates r R t*s.e	$\frac{pCO_2(1)}{x}$	<u>t*s.e</u>	<u>x</u>	<u>t</u>
Species MMJ1-001 I MMJ2-003 I	ary of stomata Age Late Eocene	<u>x</u> <u>52.5</u>	<u>ters of t</u> <u>SD (m</u> <u>a</u> <u>17.1</u>	$\frac{he adax}{m^{-2}}$ $\frac{s.e.}{3.1}$	<u>n</u> 30	$\frac{\underline{x}}{\underline{2.08}}$	<u>ssil Nag</u> <u>SI (%</u> <u>0.7</u>	<u>eia an</u> 5) <u>s.e.</u> <u>0.1</u>	<u>d pCO₂</u> <u>n</u> <u>30</u>	$\frac{[C_{(f)}] \text{ est}}{\underline{x}}$ $\frac{\underline{x}}{\underline{1.35}}$	$\frac{\text{timates r}}{\text{R}}$ $\frac{\text{t*s.e}}{0.19}$	$\frac{pCO_2(1)}{\underline{x}}$ $\frac{333.6}{\underline{x}}$	<u>t*s.e</u> <u>13.9</u>	<u>x</u> <u>412.1</u>	<u>t</u> <u>(</u> <u>1</u>
Species MMJ1-001 I MMJ2-003 I MMJ2-004 I	Age Age Late Eocene Late Eocene	<u>x</u> <u>52.5</u> <u>42.3</u>	<u>ters of t</u> <u>SD (mr</u> <u>σ</u> <u>17.1</u> <u>12.9</u>	he adax m ⁻²) <u>s.e.</u> <u>3.1</u> <u>2.4</u>	<u>n</u> <u>30</u> <u>30</u>	<u>x</u> <u>2.08</u> <u>1.80</u>	<u>ssil Nag</u> <u>SI (%</u> <u>0.7</u> <u>0.6</u>	<u>eeia an</u> 5) <u>s.e.</u> 0.1 0.1	<u>d pCO₂</u> <u>n</u> <u>30</u> <u>30</u>	$\frac{[C_{ff}] \text{ est}}{\frac{\underline{S}}{\underline{S}}}$ $\frac{\underline{X}}{\underline{1.35}}$ $\underline{1.75}$	timates r R t*s.e 0.19 0.39	$\frac{\underline{pCO_2(1)}}{\underline{x}}$ $\frac{\underline{333.6}}{\underline{356.8}}$	<u>t*s.e</u> <u>13.9</u> <u>10.5</u>	<u>x</u> <u>412.1</u> <u>536.1</u>	<u>t</u>

968 Table 6. The pCO₂ estimates proxies and corresponding references.

result based the regression approach; $C_{(i)}$ — the result based on the stomatal method.

970

972	Table 7. Sun	<u>imary of stomata</u>	<u>i parame</u>	eters of t	ne abaz	<u>ciai sur</u>	lace of tos	<u>sn na</u> g	<u>geta an</u>	<u>a pco</u> 2	$\underline{C_{(f)}}$ es	timates i	esuits.			
	Service	A ===	$\underline{SD} (mm^{-2})$			_	<u>SI (9</u>	<u>6)</u>		<u>S</u>	<u>R</u>	<u>рСО₂(р</u>	<u>omv)</u>	<u>C_(f) (pr</u>	<u>omv)</u>	
	<u>Species</u>	Age	<u>x</u>	<u>σ</u>	<u>s.e.</u>	<u>n</u>	<u>x</u>	<u>σ</u>	<u>s.e.</u>	<u>n</u>	<u>x</u>	<u>t*s.e</u>	<u>x</u>	<u>t*s.e</u>	<u>x</u>	<u>t*s.e</u>
	<u>MMJ1-001</u>	Late Eocene	<u>47.7</u>	<u>17.7</u>	<u>3.2</u>	<u>30</u>	<u>2.11</u>	<u>0.8</u>	<u>0.2</u>	<u>30</u>	<u>1.66</u>	<u>0.23</u>	<u>368.6</u>	<u>16.2</u>	<u>515.6</u>	<u>72.3</u>
	<u>MMJ2-003</u>	Late Eocene	<u>50.9</u>	<u>18.3</u>	<u>3.3</u>	<u>30</u>	<u>2.12</u>	<u>0.8</u>	<u>0.1</u>	<u>30</u>	<u>1.57</u>	<u>0.23</u>	<u>360.9</u>	<u>16.6</u>	<u>486.0</u>	<u>70.7</u>
	<u>MMJ2-004</u>	Late Eocene	<u>48.2</u>	<u>15.8</u>	<u>2.9</u>	<u>30</u>	<u>2.14</u>	<u>0.7</u>	<u>0.1</u>	<u>30</u>	<u>1.63</u>	<u>0.25</u>	<u>367.4</u>	<u>14.5</u>	<u>504.6</u>	<u>77.3</u>
	<u>MMJ3-003a</u>	Late Eocene	<u>48.9</u>	<u>12.6</u>	<u>2.7</u>	<u>22</u>	<u>1.85</u>	<u>0.5</u>	<u>0.1</u>	<u>22</u>	<u>1.52</u>	<u>0.19</u>	<u>365.4</u>	<u>13.5</u>	<u>472.3</u>	<u>59.0</u>
	Mean	Late Eocene	<u>48.9</u>	<u>16.2</u>	<u>1.5</u>	<u>112</u>	<u>2.07</u>	<u>0.7</u>	<u>0.1</u>	<u>112</u>	<u>1.60</u>	<u>0.11</u>	<u>365.6</u>	<u>7.6</u>	<u>496.1</u>	<u>35.7</u>
	<u>Note: x</u> —m	ean; σ—standard d	leviation;	s.e. —st	andard	error of	mean; n-	numbe	rs of pl	notos con	unts (400	×); t•s.e	— 95% con	fidence in	nterval. pCO	<u>2</u> — the

972 Table 7. Summary of stomatal parameters of the abaxial surface of fossil *Nageia* and pCO₂ $[C_{(f)}]$ estimates results.

result based the regression approach; $C_{(f)}$ the result based on the stomatal method.

a .		<u>SD (mm⁻²)</u>					<u>SI (</u> 9	<u>6)</u>		<u>S</u>	<u>R</u>	<u>pCO₂(</u>	<u>ppmv)</u>	<u><i>C_(f)</i> (p</u>	<u>pmv)</u>
<u>Species</u>	Age	<u>x</u>	<u>σ</u>	<u>s.e.</u>	<u>n</u>	<u>x</u>	<u></u>	<u>s.e.</u>	<u>n</u>	<u>x</u>	<u>t*s.e</u>	<u>x</u>	<u>t*s.e</u>	<u>x</u>	<u>t*s.e</u>
<u>MMJ1-001</u>	Late Eocene	<u>50.1</u>	<u>17.5</u>	<u>2.3</u>	<u>60</u>	<u>2.09</u>	<u>0.8</u>	<u>0.1</u>	<u>60</u>	<u>1.50</u>	<u>0.15</u>	349.7	<u>10.6</u>	<u>471.2</u>	<u>47.8</u>
<u>MMJ2-003</u>	Late Eocene	<u>46.5</u>	<u>16.3</u>	<u>2.1</u>	<u>60</u>	<u>1.96</u>	<u>0.7</u>	<u>0.1</u>	<u>60</u>	<u>1.67</u>	<u>0.24</u>	<u>358.3</u>	<u>9.8</u>	<u>524.1</u>	<u>75.7</u>
<u>MMJ2-004</u>	Late Eocene	<u>44.0</u>	<u>15.8</u>	<u>2.0</u>	<u>60</u>	<u>1.90</u>	<u>0.7</u>	<u>0.1</u>	<u>60</u>	<u>1.73</u>	<u>0.17</u>	<u>364.3</u>	<u>9.5</u>	<u>542.9</u>	<u>52.6</u>
<u>MMJ3-003a</u>	Late Eocene	<u>45.6</u>	<u>16.1</u>	<u>2.2</u>	<u>52</u>	<u>1.75</u>	<u>0.6</u>	<u>0.1</u>	<u>52</u>	<u>1.73</u>	<u>0.28</u>	<u>360.5</u>	<u>10.4</u>	<u>544.6</u>	<u>88.3</u>
Mean	Late Eocene	<u>46.6</u>	<u>16.4</u>	<u>1.1</u>	<u>232</u>	<u>1.93</u>	<u>0.7</u>	<u>0.1</u>	<u>232</u>	<u>1.66</u>	<u>0.11</u>	<u>358.1</u>	<u>5.0</u>	<u>519.9</u>	<u>35.0</u>
<u>Note: x</u> —m	ean; σ—standard o	deviation;	s.e. —st	andard	error of	mean; n-	- numbe	rs of pl	notos co	<u>unts (400</u>	×); t•s.e	<u>— 95% co</u>	nfidence	interval. pCC	<u>D₂— the</u>
result based	the regression app	roach; C _{(f}	<u>— the re</u>	sult bas	ed on the	e stomatal r	nethod.								

974 Table 8. Summary of stomatal parameters of the combined data of the adaxial and abaxial surfaces of fossil *Nageia* and pCO₂ [$C_{(f)}$] estimates 975 results.

977 <u>Table 9. pCO₂ estimates proxies and corresponding references.</u>

Proxies	References
Boron	<u>Pearson et al., 2009; Seki et al., 2010</u>
<u>B/Ca</u>	Tripati et al., 2009
Phytoplankton	Freeman and Hayes, 1992; Stott, 1992; Pagani et al., 1999, 2005; Henderiks and Pagani, 2008; Seki et al., 2010
Nahcolite	Lowenstein and Demicco, 2006
Liverworts	Fletcher et al., 2008
Paleosols	Cerling, 1992; Koch et al., 1992; Ekart et al., 1999; Royer et al., 2001; Nordt et al., 2002; Retallack, 2009b; Huang et al. 2013
<u>Stomata</u>	Van der Burgh et al., 1993; Kürschner et al., 1996, 2001, 2008; McElwain, 1998; Royer et al., 2001, 2003; Greenwood et al., 2003; Beerling
	et al., 2009; Retallack, 2009a; Smith et al., 2010; Doria et al., 2011; Roth-Nebelsick et al., 2012; 2014; Grein et al., 2013; Maxbauer et al.,
	<u>2014</u>