



Supplement of

Trends in south pole particle concentrations imply holocene westerly wind strengthening

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S1. Additional Methods

We use a subset of the total Abakus size bins (21 of 30 size bins) ranged continuously in 0.1 μm increments from endpoints (lower) 1.0 – 2.5 μm , followed by bins 2.7, 2.9, 3.2, 3.6, 4.0, and 4.5 μm , with an upper limit of 5.1 μm (Abakus data from the South Pole Ice Core microparticle record; Kreutz, 2022). Bin lower limit represents the upper extent of the of the previous bin size. We use this subset to match other dust concentration methodology from East Antarctica (i.e., Delmonte et al., 2020). The CFA system has an effective resolution of 3 mm with a dispersion signal of 1 cm (Breton et al., 2012; Osterberg et al., 2006; Winski et al., 2019). The implications of the CFA resolution and dispersion are negligible to the timeframe that we resample our average data to 50-year intervals. Current accumulation rates at the South Pole are $\sim 8\text{cm/year}$ (Casey et al., 2014; Winski et al., 2019). While the CFA system can resolve particle concentrations to 3 mm intervals well below the accumulation rate, the coregistration of particles in the melter is a function of melt speed, ice density, and signal symmetry (Breton et al., 2012). Because we resample our data to 50 years intervals (mean depth interval of $\sim 3.5\text{ m}$), the dispersion of particles at 1 cm is negligible. The SPC14 chronology within the top 5 – 796 m ($< 11,300\text{ years BP}$) is annually resolved and was developed using seasonal variations in Na^+ and Mg^{2+} ions constrained by volcanic tie points (Winski et al., 2019). Age uncertainty during the Holocene is $< 18\text{ years}$ except between 3100 – 1800 years BP where the uncertainty increases to $\sim 25\text{ years}$ relative to the WD2014 timescale (Buizert et

al., 2015; Sigl et al., 2016; Winski et al., 2019). To mitigate contamination from fractured sections of core that contain trace amounts of Estisol-140 and other impurities (Warming et al., 2013), we used a Python-based cleaning code (Chesler et al., 2023) to detect and remove inferred Estisol-140-related peaks in the Abakus data. While overlapping Abakus dust measurements (i.e., coincidence) can influence dust measurement counts, prior work (Chesler et al., 2023) found no significant evidence of coincidence in the SPC14 dust record.

S2. CPP sources and ^{10}Be calibration

Course particle percentage (CPP) is hypothesized to reflect local Antarctic dust source activity, where increases in CPP reflect activation of local dust sources and a decrease reflects either a more active distal mid-latitudinal source or a decrease in local sources (Aarons et al., 2016; Albani et al., 2012; Koffman et al., 2014). We therefore compare the timing of CPP increase to coastal Antarctic glacier surface elevation lowering, to assess activation of local Antarctic dust sources. Because published surface exposure ages (^{10}Be) were calculated using different calibration rates, we recalibrate ^{10}Be surface exposure ages from the Pine Island Glacier (Johnson et al., 2014), Ross Sea (Jones et al., 2015; Jones et al., 2020), and Dronning Maud Land (Kawamata et al., 2020) using the University of Washington exposure age calculator, using production rates from Putnam et al. (2010) and Kaplan et al. (2011), and using the Lm scaling factor (Balco et al., 2008).

S3. Atmospheric Distillation

Changes in aerosol particle concentration may also be induced by atmospheric distillation, which results from variability in temperature and humidity gradients (Markle

et al., 2018). If this mechanism were primarily responsible for the decrease in Holocene dust, we would expect associated changes in temperature and/or stable water isotopes, as is the case on glacial/interglacial timescales (Markle et al., 2018). However, Antarctic paleoproxy records indicate only modest Holocene temperature changes and there is no significant relationship between dust concentration and SPC14 $\delta^{18}\text{O}$ (Kahle et al., 2020; Steig et al., 2021). Therefore, atmospheric distillation processes were likely not a major driver of altering the transport of dust to the South Pole during the Holocene.

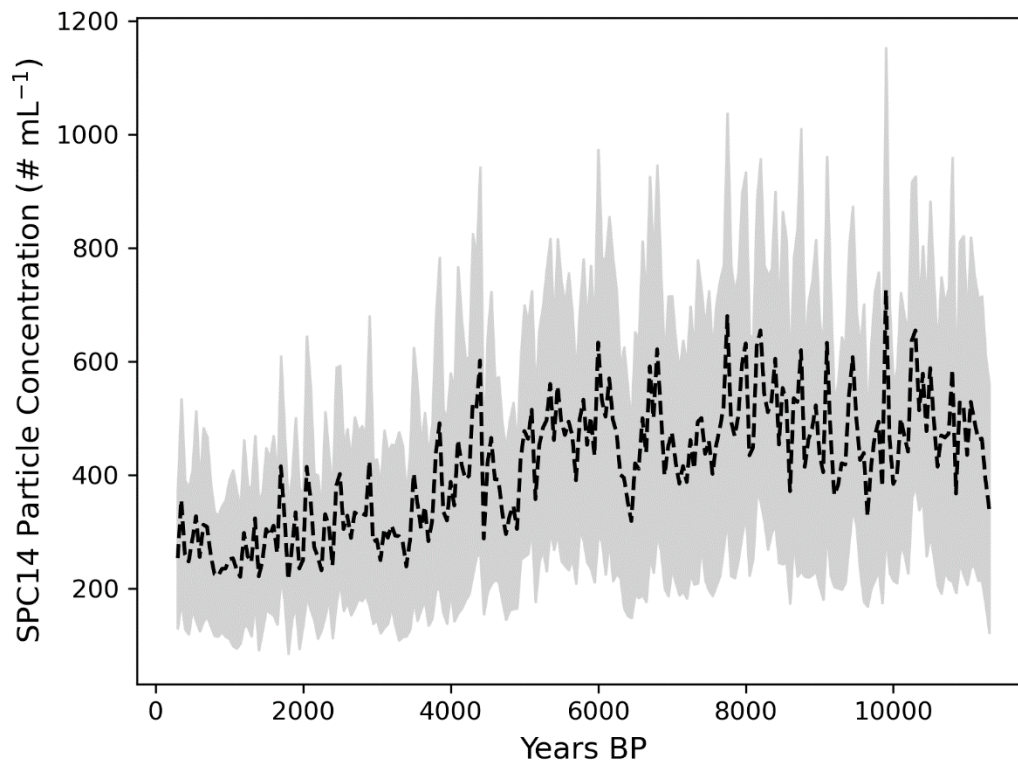


Figure S1: Median SPC14 Particle concentration in black dashed line with concentrations within median absolute deviation (MAD) shaded in the light grey.

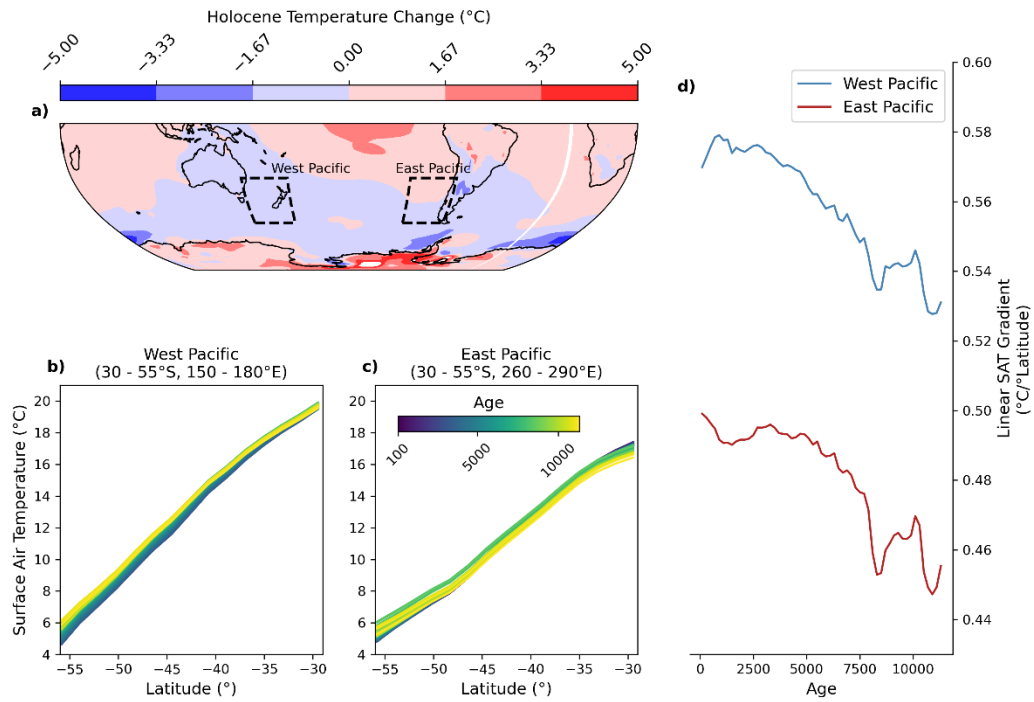


Figure S2 a – d: Surface air temperature data (11300 – 100 years ago) from Osman et al. (2021). S2a) highlights Southern Hemisphere average temperature change from the early Holocene (11300 – 7000 years ago) to pre-Industrial (300 – 100 years ago). S2b and c) are temperature gradients, colored by model age, from the eastern and western Pacific basin (highlighted in S2a). Slope values of linear regressions for each age (°C/°Latitude) are plotted in S2d.

88 **Table S1:** Median dust metrics and standard deviations for different time periods (2σ).

	Dust Concentration (# mL ⁻¹)	CPP (%)	Dust Flux (mg m ⁻² yr ⁻¹)
Average Holocene	417 ± 223	8.9 ± 1.3	0.43 ± 0.22
Early Holocene (11,300 – 7,000 years BP)	471 ± 166	8.6 ± 1.4	0.46 ± 0.17
Mid Holocene (7,000 – 3,000 years BP)	421 ± 190	8.9 ± 1.1	0.47 ± 0.22
Late Holocene (3000 – 250 years BP)	286 ± 108	9.4 ± 1.1	0.32 ± 0.11

Table S2a and b. Student t-test comparison of a) particle concentration and b) CPP distributions between the early, mid-, and late Holocene.

	Late Holocene	Middle Holocene	Early Holocene
Late Holocene	X	t-statistic = -9.1 p-value = <0.01	t-statistic = -15.6 p-value = <0.01
Middle Holocene	X	X	t-statistic = -4.9 p-value = <0.01
Early Holocene	X	X	X

	Late Holocene	Middle Holocene	Early Holocene
Late Holocene	X	t-statistic = 3.6 p-value = <0.01	t-statistic = 5.0 p-value = <0.01
Middle Holocene	X	X	t-statistic = 2.2 p-value = >0.01
Early Holocene	X	X	X

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