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Interactive comment on "Pliocene diatom and sponge spicule oxygen isotope ratios from the Bering Sea: isotopic offsets and future directions" by A. M. Snelling et al.

A. M. Snelling et al.

asnel@bgs.ac.uk

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There are insufficient quantities of sponge spicules in the samples to pick pure aliquots for sponge δ 180 by our method (fluorination).

With regards to the apparent match between the δ 18Osponge and the stacked benthic δ 18O record (Lisiecki and and Raymo, 2007), we do not have a conclusive explanation as to why the amplitude of the modelled δ 18Osponge is much greater. Whilst we speculate in the manuscript that this could be due to a mixture of non-equilibrium fractionation and/or inter-species variations in the fractionation of δ 18Owater, modern day

calibrations of δ 18Osponge are needed to address this further.

With regards to whether we expect the Bering Sea to record a "global" δ 18O signature, we would expect both surface and deep water δ 18O to be strongly dominated by local oceanographic conditions including changes in the Bering Sea gateway and regional ice-sheet dynamics. We are currently working on a longer diatom δ 18O record from which we hope to reconstruct some of these changes, which in turn may help to explain why the "relationship" between δ 18Osponge and stacked δ 18O appears to breakdown at c. 2.8 Ma. Whilst local effects may mean that it is not ideal to compare the δ 18Osponge record to the global stacked benthic δ 18O record, there are no local benthic δ 18O records available from the Bering Sea for this time period.

References Lisiecki, L. E. and Raymo, M. E.: Plio–Pleistocene climate evolution: trends and transitions in glacial cycle dynamics, Quaternary Sci. Rev., 26, 56–69, 2007.

Interactive comment on Clim. Past Discuss., 10, 2087, 2014.